| 1 | |
|----|---|
| 2 | |
| 3 | |
| 4 | |
| 5 | This document is the unedited Author's version of a Submitted Work that was |
| 6 | subsequently accepted for publication in Environmental Science and Technology, |
| 7 | copyright $\ensuremath{\mathbb{C}}$ American Chemical Society after peer review. To access the final edited |
| 8 | and published work see |
| 9 | |
| 10 | http://pubs.acs.org/articlesonrequest/AOR-jneyhAjJQIsFAmd7Bgp3 |
| 11 | |

| 13 | C and Cl isotope fractionation of 1,2-dichloroethane |
|----------------------|--|
| 14 | displays unique $\delta^{13}C/\delta^{37}Cl$ patterns for pathway |
| 15 | identification and reveals surprising C-Cl bond |
| 16 | involvement during microbial oxidation |
| 17 | |
| 18 | Jordi Palau, ^{†,*} Stefan Cretnik, [‡] Orfan Shouakar-Stash, [§] Martina Höche, [‡] Martin Elsner, [‡] Daniel |
| 19 | $Hunkeler^{\dagger}$ |
| 20 | |
| 21 | †Centre for Hydrogeology and Geothermics, University of Neuchâtel, Neuchâtel, Switzerland. |
| 22 | ‡Helmholtz Zentrum München, German Research Center for Environmental Health, Neuherberg, |
| 23 | Germany. |
| 24 | §Department of Earth and Environmental Sciences, University of Waterloo, Waterloo, Canada. |
| 25 | |
| 26 27 28 29 | |

31 TOC/Abstract art



36 ABSTRACT

37 This study investigated dual element isotope fractionation during aerobic biodegradation of 1,2dichloroethane (1,2-DCA) via oxidative cleavage of a C-H bond (Pseudomonas sp. Strain 38 39 DCA1) versus C-Cl bond cleavage by S_N2 reaction (Xanthobacter autrophicus GJ10 and 40 Ancylobacter aquaticus AD20). Compound-specific chlorine isotope analysis of 1,2-DCA was performed for the first time and isotope fractionation $\varepsilon_{\text{bulk}}^{\text{Cl}}$ was determined by measurements of 41 42 the same samples in three different laboratories using two GC-IRMS and one GC-quadrupole 43 MS. Strongly pathway-dependent slopes ($\Delta\delta^{13}C / \Delta\delta^{37}Cl$), 0.78±0.03 (Oxidation) and 7.7±0.2 44 $(S_N 2)$, delineate the potential of the dual isotope approach to identify 1,2-DCA degradation pathways in the field. In contrast to different $\varepsilon_{\text{bulk}}^{\text{C}}$ values: -3.5±0.1‰ (Oxidation), -31.9±0.7‰ 45 and -32.0±0.9‰ (S_N2), the obtained ϵ_{bulk}^{Cl} values were surprisingly similar for the two pathways: 46 47 -3.8±0.2‰ (Oxidation), -4.2±0.1‰ and -4.4±0.2‰ (S_N2). Apparent kinetic isotope effects of ¹³C-AKIE=1.0070±0.0002 ¹³C-AKIE=1.068±0.001 ³⁷Cl-48 (Oxidation), $(S_N 2)$ and 49 AKIE=1.0087 \pm 0.0002 (S_N2) fell within expected ranges. In contrast, an unexpectedly large 50 secondary ³⁷Cl-AKIE of 1.0038±0.0002 reveal a hitherto unrecognized involvement of C-Cl 51 bonds in microbial C-H bond oxidation. Our 2D isotope fractionation patterns enable for the first 52 time reliable 1,2-DCA degradation pathway identification in the field, which unlocks the full 53 potential of isotope applications for this important groundwater contaminant.

55 INTRODUCTION

Chlorinated ethanes are among the most widespread contaminants in groundwater¹ and 1,2-56 57 dichloroethane (1,2-DCA) has been found in 36% of 1,585 National Priorities List sites identified by the United States Environmental Protection Agency (2001).² The presence of 1,2-58 59 DCA - an intermediate in plastics production - in groundwater is mainly a consequence of industrial activity.² A number of laboratory³⁻⁹ and field^{10, 11} studies showed 1,2-DCA 60 biodegradation under aerobic³⁻⁶ and anaerobic^{4, 7-11} conditions via different reaction pathways.⁹ 61 Under aerobic conditions, 1,2-DCA can be degraded either via nucleophilic substitution $(S_N 2)^{5, 6}$ 62 63 or via oxidative cleavage of a C-H bond³ catalyzed by hydrolytic dehalogenase and 64 monooxygenase enzymes, respectively (Scheme 1). Initial products of both reactions, 2-65 chloroethanol (S_N2-reaction) and 1,2-dichloroethanol (Oxidation), are further degraded to 66 ubiquitous end products, which hampers a direct identification of degradation pathways from 67 metabolite analysis. Alternative approaches to detect, and identify 1,2-DCA transformation 68 pathways in the subsurface are therefore warranted. This is crucial information in environmental 69 field studies to assess 1,2-DCA natural attenuation.

70 Scheme 1



S_N2-reaction (X. autotrophicus GJ10 and A. aquaticus AD20)

71

$$\begin{array}{c} H \\ Nu \\ -C \\ -CI \\ -K_{2}CI \end{array} \rightarrow \left[\begin{array}{c} H \\ Nu \\ -C \\ -CI \\ -K_{2}CI \end{array} \right]^{+} \rightarrow \begin{array}{c} H \\ Nu \\ -C \\ -K_{2}CI \\ -K_{2}CI \end{array} \xrightarrow{KIE_{c} 1.03 - 1.09^{21}} KIE_{c} 1.03 - 1.09^{21} \\ KIE_{c} 1.006 - 1.009^{45} \end{array} \right]$$

Compound specific isotope analysis (CSIA) is an innovative tool to investigate degradation
pathways of organic contaminants.¹²⁻¹⁴ Isotope ratios of individual compounds, measured either
by gas chromatography isotope ratio mass spectrometry (GC-IRMS) or GC-quadrupole mass
spectrometry (GC-qMS), are reported using the delta notation:

77

78
$$\delta^{h}E_{sample} = \frac{R({}^{h}E/{}^{l}E)_{sample}}{R({}^{h}E/{}^{l}E)_{standard}} - 1$$
 (1)
79

80 where R is the isotope ratio of heavy (^hE) and light (^lE) isotopes of element E (e.g., ¹³C/¹²C and 81 ³⁷Cl/³⁵Cl). The isotope fractionation (ϵ) expresses by how much ^hE/^lE is smaller (negative 82 values) or larger (positive values) in the average of freshly formed products compared to the 83 substrate from which they are produced. Transformation-induced isotope fractionation is 84 generally larger than the one related to phase transfer processes such as sorption or 85 volatilization.¹⁵ Bulk (i.e. compound-average) ϵ values can be calculated using a modified form 86 of the Rayleigh distillation equation:

87

88
$$\ln \frac{R_t}{R_0} = \ln \left(\frac{\delta^{h} E_t + 1}{\delta^{h} E_0 + 1} \right) = \varepsilon_{\text{bulk}} \cdot \ln f$$
 (2)
89

90 where R_t and R_0 are the current and initial isotope ratios respectively, and f is the compound 91 remaining fraction.

Previous laboratory experiments¹⁶ showed that for 1,2-DCA different carbon ε_{bulk} values of -29.2 92 and -3.9% reflected different reaction pathways: hydrolytic dehalogenation (C-Cl bond cleavage 93 94 via S_N versus oxidation (C-H bond cleavage) respectively (Scheme 1). Knowledge of in situ 95 contaminant biodegradation reactions is essential to evaluate the fate and long term impact of 96 1,2-DCA in groundwater. For aerobic biodegradation of 1,2-DCA isotope data are particularly important as no characteristic products accumulate. However, while isotope fractionation of one 97 98 element alone can provide pathway distinction in laboratory experiments (where mass balances can be established and ε_{bulk} values be determined), this is not possible under field conditions. 99 100 Here, evidence from a second element and a dual isotope approach is necessary to distinguish 1,2-DCA degradation pathways.¹⁷ As observed experimentally,^{18, 19} for a given compound, 101 combined isotope analysis of two elements (e.g., δ^{13} C vs. δ^{37} Cl) during the course of a reaction 102 103 generally yields a linear trend in a dual element isotope plot with a slope characteristic of the reaction mechanism. The reason is that the dual element isotope slope ($\Lambda = \Delta \delta^{13}C / \Delta \delta^{37}CI$, 104 where $\Delta \delta^{13}$ C and $\Delta \delta^{37}$ Cl are changes in isotope ratios during degradation) reflects isotope 105 fractionation of both elements.¹² Therefore, different slopes are expected from distinct reaction 106 pathways involving different bonds with different elements. 107

108 Currently, there is increasing interest in dual element isotope analysis for improved 109 differentiation of transformation mechanisms and several authors pointed out that 110 complementary mechanistic insight for 1,2-DCA aerobic biodegradation reactions could be 111 achieved by the additional analysis of chlorine¹² and/or hydrogen^{16, 20} isotope ratios. The reason 112 is that isotope fractionation can be traced back to underlying kinetic isotope effects, which are 113 highly reaction-specific.²¹ During enzymatic degradation molecules containing the light isotope 114 at the reactive site (e.g., ³⁵Cl) typically exhibit a higher reaction rate (e.g., ³⁵k) than those with a heavy isotope (e.g., 37 Cl) resulting in a kinetic isotope effect of KIE = 35 k / 37 k.²² When a C-Cl bond is broken a (primary) chlorine leaving group KIE would be expected, whereas in the oxidation reaction chlorine atoms sit next to the reacting bond so that only a secondary KIE would be expected (Scheme 1). Secondary isotope effects at positions next to the reacting bond are generally much smaller than primary isotope effects.¹²

120 Until recently, Cl-CSIA of chlorinated aliphatic compounds was not feasible, however, because a 121 direct method that would produce a suitable Cl-containing measurement gas inside a 122 chromatographic separation gas was lacking. However, new analytical methods were developed 123 based on the measurement of selected isotopologue ions or isotopologue ion fragments in unconverted analyte molecules using both continuous flow GC-IRMS²³ and GC-qMS.²⁴⁻²⁶ In 124 125 addition, a theoretical framework provided the theoretical justification for such evaluation of isotope fractionation from ion-current ratios of molecular and fragment-ion multiplets.²⁷ A recent 126 interlaboratory study took the next step and demonstrated that comparable $\delta^{37}Cl$ values were 127 128 obtained when analyzing a set of pure trichlorethene (TCE) standards on eight different instruments.²⁸ Since the technology is so new, however, a comparative study would also be 129 130 desired which shows that comparable ε_{Cl} values are obtained when analyzing degradation 131 samples on different instruments and in different laboratories. No such study has been conducted 132 yet. Most notably, Cl-CSIA studies have so far been applied to only few compounds, because two isotopically distinct compound-specific standards are necessary for every new substance.²⁸ 133 This has restricted applications primarily to chlorinated ethylenes^{18, 19, 29-34} so that - to our 134 135 knowledge - dual element isotope data are currently non-existent for chlorinated ethanes.

This study, therefore, aimed (i) to establish for the first time dual element (C & Cl) isotope analysis of the chlorinated ethane 1,2-DCA; (ii) to perform the Cl isotopic analysis in three different laboratories (i.e. Waterloo, München and Neuchâtel), using two different GC-IRMS and one GC-qMS, in order to investigate the consistency of ε_{C1} values obtained with different instruments and analytical methods; (iii) to investigate carbon and chlorine isotope fractionation during aerobic biodegradation of 1,2-DCA with three pure strains to determine whether the dual isotope slopes are sufficiently different to potentially distinguish between hydrolytic dehalogenation (S_N2) and oxidation (C-H bond cleavage) in the field.

144

145 MATERIALS AND METHODS

146 **Pure cultures preparation**

Three pure strains with known initial biotransformation mechanisms were used for the batch experiments: *Pseudomonas sp.* Strain DCA1 (Oxidation), *Xanthobacter autrophicus* GJ10 and *Ancylobacter aquaticus* AD20 (S_N2 reaction). *X. autrophicus* GJ10 (DSMZ 3874) and *A. aquaticus* AD20 (DSMZ 9000) were purchased (DSMZ, Braunschweig, Germany) and *Pseudomonas sp.* Strain DCA1 was kindly provided by Dr. Elizabeth Edwards (Department of Chemical Engineering and Applied Chemistry, University of Toronto, Canada). The growth medium was prepared as described by Hunkeler and Aravena (2000).³⁵

154 Cultures and experiments were prepared in 250 mL glass bottles, which contained 185 mL of 155 medium and were capped with Mininert-Valves (Vici Precision Sampling, Baton Rouge, LA, 156 US). Cultures were amended with 1,2-DCA (Fluka, ≥ 99.5% purity) and incubated in the dark at 157 room temperature and under continuous shaking (100 rpm). Headspace 1,2-DCA concentrations 158 were monitored throughout the incubation period. Before starting the experiments the cultures were transferred three times. Each subculture was spiked with pure 1,2-DCA four times before a 15 mL aliquot was transferred to 170 mL of autoclaved fresh medium. The spike volume was 9 μ L of pure 1,2-DCA for the first and second subcultures and 22.5 μ L for the third, leading to aqueous phase concentrations of 0.6 and 1.5 mM, respectively.

163 Experiment sampling

164 All experiments were conducted in triplicate. Experiments and controls were amended with 22.5 165 μ L of pure 1,2-DCA, corresponding to the chlorine isotopic working standard (δ^{37} Cl₀-CHYN2 = 166 $+0.8 \pm 0.1\%$), to produce an initial aqueous concentration of 1.5 mM. Bottles were shaken 167 upside down to prevent leakage of gas phase through the valve. After 1h of shaking, samples representing the initial concentration were collected. For concentration and isotopic analysis, 168 aqueous samples (1.5 mL) were taken at selected time points and preserved frozen³⁶ in 2 mL 169 170 vials with NaN₃ (1 g/L). Two abiotic control bottles were prepared with 185 mL of autoclaved 171 mineral medium and samples were collected and preserved like in the experiments. For each 172 culture experiment, sample triplicates were shipped frozen to the University of Waterloo 173 (Canada) and to the Helmholtz Zentrum München (Germany) for chlorine isotope measurements.

174 Isotopic and concentration analysis

Five pure 1,2-DCA isotopic working standards, one for carbon and four for chlorine, were used for instrument monitoring and external calibration of sample raw δ^{37} Cl values to the international Standard Mean Ocean Chloride (SMOC) scale. The isotopic signature of the carbon standard ($\delta^{13}C_{V-PDB} = -29.47 \pm 0.05\%$) was determined beforehand using an elemental analyzer coupled to an IRMS. Regarding the chlorine standards, CHYN1 and CHYN2 ($\delta^{37}Cl_{SMOC} = +6.30 \pm$ 180 0.06‰ and +0.84 \pm 0.14‰, respectively) were characterized relative to SMOC in München by 181 IRMS after conversion of 1,2-DCA to methyl chloride according to Holt et al. (1997).³⁷ IT2-182 3001 and IT2-3002 ($\delta^{37}Cl_{SMOC} = +0.83 \pm 0.09\%$ and $-0.19 \pm 0.12\%$, respectively) were 183 calibrated against the CHYN standards using a GC-IRMS in Waterloo and a GC-qMS in 184 Neuchâtel, respectively.

185 A detailed description of analytical methods is available in the Supplementary Information (SI). 186 Carbon isotopes ratios were measured by GC-IRMS and precision based on the working standard δ^{13} C value reproducibility was 0.5% (1 σ). Chlorine isotopic analysis was performed using the 187 188 following instruments: 1) Waterloo - a 6890 GC (Agilent, Santa Clara, CA, US) coupled to a 189 continuous flow IsoPrime IRMS (Micromass, Manchester, UK; currently Isoprime Ltd, UK), 2) München - a TRACETM GC (Thermo Fisher Scientific, Milan, Italy) directly coupled to a 190 191 Finnigan MAT 253 IRMS (Thermo Fisher Scientific, Bremen, Germany) and, 3) Neuchâtel - a 192 7890A GC coupled to a 5975C quadrupole MS (Agilent, Santa Clara, CA, US). The instrument 193 used in Neuchâtel will be referred to in the following text as qMS and those used in Waterloo 194 and München as IRMS-1 and IRMS-2, respectively.

For analyzing 1,2-DCA, two ions of the molecular group (100, 102 m/z) were measured with the IRMS-2, whereas the two most abundant fragment ions (62, 64 m/z) were used for the IRMS-1 and the qMS. For 1,2-DCA the intensities of the most abundant fragment ion peaks are much higher than those of the parent ion peaks. Both ion couples (100, 102 and 62, 64 m/z) correspond to isotopologue pairs ($[^{37}Cl_{2}^{12}C_{2}^{1}H_{4}]^{+}$, $[^{37}Cl^{35}Cl^{12}C_{2}^{1}H_{4}]^{+}$ and $[^{37}Cl^{12}C_{2}^{1}H_{3}]^{+}$, $[^{35}Cl^{12}C_{2}^{1}H_{3}]^{+}$, respectively) that differ by one heavy chlorine isotope. The isotope ratio (R) can be obtained from the ratio of these isotopologues according to eq. 3,²⁷ 202

203
$$R = \frac{{}^{37}Cl}{{}^{35}Cl} = \frac{{}^{37}p}{{}^{35}p} = \frac{k}{(n-k+1)} \cdot \frac{{}^{37}Cl_{(k)}{}^{35}Cl_{(n-k)}}{{}^{37}Cl_{(k-1)}{}^{35}Cl_{(n-k+1)}} = 2 \cdot \frac{{}^{102}I}{{}^{100}I} = \frac{{}^{64}I}{{}^{62}I}$$
(3)

204

where ³⁷p and ³⁵p are the probabilities of encountering ³⁷Cl and ³⁵Cl, n is the number of Cl atoms, k is the number of ³⁷Cl isotopes, ³⁷Cl_(k)³⁵Cl_(n-k) and ³⁷Cl_(k-1)³⁵Cl_(n-k+1) represent the isotopologues containing k and (k-1) heavy isotopes, respectively, and "I" indicates the peak intensity of each ion.

For the qMS, isotope ratios were calculated using eq. 3 and raw δ^{37} Cl values were determined by 209 210 referencing versus an external 1.2-DCA working standard according to eq. 1. In this case the standard was dissolved in water and measured like the samples in the same sequence.²⁶ In IRMS-211 1 and IRMS-2 raw δ^{37} Cl values were determined by automatic evaluation of sample's target ion 212 213 peaks against the ion peaks of the 1,2-DCA monitoring gas, which was introduced by a dual inlet system during each sample run providing an anchor between sample measurements.²⁸ 214 Subsequently, a two point linear calibration of raw δ^{37} Cl values to SMOC scale was performed 215 216 in each laboratory using two external working standards, i.e. IT2(3001 and 3002) for IRMS-1 and CHYN(1 and 2) for IRMS-2 and qMS.²⁸ The analysis schemes applied in each laboratory are 217 218 available in SI. For the measurements with the qMS, samples and standards were diluted to a 219 similar concentration and each of them was measured ten times, leading to a precision (1σ) on 220 the analysis of standards of ±0.3‰. For IRMS-1 and IRMS-2 precision on the analysis of 221 standards was $\pm 0.1\%$ (1 σ). The δ^{37} Cl of controls remained constant (+0.7 \pm 0.2‰, n = 12) at δ^{37} Cl₀ within the analytical uncertainty throughout the experiment. 222

Concentrations of 1,2-DCA were measured by headspace analysis using a TRACETM GC-DSQII 223 224 MS (Thermo Fisher Scientific, Waltham, MA, US) in single ion mode (51, 62, 64, 98 and 100 225 m/z). Concentrations were determined using a five point calibration curve. The estimated total 226 relative error on analysis of external standards interspersed along the sequence was $\pm 4\%$. 227 Concentrations were corrected for 1,2-DCA volatilization to the bottle headspace using Henry's law and Henry coefficients.³⁸ An aqueous phase 1,2-DCA concentration decrease smaller than 228 229 5%, due to change of headspace to solution ratio during the experiment, was estimated. The 230 average concentration of controls remained constant throughout the experiments (1.55 ± 0.03) mM, $\pm 1\sigma$, n = 12) indicating that compound losses through the valves and caps during bottles 231 232 shaking and samples preservation were insignificant.

233 Calculation of apparent kinetic isotope effects (AKIE)

Kinetic isotope effects are position specific whereas ε_{bulk} values are calculated from compoundaverage isotope data. Therefore, observable ε_{bulk} values have to be converted into apparent kinetic isotope effects (AKIEs) in order to obtain information about the underlying reaction mechanisms.²¹ For the calculation of AKIEs a hypothesis about the reaction mechanism, or assumed reaction mechanism, is necessary. The effects of non-reacting positions within the molecule, as well as of intramolecular competition, are then taken into account using eqs. 4 and 5,²¹ respectively,

242
$$\varepsilon_{\rm rp} \approx \frac{n}{x} \cdot \varepsilon_{\rm bulk}$$
 (4)
243

244 AKIE_{C,Cl} =
$$\frac{1}{z \cdot \varepsilon_{rp} + 1}$$
 (5)
245

where ε_{rp} is the isotopic fractionation at the reactive position, n is the number of atoms of the 246 247 element considered, x is the number of reactive sites and z the number of identical reactive sites 248 undergoing intramolecular competition. In symmetric molecules such as 1,2-DCA all atoms are in equivalent reactive positions (n = x) and, therefore, ε_{rp} is directly obtained from eq 2.²¹ The 249 250 isotope fractionation values were quantified by least square linear regression according to eq 2 without forcing the regression to the origin.³⁹ As demonstrated Elsner and Hunkeler (2008),²⁷ the 251 252 Rayleigh equation (eq. 2) can also be applied to calculate the isotopic fractionation of chlorine despite the higher natural abundance of ³⁷Cl compared to ¹³C. 253

254

255 RESULTS AND DISCUSSION

256 Chlorine isotope fractionation values from different instruments

- 258 **Table 1**. Chlorine isotopic fractionation ($\epsilon_{\text{bulk}}^{\text{Cl}}$) values from pure cultures and p-values for paired
- t-test between instruments used.

| | | | IRMS-1 | | IRMS-2 | | qMS | | | |
|----------------|----|-------------------------|---------|-----------------------|---------|-----------------------|---------|-----------------------|-----------------------|--|
| | nª | r ^{2 b} | Average | Variance ^c | Average | Variance ^c | Average | Variance ^c | p-values ^d | |
| A. aquaticus | 3 | ≥0.96 | -4.37 | 0.01 | -4.32 | 0.02 | -4.63 | 0.04 | 0.923/0.390/0.670 | |
| X. autrophicus | 3 | ≥0.98 | -4.343 | 0.005 | -3.960 | 0.007 | -4.39 | 0.01 | 0.034/0.829/0.111 | |
| Pseudomonas | з | >0 95 | -4 08 | 0.03 | -3 64 | 0.03 | -3 650 | 0.007 | 0 085/0 249/0 946 | |
| sp. | 0 | -0.00 | 4.00 | 0.00 | 0.04 | 0.00 | 0.000 | 0.007 | 0.000/0.2+0/0.0+0 | |

^a Number of replicates. ^b Correlation coefficient of least-squares regression according to eq. 2. ^c The variance among triplicate experiments was determined from the variance of regression for each ε_i as $s_{\varepsilon}^2 = (s_{\varepsilon 1}^2 + s_{\varepsilon 2}^2 + s_{\varepsilon 3}^2)/9$.^{39 d} IRMS-1 vs. IRMS-2 / IRMS-1 vs. qMS / IRMS-2 vs. qMS.

265

Chlorine isotope fractionation values (ϵ_{bulk}^{Cl}) measured with different instruments were compared 266 in Table 1. Results show excellent regressions (i.e., $r^2 \ge 0.95$) for the data from triplicate 267 268 experiments when measured on the same instrument (entries in Table 1). In contrast, variation was greater when data from different instruments was compared. The ϵ_{bulk}^{Cl} values from different 269 instruments were compared using paired t-tests carried out with SigmaPlotTM. The p-values were 270 271 above a significant level of 0.05, with the exception of the p-value for X. autrophicus between IRMS-1 (ϵ_{bulk}^{Cl} = -4.3‰) and IRMS-2 (ϵ_{bulk}^{Cl} = -4.0‰) (Table 1). This result indicates that, in 272 general, there is no statistically significant difference in $\varepsilon_{\text{bulk}}^{\text{Cl}}$ values at the 95% confidence level. 273 274 The comparison of p-values for different instrument pairs does not show systematic differences, suggesting that ϵ_{bulk}^{Cl} variation between laboratories could be in part related to minor effects 275 276 during sample handling. The effect of scatter in data points at late stages of reaction on calculated $\epsilon_{\text{bulk}}^{\text{Cl}}$ (Figure 1) could also explain the small differences.⁴⁰ However, removing these 277 278 data points did not significantly improve the agreement.

279 C and Cl isotope fractionation and dual isotope slopes



Figure 1. For each culture C and Cl isotopes data from triplicate experiments were combined in Rayleigh plots. For chlorine, isotope data from different instruments were also combined. The uncertainty for the carbon and chlorine ε values corresponds to the \pm 95% confidence interval calculated from standard deviation of regression slope. For chlorine, dashed lines represent the 95% C.I. of regression parameters.

Degradation experiments lasted between 12h (*A. aquaticus*) and 21h (*X. autrophicus*) and 1,2-DCA transformation above 90% was reached for all the replicates. Carbon and chlorine $\varepsilon_{\text{bulk}}$ values ($r^2 \ge 0.95$) (Figure 1) were determined as indicated above (eq. 2). Transformation of 1,2-DCA by *A. aquaticus* and *X. autrophicus* by a haloalkane hydrolytic dehalogenase reaction resulted in a strong enrichment of ¹³C in the remaining substrate, showing a δ^{13} C shift of approximately 98‰ at 95% degradation. The obtained $\varepsilon_{\text{bulk}}^{\text{C}}$ and $\varepsilon_{\text{bulk}}^{\text{Cl}}$ values for both cultures are the same within 95% confidence intervals and $\varepsilon_{\text{bulk}}^{\text{C}}$ values fall within the ranges determined

from previous studies $^{16,\ 35,\ 41}$ (Table 2). As compared with carbon, a lower ϵ_{bulk}^{Cl} value of 295 approximately -4.3‰ was determined for both cultures ($\epsilon_{bulk}^{C} / \epsilon_{bulk}^{Cl} = 7.4$). Degradation of 1,2-296 DCA by Pseudomonas sp. Strain DCA1 in an enzymatic monooxygenase reaction resulted in a 297 smaller carbon ϵ_{bulk}^{C} (-3.5‰) compared to nucleophilic (S_N2) reaction (Figure 1). This result is 298 consistent with the previously reported value (Table 2).¹⁶ For Cl, a ϵ_{bulk}^{Cl} value close to those 299 associated with primary Cl isotopic effects was measured (ϵ_{bulk}^{C} / ϵ_{bulk}^{Cl} = 0.9). This value is 300 301 unusually high given that from known mechanism a secondary isotope effect is expected (see 302 further discussion in the next Section).

A linear relation between $\delta^{13}C$ and $\delta^{37}Cl$ was obtained for all the strains ($r^2 \ge 0.97$) and strongly 303 different slopes ($\Lambda = \Delta \delta^{13}C / \Delta \delta^{37}Cl \approx \epsilon_{bulk}^C / \epsilon_{bulk}^{Cl}$) were determined for the S_N2 reaction ($\Lambda = \delta^{13}C / \delta^{13}C / \delta^{13}Cl \approx \epsilon_{bulk}^C / \epsilon_{bulk}^C$) 304 305 7.7 ± 0.2) and oxidation ($\Lambda = 0.78 \pm 0.03$) (Figure 2). The slopes obtained from X. autrophicus 306 and A. aquaticus degradation experiments were the same within 95% confidence intervals. 307 Therefore, the large Λ difference among the investigated reactions enables the use of a dual 308 isotope approach to identify the different pathways in the field. In contrast, a single element 309 approach based only on carbon isotope data would lead to ambiguous interpretations because a certain extent of isotope fractionation (e.g. $\Delta \delta^{13}$ C) could have been caused by a strongly isotope 310 311 fractionating reaction that has proceeded little, or a weakly isotope fractionating reaction that has proceeded further (or an unknown combination of both). Unlike in lab experiments where ϵ^{C}_{bulk} 312 values can be determined,¹⁶ insight into pathways would therefore be elusive in the field. The 313 314 starkly contrasting trends of Figure 2 show how isotopic analysis of chlorine as second element 315 can resolve this issue. The large difference between determined slopes also enables to identify if 316 both pathways occur in the field. In ideal situations, the proportion of both competing pathways

could be estimated based on dual isotope data.^{42, 43} A dual isotope analysis can also be helpful in
microbial laboratory experiments for substantiating conclusions about prevailing mechanisms.



319

Figure 2. Carbon and chlorine δ isotope values from triplicate experiments and all used instruments (i.e. IRMS-1, IRMS-2 and qMS) were combined in a dual isotope plot. Symbols are as follows: blue (*A. aquaticus*), red (*X. autrophicus*), gren (*Pseudomonas sp.*), circles (IRMS-1), squares (IRMS-2) and diamonds (qMS). The slopes of the linear regression lines (solid lines) give the Λ values (±95% C.I.) and the dashed lines correspond to the 95% confidence intervals. Error bars of δ^{13} C values are smaller than the symbols.

326

327 Interpretation of ¹³C- and ³⁷Cl-AKIEs for oxidation and S_N2-type reactions

| 329 | Table 2. Measured carbon and chlorine isotopic fractionation values (in ‰) and apparent kinetic |
|-----|---|
| 330 | isotopic effects. |

| | | Reported | | | | |
|--------------------|-----------------------|-----------------------------|-----------------------------------|--------------------------|------------------------------------|--------------------------------------|
| | Reaction mechanism | $\epsilon^{C}_{bulk}{}^{a}$ | ¹³ C-AKIE ^b | ϵ_{bulk}^{Cl} a | ³⁷ CI-AKIE ^b | ϵ_{bulk}^{C} |
| A. aquaticus | S _N 2 | -32.0 ± 0.9 | 1.068 ± 0.002 | -4.4 ± 0.2 | 1.0089 ± 0.0004 | -31.9 to -32.4 ¹⁶ |
| X. autrophicus | S _N 2 | -31.9 ± 0.7 | 1.068 ± 0.002 | -4.2 ± 0.1 | 1.0085 ± 0.0002 | -28.7 to -33.0 ^{16, 35, 41} |
| Pseudomonas sp. | C-H bond cleavage | -3.5 ± 0.1 | 1.0070 ± 0.0002 | -3.8 ± 0.2 | 1.0038 ± 0.0002 | -3.0 ± 0.2^{16} |

^a See Rayleigh plots in Figure 1. The uncertainty corresponds to the 95% confidence interval 332 calculated from standard deviation of regression slope. ^b Calculated according eq. 5. The 333 334 uncertainty was estimated by error propagation.

335

331

336 Hydrolytic dehalogenation (S_N2) reaction

Determined ε_{bulk} values were used to estimate the AKIEs (Table 2) according to eq. 5, which 337 assumes that secondary isotopic effects can be neglected. For S_N2 reaction, ¹³C-AKIEs were 338 calculated using z = 2 since both C-Cl bonds compete for reaction. ¹³C-AKIEs agreed well with 339 the typical ¹³C-KIE range for a $S_N 2$ reaction $(1.03 - 1.09)^{21}$ (Scheme 1). Abe et al. $(2009)^{41}$ 340 determined the ¹³C-AKIE of 1,2-DCA in batch degradation experiments prepared with cell free 341 extract from X. autrophicus GJ10. These authors obtained an average value (1.0597) close to the 342 343 ¹³C-AKIE observed in this study (1.068), suggesting that there was no significant masking of the intrinsic KIE during compound transport thorough the cell membrane. This conclusion is in 344

agreement with the Streitwieser limit for ¹³C-KIE in C-Cl bonds (1.057),⁴⁴ and could explain in part the relatively narrow range of reported $\varepsilon_{\text{bulk}}^{\text{C}}$ values (from -28.7 to -33.0‰) for both pure cultures using the haloalkane hydrolytic dehalogenase reaction (Table 2). Hirschorn et al. (2007)⁹ measured a similar ¹³C-AKIE (1.05) for 1,2-DCA in laboratory biodegradation experiments under nitrate reducing conditions by an enrichment culture from a contaminated site, which was interpreted as transformation via hydrolytic dehalogenation.

Similarly to carbon isotope effects, 37 Cl-AKIE were obtained from eq. 5 with z = 2. The 351 calculated ³⁷Cl-AKIE (1.009) corresponded well to the typical ³⁷Cl-KIE range for a S_N 2 reaction 352 but very close to the upper end (1.006-1.009)⁴⁵ (Scheme 1). However, the ³⁷Cl-AKIE measured 353 in this study is above the theoretical primary isotope effect for 1,2-DCA enzymatic 354 dehalogenation reported by Lewandowicz et al. (2001) (37 Cl-KIE = 1.0065). 46 These authors also 355 measured experimentally the leaving group ³⁷Cl-AKIE for 1,2-DCA (1.0045) and 1-chlorobutane 356 357 (1.0066) dechlorination catalyzed by haloalkane hydrolytic dehalogenase (extracted from X. autrophicus GJ10). In this former study the experimental ³⁷Cl-AKIE was determined by the 358 359 isotopic analysis of the released Cl⁻ during 1,2-DCA dechlorination and, therefore, it represents the primary ³⁷Cl-AKIE. According to Lewandowicz et al. (2001)⁴⁶ and Paneth (2003)⁴⁷ an 360 explanation for the lower ³⁷Cl-AKIE of 1,2-DCA compared to 1-chlorobutane could be that the 361 362 dehalogenation step is reversible and the hydrolysis of the enzyme-bound intermediate is 363 responsible for the overall irreversibility of the reaction. In addition, in a recent study that 364 investigated 1,1-dichloroethane and 1,1,1-trichlorethane biodegradation by whole cell and cell free extract systems,⁴⁸ ¹³C-AKIEs for both chlorinated ethanes during degradation by cell free 365 366 extracts were unexpectedly lower than those determined in whole cell experiments. The higher ³⁷Cl-AKIE of 1.009 compared to the theoretical primary ³⁷Cl-KIE could be explained by the 367

368 contribution of a β -secondary isotopic effect given that, in our study, the ³⁷Cl / ³⁵Cl ratios were 369 measured in the remaining 1,2-DCA and, secondary isotopic effects were neglected in the 370 primary AKIE calculation (eq. 5). The magnitude of the β -secondary ³⁷Cl-KIE can be estimated 371 from the average of KIE_i in both Cl molecular positions according to eq. 6.¹²

372

373
$$\varepsilon_{\text{bulk}}^{\text{Cl}} \approx \frac{1}{2} \cdot \left(\frac{1}{\text{KIE}_{\text{primary}}^{\text{Cl}}} + \frac{1}{\text{KIE}_{\text{secondary}}^{\text{Cl}}} \right) - 1$$
 (6)

374

Plugging in the ϵ_{bulk}^{Cl} value measured in our study (-4.3 ‰) and the reported theoretical primary 375 Cl isotopic effect of 1.0065,^{46, 47} a β-secondary ³⁷Cl-KIE of 1.0021 was estimated. Equation 6 376 assumes that $\varepsilon_{\text{bulk}}^{\text{Cl}}$ is not significantly masked by non-fractionating rate limiting processes 377 preceding the reaction step. This is a likely assumption in our case judging by the relatively high 378 carbon and chlorine AKIEs (see discussion above and Table 2). Secondary ¹³C-KIEs are 379 380 generally much smaller than primary isotope effects and, therefore, they are usually omitted in ¹³C-AKIE calculations.²¹ However, for chlorine, secondary KIEs as large as primary isotope 381 382 effects have been recently determined theoretically during S_N2 reactions that proceed with chlorine transfer between two heavy atoms.⁴⁵ In addition, in a recent experimental study that 383 investigates TCE multi-isotope fractionation during biotic reductive dechlorination,³² daughter 384 385 products depleted in ³⁷Cl relative to their immediate parent compounds were interpreted as 386 evidence of significant secondary Cl effects related to nucleophilic substitution reaction. The conclusions of these recent studies^{32, 45} support the hypothesis that large β-secondary ³⁷Cl-KIE 387 388 occurs during this study.

390 As in the hydrolytic reaction, during oxidation both C atoms also compete for reaction and thus the ¹³C-AKIE was calculated using z = 2 in eq. 5 (Table 2). The obtained ¹³C-AKIE agrees well 391 with the typical ¹³C-KIE for C-H bond cleavage $(1.01 - 1.03)^{21}$ (Scheme 1). Primary ¹³C-KIEs 392 generally increase with increasing mass of the bonding partner (i.e. ${}^{13}C-KIE_{C-H} < {}^{13}C-KIE_{C-CI}$).²¹ 393 For Cl, the unexpectedly high $\varepsilon_{\text{bulk}}^{\text{Cl}}$ suggests the contribution of secondary ³⁷Cl-KIEs from both 394 Cl atoms. In the absence of primary chlorine isotopic effect, the secondary ³⁷Cl-AKIE can also 395 396 be evaluated using eq. 5. In this case z = 1 since no specific bond containing Cl is broken and 397 there is, therefore, no intramolecular competition for this bond. The resultant secondary ³⁷Cl-398 AKIE (1.004) represents the average secondary isotope effect of all positions. This result supports the large β -secondary ³⁷Cl-KIE estimated above for the nucleophilic reaction. This 399 finding suggests that significant secondary ³⁷Cl-KIEs are also associated with enzymatic 400 401 oxidation via C-H bond cleavage. Until now, oxidative cleavage of a C-H bond has been 402 believed to affect primarily the C-H bond and to leave chlorine substituents largely unchanged. 403 This common assumption is challenged by the observed large chlorine isotope fractionation in C-404 H bond cleavage, where involvement of a C-Cl bond would not be expected. This indicates an 405 intriguing role of the chlorine atoms which still remain to be resolved. Hypothetical contribution 406 of chlorine isotope fractionation during binding of 1,2-DCA molecules to the enzyme could be 407 an additional explanation.

408 Further insight can possibly be obtained in future studies that address ${}^{2}\text{H} / {}^{1}\text{H}$ isotope analysis. 409 Recently, Shouakar-Stash and Drimmie (2013)⁴⁹ developed and online methodology for H-CSIA 410 of TCE and 1,2-cis-dichloroethene, however, this analytical method is still not implemented for411 chlorinated ethanes.

412 Implications for application of CSIA to environmental studies

One of the main applications of CSIA to field studies is the estimation of contaminant 413 biodegradation extent and rate.^{15, 44, 50} For this purpose compound specific ϵ_{bulk} values from 414 laboratory experiments are necessary. For a given compound, different ε_{bulk} values are generally 415 416 associated with distinct biodegradation pathways, which in turn are sometimes related to 417 different subsurface redox environments. Therefore, redox conditions are usually used as criteria 418 to constrain the range of reported ε_{bulk} values. However, for 1,2-DCA different degradation 419 pathways associated with distinct ε_{bulk} values may even be active under aerobic conditions. In 420 addition, the hydrolytic dehalogenation pathway of 1,2-DCA has been observed under oxygen and nitrate reducing conditions alike (see above)⁹. In this context, the identification of the active 421 422 degradation pathway in the field is crucial to choose the appropriate ε_{bulk} value for quantification 423 of degradation. This study demonstrates that dual isotope slopes are strongly different for 424 nucleophilic substitution (S_N2) and oxidation (C-H bond cleavage) reactions, which opens the 425 possibility to identify them using a dual isotope approach. Following this approach isotopic data 426 from the field site can be directly and intuitively interpreted.

427 Significant secondary chlorine isotopic effects were determined for the investigated reactions. 428 These results indicate that primary ³⁷Cl-AKIEs derived from CSIA could be higher than reported 429 primary ³⁷Cl-KIEs (e.g. from computational studies) if secondary isotopic effects are omitted in 430 the calculations. Hence, mechanistic interpretations based on the comparison with primary ³⁷Cl-431 KIEs should be made with caution. Finally, chlorine ε_{bulk} values measured with three different instruments, two GC-IRMS and one GC-qMS, showed a fairly good agreement varying at most by $\pm 3.8\%$ (SD of the mean). Even though the agreement was not perfect these results lend confidence to the methods used and encourage the application of Cl-CSIA to investigate the fate of chlorinated compounds at contaminated sites. However, further research and methodological developments are still required to improve Cl-CSIA data agreement between different laboratories.

438

439 ASSOCIATED CONTENT

440 Supporting Information

441 A detailed description of analytical methods is available. This material is available free of charge

442 via the Internet at <u>http://pubs.acs.org</u>.

443

- 444 AUTHOR INFORMATION
- 445 Corresponding Author*
- 446 Jordi Palau
- 447 Centre d'Hydrogéologie et de Géothermie, Université de Neuchâtel, Rue Emile-Argand 11,
- 448 CH-2000 Neuchâtel, Switzerland
- 449 E-mail: jordi.palau@ub.edu

450 Author Contributions

451 The manuscript was written through contributions of all authors. All authors have given approval

452 to the final version of the manuscript.

453 ACKNOWLEDGEMENTS

454 **References**

455 1. Squillace, P. J.; Moran, M. J.; Lapham, W. W.; Price, C. V.; Clawges, R. M.; Zogorski, J.

456 S., Volatile Organic Compounds in Untreated Ambient Groundwater of the United States,
457 1985–1995. *Environ Sci Technol* 1999, *33*, (23), 4176-4187.

458 2. ATSDR Toxicological Profile for 1,2-Dichloroethane.
459 http://www.atsdr.cdc.gov/tfacts38.pdf

460 3. Hage, J. C.; Hartmans, S., Monooxygenase-mediated 1,2-dichloroethane degradation by

461 Pseudomonas sp. strain DCA1. *Appl Environ Microbiol* **1999**, *65*, (6), 2466-70.

462 4. Klecka, G. M.; Carpenter, C. L.; Gonsior, S. J., Biological transformations of 1,2463 dichloroethane in subsurface soils and groundwater. *J Contam Hydrol* 1998, *34*, (1-2), 139-154.

464 5. van den Wijngaard, A. J.; van der Kamp, K. W.; van der Ploeg, J.; Pries, F.; Kazemier,
465 B.; Janssen, D. B., Degradation of 1,2-dichloroethane by Ancylobacter aquaticus and other
466 facultative methylotrophs. *Appl Environ Microbiol* 1992, *58*, (3), 976-83.

467 6. Janssen, D. B.; Scheper, A.; Dijkhuizen, L.; Witholt, B., Degradation of halogenated
468 aliphatic compounds by Xanthobacter autotrophicus GJ10. *Appl Environ Microbiol* 1985, 49,
469 (3), 673-7.

470 7. Grostern, A.; Edwards, E. A., Characterization of a Dehalobacter Coculture That
471 Dechlorinates 1,2-Dichloroethane to Ethene and Identification of the Putative Reductive
472 Dehalogenase Gene. *Appl Environ Microbiol* 2009, 75, (9), 2684-2693.

Yu, R.; Peethambaram, H. S.; Falta, R. W.; Verce, M. F.; Henderson, J. K.; Bagwell, C.
E.; Brigmon, R. L.; Freedman, D. L., Kinetics of 1,2-Dichloroethane and 1,2-Dibromoethane
Biodegradation in Anaerobic Enrichment Cultures. *Appl Environ Microbiol* 2013, *79*, (4), 13591367.

477 9. Hirschorn, S. K.; Dinglasan-Panlilio, M. J.; Edwards, E. A.; Lacrampe-Couloume, G.;
478 Sherwood Lollar, B., Isotope analysis as a natural reaction probe to determine mechanisms of
479 biodegradation of 1,2-dichloroethane. *Environ Microbiol* 2007, *9*, (7), 1651-7.

Maes, A.; van Raemdonck, H.; Smith, K.; Ossieur, W.; Lebbe, L.; Verstraete, W.,
Transport and activity of Desulfitobacterium dichloroeliminans strain DCA1 during
bioaugmentation of 1,2-DCA-contaminated groundwater. *Environ Sci Technol* 2006, *40*, (17),
5544-5552.

Hirschorn, S. K.; Grostern, A.; Lacrampe-Couloume, G.; Edwards, E. A.; Mackinnon, L.;
Repta, C.; Major, D. W.; Sherwood Lollar, B., Quantification of biotransformation of chlorinated
hydrocarbons in a biostimulation study: added value via stable carbon isotope analysis. *J Contam Hydrol* 2007, *94*, (3-4), 249-60.

Elsner, M., Stable isotope fractionation to investigate natural transformation mechanisms
of organic contaminants: principles, prospects and limitations. *J Environ Monitor* 2010, *12*, (11),
2005-2031.

491 13. Hofstetter, T. B.; Berg, M., Assessing transformation processes of organic contaminants
492 by compound-specific stable isotope analysis. *Trac-Trend Anal Chem* 2011, *30*, (4), 618-627.

493 14. Hofstetter, T. B.; Schwarzenbach, R. P.; Bernasconi, S. M., Assessing Transformation
494 Processes of Organic Compounds Using Stable Isotope Fractionation. *Environ Sci Technol* 2008,
495 42, (21), 7737-7743.

496 15. Braeckevelt, M.; Fischer, A.; Kastner, M., Field applicability of Compound-Specific
497 Isotope Analysis (CSIA) for characterization and quantification of in situ contaminant
498 degradation in aquifers. *Appl Microbiol Biotechnol* 2012, *94*, (6), 1401-1421.

Hirschorn, S. K.; Dinglasan, M. J.; Elsner, M.; Mancini, S. A.; Lacrampe-Couloume, G.;
Edwards, E. A.; Lollar, B. S., Pathway dependent isotopic fractionation during aerobic
biodegradation of 1,2-dichloroethane. *Environ Sci Technol* 2004, *38*, (18), 4775-4781.

502 17. Zwank, L.; Berg, M.; Elsner, M.; Schmidt, T. C.; Schwarzenbach, R. P.; Haderlein, S. B.,
503 New evaluation scheme for two-dimensional isotope analysis to decipher biodegradation
504 processes: Application to groundwater contamination by MTBE. *Environ Sci Technol* 2005, *39*,
505 (4), 1018-1029.

Abe, Y.; Aravena, R.; Zopfi, J.; Shouakar-Stash, O.; Cox, E.; Roberts, J. D.; Hunkeler,
D., Carbon and Chlorine Isotope Fractionation during Aerobic Oxidation and Reductive
Dechlorination of Vinyl Chloride and cis-1,2-Dichloroethene. *Environ Sci Technol* 2009, *43*, (1),
101-107.

S10 19. Cretnik, S.; Thoreson, K. A.; Bernstein, A.; Ebert, K.; Buchner, D.; Laskov, C.;
Haderlein, S.; Shouakar-Stash, O.; Kliegman, S.; McNeill, K.; Elsner, M., Reductive
Dechlorination of TCE by Chemical Model Systems in Comparison to Dehalogenating Bacteria:
Insights from Dual Element Isotope Analysis (13C/12C, 37Cl/35Cl). *Environ Sci Technol* 2013,
47, (13), 6855-6863.

515 20. Chartrand, M. M. G.; Hirschorn, S. K.; Lacrampe-Couloume, G.; Lollar, B. S.,
516 Compound specific hydrogen isotope analysis of 1,2-dichloroethane: potential for delineating
517 source and fate of chlorinated hydrocarbon contaminants in groundwater. *Rapid Commun Mass*518 *Sp* 2007, *21*, (12), 1841-1847.

519 21. Elsner, M.; Zwank, L.; Hunkeler, D.; Schwarzenbach, R. P., A new concept linking
520 observable stable isotope fractionation to transformation pathways of organic pollutants. *Environ*521 *Sci Technol* 2005, *39*, (18), 6896-6916.

522 22. Bigeleisen, J.; Wolfsberg, M., Theoretical and Experimental Aspects of Isotope Effects in
523 Chemical Kinetics. *Adv Chem Phys* 1958, *1*, 15-76.

| 524 | 23. Shouakar-Stash, O.; Drimmie, R. J.; Zhang, M.; Frape, S. K., Compound-specific |
|-----|---|
| 525 | chlorine isotope ratios of TCE, PCE and DCE isomers by direct injection using CF-IRMS. Appl |
| 526 | Geochem 2006 , 21, (5), 766-781. |

527 24. Sakaguchi-Soder, K.; Jager, J.; Grund, H.; Matthaus, F.; Schuth, C., Monitoring and
528 evaluation of dechlorination processes using compound-specific chlorine isotope analysis. *Rapid*529 *Commun Mass Sp* 2007, *21*, (18), 3077-3084.

Jin, B.; Laskov, C.; Rolle, M.; Haderlein, S. B., Chlorine isotope analysis of organic
contaminants using GC-qMS: method optimization and comparison of different evaluation
schemes. *Environ Sci Technol* 2011, 45, (12), 5279-86.

533 26. Aeppli, C.; Holmstrand, H.; Andersson, P.; Gustafsson, O., Direct Compound-Specific
534 Stable Chlorine Isotope Analysis of Organic Compounds with Quadrupole GC/MS Using
535 Standard Isotope Bracketing. *Anal Chem* 2010, *82*, (1), 420-426.

536 27. Elsner, M.; Hunkeler, D., Evaluating chlorine isotope effects from isotope ratios and 537 mass spectra of polychlorinated molecules. *Anal Chem* **2008**, *80*, (12), 4731-4740.

Bernstein, A.; Shouakar-Stash, O.; Ebert, K.; Laskov, C.; Hunkeler, D.; Jeannottat, S.; 538 28. 539 Sakaguchi-Soder, K.; Laaks, J.; Jochmann, M. A.; Cretnik, S.; Jager, J.; Haderlein, S. B.; 540 Schmidt, T. C.; Aravena, R.; Elsner, M., Compound-Specific Chlorine Isotope Analysis: A 541 Comparison of Gas Chromatography/Isotope Ratio Mass Spectrometry and Gas 542 Chromatography/Quadrupole Mass Spectrometry Methods in an Interlaboratory Study. *Anal*543 *Chem* 2011, 83, (20), 7624-7634.

Wiegert, C.; Mandalakis, M.; Knowles, T.; Polymenakou, P. N.; Aeppli, C.;
Macháčková, J.; Holmstrand, H.; Evershed, R. P.; Pancost, R. D.; Gustafsson, Ö., Carbon and
Chlorine Isotope Fractionation During Microbial Degradation of Tetra- and Trichloroethene. *Environ Sci Technol* 2013, 47, (12), 6449-6456.

548 30. Wiegert, C.; Aeppli, C.; Knowles, T.; Holmstrand, H.; Evershed, R.; Pancost, R. D.;

549 Machackova, J.; Gustafsson, O., Dual Carbon-Chlorine Stable Isotope Investigation of Sources

and Fate of Chlorinated Ethenes in Contaminated Groundwater. *Environ Sci Technol* 2012.

31. Hunkeler, D.; Abe, Y.; Broholm, M. M.; Jeannottat, S.; Westergaard, C.; Jacobsen, C. S.;
Aravena, R.; Bjerg, P. L., Assessing chlorinated ethene degradation in a large scale contaminant
plume by dual carbon-chlorine isotope analysis and quantitative PCR. *J Contam Hydrol* 2011, *119*, (1-4), 69-79.

32. Kuder, T.; van Breukelen, B. M.; Vanderford, M.; Philp, P., 3D-CSIA: Carbon, Chlorine,
and Hydrogen Isotope Fractionation in Transformation of TCE to Ethene by a Dehalococcoides
Culture. *Environ Sci Technol* 2013.

33. Palau, J.; Marchesi, M.; Chambon, J. C.; Aravena, R.; Canals, A.; Binning, P. J.; Bjerg,
P. L.; Otero, N.; Soler, A., Multi-isotope (carbon and chlorine) analysis for fingerprinting and

site characterization at a fractured bedrock aquifer contaminated by chlorinated ethenes. *Sci Total Environ* 2014, 475C, 61-70.

562 34. Lojkasek-Lima, P.; Aravena, R.; Parker, B. L.; Cherry, J. A., Fingerprinting TCE in a
563 Bedrock Aquifer Using Compound-Specific Isotope Analysis. *Ground Water* 2012, *50*, (5), 754564 764.

565 35. Hunkeler, D.; Aravena, R., Evidence of substantial carbon isotope fractionation among 566 substrate, inorganic carbon, and biomass during aerobic mineralization of 1, 2-dichloroethane by 567 Xanthobacter autotrophicus. *Appl Environ Microbiol* **2000**, *66*, (11), 4870-6.

568 36. Elsner, M.; Couloume, G. L.; Lollar, B. S., Freezing to preserve groundwater samples 569 and improve headspace quantification limits of water-soluble organic contaminants for carbon 570 isotope analysis. *Anal Chem* **2006**, *78*, (21), 7528-7534.

37. Holt, B. D.; Sturchio, N. C.; Abrajano, T. A.; Heraty, L. J., Conversion of chlorinated
volatile organic compounds to carbon dioxide and methyl chloride for isotopic analysis of carbon
and chlorine. *Anal Chem* 1997, *69*, (14), 2727-2733.

574 38. Staudinger, J.; Roberts, P. V., A critical compilation of Henry's law constant temperature
575 dependence relations for organic compounds in dilute aqueous solutions. *Chemosphere* 2001, *44*,
576 (4), 561-576.

577 39. Scott, K. M.; Lu, X.; Cavanaugh, C. M.; Liu, J. S., Optimal methods for estimating
578 kinetic isotope effects from different forms of the Rayleigh distillation equation. *Geochim*579 *Cosmochim Ac* 2004, 68, (3), 433-442.

40. Mundle, S. O. C.; Vandersteen, A. A.; Lacrampe-Couloume, G.; Kluger, R.; Lollar, B. S., Pressure-monitored headspace analysis combined with compound-specific isotope analysis to measure isotope fractionation in gas-producing reactions. *Rapid Commun Mass Sp* **2013**, *27*, (15), 1778-1784.

41. Abe, Y.; Zopfi, J.; Hunkeler, D., Effect of molecule size on carbon isotope fractionation
during biodegradation of chlorinated alkanes by Xanthobacter autotrophicus GJ10. *Isotopes Environ Health Stud* 2009, 45, (1), 18-26.

42. van Breukelen, B. M., Extending the Rayleigh equation to allow competing isotope
fractionating pathways to improve quantification of biodegradation. *Environ Sci Technol* 2007,
41, (11), 4004-10.

590 43. Centler, F.; Hesse, F.; Thullner, M., Estimating pathway-specific contributions to
591 biodegradation in aquifers based on dual isotope analysis: Theoretical analysis and reactive
592 transport simulations. *J Contam Hydrol* 2013, *152*, 97-116.

44. Aelion, C. M.; Hohëner, P.; Hunkeler, D.; Aravena, R., *Environmental isotopes in biodegradation and bioremediation*. CRC Press: Boca Raton, 2010; p xiv, 450 p.

595 45. Swiderek, K.; Paneth, P., Extending Limits of Chlorine Kinetic Isotope Effects. J Org
596 Chem 2012, 77, (11), 5120-5124.

597 46. Lewandowicz, A.; Rudzinski, J.; Tronstad, L.; Widersten, M.; Ryberg, P.; Matsson, O.;
598 Paneth, P., Chlorine kinetic isotope effects on the haloalkane dehalogenase reaction. *J Am Chem*599 Soc 2001, 123, (19), 4550-4555.

47. Paneth, P., Chlorine kinetic isotope effects on enzymatic dehalogenations. *Acc Chem Res*2003, *36*, (2), 120-126.

48. Lollar, B. S.; Hirschorn, S.; Mundle, S. O.; Grostern, A.; Edwards, E. A.; LacrampeCouloume, G., Insights into enzyme kinetics of chloroethane biodegradation using compound
specific stable isotopes. *Environ Sci Technol* 2010, *44*, (19), 7498-503.

49. Shouakar-Stash, O.; Drimmie, R. J., Online methodology for determining compoundspecific hydrogen stable isotope ratios of trichloroethene and 1,2-cis-dichloroethene by
continuous-flow isotope ratio mass spectrometry. *Rapid Commun Mass Sp* 2013, *27*, (12), 13351344.

50. Thullner, M.; Centler, F.; Richnow, H. H.; Fischer, A., Quantification of organic
pollutant degradation in contaminated aquifers using compound specific stable isotope analysis Review of recent developments. *Org Geochem* 2012, *42*, (12), 1440-1460.