Dual-Element Isotope Analysis of Desphenylchloridazon to Investigate its Environmental Fate in a Systematic Field Study - A Long-Term Lysimeter Experiment

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14 Abstract

15 Desphenylchloridazon (DPC), the main metabolite of the herbicide chloridazon (CLZ), is more water 16 soluble and persistent than CLZ and frequently detected in water bodies. When assessing DPC 17 transformation in the environment, results can be non-conclusive if based on concentration analysis 18 alone, because estimates may be confounded by simultaneous DPC formation from CLZ. This study 19 investigated the fate of DPC by combining concentration-based methods with compound-specific C and 20 N stable isotope analysis (CSIA). Additionally, DPC formation and transformation processes were 21 experimentally deconvolved in a dedicated lysimeter study considering three scenarios. First, surface 22 application of DPC enabled studying its degradation in the absence of CLZ. Here, CSIA provided evidence of two distinct DPC transformation processes: one shows significant changes only in ${}^{13}C/{}^{12}C$, whereas the 23 other involves changes in both ¹³C/¹²C and ¹⁵N/¹⁴N isotope ratios. Second, surface application of CLZ 24 mimicked a realistic field scenario showing that during DPC formation, $^{13}C/^{12}C$ ratios of DPC were depleted 25 26 in ¹³C relative to CLZ, while ¹⁵N/¹⁴N ratios remained constant. Finally, CLZ depth injection simulated 27 preferential flow and demonstrated the importance of the topsoil for retaining DPC. The combination of 28 the lysimeter study with CSIA enabled insights into DPC transformation in the field that are superior to 29 studies of concentration trends.

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31 Introduction

32 Groundwater is one of the most important drinking water resources¹ and, therefore, constantly screened 33 for contaminants²⁻⁵. Due to their extensive application in agriculture, pesticides and their metabolites⁶ are 34 commonly detected in ground and surface water. A prominent example is desphenylchloridazon (DPC), 35 the main metabolite of the herbicide chloridazon (CLZ). CLZ is a selective systemic herbicide that is used to control broad-leaved weeds in the agricultural production of swiss chard, red beet and sugar beet⁶⁻¹¹. 36 37 The metabolite DPC is a compound of concern as it is continuously formed from CLZ. The continuous input of newly formed DPC makes it challenging to evaluate its environmental transformation from 38 39 concentration data over time. Detection of DPC has increasingly been reported exceeding concentrations of 10 µg/L in natural water bodies^{6, 11-14}. DPC can be transported into ground and surface water by 40 precipitation events as it is water-soluble (490 mg/L), and has a lower tendency to bind to the soil 41 42 (Freundlich constant K_{foc} of 50 mL/g) than CLZ (K_{foc} of 199 mL/g). Additionally, DPC has a high leaching potential, which is indicated by the groundwater ubiquity score (GUS) of 5.5, a parameter used to evaluate 43 pesticides for their potential to seep into the groundwater ^{9, 15, 16}. Thus, there is great interest in the 44 45 question whether DPC can be subject to further transformation. The fate of DPC, however, is not well understood yet ^{2, 17, 18}. It is known that DPC is a persistent and polar compound. In soil, it can be further 46 transformed into methyldesphenylchloridazon (MDPC, Figure S1)^{10, 12, 19, 20}. Whether there is a wider range 47 48 of degradation pathways, remains unclear.

49 Current attempts to quantify degradation of organic micropollutants are often based on metabolite-to-50 parent-compound ratios. This is an analytical approach based on concentration measurements. It is advantageous to quantify degradation even at low concentration ranges, and is simple to use²¹. However, 51 52 in case of DPC, which may be simultaneously formed while undergoing further transformation (Figure S1), metabolite-to-parent ratios can lead to erroneous interpretations²². An additional confounding factor is a 53 54 different drainage-dependent re-mobilization of the parent compound and the metabolite due to 55 differences in their mobility. Thus, concentrations may fluctuate in a non-trivial manner making it difficult, if not impossible, to inform about how much of the DPC has been transformed. Consequently, a 56 complementary method is needed to detect transformation if metabolite analysis alone is not conclusive. 57

58 Compound-specific stable isotope analysis (CSIA) allows to identify degradation processes by analyzing 59 variations of natural stable isotope abundances of different isotopic elements during (bio)degradation

and transformation of organic contaminants²³⁻²⁶. While CSIA of polar micropollutants has rarely been 60 performed at field scales²⁶, analytical methods for the analysis of carbon (¹³C/¹²C) and nitrogen (¹⁵N/¹⁴N) 61 isotope ratios of DPC have recently become available²⁷. So far, isotope studies of DPC have been carried 62 out neither in laboratory experiments nor in field applications, however. As illustrated in Figure S1, unique 63 insight on the formation and subsequent transformation of DPC can be expected. On the one hand, ¹³C/¹²C 64 and ¹⁵N/¹⁴N ratios of DPC are expected to show the isotopic signature of the pyridazinone ring in the 65 66 precursor CLZ. When CLZ is transformed, its phenyl-ring is first oxidized and then cleaved off. Thus, any isotope effect-induced changes in ¹³C/¹²C and ¹⁵N/¹⁴N ratios will be manifested in the molecular average 67 of CLZ and in the oxidized phenyl-part that is cleaved off. In contrast, none of the molecular positions of 68 69 the pyridazinone-ring are involved in the reaction, meaning that only secondary kinetic isotope effects 70 occur so that the isotope ratios within the pyridazinone-ring remain mainly unaffected when they end up 71 in DPC (Figure S1). If, however, further transformation of DPC takes place, this process is expected to 72 result in pronounced changes in isotope ratios in DPC, because now, carbon and nitrogen atoms are 73 directly involved (primary isotope effect). This would lead to carbon and nitrogen isotope fractionation in DPC giving a strong indication of further DPC transformation²⁸. CSIA of DPC, therefore, holds promise to 74 identify both processes, formation of DPC from CLZ, as well as independent further transformation of DPC. 75 76 According to the current mechanistic picture, DPC is only formed from CLZ and transformed through Nmethylation^{19, 20, 29}. Thus, the combined analysis of carbon and nitrogen isotope ratios of DPC may offer 77 78 new insights into its fate in soil leachate.

Evidence from CSIA may be inconclusive, however, if physical processes (e.g., multiple sorption-79 80 desorption steps, dissolution from non-aqueous phase, volatilization/diffusion, dispersion) or the 81 heterogeneity of the system, the soil in this case, affect degradation-induced changes in isotope ratios. 82 For example, a freshly dissolved compound, which has not been transformed yet, can mix with water 83 containing the contaminant that has already undergone varying degrees of degradation and thus isotope fractionation³⁰⁻³². Consequently, the transformation-induced isotope ratios in the degraded fraction might 84 85 not be discernible any longer^{33, 34}. When applying CSIA to a field site either for the interpretation of a 86 compound's environmental fate or to monitor the success of remediation processes, it is therefore 87 suggested to combine it with complementary approaches in order to obtain as many lines of evidence as 88 possible^{30, 35, 36}.

Thus, the aim of this study was to explore different complementary and innovative approaches for assessing the environmental long-term fate of DPC in drainage water after agricultural application over a

91 period of 3 years. To that end, we combined concentration measurements with the analysis of carbon and nitrogen isotope ratios in a comprehensive and systematic study in a well-characterized model lysimeter 92 93 system. This lysimeter system mimics pesticides fate in natural soil environment under high control over 94 environmental and hydrological factors (i.e. soil type and humidity, precipitation levels, temperature, 95 evapotranspiration, etc.). In order to separate the relevant transport and transformation processes, these 96 complementary approaches were integrated into a dedicated experimental design where CLZ and DPC 97 were applied in three different scenarios (Figure S2): (i) DPC was applied to the lysimeter directly, without the presence of CLZ, to investigate whether further DPC transformation is observable in drainage water 98 and whether this transformation is detectable from analyzing carbon and nitrogen isotope signatures of 99 100 DPC when interfering simultaneous formation of DPC can be excluded. (ii) The concurrent formation of 101 DPC from CLZ and potential DPC transformation were evaluated through surface application of CLZ to the 102 lysimeters. (iii) To simulate the preferential flow and to study whether DPC formation and transformation 103 is also occurring below the top soil, CLZ was injected below the root zone. For each scenario, these 104 complementary approaches were tested with two different soil types through a replication of the 105 lysimeter studies with moraine and gravel soil, respectively.

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107 Experimental / Methods

108 Experimental Set-up of Lysimeter Experiments. For this study, the lysimeter facility from Agroscope was used, located in Zurich-Reckenholz, Switzerland. The facility itself and the characteristics of the lysimeters 109 110 are described in detail by Torrentó et al.³⁷. Briefly, the site consisted of 12 gravitation lysimeters (L) (3.14m² surface area, 2.5 m depth, approximately 14 000 kg of soil in each) filled with two soil types 111 112 (gravel/moraine). Both soil types consisted of repacked Cambisol. Cambisols, widely and intensively used as agricultural land, are among the most extensive soil types on earth, extending over about 11 % of the 113 global land surface³⁸. The soils used in this study differed in the B horizon and the draining properties of 114 115 the parent material, and thus they were expected to show a different extent of preferential flow³⁷. Gravel 116 soil was represented by well-drained sandy loamy Cambisol (L1-L6), while moraine soil consisted of a 117 poorly drained loamy Cambisol (L7-L12) (Table S1). Six of these lysimeters were used for this study (three of each soil type). The lysimeters were planted in 2014 with corn (Zea mays L.) followed by sugar beet 118 (Beta vulgaris ssp. vulgaris var. altissima Doel) in 2015, with corn (Zea mays L.) again in 2016 and finally 119 120 with broccoli, Chinese cabbage, lettuce and leek in 2017. 3.0 kg/ha (0.96 g/lysimeter) of CLZ were applied

121 on the surface of two lysimeters (L4 and L8) simulating the scenario of pesticide application at the threeto four-leaf stage in the field¹⁰. To simulate preferential transport through topsoil, two additional 122 123 lysimeters were used (L6 and L7), where 2.0 g of CLZ were injected in each lysimeter at a depth of 40 cm 124 at eleven injection points uniformly distributed over the area of each lysimeter by using a metal rod 125 connected to a gear pump through a Teflon tube. Additionally, 3.2 kg/ha (1.0 g/lysimeter) DPC was applied 126 on the surface of two lysimeters (L1 and L12). In addition to CLZ or DPC, the following tracers were applied 127 at the same time as the pesticides: uranine (1.3 kg/ha) and NaBr (500 kg/ha) to lysimeters L1 and L12, uranine (1.3 kg/ha) to lysimeters L4 and L8, and uranine (0.4 g injected in each lysimeter) to lysimeters L6 128 129 and L7. Bromide was used as conservative tracer and uranine (K_{foc} of 120 mL/g) as a marker for preferential 130 leaching shortly after pesticide application³⁷. A detailed set-up is shown in the Supporting Information 131 (sections II.2 and II.3). Details about application methods can be found in Torrentó et al.³⁷. All lysimeters 132 were irrigated artificially and the seepage water was collected for analysis over a time period of 3 years 133 (Table S2).

Concentration Measurements of CLZ, DPC and MDPC. For concentration measurements of CLZ, DPC and 134 135 MDPC, an Ultimate® 3000 RS high-pressure liquid chromatography (HPLC) (Dionex, Thermo Fisher 136 Scientific, Waltham, MA, USA) coupled to a 4000-hybrid triple quadrupole-linear ion trap mass 137 spectrometer (QTRAP[®], ABSciex, Framingham, MA, USA) was used. Five microliters were injected on an 138 Acquity UPLC BEH Shield RP18 column (100 × 2.1 mm, 1.7 μm, Waters, Milford, MA, USA) maintained at 139 25 °C. The separation was performed at a flow rate of 0.4 mL/min using a binary mobile phase system 140 consisting of 0.05% formic acid in water (mobile phase A) and 0.05% formic acid in acetonitrile (mobile 141 phase B) according to the following gradient program: 5-15 % phase B in 2 min, 15-100 % phase B in 4 min, 142 holding at 100 % phase B for 2 min, and re-equilibration at 2 % phase B for 6 min. Detection was 143 performed in electrospray positive ionization (ESI+) using the multiple reaction monitoring (MRM) mode 144 by monitoring both a quantifier (Q) and a qualifier (q) transition ion for each compound. Precursor and 145 fragment ions (m/z) were 222.1 and 104.0 (Q) or 77.0 (q) for CLZ, 146.0 and 117.0 (Q) or 66.0 (q) for DPC, 146 160.0 and 117.0 (Q) or 88.0 (q) for MDPC, and 227.0 and 108.0 (Q) or 81.0 (q) for CLZ-d₅, respectively 147 (Table S3). Quantification was performed using standard curves calculated from standard solutions of CLZ, DPC and MDPC at 0.25, 0.5, 1, 3, 5 and 10 ng/mL, each containing deuterated CLZ-d₅ as internal standard 148 149 at a constant concentration of 2 ng/mL. The limits of quantification were 0.05 μ g/L for CLZ, 0.4 μ g/L for 150 DPC and 0.1 μ g/L for M-DPC. For those drainage water samples with CLZ, DPC and MDPC concentrations 151 lower than 0.2 µg/L, solid-phase extraction (SPE) of 20-mL samples was performed using 6 mL cartridges 152 packed with 0.2 g of Bakerbond SDB-1 sorbent and 0.2 g of Sepra ZT sorbent, as described by Torrentó et al.³⁹. After SPE, the extracts were analyzed by UHPLC-QTOF-MS. The method is briefly described in the
 Supporting Information (II.5.).

Large Volume Solid-Phase Extraction. For isotope analysis, all lysimeter samples were filtered through
 0.7-μm glass fiber filters and were concentrated by SPE using the method described in Torrentó et al.³⁹,
 as detailed in the Supporting Information (II.6.).

Elemental Analyzer-Isotope Ratio Mass Spectrometry Measurement for Determination of Reference Values. Carbon and nitrogen isotope reference values of our in-house standards of CLZ, DPC and MDPC were determined by elemental analysis – isotope ratio mass spectrometry (EA-IRMS) according to the method of Meyer et al.⁴⁰. The system consisted of an EuroEA (Euro Vector, Milano, Italy) coupled with a Finnigan MAT 253 IRMS via a FinniganTM ConFlow III interface (Thermo Fisher Scientific, Bremen, Germany). For calibration, USG 40, USG 41 (L-glutamic acid) and IAEA 600 (caffeine), supplied by the International Atomic Agency (IAEA), were used as organic reference materials.

165 Carbon (δ^{13} C) and nitrogen (δ^{15} N) isotope signatures are usually expressed using the Delta notation in per 166 mille as described in equation 1 and 2. There, the isotope ratios (13 C/ 12 C_{sample} and 15 N/ 14 N_{sample}) are stated 167 relative to the international references PeeDee Belemnite (V-PDB) for carbon and air for nitrogen.

$$\delta^{13} C = \frac{{}^{13} C / {}^{12} C_{\text{Sample}} {}^{-13} C / {}^{12} C_{\text{Reference}}}{{}^{13} C / {}^{12} C_{\text{Reference}}}$$
(1)

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$$\delta^{15} N = \frac{{}^{15} N / {}^{14} N_{\text{Sample}} {}^{15} N / {}^{14} N_{\text{Reference}}}{{}^{15} N / {}^{14} N_{\text{Reference}}}$$
(2)

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Carbon Isotope Analysis of DPC by LC- IRMS. For carbon isotope analysis of DPC we applied the method 170 of Melsbach et al.²⁷. Briefly, 10 to 100 µL of SPE extracts reconstituted in ultrapure water were injected 171 172 into an LC-IRMS Dionex system consisting of an Ultimate 3000 HPLC pump and an Ultimate 3000 173 autosampler (Thermo Fisher Scientific) coupled via an LC-Isolink interface with a Delta V Advantage IRMS 174 (Thermo Fisher Scientific). Chromatography was accomplished using a Sentry guard column (3 µm, 175 20 mm) and an Atlantis T3 column (3 μ m, 100 mm, Waters) at a flow rate of 500 μ L/min. Phosphoric acid at pH 2 was chosen as mobile phase. The method was run isocratically at room temperature. The analytes 176 177 were converted by wet oxidation at a temperature of 99.9 °C after the separation unit. Thereto, 90 g/L 178 Na₂S₂O₈ and phosphoric acid (1.5 M H₃PO₄) were introduced at a flow rate of 30 μ L/min. The vacuum 179 inside the IRMS was 2×10⁻⁶ mbar. Its ion source was set to an accelerating voltage of 3 kV and an electron 180 ionization energy of 124 eV. The isotope ratios were calibrated using our laboratory monitoring gas (CO₂), 181 which had previously been calibrated against the international standard RM8563 (CO₂), supplied by the 182 IAEA.

183 Derivatization of DPC for Nitrogen Isotope Analysis. Nitrogen isotope analysis was conducted using the derivatization procedure proposed by Melsbach et al.²⁷. Briefly, DPC was methylated to MDPC by adding 184 an excess of greater than 160 $n_{analyte}/n_{TMSD}$ (140 μ L of a 2 M TMSD solution) into a vial containing a 185 186 standard or a SPE extract reconstituted in 1 mL methanol. The vial was crimped tightly before putting it 187 into a 70°C water bath for 2 h. For samples from lysimeters with CLZ depth injection, the volume of the 188 2 M TMSD solution added to the reconstituted SPE extracts was increased to 200 µL to ensure complete 189 derivatization, as concentrations of DPC were up to an order of a magnitude higher compared to the other 190 lysimeter samples. Afterwards, the solvent was evaporated to dryness. The sample was then reconstituted 191 in 50 µL acetone.

192 Separation of Drainage Sample Fractions for Analysis of DPC and MDPC. For drainage water samples 193 from the lysimeters where CLZ was applied on the surface and for which the ratio of DPC to naturally 194 formed MDPC was greater than 10 %, preparative HPLC was used prior to derivatization to isolate this 195 naturally formed MDPC and thus to avoid interferences in the isotopic signature of DPC when subjected 196 to methylation in the derivatization procedure. The method is briefly summarized in the Supporting Information (II.7.)²⁷. Additionally, both DPC and MDPC fractions were used for δ^{15} N isotope analysis when 197 possible. For samples with an MDPC to DPC ratio <10 %, no preparative HPLC method was applied prior 198 199 to derivatization, as the influence of the isotope ratio of MDPC on the isotope ratio of derivatized DPC is 200 negligible and lies within the measurement error for nitrogen CSIA (±1 ‰) of the developed ¹⁵N GC-IRMS method²⁷. 201

Nitrogen Isotope Analysis of DPC and MDPC. The method is described by Melsbach et al.²⁷ Briefly, a TRACE GC Ultra gas chromatograph (Thermo Fisher Scientific, Milan, Italy) coupled with a Finnigan MAT 253 IRMS (Thermo Fisher Scentific, Bremen, Germany) was used. A Finnigan Combustion III interface (Thermo Fisher Scientific) connected both instruments. The analytes were combusted at a temperature of 1030 °C with a NiO tube/CuO-NiO reactor (Thermo Fisher Scientific). The gas chromatograph contained a DB-1701 column (30 m × 0.25 mm × 1 µm, J&W Scientific, Santa Clara, CA). Helium (grade 5.0) at a flow rate of 1.4 mL/min was used as carrier gas. Injection was carried out with a GC Pal autosampler (CTC, Zwingen, Switzerland). A sample volume ranging between 1 and 3 μL was injected into a splitless liner
(Thermo Fischer Scientific, Australia) at a temperature of 250 °C. The GC oven was programmed to start
at a temperature of 100 °C (held for 1 min), ramped with 25 °C/min to 240 °C, and with 10 °C/min to 280 °C
(held for 5 min). The isotope ratios were calibrated using our laboratory monitoring gas (N₂), which had
previously been calibrated against the international standard NSVEC (N₂), supplied by the IAEA.

214 Correction Procedure for Isotope Analysis. Analogous to the correction procedure described by Melsbach 215 et al.²⁷, all samples and standards were measured in triplicate and their isotope ratios are reported as the 216 arithmetic means with their respective estimated standard deviations ($\pm \sigma$). In addition to the calibration 217 of the measurement gas, samples are bracketed within the sequences by in-house standards of DPC and 218 MDPC, whose isotopic signature had been determined with EA-IRMS (Table S4). Here, the principle of 219 identical treatment by Werner and Brand⁴¹ was applied to correct for trueness by identifying drifts and off-sets, caused by different combustion efficiency. $\delta^{15}N$ correction was performed using MDPC 220 221 synthesized by LGC Standards GmbH, while an authentic DPC standard was used for δ^{13} C correction of the 222 LC-IRMS method.

223 **Concentration Measurement of CLZ and DPC from Soil Samples.** CLZ and DPC residues were measured 224 within the first soil layers (0 to 10 cm) approximately one year after herbicide/metabolite application. To 225 obtain a representative and homogenous sample, subsamples for soil analysis were collected in 226 quadruplets and combined afterwards. The total amount was at least 100 g soil per sample. Sample 227 extraction and analysis were carried out by Eurofins Sofia GmbH using LC-MS/MS.

228 Statistical Analyses. Pearson correlation analysis and one-way analyses of variance (ANOVA) tests were 229 performed to identify patterns and to measure the statistical significance of the relationship between 230 variables. ANOVA tests were performed to assess the differences between soil types and pesticide 231 application methods regarding total accumulated drainage, total DPC mass leached, maximum change of 232 carbon and nitrogen isotope signatures 900 days after pesticide application/injection. Separate Pearson 233 linear correlations were performed to evaluate the relationship between irrigation and drainage, between 234 soil humidity and drainage, between drainage and DPC mass leached, and between evapotranspiration 235 and DPC mass leached. All tests were performed using the statistical package Minitab 13.31 (Minitab Inc., 236 State College, PA). All statistical differences were set to the α = 0.05 significance level (p ≤ 0.05).

237 Results and Discussion

238 Water Dynamics. Total accumulated drainage 900 days after CLZ or DPC application/injection was 239 between 488 to 656 mm for gravel soil and between 337 and 502 mm for moraine soil. In relation to the 240 water input, drainage represented 25-39 % and 18-27 % of the total irrigation, respectively. Increased 241 drainage coincided with periods of high irrigation intensity and high soil water content. A significant 242 positive correlation (Pearson's correlation coefficient - r - from 0.30 to 0.49, p < 0.0001) between 243 intensity of daily irrigation and daily drainage was observed for the six lysimeters. As detailed by Torrentó et al.³⁷, who used the same lysimeters to assess the fate of the herbicide atrazine and its metabolites, soil 244 humidity data revealed that large irrigation events resulted in a greater contribution of preferential flow 245 to drainage, and that this effect was more significant for the moraine than for the gravel soil. A statistically 246 247 significant (p < 0.05) correlation was observed between soil humidity and drainage for both gravel and 248 moraine soil at all depths where capacitance sensors were installed (at 16, 36, 56, 76, and 96 cm for moraine soil and at 11, 51, and 71 cm for gravel soil)³⁷. This correlation was stronger for moraine (r 249 250 between 0.15 and 0.22, except for one depth with r = 0.08) than for gravel soil (r between 0.06 and 0.16), 251 and is in accordance with the fact that fluctuations in the soil water content were smaller for the latter, especially at greater depths³⁷. The total accumulated drainage after 900 days was influenced by the 252 application method (higher drainage for depth injection, p = 0.331) and by the soil type (higher for gravel 253 254 soil, p = 0.426). Large amounts of drainage from the gravel soil are probably a consequence of the higher water permeability and low water content at field capacity of this soil³⁷. 255

256 The average monthly and annual irrigation, drainage, and evapotranspiration values for the lysimeters 257 used in this study are shown in Table S5. Annual evapotranspiration, estimated by the water balance computation as explained by Torrentó et al.³⁷, was for the four years of study (2014 to 2017) higher for 258 259 moraine (315 to 633 mm) than for gravel soil (266 to 585 mm), although the effect was not statistically 260 significant (p = 0.718). A significant effect (p = 0.002) on annual evapotranspiration was however observed 261 for crop type: evapotranspiration was higher for sugar beet and corn than for broccoli, Chinese cabbage, 262 lettuce and leek. The effects of soil type and pesticide application method on evapotranspiration 900 days after pesticide application were not statistically significant (p = 0.093 and p = 0.579, respectively). The 263 264 influence of the cover vegetation on drainage and pesticides fate was not assessed, since no significant differences in the plants development were observed between lysimeters. For details, see Supporting 265 266 Information section III.2.

Trends in Compound Concentrations after DPC Surface Application. Neither CLZ, nor DPC had ever been
 applied to any of the lysimeters prior that study so that trends for CLZ and DPC concentrations could be

269 uniquely attributed to our experimental design. Through application of the metabolite DPC to the surface 270 of the lysimeters, it was possible to investigate the fate of DPC separately, in the absence of CLZ and 271 without interference of constantly formed DPC. The breakthrough of DPC in the seepage water differed 272 between the soil types (Figure 1b). In the lysimeter with moraine soil (L12), concentrations changed more 273 rapidly in relation with drainage events than for gravel soil (L1). For gravel soil (L1), DPC was detected in 274 the drainage water for the first time after 137 days, while it broke through only 15 days after application 275 in moraine soil (L12). In these lysimeters, a positive correlation was observed between drainage and DPC mass leached, being more significant for gravel (r = 0.36, p = 0.029) than for moraine soil (r = 0.31, p = 276 277 0.113). The observed dependency of the drainage response, and the analytes' concentration therein, on the irrigation agrees with Torrentó et al.³⁷ for the fate of the herbicide atrazine and its metabolites in 278 279 these lysimeters. Table S6 summarizes the observed breakthrough parameters for each lysimeter. Two 280 main DPC concentration peaks were detected in the drainage water of these two lysimeters after 281 approximately 550 and 850 days (Figure 1e). They coincided with two intense irrigation events (November 282 2016 and September 2017, Table S2). In moraine soil (L12, 303 and 441 mm), less accumulated drainage 283 had occurred at peak concentration of DPC than in the gravel soil (L1, 458 and 852 mm). Concentrations 284 in the gravel soil were approximately one order of magnitude higher than in moraine soil. In contrast to 285 our previous study³⁷, no rapid breakthrough peak was observed shortly after application, neither for DPC 286 nor for uranine (Figure S4). Bromide mass recovery curves (Figure S5) showed an asymmetric sigmoidal 287 shape, which is characteristic for transport through a porous matrix with some retardation. Smoother 288 trends for DPC compared to the tracers indicate retardation by sorption and/or attenuation by 289 degradation. DPC leaching was therefore mainly driven by porous matrix flow, although intense irrigation 290 events resulted in a greater contribution of preferential flow. This was observed mainly in moraine soil. 291 For example, after 425 and 670 days, sharp increases in DPC concentrations were measured (Figure S4). 292 This might be a consequence of transport by preferential flow induced by intense irrigation events (July 293 2016 and March 2018, respectively, Table S2).

The transformation product of DPC, MDPC, was first detected after 256 days and 425 days for gravel and moraine soil, respectively. At the end of the monitoring period (950 days after DPC application), 6.0 % of the DPC mass was recovered in the drainage water of the gravel soil and only 0.3 % in the moraine soil (details about the calculation of analyte recovery can be found in the Supporting Information section II.9). MDPC accounted for 0.55 % and 0.06 % of the applied DPC, respectively. One year after application, a DPC residue of approximately 3 % and 7 % of the applied DPC was quantified within the first soil layers (0 to 10 cm) of gravel and moraine soil, respectively (Table S7). Thus, an incomplete mass balance was 301 observed. Here, possible explanations might be: (i) sorption of DPC to lower soil layers within the root zone, where further sampling was not possible without disturbing the lysimeter, (ii) the uptake and 302 metabolism of DPC by plants ^{42,43}, and (iii) the presence of DPC-fulvic acid complexes, as their functional 303 groups can bind DPC. This has been demonstrated by Gatzweiler⁴⁴, who conducted lysimeter experiments 304 with ¹⁴C-labelled CLZ. Using thin-layer chromatography and analyzing the radioactivity, Gatzweiler⁴⁴ 305 306 detected DPC in fulvic acid fractions verifying the existence of these DPC-fulvic acid complexes. 307 Nevertheless, the MDPC/DPC concentration ratio suggests that further DPC degradation to MDPC occurred in both soils, mainly after 425 days. (As both DPC and MDPC have a similar GUS leaching 308 potential and show only minor differences in their mobility, no major retardation effect on the transport 309 310 of either compound is expected so that the use of metabolite-to-parent compound ratios appears justified in this case)⁹. This further degradation agrees with the findings of Schuhmann et al.⁴² and the 311 environmental degradation pathway predicted by Roberts et al.¹⁹. This demonstrates that transformation 312 313 of DPC is occurring only slowly. For the moraine soil, a local maximum for the MDPC/DPC concentration 314 ratio was reached after 750 days (Figure 1, L12e). To obtain additional insight into DPC transformation, 315 we, therefore, evaluated the results from CSIA of the lysimeter experiment.

Insights into DPC Transformation by Isotope Analysis of DPC from Surface Application. Initially, the δ^{13} C 316 317 and $\delta^{15}N$ values of the leached DPC were close to the original isotope signature of the applied DPC 318 (Figures 1, L1d and L12d). Over the course of the observation period carbon isotope signatures of DPC showed significant enrichment in ${}^{13}C$ ($\Delta\delta^{13}C_{DPC}$) of approximately +4 ‰ in both soil types. The heavy 319 320 irrigation event 672 days after DPC surface application (March 2017, Table S2) caused a new small DPC 321 breakthrough peak, in which DPC isotope values returned to the original isotopic composition, most likely 322 because new DPC was mobilized, which had not yet been subject to transformation. This effect was more 323 significant in moraine soil, where a greater contribution of preferential flow in response to this heavy 324 irrigation event was observed, resulting in a recovery of up to 20 % of the total mass of DPC leached in 325 the drainage water after the monitoring period. Additionally, significant changes of nitrogen isotope signatures ($\Delta \delta^{15} N_{DPC}$) of +2 ‰ to +3 ‰ were observed – however, mainly in the gravel soil (L1). 326 Furthermore, these shifts were observed at a later time point than the enrichment in ¹³C, approximately 327 450 days after application. The fact that during the first 450 days DPC was only becoming enriched in 13 C, 328 329 and then in both ¹³C and ¹⁵N, suggests that DPC was transformed by two distinct processes and that only 330 the latter one starting after 450 days involved a reaction of a nitrogen atom. The transition between the 331 two trends coincides with an increase in the MDPC/DPC concentration ratio (Figure 1, L1e). As there had 332 never been any application of CLZ or DPC to these lysimeters, the carbon and nitrogen isotope values of 333 DPC can be uniquely attributed to the substance applied in this study, and changes in these isotope signatures are attributable to its further degradation. Interestingly, due to the high concentrations of 334 MDPC in the drainage water, it was possible to measure the $\delta^{15}N$ of formed MDPC after purification by 335 preparative HPLC (Tables S9 and S10). In both lysimeters, the $\delta^{15}N$ of MDPC was significantly more 336 negative (approximately by 4‰) compared to the δ^{15} N value of the DPC at that time (Figure 1d). Since 337 338 DPC contains three nitrogen atoms out of which only one is methylated, it can be estimated that the methylation of DPC causes a nitrogen isotope effect of approximately 3×4 ‰ = +12 ‰ at the reactive 339 atom. Our data for the DPC surface application show an enrichment in ¹³C and, to a lesser extent, in ¹⁵N 340 for DPC in both soils, which was significantly masked in the moraine soil due to the leaching of fresh DPC 341 342 after heavy irrigation events. Transformation extent can thus be underestimated. Here, transformation of 343 DPC may be easier to detect using the metabolite-to-parent concentration ratio, at least for the pathway 344 involving MDPC formation. On the other hand, using the metabolite-to-parent concentration ratio only to 345 investigate the transformation of DPC, the evidence of an additional transformation mechanism would 346 have remained undetected. Additionally, CSIA appears to be more robust as the integrated isotope signal, which indicates degradation remains measurable, even if the metabolite might be subject to sorption or 347 348 further transformation.



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350 Figure 1. Lysimeters with DPC application on surface (a single application in May 2015): L1 in gravel soil (left panels) and L12 351 in moraine soil (right panels). a) Daily irrigation (black bars) and cumulative drainage (grey line); b)-c) Concentration of DPC 352 (blue diamonds) and MDPC (black triangles), note that different scales are used for both soil types; d) Carbon (black diamonds) 353 and nitrogen (red diamonds) isotope ratios of DPC and nitrogen isotope values of MDPC (petrol triangles), error bars show the 354 associated uncertainties (±0.5 ‰ for carbon, ±1.0 ‰ for nitrogen isotope analysis; or when exceeding this uncertainty, 355 standard deviations of triplicate measurements are given , EA isotope values of the applied DPC are shown as lines, whereas 356 associated uncertainties (±0.5 ‰ for carbon, ±1.0 ‰ for nitrogen isotope analysis) are shown as dashed lines in the 357 corresponding color, respectively; e) metabolite-to-parent compound molar ratio of MDPC/DPC (black diamonds); f) season 358 corresponding to the time since application - spring (green horizontal lines), summer (red vertical lines), autumn (yellow dots), 359 winter (blue diagonal lines); the grey dashed lines repeated in each sub-figure represent the start of a new year.

360

361 CLZ Surface Application Mimicking A Realistic Field Scenario. For the surface application of CLZ (Figure 2,
 362 L4 and L8), the metabolites DPC and MDPC were detected in the seepage water 425 days after CLZ
 363 application, coinciding with a heavy irrigation event (July 2016, Table S2), while the applied parent

364 compound remained below or close to the limit of detection of $0.05 \,\mu g/L$ during the time of monitoring 365 (970 days). Analytes breakthrough curves and concentrations differed between the soil types. For uranine, 366 a rapid breakthrough shortly after application was detected in moraine soil (Figure S4). During the 367 monitoring period, the maximum uranine concentration was measured within the first day, after only 4 368 mm of accumulated drainage (Table S6), suggesting that it was mainly transported through preferential 369 flow, bypassing large fractions of the soil matrix. Furthermore, a pronounced uranine peak tailing was 370 observed, which is typical for preferential flow (Figure S4). Furthermore, the DPC mass recovery curves 371 were significantly different for the two soils (Figure S5), giving further evidence of a greater contribution 372 of transport through preferential flow for moraine soil. This difference in soil type agrees with the results 373 of the lysimeters with surface application of DPC as well as well as with the findings for other compounds 374 described in Torrentó et al.³⁷.

375 Approximately 0.5 % and 0.13 % of the applied CLZ was leached as DPC after 950 days in gravel and 376 moraine soil, respectively. When analyzing the CLZ and DPC content in the upper soil, for none of the 377 lysimeters a closed mass balance was obtained. While no CLZ was detected in the first soil layer (0 to 378 10 cm) approximately 1 year after CLZ application to the lysimeter surface (consistent with Pestemer & 379 Malkomes⁴⁵), DPC amounts corresponding to 5 to 9 % of the applied CLZ were found (Table S7). CLZ and DPC are expected to be incorporated into maize plants based on the findings of Schuhmann et al.⁴² and 380 381 Stephenson & Ries⁴³. In addition, Barra et al.⁴⁶ showed that during the first 90 days after CLZ application, 382 CLZ dissipation was mainly due to volatilization and degradation, whereas later on, when CLZ was already 383 in the subsurface, its disappearance from soil occurred mainly due to degradation. Higher DPC/CLZ 384 concentration values were measured in the drainage water of the gravel soil compared to moraine soil. 385 These results suggest that either DPC leached more rapidly through the soil matrix in the gravel soil 386 because of higher permeability. Or, alternatively, the extent of CLZ degradation was higher for the gravel 387 soil compared to moraine soil, as there is a greater contribution of preferential flow in moraine soil, which 388 bypasses the top layer where degradation is mostly expected to take place. When preferential flow occurs, 389 pesticides bypass large fractions of the soil matrix, reducing the degradation and sorption potential, as 390 the topsoil is microbiologically more active and with higher organic matter content. CSIA results provide 391 additional insights about these two hypotheses (see below). Concentration ratios and isotope results 392 point to a higher extent of CLZ degradation in gravel than in moraine soil. Nevertheless, some metabolite-393 to-parent ratio values may be underestimated, because CLZ was below the limit of detection and, 394 therefore, CLZ concentrations corresponding to the detection limit were chosen for calculation, resulting 395 in a minimum estimated ratio in that case.

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397 Figure 2. Lysimeters with CLZ application on surface (a single application in May 2015), L4 (left panels) and L8 (right panels). a) 398 Daily irrigation (black bars) and cumulative drainage (grey line), b)-d) Concentration of CLZ (green circles), DPC (blue diamonds) 399 and MDPC (black triangles) over time, e) metabolite-to-parent compound molar ratio of DPC/CLZ (black hexagon), f) carbon 400 (black diamonds) and nitrogen (red diamonds) isotope ratios of DPC, error bars show the associated uncertainties (±0.5 ‰ for 401 carbon, ±1.0 ‰ for nitrogen isotope analysis; or when exceeding this uncertainty, standard deviations of triplicate 402 measurements are given, EA isotope values of the applied CLZ are shown as lines, whereas associated uncertainties (±0.5 ‰ 403 for carbon, ±1.0 ‰ for nitrogen isotope analysis) are shown as dashed lines in the corresponding color, respectively; g) 404 metabolite-to-parent compound molar ratio of MDPC/DPC (black diamonds), h) season corresponding to the time since 405 application - spring (green horizontal lines), summer (red vertical lines), autumn (yellow dots), winter (blue diagonal lines); 406 the grey dashed lines repeated in each sub-figure represent the start of a new year.

407 In lysimeters with CLZ surface application (Figure 2, L4f and L8f), some carbon isotope values of DPC show a shift to more negative δ^{13} C values compared to the carbon isotope signature of the applied CLZ (Table 408 409 S4). This behavior is observable in both lysimeters, especially after heavy rain events such as that one performed 550 days after CLZ application (November 2016, Table S2), which resulted in a depletion in ¹³C 410 411 by 3.4 ‰ for gravel soil (L4). This shift may be attributed to the mobilization of freshly formed DPC, which 412 is formed from CLZ by loss of the aromatic moiety through C–N bond cleavage. Presuming that the phenylring contains more ¹³C atoms than the average molecule (Figure S1), which may have been introduced by 413 414 the synthesis process, this would result on a ¹³C-depletion. Alternatively, the shift may be due to 415 secondary normal carbon isotope effects. Once transformation of DPC starts – as evidenced by the 416 detection of MDPC – this ¹³C-depletion may be masked compared to the carbon isotope composition of the applied CLZ, as an enrichment in ¹³C in DPC is expected. Consistently, observed $\delta^{13}C_{DPC}$ values are close 417 418 to or higher than the EA-IRMS value of the applied CLZ.

419 In moraine soil (L8), no evidence of DPC degradation was obtained based on carbon isotope values, as changes of δ^{13} C values were within the uncertainty of the method (Figure 2, L8f). In contrast, carbon 420 isotope values of DPC in gravel soil (L4) showed an enrichment in ¹³C by up to +8.4 ‰ (Figure 2, L4f) 421 422 indicating that DPC was further transformed. At a subsequent time point (930 days after application), 423 however, the $\delta^{13}C_{DPC}$ value changed back close to the original isotopic signature detected at the beginning 424 of monitoring. This indicates that the change in δ^{13} C DPC values was "diluted" by the input of newly 425 mobilized DPC, as supported by a concomitant increase of the DPC/CLZ concentration ratio (Figure 2, L4e). 426 Hence, the two lines of evidence (isotope and DPC/CLZ concentration ratios) were found to complement 427 each other in the assessment of DPC degradation - when one line of evidence was about to fail, the other 428 was able to provide conclusive evidence.

The more substantial changes in both $\delta^{13}C_{DPC}$ values and DPC/CLZ concentration ratios indicate that DPC 429 430 degradation was higher in L4 (gravel soil) than in L8 (moraine soil) leading to the hypothesis that 431 differences in the transformation rate of CLZ to DPC existed. This is supported by the findings of Capri et al.⁴⁷, who reported that the extent of CLZ degradation is influenced by the moisture content of the soil. 432 As described by Torrentó et al.³⁷, there is a higher soil water content and less fluctuation of the water 433 434 content in the gravel soil than in moraine soil. On the other hand, for both moraine and gravel soil, $\delta^{15}N$ 435 values of the DPC formed are, as hypothesized in Figure S1, close to the nitrogen isotope signature of the 436 applied CLZ. Based on the findings of Lingens et al.²⁰, the pyridazinone-ring of the CLZ molecule is not 437 involved in the first transformation steps (dioxygenation of the phenyl-ring, Figure S1) so that no

438 significant nitrogen isotope fractionation is expected during CLZ transformation to DPC ^{27, 29, 48, 49}. As the 439 isotope effect during multi-step reactions is reflected by the rate-limiting steps, our results indicate that 440 the amidase-driven cleavage of the moiety (2-hydroxymuconate) at the C–N bond, may be not rate-441 limiting. As a result, changes in nitrogen isotope values of DPC can be uniquely attributed to its further 442 degradation.

443

444 Transformation-Potential after Herbicide Injection Below the Root Zone. Finally, two lysimeters (L6 and 445 L7) were chosen to simulate the preferential flow after a heavy irrigation event by injecting CLZ into a depth of 40 cm, following the approach described by Torrentó et al.³⁷. In contrast to surface application 446 447 observations, CLZ and DPC broke through a few days after CLZ was injected (Figure S6). The second 448 metabolite MDPC was detected in the drainage water after 130 days. The detection of the metabolites 449 indicated that CLZ degradation occurred, even when it was injected below the root zone. Additionally, 450 significantly greater concentrations of CLZ, DPC and MDPC (1 to 2 orders of magnitude higher) were 451 measured in the drainage water of the lysimeter with CLZ depth injection compared to the CLZ surface 452 applications. In contrast to surface application observations, early breakthrough of injected uranine and 453 CLZ occurred for the two soil types within a few days (< 11 days for gravel and 6 hours for moraine soil) 454 and after a small amount of accumulated drainage (< 55 mm and 8 mm, respectively). This rapid response 455 and the peak tailing for both solutes are typical for preferential flow. More than 80 % of the total uranine 456 recovered mass was received during this early breakthrough. These results confirm that preferential flow 457 was enhanced by depth injection. In agreement with Torrentó et al.³⁷, the response to intense irrigation 458 events was more significant than for surface applications. It results in several fluctuations of CLZ and DPC 459 concentrations in the drainage water during the first 370 days for both soils (Figure S4). A great increase 460 in CLZ and DPC concentrations occurred in both lysimeters after 330-345 days (at 225-320 mm of 461 accumulated drainage), coinciding with the heavy irrigation events in May 2015 (Table S2). After this 462 pulse, no CLZ was recovered, while a steady increment in accumulated mass recovery was observed for 463 DPC for both soils (Figure S5).

At the end of the monitoring period (1250 days after CLZ injection), total leached analytes accounted for 24 and 22 % of the injected CLZ mass, respectively. Even though comparison between the two application methods may be limited (eleven uniformly distributed CLZ injections versus broad surface application), higher recoveries were obtained for CLZ injection after the same time of monitoring (950 days): from 2.0 to 3.4 % of CLZ, between 16.4 and 17.2 % of DPC and from 0.2 to 0.4 % of MDPC compared to no CLZ 469 leaching, 0.13 to 0.15 % of DCP and below 0.02 % of MDPC with surface application. As the mass balance 470 remains incomplete for CLZ injection, there is evidence that additional processes occurred. With surface application, processes such as volatilization⁴⁶, additional transformation pathways¹⁹ and uptake by 471 472 plants⁴² likely accounted for the mass losses. Additional influences on the low recovery, which might also occur after CLZ depth injection, might be the low mobility for CLZ⁵⁰ and the formation of putative fulvic 473 acid complexes of DPC⁴⁴. The DPC/CLZ concentration ratio in these lysimeters with CLZ depth injection 474 475 shows that the main fraction of DPC seems not to be involved in sorption as this concentration ratio has 476 a single global maximum starting approximately 600 days after CLZ injection (Figure S6). This global 477 concentration maximum is two orders of magnitude greater than DPC/CLZ concentration ratios observed 478 for CLZ surface application. It shows the importance of the topsoil to retain DPC. As indicated by the 479 MDPC/DPC concentration ratio, further transformation of DPC occurred, although its extent and nature 480 is unknown.

ANOVA tests were performed to assess the differences between the two soil types and the CLZ application method (i.e. surface application vs. depth injection) regarding DPC leaching and its carbon and nitrogen isotope fractionation. The results showed that the DPC mass leached after 900 days was significantly influenced by the CLZ application method (p < 0.0001). A 90- to 260-fold increase in DPC leaching was observed for depth injection compared to surface application. Although the effect of soil type was not statistically significant (p = 0.998), CLZ surface application resulted in higher DPC mass leached for gravel than for moraine soil.

488 Similar to observations in lysimeters with CLZ surface application, carbon isotope data of DPC show an 489 enrichment in δ^{13} C of 3.8 ‰ after 648 days of herbicide injection below the root zone in the gravel soil, 490 while no significant change is observed in moraine soil (Figure S6). There, up to 648 days, no significant changes in ${}^{13}C/{}^{12}C$ and ${}^{15}N/{}^{14}N$ ratios were measured. The $\delta^{15}N$ value of DPC shows the initial isotope 491 composition of the CLZ applied to the lysimeter. In very few cases, it was possible to measure δ^{15} N values 492 493 of MDPC formed from DPC (Figure S6 and Table S11). Nitrogen isotope values of MDPC were by approximately 6 % more negative than δ^{15} N signature of its parent compound DPC. This shift agrees with 494 495 the findings in DPC transformation experiments (Figure 1, L1 and L12) and, thus, supports the nitrogen 496 isotope effect of DPC methylation of approximately +12 ‰ as estimated above (Table S9 and S10). 497 According to the ANOVA results, isotope fractionation was mainly influenced by the soil type (higher ¹³C 498 and ¹⁵N enrichment for gravel soil) rather than by the CLZ application method.

499

500 Dual-Element Isotope Plot to Identify DPC Formation and Transformation. A dual-element plot was used 501 for an overview of observed trends in carbon and nitrogen isotope signatures of DPC (i) either from 502 formation from CLZ, or (ii) when DPC was further transformed. In Figure 3a, isotope data of all lysimeters 503 with DPC surface application are combined, whereas in Figure 3b, data of all lysimeters with CLZ 504 application/injection are shown. In Figure 3a, where DPC represents the original applied compound, a 505 general trend towards more positive δ^{15} N and δ^{13} C values is observable. This observation is consistent 506 with the well-established phenomenon that, in most cases, heavy isotopes become enriched in the 507 remaining substrate during (bio)degradation. As detailed above, DPC in first drainage samples (first 450 days) of the gravel soil showed a significant enrichment in ¹³C but not in ¹⁵N, indicating that two 508 509 distinct processes for DPC transformation occurred. In contrast, Figure 3b shows two opposing trends 510 pointing to the occurrence of both DPC formation and transformation. On the one hand, similar to the lysimeters with DPC application, a trend is observed towards more positive δ^{13} C and δ^{15} N values during 511 the transformation of DPC. On the other hand, numerous data points show more negative δ^{13} C and δ^{15} N 512 513 isotope values. As this trend is only observable for lysimeters with CLZ application and injection, we 514 attribute it to the formation of DPC. As discussed above, possible explanations for the observed depletion in ¹³C (more negative δ^{13} C values) is (i) an artefact of an uneven ¹³C isotope distribution in the cleaved 515 516 phenyl-ring during DPC formation; or (ii) that the formation of DPC from CLZ (Figure S1) may be 517 accompanied by a small and normal secondary carbon isotope effect.



519Figure 3. Dual-element isotope plot of a) DPC degradation in lysimeters L1 and L12, where DPC was applied on the surface, and520b) formation and degradation of DPC in the lysimeters where CLZ was either applied (L4 and L8) or injected in a depth of 40 cm521(L6 and L7); the red circles represent the isotopic signature of the applied/injected a) DPC and b) CLZ – position of CLZ and DPC522differ within the dual-element isotope plots due to their different isotopic source signatures (Table S4).

523

524 Environmental Significance and Outlook

525 The isotope fractionation in DPC observed for the three tested scenarios is particularly important because 526 (i) the change in carbon and nitrogen isotopic signature of DPC evidenced transformation of an apparently 527 persistent metabolite, and (ii) these changes provide evidence that likely more than one transformation 528 pathway is involved in DPC transformation. In soil, only methylation of DPC to MDPC is known and thus 529 our data suggest the need for further laboratory experiments and mechanistic studies on DPC 530 (bio)degradation to gain further insight into possible additional transformation pathways. (iii) Formed 531 DPC, which had not been subject to further transformation yet, showed the same nitrogen isotope signature as its precursor CLZ. Hence, δ^{15} N values may serve as isotopic fingerprints to identify the origin 532 533 of such compounds in groundwater.

534 When applying CSIA, the combination with conventional methods was found to be complementary and 535 advantageous, especially when formation and transformation of the metabolite was occurring 536 simultaneously. Once introduction of newly formed metabolite dominated, evidence from CSIA was not 537 necessarily conclusive because transformation-related changes in isotope ratios were masked by the 538 continuous input of DPC. Here, additional information was gained by metabolite-to-parent concentration 539 ratios, which became greatest and could provide evidence of DPC formation. Vice versa, when metabolite-540 to-parent-ratios were small because DPC was further transformed, it was the changes in isotope ratios of 541 DPC which still carried the isotopic imprint of the reaction and, hence, made transformation visible. For further understanding of the environmental fate of DPC, reference experiments focusing on the 542 543 determination of stable isotope fractionation factors as well as microbial processes during DPC 544 transformation are required in order to identify transformation mechanisms and quantify them.

545 For the future, our approach with CSIA in combination with concentration measurements and systematic 546 long-term lysimeter experiments holds promise to answer questions about transformation pathways and 547 the extent of soil / vadose zone (bio)transformation not only for DPC - one of the most widely detected 548 substances – in groundwater, but also for other micropollutants of concern and their metabolites. 549 Additionally, this study confirmed that the application of CSIA in combination with solid-phase 550 extraction^{27, 39} is feasible for the analysis of polar micropollutants in drainage water at environmentally 551 relevant concentrations. Thus, it can be also applied to studies in agricultural soil and groundwater from 552 common unconsolidated sand and/or gravel aquifers with catchment areas within agricultural production.

553 ASSOCIATED CONTENT

554	* Supporting Information
555	Further details on chemicals, lysimeters, analytical methods, analyte recovery calculations and results of
556	water balance computation, vegetation cover evolution, breakthrough trends, EA-IRMS analyses, soil
557	analyses, CLZ depth injection and nitrogen isotope ratios.
558	
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575	
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578	
579	ABBREVIATIONS
580	CLZ - 5-Amino-4-chloro-2-phenyl-2 <i>H</i> -pyridazin-3-one
581	DPC - 5-Amino-4-chloro-3(2H)-pyridazinone

582 583 MDPC - 5-Amino-4-chloro-2-methyl-3(2H)-pyridazinone

584 **References**

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