

1 **Pluridisciplinary analysis and multi-archive reconstruction of**  
2 **paleofloods: societal demand, challenges and progress**

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12  
13 **Abstract.**

14 Floods are one of the gravest natural hazards for societies, worsened by  
15 population growth, unchecked development, and climate change. From a Global  
16 Change perspective, past extreme events merit particular interest because they  
17 can be linked to wider climate and environmental changes, introduce  
18 perturbations. During the last decade, knowledge of long-term flood frequency  
19 and magnitude has been improved by extracting data from different types of  
20 archive. But, despite advances in dating methods, proxies and statistical  
21 techniques and efforts to identify atmospheric drivers, some fundamental  
22 questions remain unresolved. The Special Issue entitled “Pluridisciplinary  
23 analysis and multi-archive reconstruction of paleofloods” in the journal Global and  
24 Planetary Change addresses these uncertainties and complexities by assembling  
25 a selection of studies, which were first presented at the Past Climate Changes  
26 (PAGES) Open Scientific Meeting held at Zaragoza in 2017. In this introductory

27 paper, the guest editors outline the 17 research contributions and meta-data from  
28 the 17 paleoflood studies were systematically analyzed in terms of i) geographical  
29 distribution; ii) methodologies applied; iii) types of archives; iii) numbers of flood  
30 series compiled and iv) spatial and temporal resolution of paleoflood data. The  
31 data indicate that paleoflood studies focused on fluvial depositional environments  
32 show a higher rate of integration with other types of paleoflood archive (mean of  
33 4.5 types of archive) than studies focused on documentary sources (mean of 3.5)  
34 and lake sediments (mean of 2.4). We suggest that this strategy of archive  
35 integration has been adapted to effectively compensate for the higher  
36 uncertainties of fluvial deposition in floodplains. Statistical processing of the  
37 meta-data shows quantitative associations between specific types of flood  
38 archive and offers a solid platform for designing the optimal approach for multi-  
39 archive paleoflood research. A qualitative review and visual comparison of the 17  
40 paleoflood series shows some consistent trends and breaks but also notable  
41 differences within and between regions. While a trend of increased flooding since  
42 4-5 ka BP is evident, the lack of synchronicity between breaks and the coeval  
43 increases and decreases in fluvial activity is manifest. The majority of studies in  
44 the Special Issue do denote the 19<sup>th</sup> century - including the youngest cool climate  
45 pulses during the Little Ice Age - as a particularly flood-rich period. It is more  
46 difficult to assess the 20<sup>th</sup> century because of social changes, population growth  
47 and extensive river modification. Despite the mentioned uncertainties, 10 of 14  
48 papers do not record the 20<sup>th</sup> century as an exceptional flood period. Assessing  
49 the effects of human impact on paleoflood calendars and disentangling  
50 anthropogenic from natural drivers are major challenges in integrated paleoflood  
51 analysis.

52 It is concluded that the interpretation of flood series is complex as landscapes  
53 and flood drivers are heterogeneous and systems show different sensitivities to  
54 flood control and drivers. Thus, the study of past floods, from historical and  
55 natural archives, is challenging but also offers unparalleled opportunities to  
56 document low-frequency, large-magnitude flood events, which occurred under a  
57 broad range of climate and/or environmental scenarios, and, probably, the only  
58 way to reconstruct robust paleoflood series.

59 **Keywords:** Paleoflood hydrology; flood hazard; natural archives; documentary  
60 sources; multi-archive reconstruction; integration model.

61

## 62 **1. The motivation of this Special Issue**

63 The integration of multi-archive flood proxies to reconstruct flooding hundreds or  
64 thousands of years ago is like putting together a puzzle. In the beginning, the  
65 numerous pieces appear chaotic and confusing. But after struggling for some  
66 years or even decades, structure and eventually a diffuse picture becomes  
67 recognizable. This differs from an ordinary puzzle, however, because the number  
68 of pieces is not finite and the puzzle will never be finished. Although this might  
69 produce at a first glance a certain frustration, there are many other aspects which  
70 fully satisfy the expectations of scientists.

71 The first aspect is that the development of multi-archive flood records is a  
72 relatively modern approach. Over the last two decades, a growing number of  
73 studies have reconstructed flood records spanning centuries and millennia from  
74 fluvial sediments, lake deposits, speleothems or tree-rings (Baker, 1987; Benito  
75 et al., 2004; Schulte et al., 2008, 2009b; Wilhelm et al., 2012; 2019; Díez-Herrero

76 et al., 2013; Wirth et al., 2013, Schillereff et al., 2014; Ballesteros-Cánovas et al.,  
77 2014; Santisteban et al., 2017; Denniston and Lüscher, 2017). These data series  
78 were largely confined to comparisons of paleoflood series with historical sources  
79 and instrumental measurement of discharge and precipitation. More recently,  
80 efforts were made to depict flood patterns across larger regions based on  
81 instrumental data covering half a century (Blöschel et al., 2017). At the same  
82 time, historians and geographers produced regional compilations also from  
83 historical sources (e.g. Röthlisberger, 1991; Glaser, 2001; Brázdil et al., 2005a,  
84 2005b; Wetter, 2011; Macdonald and Sangster, 2017; Paprotny et al., 2018).  
85 Historical records and long instrumental records are mostly restricted to larger  
86 river towns (Pfister, 1999; Barriendos et al., 2014; Elleder et al., 2015; Wetter,  
87 2017) whereas catchments in more remote regions, particularly mountain basins,  
88 are often ungauged and historical sources may be scarce (Schulte et al., 2009a,  
89 2015). Since flood archives are embedded in different geographical and  
90 environmental settings, and their “perfect” study sites do not coincide  
91 geographically, paleoflood information is often fragmented.

92 Multi-proxy approaches have become standard in paleoenvironmental and  
93 paleoflood research (Santisteban et al., 2017; Wilhelm et al., 2019), whereas  
94 multi-archive studies *in sensu strictu* which integrate more than three different  
95 types of flood archives are extremely rare (Schulte et al., 2015). This presents  
96 opportunities for creative researchers to open doors to a fascinating world where  
97 they can explore, combine, disentangle and test several combinations of flood  
98 archives.

99 A second motivation is the attraction of multi-disciplinary research. It is exciting  
100 to meet researchers from other fields at conferences, workshops, field

101 excursions, or on interdisciplinary field work and listen to them describe their  
102 approaches to researching floods. It is remarkable that they all look at the same  
103 physical process but use other archives, proxies, markers, thresholds and so on.  
104 For example, who could imagine that bioindicators such as algae and lichens in  
105 cm-small alveoli in canyon rock walls could provide information about floods?  
106 Therefore, it is vital for the paleoflood community to test reconstructed past floods  
107 through different techniques, methods, and scientific views.

108 A third aspect is the spatial dimension of the flood phenomenon. Different  
109 archives allow flood information to be obtained that better reflects the diversity of  
110 landscapes that experience flooding compared to studies that focus on only one  
111 type of archive. For example, in mountain regions, flood information can be  
112 obtained from high-altitude lakes, tree-rings and lichen colonization of river  
113 banks, gorge rock surfaces, and alluvial fan deposits at mid-altitudes, and from  
114 low-altitude floodplains and deltas (alluvial sediments, historical and  
115 archaeological evidence, pollen, etc.). In some basins, flood data can be obtained  
116 from multiple sites that differ in elevation by 1500 m or more within only a few  
117 kilometers (Schulte et al., this issue; Zaginaev et al., this issue). Another  
118 promising approach is the reconstruction of single flood events in terms of total  
119 flooded area, the propagation of the flood wave and the path of the precipitation  
120 field (Kiss, 2009; Elleder et al., this issue). Furthermore, where a high density of  
121 paleohydrological data is available, the production of paleoflood maps can  
122 improve our spatial understanding of flood dynamics (Röthlisberger, 1991;  
123 Schmocker-Fackel and Naef, 2010; Barriendos et al., this issue; Schulte et al.,  
124 this issue). In this context, a further methodological innovation is the  
125 reconstruction, reanalysis and modelling of synoptic sea level pressure maps of

126 extreme flood episodes, which improve our knowledge about atmospheric  
127 variability as a flood driver (Ortega et al., this issue; Peña and Schulte, this issue;  
128 Sánchez-García et al., this issue, Schulte et al., this issue).

129

## 130 **2. Foci of the PAGES Floods Working Group**

131 The exposed range of opportunities is one of the reasons why multidisciplinary  
132 analysis and multi-archive reconstruction of paleofloods define one of the three  
133 core activities of the Past Climate Changes (PAGES) Floods Working Group  
134 since its founding in 2015. According to the White Paper (PAGES - Floods  
135 Working Group, 2017), the Working Group “aims to bring together all the scientific  
136 communities reconstructing past floods (historians, geologists, geographers, etc.)  
137 and those studying current and future floods (hydrologists, modelers,  
138 statisticians, etc.) to coordinate, synthesize and promote data and results on the  
139 natural variability of floods”. Also in 2017, the Floods Working Group launched  
140 three work packages (WP): WP1 focuses on collecting, storing and sharing of  
141 global paleoflood data, WP2 on integrating and analyzing paleoflood data and  
142 WP3 on communicating and disseminating paleoflood science and data at  
143 different levels, including stakeholders. Conference sessions of the WP2 topic  
144 were organized in 2016 during the first Floods Working Group Workshop in  
145 Grenoble and in 2017 on the PAGES Open Scientific Meeting held at Zaragoza.  
146 Also in 2016 several members of WP2 launched a pilot project of paleoflood data  
147 integration in the Swiss Alps. The results of the first phase of this research of  
148 geographers, historians, geologists, and geochemists are presented in the paper  
149 of Schulte et al. entitled “Integration of multi-archive datasets towards the  
150 development of a four-dimensional paleoflood model in alpine catchments” in this

151 Special Issue. These research activities will continue during the second phase  
152 (2019-2021) of the FWG Program.

153

### 154 **3. Societal demand for multi-archive reconstruction of paleofloods**

155 Floods are one of the gravest natural hazards for societies, worsened by  
156 population growth, unchecked development and climate change (UNISDR,  
157 2015). So, the transfer of long flood series to public agencies is crucial for  
158 producing reliable evaluations of floods and societal risk. However, although  
159 policies have been developed (e.g. Directive 2007/60/EC of the European  
160 Parliament and of the Council of 23 October 2007 on the assessment and  
161 management of flood risks; Real Decreto 903/2010, of July 9<sup>th</sup>, Assessment and  
162 management of floods), the integration of paleoflood studies into spatial planning  
163 and flood risk assessment is not sufficiently applied. From a Global Change  
164 perspective (Baker, 2006), past extreme events are of interest as they can be  
165 linked to climate and environmental changes, introduce perturbations in natural  
166 systems and can be traced in paleoenvironmental archives. Hence, paleoflood  
167 research is a rapidly developing approach through which insight from multiple  
168 disciplines (hydrology, geomorphology, climatology, paleolimnology) has  
169 implications for human life as its goal is to understand and quantify flood risk over  
170 extended periods of time.

171 During the last decade, knowledge of flood frequency and magnitude has  
172 improved through data coming from different types of archives (Baker, 1987;  
173 Benito et al., 2004; Schulte et al., 2015; Schillereff et al., 2016; Wilhelm et al.,  
174 2019). But, despite advances in dating methods, proxies and statistical  
175 techniques and efforts to identify atmospheric drivers, some fundamental

176 questions remain unresolved. The interpretation of flood series is complex as  
177 landscapes and flood drivers are heterogeneous and systems show different  
178 sensitivities to hydrometeorological forcings. Thus, the study of past floods using  
179 historical and natural archives is challenging but also a rare opportunity to  
180 document low-frequency, large-magnitude flood events. Long-term studies also  
181 allow trends in flooding that occurred under a broad range of climate and/or  
182 environmental scenarios to be explored, which is probably the only way to  
183 reconstruct robust paleoflood series. This issue addresses these uncertainties  
184 and complexities by assembling a selection of studies with a global geographical  
185 distribution (high to low latitude, from mountains to lowlands) and provide an  
186 insight about present state on multi-source data (lakes, floodplains,  
187 geomorphology, tree-rings, historical and archaeological sources, soils, marine  
188 sediments, etc.), controls/drivers and time-scales integration (from Pleistocene to  
189 present time) plus methodological and societal issues in paleoflood research.

190

191 **4. Contents of contributions to multi-archive paleoflood reconstruction**

192 The compilation of this Special Issue, which includes 17 research papers, is the  
193 outcome of the PAGES OSM Conference session entitled “Multidisciplinary  
194 reconstruction of paleofloods”. Sixteen oral contributions and 18 posters from  
195 most continents were presented and lively discussed. The papers showcase  
196 substantial progress in the analysis and interpretation of flood archives, important  
197 methodological advancements, including innovative approaches to integrate and  
198 model diverse archives and flood series, and a focus on remote regions with  
199 difficult access.



200 The research papers of Santisteban et al. (this issue) and Fuller et al. (this issue)  
201 are case studies from Central Spain and New Zealand that demonstrate how  
202 high-resolution, continuous geochemical flood proxies can be inferred from  
203 alluvial sediments that span most of the Holocene. Santisteban et al. (this issue)  
204 use several geochemical ratios as proxies for water competence, water level, and  
205 sediment discharge to reconstruct flood pulses. Similarly, Fuller et al. (this issue)  
206 estimated the flood recurrence interval using normalized Zr/Rb measurements  
207 and a tight age-depth model in a volcanically-reset catchment.

208 The studies of Agatova et al. (this issue) and Lombardo et al. (this issue) focus  
209 on large-scale flood areas in Asia and South America which are difficult to access.  
210 In south-western Amazonia, Lombardo et al. (this issue) combined proxies such  
211 as phytoliths and stable carbon isotopes from sedimentary flood archives and  
212 soils to provide a solid reconstruction of past Holocene land cover change and  
213 periods of low or modest flooding. Agatova et al. (this issue) used  
214 geomorphological, geological and geoarchaeological data to reconstruct the  
215 presence of Late Pleistocene ice-dammed lakes and cataclysmic outburst floods  
216 in the Mongolian Inland Drainage Basin. A multi-century dataset of regional  
217 glacial outburst floods (GLOF) is presented by Zaginaev et al. for the Tien Shan  
218 (Central Asia). These high discharge flash-floods were reconstructed by tree-ring  
219 analyses from six different torrential fans providing insights on regional process  
220 activity.

221 A different approach is adopted by the following four papers: extracting evidence  
222 from documentary archives to produce regional centennial flood series.  
223 Barriendos et al. (this issue) provide 18 extensive flood event chronologies for  
224 the Spanish Mediterranean coast from 14<sup>th</sup> to 20<sup>th</sup> centuries. They discuss the

225 profound influence of social factors on historical flood data series and evaluate  
226 methods of integrating multi-source information such as population and flood  
227 protection measures. This human component also affects the 450-year  
228 reconstruction of historical discharges performed by Sánchez-García et al. (this  
229 issue) from semi-arid South-eastern Spain. Furthermore, the synoptic  
230 atmospheric configurations of four catastrophic flood events were investigated.  
231 In the River Jing catchment, southern Chinese Loess Plateau, Yu et al. (this  
232 issue) identified decadal solar activities as an important driver for floods and  
233 droughts. Multiple documentary sources and a precipitation-runoff model were  
234 used by Elleder et al. (this issue) to explore the spatial imprint of the 1872 flash-  
235 flood in central Bohemia and model the river's runoff response.

236 Another five papers deal with paleoflood reconstruction and flood frequency  
237 analysis using lake records. Evin et al. (this issue) propose a novel statistical  
238 approach that combines a classic series of paleoflood observations for the Rhône  
239 River reconstructed from lake sediments (Lake Bourget, Northwestern Alps,  
240 France) and disseminates uncertainties related to the reconstruction method  
241 during the estimation of extreme quantiles. Albrecher et al. (this issue) applied a  
242 change-point analysis to sedimentary flood frequency data from six large alpine  
243 lakes. This enabled a comparison to be made with other flood records and  
244 possible links to be drawn between event frequencies and climatic conditions.  
245 Corella et al. (this issue) present a new method for estimating seasonally-  
246 resolved flood erosion rates using millennium-long varved lake sediments. Their  
247 use of high-precision, multi-proxy data also sheds light on the main environmental  
248 drivers (climatic or anthropogenic) controlling sediment yield in a mountainous  
249 Mediterranean watershed. The respective roles of human and climate forcings on

250 Holocene flood frequency were also investigated by Rapuc et al. (this issue) in  
251 Lake Iseo (Southern Alps). Similarly, Schillereff et al. (this issue) showed that  
252 detailed sub-sampling and proxy analysis based on particle size data, coupled  
253 with careful evaluation against independent hydrological data and accounting for  
254 variations in external sediment supply potentially driven by anthropogenic  
255 landscape modification, is an appropriate methodology to extract paleoflood  
256 records from temperate lakes.

257 To explore climatic forcing of floods Peña and Schulte (this issue) performed a  
258 paleoclimate modeling experiment of the atmospheric variability related to large  
259 summer floods in the Hasli-Aare (Swiss Alps) from the AD 1300 to 2010. They  
260 propose the name of paleo-SNAO to define this decadal atmospheric variability  
261 related to summer floods in the alpine catchment.

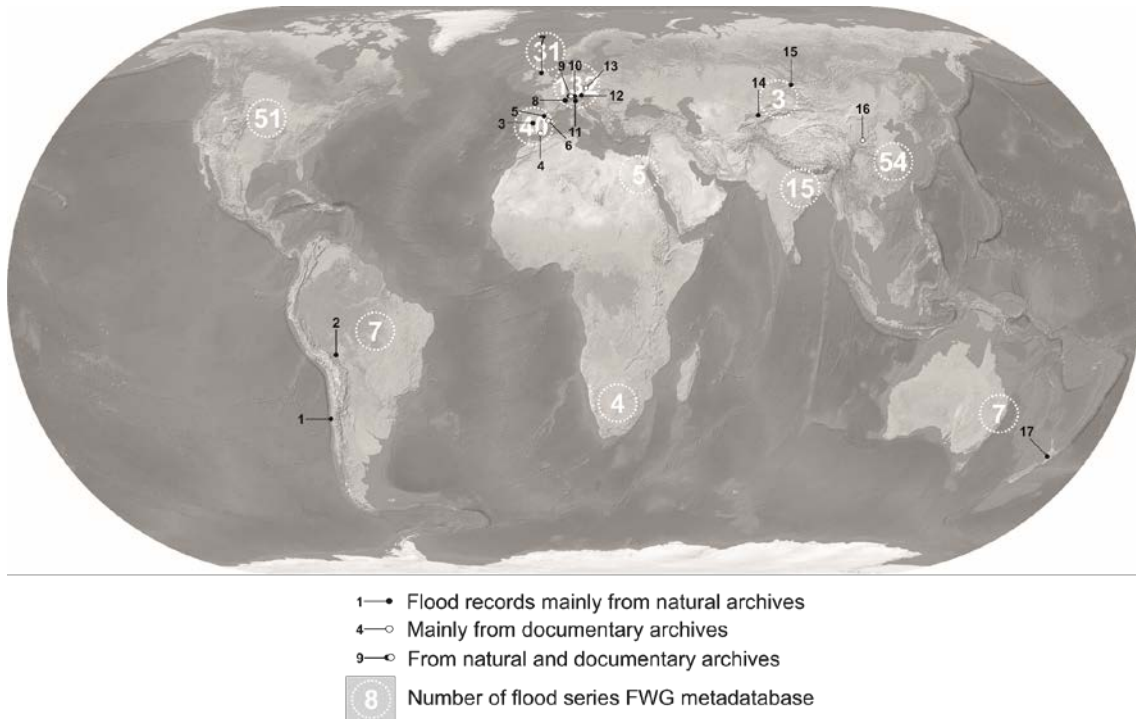
262 Schulte et al. (this issue) designed an innovative methodology that integrates  
263 multi-archive datasets towards the development of a spatial-temporal (four-  
264 dimensional) paleoflood model in alpine catchments. The most continuous and  
265 accurate series from natural and anthropogenic flood archives were integrated  
266 over the period from AD 1400 to 2005 into a synthetic flood master curve for the  
267 Bernese Alps.

268 Ortega et al. (this issue) analyzed extreme ENSO-driven torrential rainfalls at the  
269 southern edge of the Atacama Desert during the Late Holocene and their  
270 projection into the 21<sup>st</sup> century. The integration of marine paleoclimate proxies,  
271 historical data, and the future projection helps to understand oceanic and climatic  
272 factors conditioning the variability of extreme rainfall events.

273

## 274 **5. Geographical location and meta-data of paleoflood records**

275 The 17 papers of the Special Issue present a diverse body of work in terms of i)  
 276 geographical distribution (four continents, northern and southern hemisphere;  
 277 Figure 1); ii) types of paleoflood archives; iii) numbers of flood series compiled;  
 278 iv) spatial and temporal resolution and v) methodologies applied for the  
 279 integration of flood data (Table 1).



280

281 Figure 1. Location of the case studies (black numbers) presented in the research papers of this  
 282 Special Issue. The ID of each paper is listed in Table 1. White numbers encircled presents the  
 283 number of studies recorded in the paleoflood metadata base of the PAGES Floods Working Group  
 284 for each region (PAGES-FWG, <http://www.pages-igbp.org/ini/wg/floods/wp1/data>, date of access  
 285 January 10<sup>th</sup>, 2019).

286

287 Figure 1 reflects two types of data: the papers presented in this Special Issue  
 288 (black numbers) and the regional distribution of studies recorded in the paleoflood  
 289 metadata base of the PAGES Floods Working Group (FWG; <http://www.pages-igbp.org/ini/wg/floods/wp1/data>, white circled numbers). More than half of the  
 290

291 studies archived in the FWG databank are located in Europe. More than 50  
292 studies were carried out in North America as well as in China, 15 in India, whereas  
293 low numbers are recorded in Australia/New Zealand, Central Asia, Africa, and  
294 South America. 36% of these records were obtained from historical documents,  
295 33% from riverine sediments, 29% from studies of lake sediments and tree rings  
296 and, finally 2% from speleothems (PAGES Floods Working Group, 2019). These  
297 numbers and consequent distribution do not include all worldwide published  
298 works but may reflect some general trends. The high numbers obtained in Europe  
299 reflect the intense research activities in the field of paleoflood reconstruction  
300 across the continent, but, on the other hand, these numbers are also influenced  
301 by the location of organized workshops and annual meetings, the hosting of the  
302 PAGES office and members of the FWG steering committee and the lower level  
303 of cooperation with researchers from other continents. For example, in recent  
304 years several European researchers (many active in the FWG network) attended  
305 paleoflood conferences organized by US researchers (e.g. Rapid City 2016) and  
306 vice versa, but no joint conference has yet been organized by both communities.  
307 The paleoflood community is in a similar situation with regard to links with Asia.  
308 In addition to the innovative topics and methodological progress of the presented  
309 research papers, the metadata of these case studies presented in Table 1  
310 provides interesting insight into the structure of multi-archive paleoflood  
311 approaches. According to Table 1 and Figure 3, the papers presented in this  
312 special issue can be subdivided into different groups. The first group includes  
313 papers which present flood series from alluvial and fluvial depositional  
314 environments and landscapes; the second group reconstructs flood calendars  
315 from historical sources, and the third investigates past floods from lake deposits.

316 Finally, there are contributions that focus on flood records from tree-rings and  
317 marine sediments or integrate numerous types of flood archives.

318

| Basic meta data                                  | Research papers of Special Issue |  | Multi                            | Botanical        | Marine                                     | Alluvial sediments and soils |                    | Documentary sources |                    |  |   | Lake sediments      |                    |                   |                   |                    | Model             |              |                    |
|--|----------------------------------|--|----------------------------------|------------------|--|------------------------------|--------------------|---------------------|--------------------|--|---|---------------------|--------------------|-------------------|-------------------|--------------------|-------------------|--------------|--------------------|
|  | Schulte et al.                   | Zaginaev et al.                            |                                  |                  |  | Ortega et al.                | Santisteban et al. | Agatova et al.      | Lombardo et al.    | Fuller et al.                            | Elleder et al.                              | Sánchez-G. et al.   | Barreros et al.    | Yu et al.         | Evin et al.       | Corella et al.     |                   | Rapuc et al. | Schillereff et al. |
| Reference in Figure 1                            | 9                                | 14   | 1                                | 1                | 1  | 3                            | 15                 | 2                   | 17                 | 13                                       | 4   | 6                   | 16                 | 8                 | 5                 | 11                 | 7                 | 12           | 10                 |
| Location of study area                           | 46°41'N<br>6°04'E                | 74°33'E-<br>74°48'E<br>42°25'N-<br>42°42'N | 32°S-26°S<br>71°43'W-<br>69°28'W | 39°5'N<br>3°45'W | 50°15'N-<br>49°45'N<br>89°45'E-<br>90°10'E | 14°30'S<br>65°00'W           | 39°43'S<br>17°59'E | 49°57'N<br>14°04'E  | 37°12'N<br>19°45'W | 35°50'N-<br>43°34'N<br>5°58'W-<br>4°30'E | 34°48'N-<br>36°12'N<br>107°14E-<br>108°40'E | 45° 43'N<br>5° 52'E | 42°19' N<br>0959 E | 45°44'N<br>10°4'E | 54°30'N<br>2°55'W | 47°48'N<br>13°23'E | 46°41'N<br>6°04'E |              |                    |
| Total catchment area (km2)                       | 2117                             | 145  | no data                          | 26232            | no data                                    | 78000                        | 7380               | 8286                | 2611               | 156930                                   | 16057                                       | 4000                | 1.39               | 45                | 13.01             | 241                | 596               |              |                    |
| Sediments (Stratigraphy, geochemistry)           | 4                                | 6  |                                  | 3 (12)           | 2 (15 sect.)                               | 4                            | 1 (3)              |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              | 1                  |
| Landforms: flood plains, terraces, alluvial fans | 3                                |  |                                  | 2 (basins)       | 4  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Soil sequences                                   | 4                                | 6  |                                  | 2 (15 sect.)     | 4 (37)                                     |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Mapping, aerial photographs, GIS                 | 3                                |  |                                  | 4                |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| 10Be erosion rate data                           |                                  |  |                                  |                  |  |                              | 1                  |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Lichenometry                                     | 4                                |  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Dendrochronology/Dendromorphology                | [1]                              | 6  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Palynology                                       | [1]                              |  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Documentary sources (flood series)               | 6                                | 6  | 1                                | 2                | 1  | 1                            | 1                  | 2                   | 1 (4)              | 18                                       | 1 (10)                                      |                     |                    |                   | 1                 |                    | [1]               | 1            |                    |
| Documentary sources (drought series)             |                                  |  |                                  | 1                |  |                              |                    | 10                  |                    |  | 1 (10)                                      |                     |                    |                   |                   |                    |                   |              |                    |
| Flood series from flood marks                    | [1]                              |  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   | [1]                |                   |              |                    |
| Archeological sites                              | 1                                |  |                                  | 1                | 2  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Precipitation records (meteorological stations)  | [4]                              |  | 1                                | [11]             |  |                              |                    | 12                  | 2                  | 18                                       |   |                     |                    |                   | 1                 |                    |                   |              |                    |
| Instrumental records (gauging stations)          | 7                                |  |                                  | 4                |  |                              | 1                  | 9                   | 3                  | 18                                       |   |                     |                    |                   | 1                 |                    |                   |              |                    |
| Lake sediments                                   | 4                                |  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   | 1 (32)            | 1                  | 1                 | 1            | 1 [5]              |
| Marine sediments                                 |                                  |  | 1                                |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Geophysical sections                             |                                  |  |                                  | 1                |  |                              |                    |                     |                    |  |   |                     |                    |                   | 1                 | 1                  |                   |              |                    |
| Variability of glaciers                          | [2]                              | 6  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Climate reanalysis / modeling (CESM at NCAR)     | 1                                |  | 1                                |                  |  |                              |                    |                     | 1                  |  |   |                     |                    |                   |                   |                    |                   |              | 1                  |
| Hydraulic / Hydrological modeling                | [1]                              |  |                                  |                  |  |                              |                    | 1                   |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Total of systematically analysed archive         | 10                               | 5  | 4                                | 6                | 4  | 4                            | 4                  | 5                   | 4                  | 3  | 2   | 2                   | 4                  | 2                 | 3                 | 3                  | 1                 | 3            |                    |
| Max. range of flood series (cal yr AD)           | 1400                             | 1874                                       | 1301                             | 9500             | >11600                                     | 10000                        | 1770               | 1872                | 1550               | 1301                                     | 1646  | 1650                | 2775               | 12000             | 450               | 7096               | 1300              |              |                    |
| Max. range of flood series (cal yr BP)           | 3600                             |  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| Resolution of flood series                       | 1                                | 1  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| exact day  |                                  |  |                                  |                  |  |                              |                    |                     |                    |  |   |                     |                    |                   |                   |                    |                   |              |                    |
| seasonal   | 1                                | 1  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| annual   | 1                                | 1  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| intra-decadal                                    | 1                                | 1  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| decadal  | 1                                | 1  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| multidecadal                                     | 1                                | 1  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| centennial                                       |                                  |  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| multicentennial                                  |                                  |  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |
| millennial                                       |                                  |  |                                  | 1                |  |                              |                    | 1                   | 1                  | 1  | 1   | 1                   | 1                  | 1                 | 1                 | 1                  | 1                 | 1            |                    |

320 Table 1: Metadata of Paleoflood case studies published in the Special Issue. A) Numbers of  
321 different types of archive analyzed to compile regional flood series. Legend: 2 = number of  
322 analyzed flood series (one series per type of archive and catchment, region or area). If data is  
323 available: (2) = number of records; [1] = punctual data record or data not explicitly discussed in  
324 paper

325

326 The number of systematically analyzed types of flood archives listed in Table 1  
327 varies significantly between papers. Seven studies - primarily investigating  
328 historical sources and lake sediments - draw on one to three types of archives;  
329 nine papers incorporate four to six types of archive and one paper utilizes up to  
330 ten. The meta-data indicates that paleoflood studies focused on fluvial  
331 depositional environments show a higher rate of integration of different types of  
332 paleoflood archives (mean of 4.5 types of archives) than studies focused on  
333 documentary sources (mean of 3.5) and lake sediments (mean of 2.4).

334 Papers in this Special Issue present paleoflood reconstructions over variable time  
335 periods, ranging from one and a half centuries (tree rings; Zaginaev et al., this  
336 issue) to the Early Holocene (lake and alluvial flood records; Rapuc et al., this  
337 issue; Santisteban et al., this issue) while the reconstruction of catastrophic floods  
338 from the Mongolian Great Lakes Basin reaches back to the Late Pleistocene  
339 (Agatova et al., this issue). The highest temporal resolution (exact dates,  
340 seasonal and annual flood information) were obtained by studies using  
341 documentary sources or papers that combine natural flood archives (e.g. tree  
342 rings and varved lake sediments) with documentary and instrumental data. It is  
343 striking that most approaches exploiting natural archives (except soils and fluvial  
344 landforms) achieve a temporal accuracy of decadal or better.



345 With regard to the spatial scale, large differences in the size of study areas and  
346 catchments are noticeable in Table 1. Catchments smaller than 100 km<sup>2</sup> are  
347 associated with lake reconstructions, whereas larger areas of more than 50,000  
348 km<sup>2</sup> were studied by papers focused on landscape, landform and soil  
349 development. Barriendos et al. (this issue) submitted the paper with a total area  
350 of 156,930 km<sup>2</sup>, which presents documentary flood records from 18 catchments  
351 spanning the Mediterranean slope of the Iberian Peninsula.

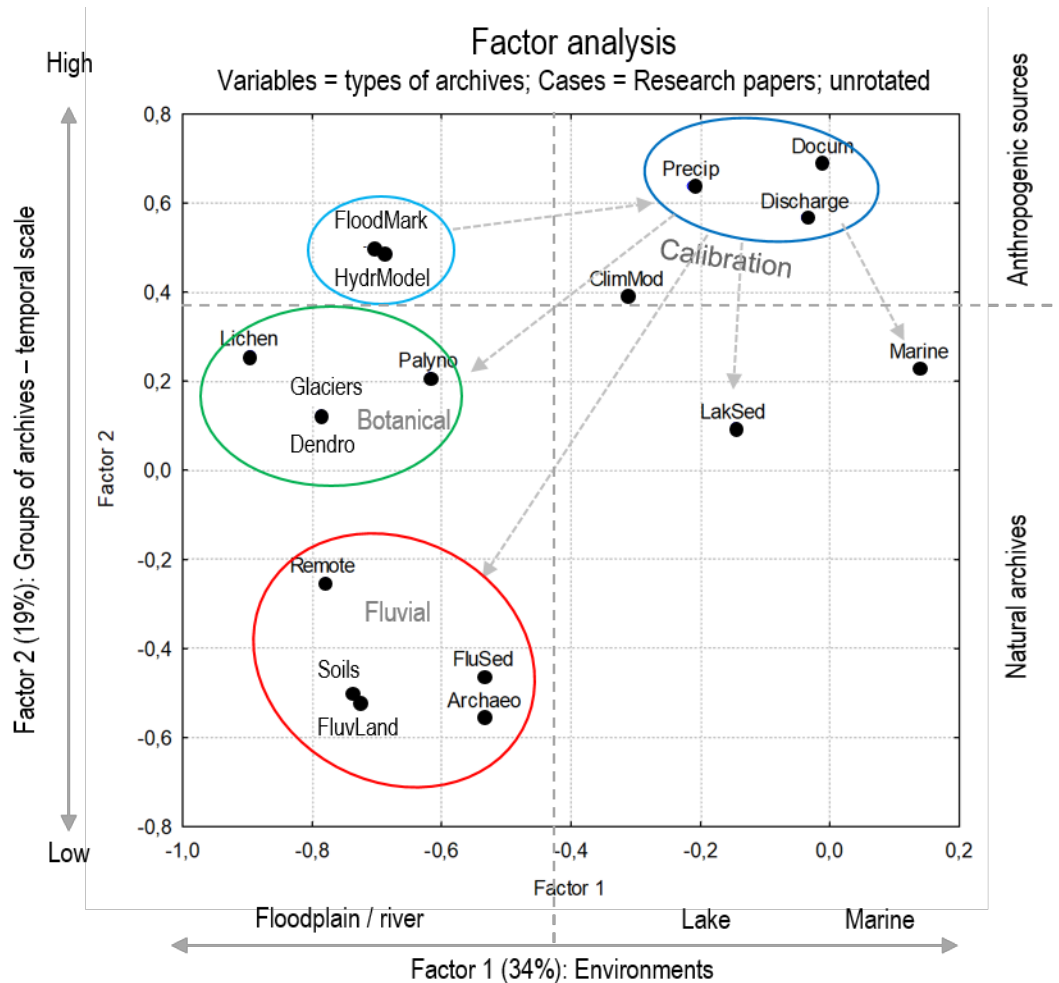
352

## 353 **6. How are paleoflood archives combined and integrated?**

354 A possibility to develop a conceptual framework for multiple-archive paleoflood  
355 integration is the performance of a qualitative approach (PAGES - Floods  
356 Working Group, 2017). However, statistical processing of the meta-data from  
357 paleoflood studies published in this Special Issue could provide valuable insight.  
358 This testing shows associations between different types of flood archive,  
359 achieving a solid background for the design of an integrated multi-archive  
360 paleoflood approach.

361 The variables (number of flood archive types analyzed for generating paleoflood  
362 series) presented in Table 1 were transformed into a binary system and  
363 introduced into a matrix. Factor analysis (FA) was performed to explore the  
364 variability of flood archives across the 17 studies. Figure 2 shows the 2-  
365 dimensional plot of the first two factors explaining 34% (F1) and 19% (F2) of the  
366 variability. The following groups are identified: (i) fluvial and terrestrial archives  
367 (red circle); (ii) botanical archives (green circle); hydrological archives (light blue  
368 circle) and (iv) precipitation, discharge measurements and documentary sources

369 (blue circle). The variables climate modeling, lake sediments, and marine  
 370 sediments show a more scattered distribution.



371

372 Figure 2: 2-D plot of the first two factors explaining the variability of types of archives analyzed by  
 373 the research papers in this Special Issue.

374

375 The factor F1 is interpreted as the range of environments of natural flood  
 376 archives: floodplain and river environments show negative loadings, whereas  
 377 subaquatic (lakes and marine) archives show positive loadings. The second  
 378 factor reflects the temporal resolution of archives from low (millennial-scale  
 379 resolution of fluvial landforms and soils) to high (exact hour and/or day of river  
 380 discharge, documentary sources, flood marks and precipitation records). This

381 distribution is similar to the 2-D plot (not shown) where temporal resolution (Table  
382 1) is included in the FA matrix as an additional binary variable. Another interesting  
383 outcome is the clear division (0.3 F2 loading) between natural archives and  
384 anthropogenic sources (Figure 2).

385 Our explanation for this variability is that high-resolution lake records are mostly  
386 calibrated against instrumental records of discharge and precipitation as well as  
387 documentary sources at annual and, in the best cases, seasonal resolution  
388 (Corella et al., this issue; Evin et al., this issue). Accurate calibration can also be  
389 applied in the studies of tree-rings and lichens. However, the botanical archives  
390 show a closer relationship with the group of fluvial archives, landforms, and soils  
391 because they are also used when dating flood deposits, flood levels, and impacts  
392 (Schulte et al., this issue; Zaginaev et al., this issue).

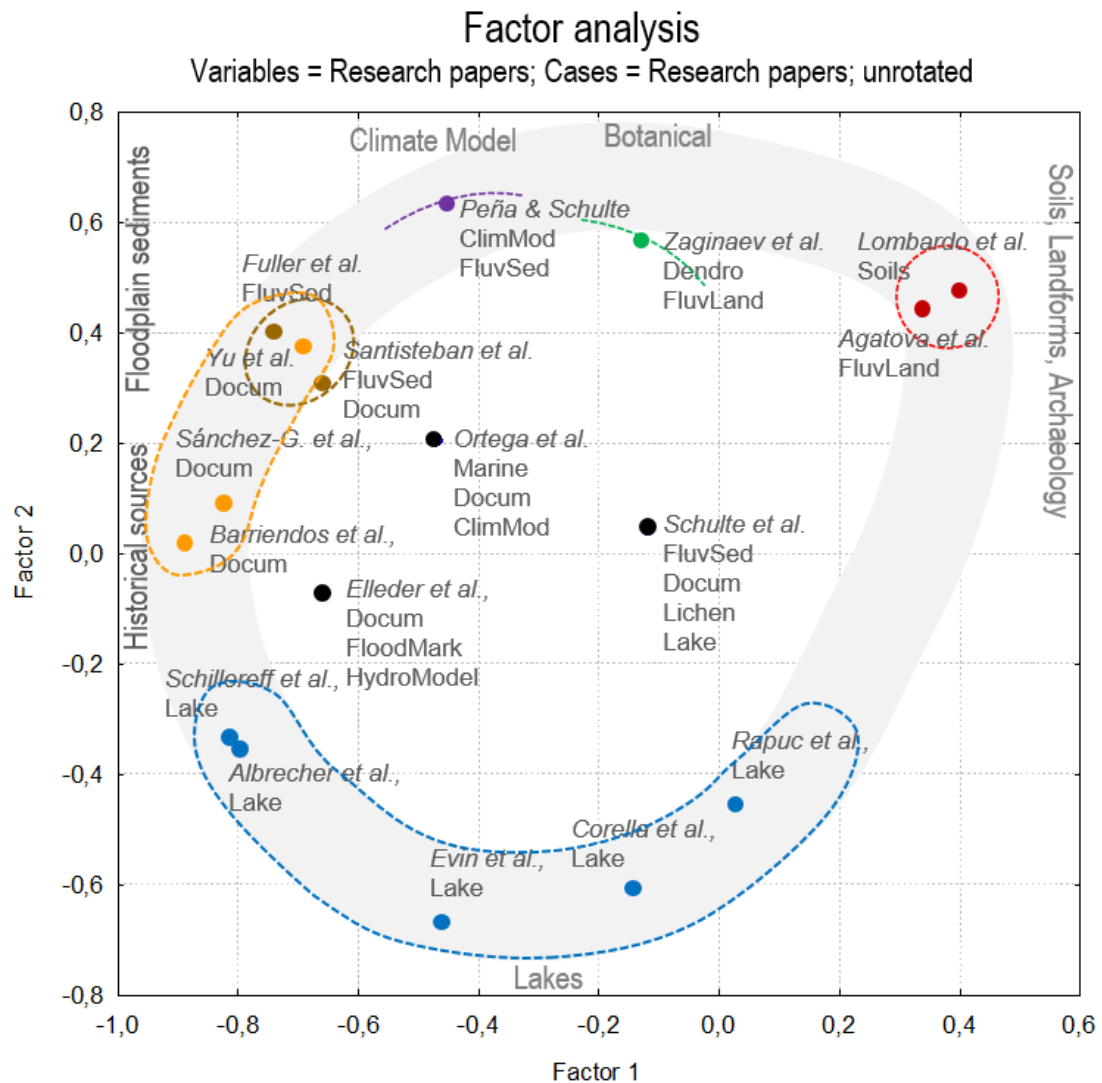
393 Terrestrial archives such as fluvial landforms, deposits, soils, and archeological  
394 sites provide flood information at lower temporal resolution than e.g. lakes but  
395 they can explain the spatial scale of flooding more accurately (Agatova et al., this  
396 issue; Lombardo et al., this issue). With regard to the fluvial sediments in  
397 floodplains, the studies of Fuller et al. (this issue), Santisteban et al. (this issue)  
398 and Schulte et al. (this issue) demonstrate that fluvial deposition mirrors  
399 sensitively severe and medium-magnitude floods.

400

## 401 **7. Perspectives of the integration of paleoflood archives**

402 To understand the epistemic concepts of paleoflood research, the thematic  
403 relationships relationship between the 17 research papers of the Special Issue  
404 were explored. Factor analysis (FA) was performed from the binary matrix  
405 (chapter 3), where research papers are variables and flood archives analyzed by

406 each study are considered as cases. The 2-dimensional plot of the first two  
 407 factors shows the distribution of the research papers. Below the citations (values),  
 408 the types of flood series, generated by each study, are listed additionally.  
 409



410  
 411 Figure 3: 2-D plot of the first two factors explaining the variations and associations of the 17  
 412 paleoflood approaches presented in the Special Issue. The matrix is defined by binary data of  
 413 analyzed types of archives. Below the references, the types of flood series generated by each  
 414 study are listed. Note that the number of generated flood series is lower than the total of flood  
 415 archives used for the compilation of flood series.

416

417 The 2D-plot in figure 3 presents a very clear structure. Papers that used data  
418 from 2 to 4 different types of flood archive are located around the periphery,  
419 essentially defining a circle (grey shading). Those papers which focus mainly on  
420 lake records are located at the bottom (negative loadings of F2). The different  
421 factor F1 loadings of these studies result from the fact that Schillereff et al. (this  
422 issue) and Albrecher et al. (this issue) compare their records with historical  
423 sources and instrumental discharges (also Evin et al., this issue), whereas  
424 Corella et al. (this issue) consider palynological data and Rapuc et al. (this issue)  
425 provide a calibration using precipitation records.

426 Papers that primarily explore historical sources of flood information (Barriendos  
427 et al., this issue; Sánchez-García et al., this issue; Yu et al., this issue) are  
428 situated on the left (strong negative loadings of F1) and show a close relationship  
429 to the studies of floodplain sediments (Fuller et al., this issue; Santisteban et al.,  
430 this issue). This association arises from the fact that both types of archives -  
431 documentary evidence of damage and flooding of settlements and infrastructure  
432 on one hand, and aggradation of overbank deposits, on the other - are sourced  
433 from similar areas of a catchment, e.g. in the floodplains of river valleys and deltas  
434 (Schulte et al., this issue).

435 The loadings and associations of the only paper exclusively dedicated to  
436 paleoclimate modeling (Peña and Schulte, this issue; strong positive loading of  
437 F2) result from the fact that the flood periods of this model were inferred from  
438 geochemical floodplain proxies. The dendromorphological paper presented by  
439 Zaginaev et al. (this issue) is positioned a relatively short distance from the two  
440 papers dedicated to landforms, soils and archaeology (Lombardo et al., this  
441 issue; Agatova et al., this issue; positive loadings of F1 and F2), since trees were

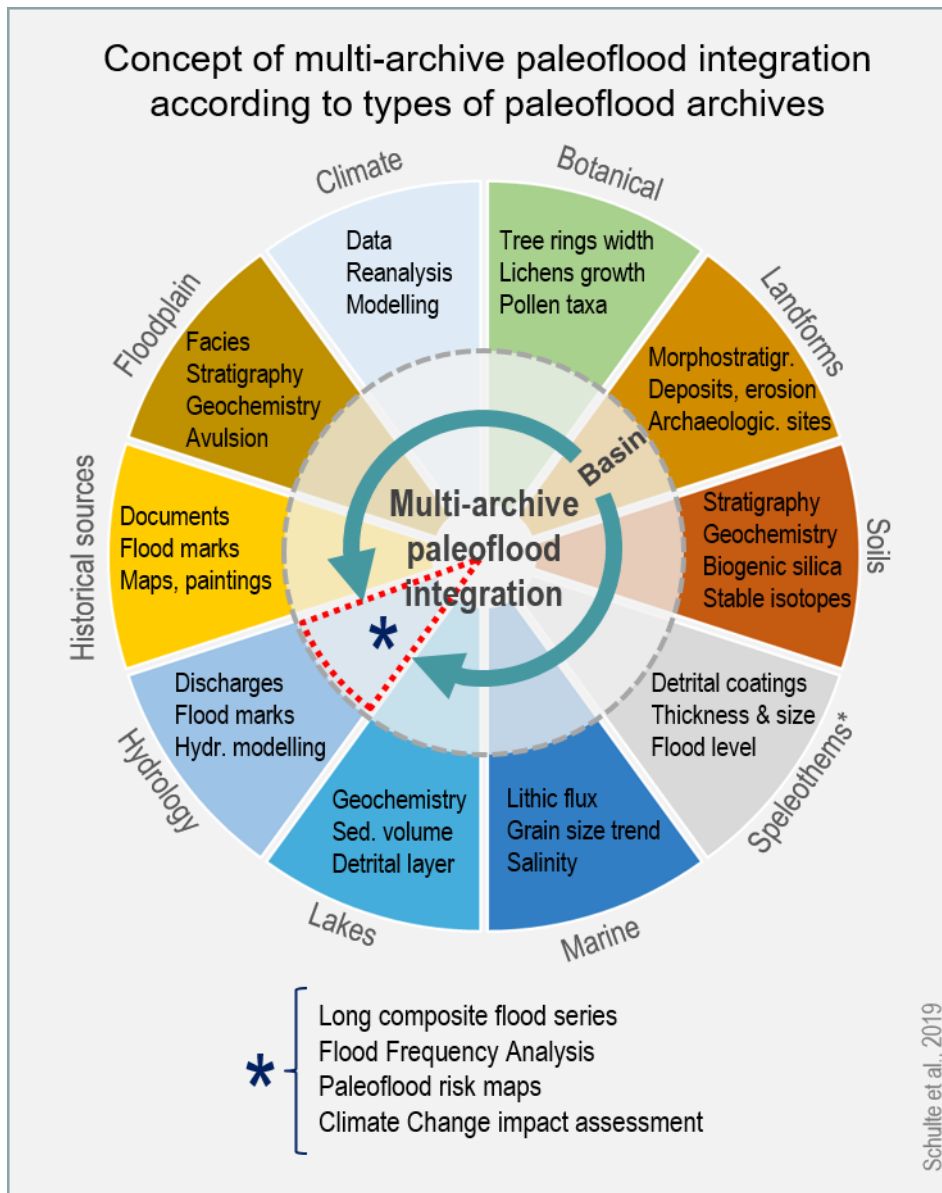
442 sampled on alluvial cones for the reconstruction of flash-floods and debris flow  
443 dynamics. The works of Agatova et al. (this issue) and Lombardo et al. (this issue)  
444 take slightly eccentric positions because these approaches are not strictly related  
445 to the analysis of flood records but rather on fluvial landscape and soil  
446 development.

447 Those papers that show the strongest degree of multi-disciplinary research and  
448 the highest number of integrated flood series are located in the center of the circle  
449 (Figure 3). Furthermore, they performed climatological or hydrological modeling.  
450 In the case of the study of Elleder et al. (this issue), the authors analyzed one  
451 single flood episode reconstructing the propagation of the 1872 flooding by  
452 different documentary and instrumental archives. The study of Ortega et al. (this  
453 issue) focuses on marine, documentary and instrumental archives. Finally, the  
454 widest range of methodologies and archives are presented by the research  
455 cluster of the Bernese Alps (Schulte et al., this issue), where fluvial, lake,  
456 documentary and lichenometric flood series are integrated (Table 1).

457 Despite the rather modest number of cases (defined by the papers published in  
458 this special issue), this factor analysis helps elucidate the different yet  
459 complementary approaches of palaeoflood research.

460 Thus, the distribution of papers and type of flood series (variables in Figure 3)  
461 were integrated into the conceptual model presented in Figure 4.

462



463

464 Figure 4: Concept of multi-archive paleoflood integration according to the type of flood archive.

465 The assembly of the type of flood series mirrors the distribution of papers and constructed flood

466 series shown by the 2D-plot of the Factor Analysis of figure 3. Thus, the concept is not only based

467 on a qualitative background but also on an empirical background inferred from the metadata of

468 Table 1. Paleoflood records from speleothems\* were not presented by any case study of the

469 Special Issue, but they were included to complete the concept.

470

471 The conceptual model is structured in the form of a pie chart, in which each slice

472 represents a type of paleoflood archive. Their arrangement is guided by their

473 nature and statistical association (Figure 3). Furthermore, each slice assembles  
474 the most commonly used paleoflood proxies or techniques applied to that type of  
475 archive. These archive types (external boundary of the chart) largely correspond  
476 to scientific disciplines focused on analyzing past floods. The (sub-) horizontal  
477 slices in the middle (vertical order) represent terrestrial archives (geosphere),  
478 including floodplains, soils, and landforms as well as historical sources. On top of  
479 the geosphere are located botanical archives (biosphere) and climatological data  
480 series (atmosphere). Slices at the bottom of the conceptual model are associated  
481 with the hydrosphere and subaquatic archives: hydrological and hydraulic data  
482 from rivers, sedimentary and environmental archives from lakes and oceans and  
483 speleothem proxies from subsurface flooding.

484 All these flood archives can be integrated by means of statistical processing  
485 (inner circle) to compose synthetic regional flood records that reflect flooding up  
486 to basin-scale. At these points, it has to be stressed again that the combination  
487 of proxies (and archives) and their statistical processing are presented in Table  
488 1 and Figures 2 and 3.

489 Finally, robust multi-archive flood records can provide accurate information for  
490 fundamental concerns of society (bottom of Figure 4):

491 i) Centennial and millennia-long flood calendars allow clusters of extreme  
492 events to be detected as well as changes in trends of flood frequency and  
493 magnitude during periods of changing cold/warm and dry/wet climate  
494 pulses and periods (Schulte et al., 2008, 2009a, 2015; Wilhelm et al.,  
495 2012);

496 ii) Flood Frequency Analyses (FFA) based on long time series of field  
497 evidence (“real flood evidence”) of extreme floods can account for



498 changes in the pattern of flooding during different climate conditions and  
499 cycles (i.e., non-stationarity; Knox, 2000; Mudelsee et al., 2003);

500 iii) Compiling spatial information of paleofloods in thematic maps contributes  
501 fundamental information on local hazard and risk, thus improving river and  
502 flood management and ensuring appropriate spatial planning  
503 (Röthlisberger, 1992; Schulte et al., 2009b, this issue; Geoportal des  
504 Kanton Bern, 2018);

505 iv) The holistic knowledge of flood dynamics during past climate periods (e.g.  
506 RWP, DA, MCA, LIA) is critical for the assessment of the impacts of  
507 flooding in the context of Climate Change (Global Warming).

508 To conclude, the strong arguments in favor of integrated paleoflood approaches  
509 are the possibility of cross-calibrating independent proxies of past floods from  
510 different archives and bringing to light the flood phenomena from different  
511 perspectives. This diversity of foci on past floods and internal validation should  
512 reduce uncertainties and help to identify unusual data in flood series.

513

## 514 **8. Major trends and drivers of paleoflood series**

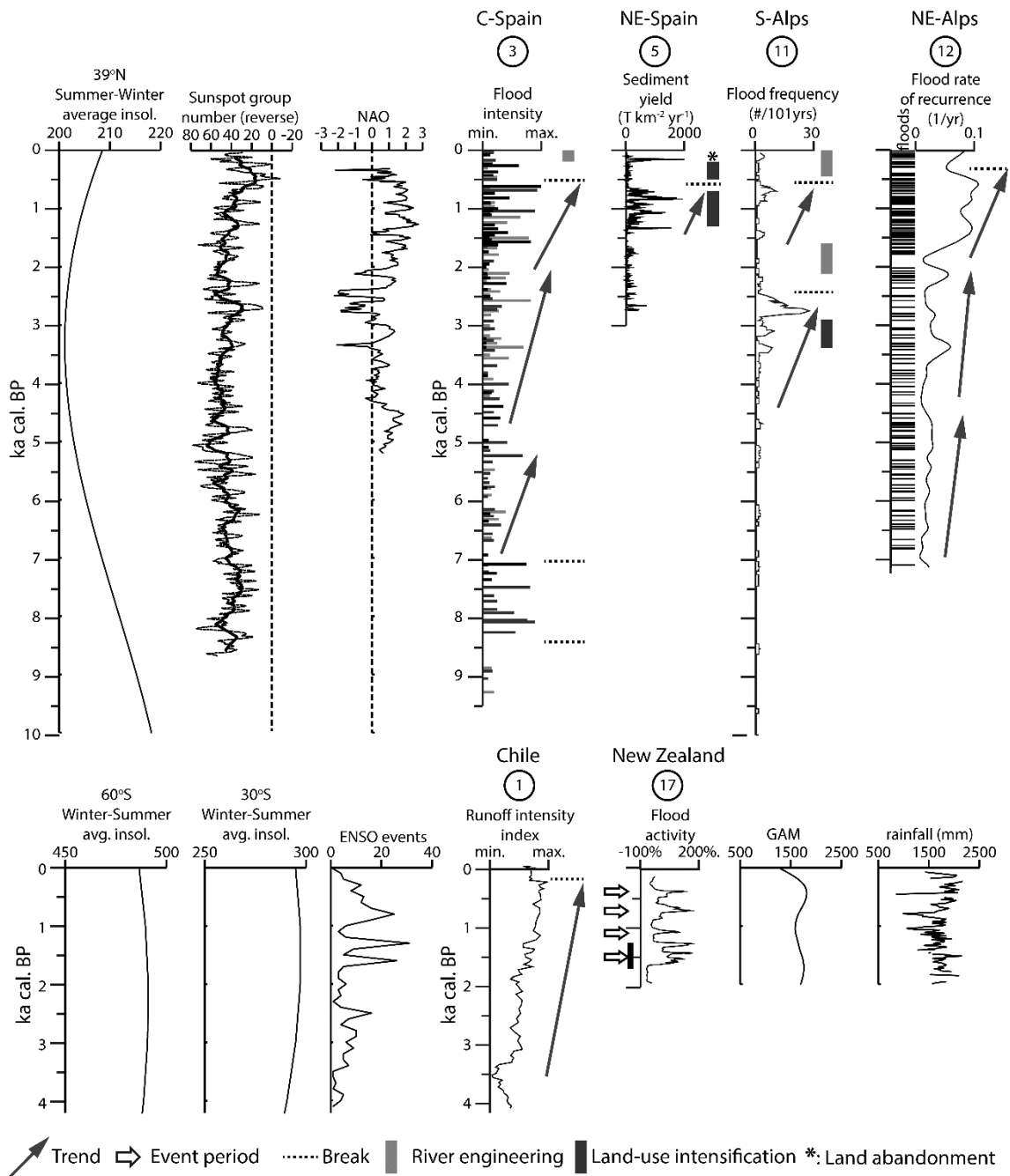
515 Despite the heterogeneity of the presented records in terms of sources, settings  
516 and timescales, a qualitative comparison between the flood series reveals some  
517 consistent patterns. The most evident structural elements in Figures 5 and 6 are:

518 i) similar trends in frequencies and magnitudes of flooding; ii) periods of  
519 noticeable higher activity (“event periods”); and iii) “breaks” at which an abrupt  
520 fall in values occurs where thresholds in fluvial and erosional systems have been  
521 exceeded.

522 Relatively few records of multi-millennial timescale (Fig. 5) are presented, and  
523 they are all based on sedimentary reconstructions. Most long records show an  
524 overall trend of increasing flood frequency/magnitude from ca. 4 to 5 ka BP, with  
525 some earlier episodes of high fluvial activity in Amazonia (before 8 ka BP;  
526 Lombardo et al., this issue) and central Spain (from 8.5 to 7 ka BP; Santisteban  
527 et al., this issue). Patterns of flooding around the world during the last five  
528 millennia show more complexity, which could relate to: i) the higher temporal  
529 resolution of available archives; ii) the higher number of shorter flood records; iii)  
530 the use of documentary sources and a wider range of natural archives spanning  
531 the last millennium; and iv) the progressive intensification of human impact on the  
532 landscape and river systems (Fig. 6). Flood regimes during the last few millennia  
533 are typically characterized by longer (many decades to centuries) periods of  
534 increased flood activity punctuated by short or abrupt drops in flood occurrence  
535 (e.g. 2.2 ka BP at site 3, 1.5 ka BP at site 5, 1.8 ka BP at site 11 and 12 in figure  
536 5). However, it is important to note that these gaps do not occur synchronously  
537 (Fig. 5).

538 Whereas flood activity increases through the mid-Holocene in most records, (i.e.  
539 Albrecher et al., this issue; Rapuc et al., this issue; Santisteban et al., this issue),  
540 the picture of flood trends for the last 2500 years is much more diverse. Rapuc et  
541 al. (this issue) show an overall decrease in flooding activity while Albrecher et al.  
542 (this issue), Ortega et al., (this issue) and Santisteban et al. (this issue) show an  
543 increase until the last centuries.

544



545

① Numbers of study area as in text and table 1 For ③: — Regional event — Sub-basin event — Local event

546

Figure 5.- Millennial-scale flooding episodes according to research papers in the Special Issue

547

and their relation to solar activity (sunspot group number and 21-points [210 years] moving

548

average: Wu et al., 2018; 39°N summer-winter difference calculated using the R package

549

'palinsol', Crucifix, 2016, using the calculations of Berger and Loutre, 1991) and NAO index (Olsen

550

et al., 2012) for European records and insolation at 60°S and 30°S (Berger and Loutre, 1991),

551

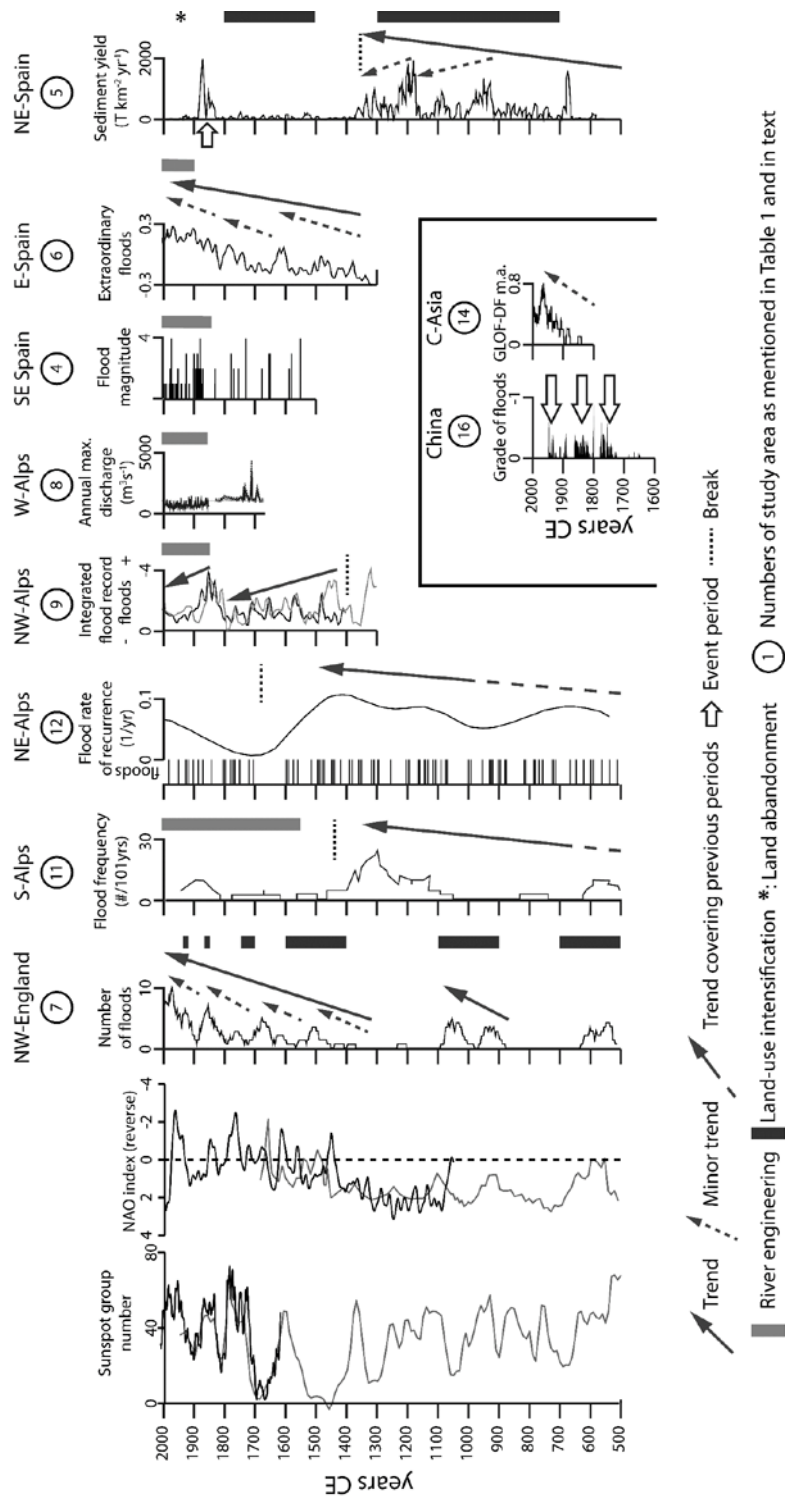
ENSO events (Moy et al., 2002) and ENSO-related GAM and rainfall episodes in subtropical

552

Australia (Barr et al., 2019) for the Pacific records. References of paleoflood studies: site 3 =

553 Santisteban et al. (this issue); site 5 = Corella et al. (this issue); site 11 = Rapuc et al. (this issue);  
 554 site 12 = Albrecher et al. (this issue); site 1 = Ortega et al. (this issue); site 17 = Fuller et al. (this  
 555 issue).

556



557

558 Figure 6.- European flood records spanning the last 2000 years, NAO (in reverse scale) (grey:  
559 Olsen et al., 2012; black: Trouet et al., 2009) and sunspot reconstructions (grey: Usoskin et al.,  
560 2014; black: WDC-SILSO, Royal Observatory of Belgium, Brussels <http://sidc.be/silso/home>).  
561 Similarities emerge in trends and periods of sunspot, NAO and flooding activity. There is a  
562 stronger link between NAO and flooding than at the millennial time-scale. References of  
563 paleoflood studies: site 7 = Schillereff et al. (this issue); site 11 = Rapuc et al. (this issue); site 12  
564 = Albrecher et al. (this issue); site 9 = Schulte et al. (this issue); site 8 = Evin et al. (this issue);  
565 site 4 = Sánchez-García et al. (this issue); site 6 = Barriendos et al. (this issue); site 5 = Corella  
566 et al. (this issue); site 16 = Yu et al. (this issue); site 14 = Zaginaev et al. (this issue).

567

568 Figure 6 presents ten data-rich paleoflood records that span the last 1.5 ka and  
569 show similar variability (Fig. 6). Whilst most records show an increasing trend  
570 over this period, there are exceptions.

571 This heterogeneous pattern continues towards the present, when human action  
572 has become an increasingly important driver. Records showing sudden  
573 decreases (i.e. Corella et al., this issue; Ortega et al., this issue; Schulte et al.,  
574 this issue) coincide at times with periods of increasing activity elsewhere (i.e.  
575 Barriendos et al., this issue; Schillereff et al., this issue; Zaginaev et al., this  
576 issue). These changes have been attributed by the authors to different natural  
577 forcings, depending on the period and timescale. For the millennial timescale,  
578 Rapuc et al. (this issue) and Santisteban et al. (this issue) relate the long-term  
579 trend to changes in insolation that could have affected seasonality and the  
580 persistence of atmospheric patterns (Fig. 5). Over shorter intervals, a number of  
581 regional ocean-atmosphere processes have been invoked as important natural  
582 forcings. Most presented studies for Europe relate flooding activity to negative  
583 NAO phases (e.g. Rapuc et al., this issue; Santisteban et al., this issue; Schillereff

584 et al., this issue), (positive and negative) phases of summer NAO (Peña and  
585 Schulte, this issue; Schulte et al., this issue), changes in solar activity (Peña and  
586 Schulte, this issue; Schillereff et al., this issue) or cold phases linked to Atlantic  
587 multidecadal variations (Barriendos et al., this issue). For the Pacific domain,  
588 floods have been correlated to displacement of the westerlies/monsoon systems  
589 (Pacific Decadal Oscillation, PDO; Southern Annular Mode, SAM) and ENSO  
590 (Fuller et al., this issue; Ortega et al., this issue; Yu et al., this issue).

591 However, the frequency and intensity of flood events are the result of diverse and  
592 interacting factors operating at the local, regional and global scales. This  
593 produces complex records that are challenging to interpret. For example, Corella  
594 et al. (this issue) show a lake sediment record that responds to the seasonal  
595 distribution of storms and longer-term changes in soil properties (resulting from  
596 climate and land-use change).

597 The comparisons presented in Figures 5 and 6 (and in the papers mentioned)  
598 highlight that temporal correlations between regional forcings and flood  
599 reconstructions can rarely be drawn precisely. However, numerous studies in the  
600 Special Issue emphasize that the 19<sup>th</sup> century - including the most recent cool  
601 climate pulses during the Little Ice Age - is a particularly flood-rich period  
602 (Barriendos et al., this issue; Rapuc et al., this issue; Schillereff et al., this issue;  
603 Sánchez-García et al., this issue; Yu et al., this issue) and the period with highest  
604 flood intensity in some regions (Corella et al., this issue; Schulte et al., this issue).

605 This could be a consequence of synergetic effects between climate forcing and  
606 human factors (land-use, river management). With regard to the 20<sup>th</sup> century, it  
607 is difficult to assess the influence of global warming because of intensifying social  
608 factors, the effect of hydraulic infrastructure and management, and demographic

609 and urban growth (Barriendos et al., this issue; Sánchez-García et al., this issue).  
610 Research drawing on historical sources might also be affected by the increased  
611 availability of flood information about smaller and moderate floods since the  
612 second half of the 19<sup>th</sup> century. In addition, some flood types such as GLOFs and  
613 debris flows can be favored by particular physiographic settings like the formation  
614 of new glacier lakes in the Tien Shan mountains (Zaginaev et al., this issue). It is,  
615 however, noteworthy that 10 of the 14 papers displayed in Figures 5 and 6 do not  
616 record the 20<sup>th</sup> century as the exceptional flood period.

617 Based on this synthesis of current research, palaeoflood data are optimally  
618 explored at decennial to centennial time scales (trends or periods). Finer-  
619 resolution data are highly desirable but we must be conscious of multiple limiting  
620 factors. For example, the mixing depth of sedimentary records may limit data  
621 resolution. In addition, the timing of flooding in adjacent catchments can differ  
622 considerably, as shown by Schulte et al. (this issue) in the Bernese Alps,  
623 Santisteban et al. (this issue) in Central Spain and Barriendos et al. (this issue)  
624 in eighteen catchments of eastern Spain. This is unsurprising because  
625 hydrometeorological processes, sensitivity to climate variability or land-use  
626 change and thresholds may be site-specific. Palaeoflood research must account  
627 for this diversity. There are many potential pathways towards improvement: new  
628 chronological tools that circumvent technical limitations, such as <sup>14</sup>C  
629 plateaus/anomalies coupled to higher-resolution, multi-proxy studies, basin-scale  
630 studies, these need to be coupled with improved reconstructions of local factors,  
631 especially human activity, and climate forcing. For example, evaluating NAO or  
632 ENSO as a long-term flood driver is limited by the temporal resolution of NAO

633 reconstructions. These requirements should guide future research and will be  
634 best achieved by expanding collaborative efforts.

635

## 636 **9. Human impact and disentangle anthropogenic from natural drivers**

637 Anthropogenic modifications to the landscape can dramatically alter flood  
638 regimes. Land-use change, in particular, the removal of mature vegetation covers  
639 and intensification of agriculture, destabilizes hillslopes and increases surface  
640 runoff and soil erosion potential (Dotterweich, 2008; Hoffmann et al., 2010).  
641 Structural interventions in rivers also affect the flood hazard (Schulte et al., 2015;  
642 Wetter et al. 2017; Munoz et al., 2018). Disentangling climatic and anthropogenic  
643 forcings in palaeoenvironmental data is a persistent challenge, however (Mills et  
644 al., 2017). Palaeoflood researchers are cognizant of these difficulties (Brisset et  
645 al., 2017; Wilhelm et al., 2019) and this set of papers presents an opportunity to  
646 explore current approaches and limitations when evaluating the human influence  
647 on long-term flood risk.

648 Some common approaches are evident, in general, but also in this Paleoflood  
649 Special Issue. There is common agreement that regional consistency across  
650 multiple proxies from independent archives denotes a climate signal whereas  
651 localized shifts probably point towards human disturbance (e.g. Barriendos et al.,  
652 this issue; Fuller et al., this issue; Rapuc et al., this issue; Sánchez-García et al.,  
653 this issue; Schulte et al., this issue). There is also evidence that interactive effects  
654 characterize the climate-human-flood nexus. For example, flood occurrence in  
655 northern Britain correlates with solar activity and NAO dynamics but concurrent  
656 woodland clearance for pastoralism appears to amplify flood frequency and  
657 magnitude (Schillereff et al., this issue). Similarly, Rapuc et al. (this issue) show



658 a striking increase in flood frequency at Lake Iseo (Italy) around 3000 yr BP that  
659 coincides with forest clearance indicators but lags the onset of climate-driven  
660 catchment erosion. Subsequent channel diversion, however, reduced  
661 sedimentation and discontinued the depositional flood record in some areas of  
662 lake basins (Rapuc et al., this issue) and floodplains (Carvalho and Schulte,  
663 2013). Researchers must be aware that channelization can equally produce  
664 areas of aggradation and delta progradation (Schulte et al., 2009a; Wirth et al.,  
665 2011; Santisteban et al., this issue). Crucially, the effects of human pressure on  
666 flood regimes are often irregular through time. Corella et al. (this issue) infer from  
667 high-resolution, multi-proxy palaeoecological data that burning, grazing and  
668 cultivation triggered a prolonged period of frequent flooding around 700-1300 AD  
669 in northeast Spain. Equivalent activity during later centuries produced a more  
670 muted flood response, suggesting the system rebalanced to accommodate  
671 disturbance.

672 Hydraulic infrastructure also induces complex effects. Flood risk declined on  
673 Spanish rivers after expansive 20<sup>th</sup>-century dam building (Barriendos et al., this  
674 issue). Similarly, in several catchments of the Alps the combination of river  
675 correction and diversion into large alpine lakes lower peak downstream  
676 discharges because these lakes regulate flood waters like large retention areas  
677 (Wetter et al., 2011, Schulte et al., this issue). Conversely, channelization  
678 amplified flood magnitudes on the Mississippi (Munoz et al., 2018) while Elleder  
679 et al. (this issue) show the Mladotice dam collapse (Czech Republic) was a key  
680 factor behind the destructiveness of the 1872 flood. Sánchez-García et al. (this  
681 issue) note more recent dam construction, such as on the Almanzora River, Spain

682 (Sánchez-García et al., this issue), has shifted the hydrometeorological baseline.  
683 This hinders efforts to evaluate flood risk under 21<sup>st</sup>-century climatic warming.  
684 This compilation of evidence stresses that the role of human activity as a flood  
685 driver and proxy evidence used to characterize an anthropogenic signature must  
686 be established on a site-by-site basis. This presents a number of challenges: first,  
687 while the multi-proxy approach of Corella et al. (this issue) that distinguishes the  
688 impacts of grazing, burning and cultivation is a powerful tool and Schulte et al.  
689 (this issue) sift their synthesis of the European Alps for sites least affected by  
690 human presence, such bountiful paleoflood data is rare. Second, we must keep  
691 in mind that vegetation and fire dynamics can respond independently to climate  
692 so evidence of human-induced changes is crucial (Corella et al., this issue). Third,  
693 inter-site comparisons will widen chronological uncertainty. For example, most  
694 periods of frequent flooding in northern Britain (Schillereff et al., this issue) align  
695 with geomorphic evidence of hillslope destabilization across the region but both  
696 datasets depend on radiocarbon dates with multi-decadal uncertainties.  
697 Communicating the temporal uncertainty for comparator datasets should,  
698 therefore, be encouraged. Similarly, the approach of Rapuc et al. (this issue) to  
699 re-examine local archaeological evidence in a flooding context builds confidence  
700 in the time-transgressive role of human activity. Lastly, we must be wary of  
701 circular reasoning: catchment destabilization usually alters the rainfall-runoff  
702 relationship and soil erodibility. In this scenario, an anomalous lamination need  
703 not reflect a major flood because of modest rainfall on recently exposed, fragile  
704 soils could produce a similar signature. Mechanistic interpretations that consider  
705 transport capacity, for example, become crucial (Evin et al., this issue; Schillereff  
706 et al., this issue). Similarly, Corella et al. (this issue) convincingly attribute a shift

707 in flood seasonality to human influence because known advancements in tilling  
708 practice would create the observed response. Site selection is also important. For  
709 example, Schulte et al. (this issue) show a divergent flood response on either  
710 flank of the Bernese Alps.

711 Isolating the anthropogenic component can be a formidable challenge. This  
712 collection of papers showcase how a coupled multi-archive, multi-proxy approach  
713 is imperative. In sedimentary systems, for example, do different geochemical  
714 proxies reflect rates of erosion and the carrying capacity of the system (grain  
715 size). Determining the nature of human modifications must also occur on a  
716 specific basis. Channel diversion and hillslope vegetation clearance may exert  
717 quite different effects on the flood regime. This could draw on archaeological  
718 evidence, for example (Rapuc et al., this issue). As human and climatic modifiers  
719 may coincide, less equivocal proxy evidence of human presence is needed.  
720 Indicator pollen (Corella et al., this issue) or pastoral DNA (Giguet-Covex et al.,  
721 2014) show real promise but may require deeper collaboration amongst  
722 international research groups in the future. Integrating sedimentary data with an  
723 independent archive, such as historical documents or tree rings would also be  
724 wise.

725

## 726 **10. Outlook: scoop and limitations of multi-archive paleoflood integration**

727 Taking a long-term perspective on flooding is fundamental for adequate hazard  
728 and risk assessment (e.g. flood-frequency analysis). The research papers in this  
729 Special Issue demonstrate that integrating field-data on “real” past floods derived  
730 from multiple historical and natural archives provides excellent flood data series.  
731 These datasets are uniquely positioned to document low-frequency, large-

732 magnitude flood events that vary under different climate regimes (cooler, warmer  
733 and transitional climate periods) and/or environmental conditions (changes in  
734 land cover, land use, and river management).

735 However, reliable and accurate integration of paleoflood data relies on  
736 accounting for several critical issues:

737 i) Although this paper does not focus on the dating of flood records, we  
738 would like to stress first that considering dating uncertainties within the  
739 time series is vital prior to perform statistical analysis. The temporal  
740 resolution of records can vary significantly, however. Before flood data  
741 series from different natural and anthropogenic archives can be  
742 integrated into a regional model, a critical assessment of chronological  
743 models and homogenization of flood data is needed.

744 ii) The comparability of flood series from heterogeneous catchments and  
745 landscapes is often complex because the controlling factors and  
746 system sensitivity (e.g. of erosion or aggradation) to climatological  
747 conditions and hydrological extremes varies greatly across diverse  
748 hydrological and environmental settings.

749 iii) Indirect flood proxies recording climatic-environmental signals (e.g.  
750 sediments that are deposited by surface runoff in a small sub-  
751 catchment) are different from records that are directly involved in the  
752 process of river flooding. Thus a precise understanding and a careful  
753 interpretation and/or calibration of physical processes are mandatory.

754 iv) Due to the heterogeneous natural response of different subsystems to  
755 flood drivers, not all flood series from a basin or a region can be  
756 integrated into a regional synthetic paleoflood master curve. The

757 criteria for selection or rejection of individual flood series must not only  
758 follow statistical protocols but also consider process-based arguments.  
759 To identify “false alarms” and “missed” floods, data series should be  
760 tested against known regional hydrological extreme events that are  
761 documented by several records.

762 v) Human modification to many river systems has had major effects on  
763 the flood regime. These non-stationary conditions impose challenges  
764 when performing flood frequency analysis and evaluating flood risk  
765 under future climate change projections. Effort should be invested to  
766 isolate the anthropogenic component, ideally through a coupled multi-  
767 archive, multi-proxy approach.

768 vi) There is good evidence that some regional/global factors can  
769 systematically affect the dynamics of flooding over long timescales.  
770 There is now a need to achieve wider spatial and temporal coverage,  
771 leading to better understand the factors affecting intra-basin variability,  
772 especially interactions between natural and anthropogenic forcings.

773 When critical points i) – vi) are carefully taken into consideration, it seems clear  
774 that using a variety of paleoflood data from multiple archives and methodologies  
775 from different scientific fields is the best approach. Such multi-dimensional  
776 investigations can better account for limitations in individual records and more  
777 effectively analyze the spatial distribution (horizontal and vertical) of flood records  
778 in order to capture the physiographic and environmental diversity of a catchment.  
779 Schulte et al. (this issue) conclude from their integrated paleoflood pilot project in  
780 the Alps that such a multi-archive methodology can be applied in many regions.  
781 They do recommend, however, that this approach will be most effective in

782 catchments where a high number of paleoflood records already exist and a  
783 profound understanding of the different flood proxies and flood generating  
784 mechanisms has been built up.

785 The meta-data of the case studies presented in this Special Issue in Table 1  
786 suggest that paleoflood studies focused on fluvial depositional environments  
787 show a higher rate of integration with other types of paleoflood archive (mean of  
788 4.5 types of archive) than studies focused on documentary sources (mean of 3.5)  
789 and lake sediments (mean of 2.4). We suggest that this adopted strategy of cross-  
790 correlation is an effective method to compensate for the higher uncertainties of  
791 fluvial deposition in floodplains due to lower temporal sample resolution and  
792 possible effects of unconformities (possible gaps of flood information). The  
793 apparently more precisely dated lake and documentary flood records focus  
794 predominantly on instrumental calibration instead of multi-archive integration.  
795 However, several studies showed that neither of these series always record  
796 continuous flood information and they should be completed by other archives. In  
797 addition, spatial accuracy of flooding processes is a weak point of studies where  
798 flood information is obtained from single locations such as lakes, flood marks or  
799 settlements instead of larger flood-prone areas. In this latter case, terrestrial  
800 natural flood archives contribute highly valuable information.

801 Based on the gathered experience from the Special Issue, the activities of the  
802 work package WP2 and the FWG pilot multi-archive project, we suggest that over  
803 the next few years the agenda of regional paleoflood research might include the  
804 following trends:

- 805 i) Design of methodological approaches integrating paleoflood datasets  
806 through numerous regional pilot studies in different environments,

- 807 ii) Improvements in flood frequency analysis and spatial flood risk  
808 assessment using multi-archive analysis,
- 809 iii) The progress of methodological and statistical approaches to analyze  
810 paleoclimate models in relation to the flood variability,
- 811 iv) Assessment of the changes of flood pattern due to the effect of land-use  
812 changes in basins,
- 813 v) Examination of changes in external forcing and atmospheric variability of  
814 the flood periods by paleoclimate modeling. Furthermore, these  
815 comparisons help to predict future climate change related to flooding.
- 816 vi) Detailed consideration of the role of anthropogenic landscape modification  
817 in controlling flood dynamics and develop and apply multi-proxy  
818 approaches to disentangle the effects of human activity from climatic  
819 drivers of flooding. This will be most effectively achieved through wide  
820 collaborations.

821 We, the invited editors of the present Special Issue and authors of this  
822 introductory article, hope that the collection of papers presented in the Paleoflood  
823 Special Issue and at the PAGES Open Scientific Meeting in 2017 contribute to  
824 the progress of paleoflood research and inspire interested readers to promote the  
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830

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851

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