

Manuscript Details

Manuscript number	HUMEV_2018_236_R1
Title	Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of Castell de Barberà to the earliest Vallesian
Article type	Full Length Article

Abstract

Castell de Barberà, located in the Vallès-Penedès Basin (NE Iberian Peninsula), is one of the few European sites where pliopithecoids (*Barberapithecus*) and hominoids (cf. *Dryopithecus*) co-occur. The dating of this Miocene site has proven controversial. A latest Aragonian (MN7+8, ca. 11.88–11.18 Ma) age was long accepted by most authors, despite subsequent reports of hipparionin remains that signaled a Vallesian age. On the latter basis, Castell de Barberà was recently correlated to the early Vallesian (MN9, ca. 11.18–10.3 Ma) on tentative grounds. Uncertainties about the provenance of the *Hippotherium* material and the lack of magnetostratigraphic data precluded more accurate dating. After decades of inactivity, fieldwork was resumed in 2014–2015 at Castell de Barberà, including the original layer (CB-D) that in the past delivered most of the fossils. Here we report magnetostratigraphic results for the original outcrop and another nearby section. Our results indicate that CB-D is located in a normal polarity magnetozone at about midheight of a short (~20 m-thick) stratigraphic section. The composite magnetostratigraphic section (~50 m) has as many as four to six magnetozones. These multiple reversals, coupled with the in situ recovery of a *Hippotherium* humerus from CB-D in 2015, make it very unlikely the correlation of any of the sampled normal polarity magnetozones with the long normal polarity subchron C5n.2n (11.056–9.984 Ma), which is characteristic of the early Vallesian. Our results support instead a correlation of CB-D with C5r.1n (11.188–11.146 Ma), where the Aragonian/Vallesian boundary is situated, and therefore indicate an earliest Vallesian age of ~11.2 Ma for Castell de Barberà. Our results settle the longstanding debate about the Aragonian vs. Vallesian age of this site, which appears roughly coeval with the Creu de Conill 20 locality (11.18 Ma), where hipparionins are first recorded in the Vallès-Penedès Basin.

Keywords	Hominoidea; Pliopithecoidea; <i>Hippotherium</i> ; Late Miocene; Magnetostratigraphy; Paleomagnetism
Corresponding Author	David Alba
Corresponding Author's Institution	Institut Català de Paleontologia Miquel Crusafont
Order of Authors	David Alba, Miguel Garces, Isaac Casanovas Vilar, Josep M. Robles, Marta Pina, Salvador Moya-Sola, Sergio Almécija

Submission Files Included in this PDF

File Name [File Type]

Cover letter.R1.pdf [Cover Letter]

Response to reviewers.R1.docx [Response to Reviewers (without Author Details)]

Text R1 with changes.docx [Revised Manuscript with Changes Marked (without Author Details)]

Title page R1 with changes.docx [Title Page (with Author Details)]

Text R1 clean.docx [Manuscript (without Author Details)]

Figure 1.tif [Figure]

Figure 2.tif [Figure]

Figure 3.tif [Figure]

Figure 4 R1.eps [Figure]

Figure 5 R1.eps [Figure]

Figure 6 R1.eps [Figure]

Figure 7 R1.tif [Figure]

Table 1 R1 with changes.docx [Table]

SOM R1.docx [e-Component]

To view all the submission files, including those not included in the PDF, click on the manuscript title on your EVISE Homepage, then click 'Download zip file'.



Dr. David M. Alba
Director, Institut Català de Paleontologia Miquel Crusafont,
Universitat Autònoma de Barcelona.
Co-Editor-in-Chief, *Journal of Human Evolution*.
Edifici ICTA-ICP, C/ Columnes s/n, Campus de la UAB,
08193 Cerdanyola del Vallès, Barcelona, Spain.
E-mail: david.alba@icp.cat

Barcelona, 12 February 2019

Dr. Mike Plavcan
Editor-in-Chief, *Journal of Human Evolution*

Dear Mike,

Please receive a revised version of manuscript HUMEV_2018_236 (“Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of Castell de Barberà to the earliest Vallesian”) by Alba and coauthors, to be considered for publication in the *Journal of Human Evolution*.

We are thankful to the reviewers for their detailed and constructive input and have substantially revised the manuscript accordingly. A detailed response to the reviewer comments (with emphasis on those instances in which we were not able to implement their suggestions) has been attached as a separate ‘response to reviewers’ file.

The submission consists of the revised title page, main text, and table files (all in Word format), as well as a newly added SOM file (also in Word) and seven figure files (of which one is new and three have been revised to some extent). The SOM includes a brief text, two new figures and a new table, all of which included in response to the reviewer comments. The main text, in turn, has been extensively edited (including the title) and, to some extent, reorganized and rewritten (with multiple additions to adequately address the reviewer comments).

We think that the revised version has significantly improved thanks to the reviewers’ input and hope that you will find it suitable for publication. We look forward to hearing back from you at your earliest convenience.

All the best,

A handwritten signature in blue ink, consisting of several overlapping loops and a long horizontal stroke at the bottom.

David M. Alba

RESPONSE TO REVIEWERS

Ref: HUMEV_2018_236

Title: Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of Castell de Barberà to the earliest Vallesian: End of the controversy

Journal: Journal of Human Evolution

Editor-in-Chief

Thank you very much for submitting your paper to the Journal of Human Evolution. I have now received comments from two reviewers as well as a recommendation of “revise” from the Associate Editor. All agree that this is a well-written and engaging piece that is highly suitable for the JHE, but that will need revision before acceptance.

Both reviewers have provided detailed comments with ample explanation, requiring little elaboration from me. As you will see, reviewer #1 (Roberts, who provides his review in an attached document, as well as an annotated version of the manuscript) is highly complementary of the piece, but questions the claimed certainty of the dating. Assuming that this is true (I am not a dating person, of course), I have to agree that you need to drop “: End of the Controversy” from the title and re-tool the text to address the uncertainties in the dates. The reviewer provides very helpful and detailed explanation of his reasoning and recommendations for revision.

Reviewer 2 likewise provides a series of detailed and helpful comments that should greatly clarify the text. Most of these are very straightforward, and include a lot of grammatical and usage comments, all of which are quite helpful. Apart from this, I have nothing on my own part at this point to add to the text. Please look over the reviews and revise appropriately. When I receive the revision, I will return it to the AE for a second round of reviews.

We thank the reviewers for their comments, which we found constructive and helpful to improve the original manuscript. Below we provide detailed explanations as to how we addressed these comments and outline the changes introduced in the text, which are also marked in a tracked version of the revised manuscript. With regard to the title, we removed “End of the controversy” at the end as per editor and reviewer’s request. However, note that the alluded controversy referred to the Aragonian vs. Vallesian dating of the site (as implicit from the title and more explicitly stated elsewhere in the original manuscript). We discuss at greater length below the more detailed criticisms by reviewer 1 with regard to the dating.

Reviewer 1 (Andrew P. Roberts)

Please see attached review and annotated manuscript.

First of all, we would like to thank the reviewer for his very thorough review of the manuscript and the multiple edits and suggestions provided. We will discuss first the comments kindly provided by the reviewer in the annotated manuscript, and subsequently the more detailed comments sent by him separately.

Annotated manuscript:

Most of the edits provided by the reviewer have been incorporated to the revised version without further discussion, even if in many instances they were largely stylistic; these changes are marked in the tracked version of the revised manuscript. We only comment below (using line numbers of the original manuscript) those instances in which we were unable, or considered inappropriate, to implement the reviewer's edits and suggestions, along with other instances that in our opinion required further clarification. Note that the list below appears very long (ca. 80 edits commented). However, the Editor should be aware that the reviewer provided ca. 550 edits to the original manuscript, implying that between 80 and 90% of them were incorporated as proposed. The remaining ones were rebutted, no longer applicable, or applied with some modifications (see below for further details). Please accept our apologies in advance if we inadvertently failed to incorporate any edit not discussed below—we would gladly do so after the next round of review.

L14, 338, 422, 427, 431, 433, 507: the reviewer suggests to use “ca” instead of “ca.” as an abbreviation of “circa”. Both options are acceptable, but the latter has been maintained as per the journal's style.

L23, 597: the reviewer suggests to underline “in situ”. Indeed, this is an English word (even directly taken from Latin) that needs not be italicized (e.g., see New Oxford American Dictionary). In any case, the journal's style does not allow italics except for genus and species names anyway (“words of Latin origin that are not abbreviated should not be italicized either” according to the Guide for Authors), so we could not implement this edit.

L30: the reviewer suggests to change “site of Creu de Conill 20” into “Creu de Conill 20 site”. This was implemented, except that “locality” was used instead of “site”, since Creu de Conill 20 is best considered one among various localities of the Creu de Conill site.

L41, 56, 60, 125, 127, 141, 142, 150, 156, 159, 173, 185, 213, 215, 225, 275, 278, 279: the reviewer suggests to add a hyphen in “Crusafont Pairó” and other composite surnames, but this is at odds with the criterion followed in the manuscript to spell the surnames

- of authors as originally stated in their publications, even if they were not consistent in this regard (see also the response to reviewer 2 in this regard below).
- L46-47, 182, 214, 321: the reviewer suggests to add a space between “author” and “#”, but this is irrelevant as these expressions are only used to anonymize the manuscript for the purposes of double-blind review, and will be substituted by actual authors’ initials upon final submission if the manuscript is eventually accepted. We prefer to keep these expressions as in the original manuscript so as not to overlook them when providing the definitive version of the file with authors’ details at a later stage.
- L81: the reviewer suggests to change “considered” into “consider”. However, we prefer to maintain the past tense for all these sentences describing what previous authors published, in further agreement with “advocated” in the preceding line (not edited by the reviewer).
- L86: the reviewer suggests to change “erected” into “proposed” regarding the new species described by Alba and Moyà-Solà (2012). We consider the former a frequent term in taxonomic contexts, but following the reviewer’s concern we have substituted it by “described”, which appears to us more suitable than the alternative proposed by the reviewer (except when the taxa are conditionally proposed, in which case they are not nomenclaturally valid according to the Code).
- L91: the reviewer suggests to add “which is” but we prefer to add “which was”, since we are referring to what Alba and Moyà-Solà (2012) did (see response to L81 edit above).
- L138: the reviewer suggests to italicize “ad hoc”, but this does not apply because of the same reasons as discussed above for “in situ” (response to L23 edit).
- L139: the reviewer suggests to change “at odds with the very same definition of the Vallesian proposed by Crusafont himself” into “at odds with the same definition of the Vallesian proposed by Crusafont”. Changed into “at odds with the original definition of the Vallesian proposed by Crusafont” not to alter our original intended meaning.
- L154: the reviewer suggests to change “would have been washed down” into “was washed down”. Changed into “was purportedly washed down” in order to highlight the uncertainty in this regard.
- L222-223: the reviewer questions the need to capitalize “Alluvial Fan System” after “Castellar del Vallès” and “Upper Continental Complexes” by remarking “Are these formal names? If not, do not capitalize”. In the revised version we have used lowercase because these terms are not formally defined in the sense alluded by the reviewer (however, we added three citations, including a new reference, where these complexes and alluvial fan systems are described in greater detail).
- L236: we are unsure about one of the edits provided by the reviewer, but anyway the sentence was rephrased following also the advice from reviewer 2.
- L237-238: regarding our fragment “based on the assumption that there is no major fault located between the two sections”, the reviewer comments “perhaps not ideal

considering likely lateral variation of alluvial fan environment”. We agree, but uncertainties in the correlation because of this fact were already explicitly discussed later in the manuscript. In the revised version we have further specified that no direct correlation was possible because of dense vegetation cover (this was implicit from the fact that we had to use heavy machinery to expose the sections, but better to explicitly note it, given the reviewer’s concern in this regard; see also our response to the following comment).

L242: the reviewer asserts “assumes also that lithology is laterally the same—which is less secure a conclusion that lack of faulting”. We agree that is an added problem and we have explicitly noted it in the revised version. However, we disagree that this problem is potentially more serious than major faulting, for the latter might cause the two sections not to overlap at all, whereas this is unlikely to be the case due to local variations in accumulation rates, at least according to our experience in this and other areas of the basin with the same type of depositional setting. Indeed, in the previous version we already discussed this problem (L343), and concluded that “Given the close distance between the two sections, local differences in accumulation rates are negligible”. The sentences added in the revised version to acknowledge this and the previous concerns by the reviewer read as follows: “Such correlation methodology is far from ideal, not only because it has to assume the lack of major faulting, but also because it does not take into account lateral changes in lithology and local accumulation rates. However, such an approach was unavoidable given the dense vegetation cover between the two sections and the impossibility to deforest the whole riverbank in between.”

L252-253: The reviewer suggests to rephrase “applied up to complete demagnetization of the NRM” as “used until samples were completely demagnetized”, but we prefer a wording more similar to our original (“applied up to complete demagnetization of the samples”; see also our response to comments by reviewer 2 in this regard). The reviewer also suggest to rephrase “or observation of unstable behavior caused by neoformation of magnetite upon heating” as “or to temperatures at which unstable behavior was caused by neoformation of magnetite upon heating”, but we have rewritten it a bit differently (“or to temperatures at which acquisition of spurious magnetization caused unstable behavior”).

L262: The reference suggested by the reviewer (Kirschvink, 1980) has been added where indicated. The reviewer further asserts “see Heslop & Roberts (2016a,b) for improved way to calculate error.” In this study a large number of samples did not yield good demagnetization data at temperatures above 400 °C, as acquisition of spurious magnetization at high temperatures prevented us from observing a clean decay of the NRM. Thus, calculation of magnetic components was done by anchoring directions to the origin. The authors are aware that this procedure yields errors that lack statistical

significance. Resulting MAD values are shown in the newly added SOM Table S1, but were not used to assess the quality of the paleomagnetic directions. Quality assessment is based on visual inspection of demagnetization plots as explained in the revised text, which has been significantly expanded in this portion to give additional methodological details. Several demagnetization plots have been provided in a new figure for the main text.

L267: the reviewer suggests to add “from this temperature range” at the end of the sentence and adds “is this correct? Why not use PCA over this temperature range? If the ChRM is not directed to the origin see Heslop & Roberts (2016b) for a suitable way to assess the need (or not) to anchor solutions”. We did use standard PCA to calculate directions. In the revised version we have rephrased this sentence and, as explained above, expanded the whole paragraph to be clearer in this regard.

L280: the reviewer suggests to change “revision of the material” into “re-evaluation of the material”. This is not suitable within a taxonomic framework, so we changed it into “taxonomic revision of the material” to be more specific.

L290: the reviewer suggests to change “during a recent revision of the collections” into “in our re-evaluation”. With all due respect, collections are not re-evaluated, but revised, re-examined or restudied. We changed it into “in our recent re-examination of the large mammal collections”, because it is relevant to specify that this is not the same as the taxonomic revision of the rodent remains alluded above.

L300: the reviewer suggests to change “broken in several fragments” into “broken into several fragments”. However, this was changed into “fragmentary” following the advice of reviewer 2.

L322: the reviewer suggests to delete the comma after “fossa”, but that edit would change our original meaning, since “fossa” is the last word of a clarifying statement within commas, and the referent for the “that” that follows is not “fossa” but “deep groove” in the preceding line. In any case we have rephrased the sentence to be more clear-cut.

L359: the reviewer comments “magnetostratigraphy is a binary signal so independent constraints are critical”. We obviously agree, but are unsure as to whether the reviewer is requesting any changes in this passage. To be more clear-cut, we added a sentence to explicitly highlight the point raised by the reviewer: “This is because paleomagnetism only provides a binary signal, whose interpretation critically relies on independent constraints”.

L361: the reviewer suggests to change “being mostly based on rodents” into “and is mostly based on rodents”. We do not particularly like this alternative and rephrased it differently by providing such information earlier in the sentence and by further taking into account the suggestions by reviewer 2 in this regard (“largely” instead of “mostly”).

L364: the reviewer suggests to change “changes in the rodent faunas” into “rodent faunal changes”. However, the latter expression sounds a little weird to us and we think it is not customarily employed in the paleontological literature, so we did not implement this suggestion.

L369-370: the same as above applies to this reviewer’s suggestion to change “datum provided by the rodent assemblage” by “rodent assemblage datum”. We appreciate the reviewer’s efforts to condense our original text as much as possible, and we generally followed his suggestions, but this should not be at the expense of intelligibility. We consider the reviewer’s suggestion would alter our intended meaning and changed it into “datum provided by rodents”.

L371: the reviewer suggests to change “and thus” into “which”, but that would imply a second order subordinate sentence, both beginning with “which” (“...which precludes a correlation with the *Cricetulodon hartenbergeri* range subzone (MN9, 10.3-9.98 Ma), which indicates conclusively...”). We preferred to simply delete the comma and “thus”, i.e., “which precludes [...] and indicates...”.

L376 (and others): the reviewer suggests to abbreviate “Castell de Barberà” into “CB”, but the use of the abbreviation throughout the manuscript was dropped following the alternate advice by reviewer 2 (see later in this document).

L378: the reviewer suggests to change “even if tentatively” into “which is tentative”, but this changes our intended meaning, because we were referring to what Casanovas-Vilar et al. did, not asserting that in our current opinion the correlation is tentative. To be more clear-cut in this regard, we changed it into “even if they did so tentatively”.

L380: the reviewer suggests to change “the find of *Hippotherium* remains at CB was” into “the *Hippotherium* remains at CB were”, but we prefer “the record of *Hippotherium* at Castell de Barberà was”.

L381: the reviewer suggests to change “dismissed” into “dismiss”, but as explained above it is preferable to use past tense to describe what previous authors did or said.

L415: the reviewer suggests to change “to ascertain” into “discrimination of” when referring to the derived distinguishing features of *Hippotherium primigenium*. However, this change does not seem advisable in this context, because the problem is that available remains do not preserve the relevant anatomical regions. We therefore changed it into “to assess”.

L416: the reviewer suggests to change “a very elongate” into “an elongated” in reference to the preorbital fossa of *Hippotherium primigenium*. However, this refers to the diagnostic criteria provided in the paper cited in this sentence: Bernor et al. (1996, p. 316) used the expression “very elongate” when referring to this anatomical structure. From a taxonomical viewpoint, the difference between “elongate” and “very elongate” is relevant, even if expressed qualitatively instead of quantitatively. Note as well that Bernor et al., like our original sentence, used “elongate” instead of “elongated”. Both

spellings are correct and have indeed slightly different meanings (e.g., see New Oxford American Dictionary): “elongate” means “long in relation to width” (and has no evolutionary implications), whereas “elongated” means “unusually long in relation to its width” (and might further be interpreted as indicating an evolutionary polarity, which is not implicit in “elongate”). Given these problems, we maintained our original wording.

L441: the reviewer suggests to place “therefore” between commas, but we prefer to move it to the beginning of the sentence (“Therefore, ...”).

L447: we accepted all of the reviewer’s edits in this line, but added “this reasoning” between “and” and “has been”, because otherwise the sentence was ambiguous.

L459-460: The reviewer asserts “You need to show the stratigraphy between the sections to convince readers. You are asking readers to take a lot by faith”. We disagree from the latter assertion, as our rationale to perform the correlation is explicitly stated. In any case, in the revised version we have more explicitly recognized the limitations of this approach and further reported the stratigraphy in greater detail (see more detailed comments below). Note, however, that the latter has only been possible regarding the two studied sections, which were exposed by us with a digger machine. As explained above (and also in the revised version), it is not possible to expose the Miocene sediments located between the two sampled sections, which extend a few hundreds of meters along the riverbank (which, by the way, happens to be located within a riverine park that is densely frequented by the local inhabitants). We could not afford it financially and we would have never gotten the permits from the Catalan government and from the land owner (the city council). Therefore, we are not asking readers to take a faith leap, we are merely trying to report and interpret our results as honestly as possible. We acknowledge they might be suboptimal in some regards, and are willing to openly and explicitly note the limitations of our study in the manuscript, but the reviewer should be more understanding about the existing constraints, particularly since our paper reports the first magnetostratigraphic results for this site ever. Nevertheless, in response to the reviewer’s concerns, in the revised version we have provided more details (including a new supplementary figure) about the stratigraphy (see also our response later in this document to the more detailed comment by the reviewer in this regard).

L470-472: the reviewer asserts “OK – but since time is short & sedimentation rates are high, you could be reading polarity manifestations of tiny wiggles”. In the revised version, this possibility suggested by the reviewer has been explicitly discussed (see further details below in response to the more detailed comments provided by the reviewer).

L471: the reviewer suggests to add “lower” before “sampled normal magnetozones”, but this is unnecessary since the fact that the uppermost part of the sample sequence is excluded was already noted later within the same sentence.

L488-490: the reviewer asserts “this text is not perfectly clear—please improve”. We agree that this sentence was unnecessarily complicated and indeed it had a few mistakes that hindered interpretation. In the revised version we have deleted and rephrased this portion of the paragraph to be more clear-cut.

L503: the reviewer asserts “not really – see separate comments”, in allusion to our sentence excluding correlation with the long normal chron that is characteristic of the earliest Vallesian. This relates to the above-mentioned tiny wiggles, and has been replied in greater detail later in this document.

L507: the reviewer asserts “I am not sure you can resolve to within 100 kyr in these sediments”. We disagree, previous magnetostratigraphic work at the Vallès-Penedès Basin indicates that this is possible, particularly for time intervals where multiple reversals occur over a short time span, as it is the case of the earliest Vallesian. Previous work in similar depositional settings from the late Aragonian of the same basin further suggests that magnetostratigraphically interpolated dates are accurate to the nearest 0.1 Ma (see Alba et al., 2017, for a more elaborate reasoning). All in all, we consider sufficient to explicitly discuss the additional sources of uncertainty (the tiny wiggles) noted by the reviewer above, and refrain from further modifying this sentence, which merely explains that we did not favor this particular correlation because it would imply an earlier record of *Hippotherium* than elsewhere.

L509: the reviewer suggests changing “representing” by “which represents”, but that would change our intended meaning, for “which” would refer to Creu de Conill 20 only, whereas we referred to both Castell de Barberà and Creu de Conill 20. It has been changed into “and both sites would represent”

L531: the reviewer suggests to change “castorids” into “castorid rodents” because he deleted “As far as rodents are concerned” from the beginning of the sentence. Although we adopted the latter change, specifying that castorids are rodents is not really necessary.

L535: the reviewer suggests to change “overall indicating” by “which overall indicate”, but this would change our intended meaning by restricting the subordinate sentence to *Chalicomys* instead of both *Euroxenomys* and *Chalicomys*. We therefore opted to make a new sentence: “These taxa indicate...”.

L536: the reviewer suggests to change “permanent water bodies at the area of the site” into “permanent water bodies at the site”, but the latter is too restrictive, as we are unsure whether these water bodies were located exactly at the location of the site or somewhere nearby. We have changed it into “permanent water bodies nearby”.

L554: the reviewer suggests to change “remains of *Euroxenomys minutus*” into “*Euroxenomys minutus* remains”. While the use of genus names as adjectivated nouns is very frequent (e.g., *Euroxenomys* remains) and we have adopted all of the reviewer’s suggestions in this regard, the use of species binomina as adjectivated nouns in written form is generally advised against. Therefore, we have not implemented this suggestion.

L561: the reviewer suggests to change “provide an accurate dating” into “provide accurate dating”, but this fragment was changed instead into “provide an accurate date” as per reviewer’s 2 request. Reviewer 1 further adds: “see my separate comments about chronology”. Our response to such comments is provided later in this document.

L597-602: the reviewer asserts “I am not convinced” regarding our preferred magnetostratigraphic correlation. This has already been discussed above and more detailed responses are given below. After explicitly discussing in the revised version alternate possibilities put forward by the reviewer, we feel there is nothing much that we can do to be more convincing. We hope to have openly discussed the pros and cons of the various possibilities, and acknowledge that the reviewer comments have been most useful in this regard, but invoke our right to explicitly favor the interpretation that in our opinion is more likely in the light of currently available data. The revised version is more explicit in this regard.

L606: the two last edits provided by this reviewer in this line no longer apply after rephrasing of the sentence following the suggestions by reviewer 2.

L995-997 (caption of Fig. 5): the reviewer asserts “c, s, & l for lithology? Please spell out. What is the meaning of black circles & white circles for paleomagnetic data? Presumably sideways triangles on the strat columns indicate sampling horizons – please specify this in the caption”. This has been explained in the revised version, please see below for more detailed explanations as this comment was also provided separately by the reviewer.

With regard to the reviewer’s more detailed comments, we have copied them below in order to provide a response.

This is a clearly articulated and beautifully written paper that claims to lay to rest debates about the age of fossil finds at the Castel de Barberà (CB) location in Catalonia. I found the paper so convincing as I read it that I did not start to worry about the final result until the punchline arrived. I expected to provide a more positive evaluation than the one below. The sense is given in various places that a magnetostratigraphy provides an unambiguous age for these deposits even though the authors acknowledge that different interpretations of the polarity zonation are possible. Geomagnetic polarity is a binary signal, so independent age constraints are usually needed to tie down a set of identified polarity zones that are otherwise “floating” in time. The authors obviously understand this and argue that a clear

interpretation is still possible. I am not convinced, as argued below. At minimum, further information needs to be presented to make the case more rigorously and convincingly, but I argue below that there are ambiguities in this part of the timescale that may make the issue difficult to finally resolve with the level of certainty claimed by the authors. I have annotated the manuscript to help the authors to improve or condense the writing. Many of my annotated comments are not reproduced below, so careful attention to both will be needed. The results presented are valuable and deserve to be published in the *Journal of Human Evolution*. My reservations relate to the final chronological interpretation, which I think should be presented so that the ambiguities that I suggest below are recognized and discussed in the final interpretation.

We appreciate the reviewer's constructive criticism, but as noted above we need to highlight that unambiguity of the correlation depends on whether we are talking about land mammal ages (Aragonian vs. Vallesian), (sub)chrons, or hundreds of ka. In our original manuscript, we did recognize that several correlations where possible, yielding slightly different ages, but based on biostratigraphic data we argued that a Vallesian age was most conclusive (as explained above, this is to what we alluded in the original title). We further provided arguments to favor one of the possible correlations with the earliest Vallesian, without omitting other possible correlations. We still adhere to our previous interpretation that the site can be securely correlated to the Vallesian (instead of the Aragonian, as previously favored by most authors) and that one of the discussed correlations is more likely to alternative ones. However, following the reviewer's concerns we have discussed in greater detail the various alternate possibilities. Below we provide more detailed responses in this regard.

1. Even if the conclusions of the paper were totally sound, I would find statements such as "End of the controversy" in the title of the paper to be grandiose. It is better to let readers decide for themselves rather than to use hyperbole. I recommend removal of such statements (my annotations on the manuscript have done this).

Done.

2. The paleomagnetic results are presented a little too sparsely to enable rigorous assessment of data quality. It is usual to include a number of demagnetization diagrams with an assessment of the magnetic mineralogy responsible for the signal and brief discussion of the remanence acquisition mechanism to provide readers with the sense that this has been worked through carefully and that the reported signals are not due to remagnetizations of various types. This level of information is missing and should be added. This is what I refer to above as the minimum further information that should be added to strengthen the paper. My annotations on p. 11 are relevant to this issue, which include questions about how the characteristic remanent magnetization was determined and

whether principal component analysis solutions were anchored or not to the origin of demagnetization plots.

The reviewer annotations have already been commented above. In the revised version we have added the demagnetization diagrams requested by the reviewer as a new figure for the main text, which has also been expanded to provide additional details.

3. In addition to the above, and while extensive effort has been made to present and justify the main paleontological finds, an equivalent level of care has not been made to illustrate the stratigraphic context of the CB site. With two locations sampled, arguments are made about lithological correlation and the dip of beds, including potential faulting — but with no lithostratigraphic correlation presented. I consider this to be a critical part of the story that needs to be documented and illustrated clearly in new figures in a way that convinces readers. Most readers, including myself, will not know the local lithological units and whether they are laterally extensive or not. From the description provided, alluvial fan environments would be expected to be highly variable laterally, with local channels and variable interbeds between fluvial gravels and finer-grained units. I am left wondering how the two sampled locations might relate to each other in such a context. It may be that these concerns are not important, but I would like to be convinced of this. It is a fundamentally important issue for the chronological interpretation. I thought the discussion of possible and undocumented faulting is less important than this issue. Likewise, I would have thought that the potential for erosion of underlying strata by migrating river channels could also be important. Without a clear diagram that illustrates the stratigraphic context, stratal dip, and elevation and lithological differences between the sampled sites, the authors are asking readers to take a lot by faith. The authors do a good job of demonstrating that past work has not been adequate for various reasons, which makes it reasonable to request full documentation of site stratigraphy and inter-site lithological correlation.

As per reviewer's request, in the revised version we have added a supplementary figure with the two lithostratigraphic profiles and their suggested correlation. However, as explained above, the reviewer should be aware that the whole area is densely covered by vegetation, and that the two sections studied had to be dug with heavy machinery in order to be exposed. This severely limits the possibility to directly assess the correlation between the two sections, which are nevertheless quite close to one another. Extensive exposures in other areas of the basin (Abocador de Can Mata) with a comparable depositional context have shown that most stratigraphic levels have a high lateral continuity (see Alba et al., 2017, cited in the ms.). This has been briefly discussed in the revised version.

4. While the above issues are important, my main comment on this paper concerns the polarity interpretation. When dealing with the binary geomagnetic polarity signal, it is nice

to have a long polarity interval that can be excluded for correlation because more frequent reversals do not fit with such a pattern. Thus, interpretation of the late Aragonian and early Vallesian seems to be a cast-iron case for such a situation ... but it is not. The authors have plotted the locations of three “cryptochrons” in the long normal polarity Chron C5n.2n. The origin of cryptochrons is indicated in their name —they are cryptic. The Cenozoic polarity timescale is based on the polarity pattern observed in marine magnetic anomalies (see Cande and Kent (1992a; *JGR*) on the GPTS). The resolution of the marine magnetic anomaly record is limited by the smallest resolvable anomalies, which are referred to as ‘tiny wiggles’ (see Cande and Kent (1992b; *JGR*) on ‘tiny wiggles’). These small anomalies with durations <30 kyr are referred to as cryptochrons because their origin is less clear. Three origins have been proposed in the literature for cryptochrons, where they could represent: (1) short-duration polarity zones; (2) shorter geomagnetic excursions (typically a few kyr in duration); and (3) largescale changes in geomagnetic dipole field intensity without an accompanying field reversal. Relatively little work has been done on the origin of tiny wiggles. I have suggested that detailed magnetostratigraphy of rapidly deposited sediments is the best way to test these alternatives (Roberts and Lewin-Harris, 2000; *EPSL*), and provided a direct magnetostratigraphic test of the origin of tiny wiggles in the same time interval of relevance to the present study, Chron C5n.2n, and concluded that some are likely to represent short polarity chrons and some probably represent geomagnetic excursions. This interpretation has been contested, but it is not unique. Several papers that I cited in 2000 reached similar conclusions for Chron C5n.2n and more have probably been produced in the succeeding years (Garcés et al., 1996; Li et al., 1997; Rössler and Appel, 1998; Roperch et al., 1999). Garcés et al. (1996) studied sediments of this age in the Vallès-Penedès Basin, as cited in the reference list, so this possible interpretation must be known by the authors and has been dismissed without discussion. My difficulty in accepting the proposed interpretation is that the time represented by the CB section is short and sedimentation rates are high, so the sediments could easily be recording short polarity intervals or geomagnetic excursions associated with the “tiny wiggles” in Chron C5n.2n. All of this adds ambiguity rather than clarity to the interpretation, which makes me sceptical of the authors’ claims that they can resolve the chronology of the CB section to within 100 kyr. I wish it were so, but I am not convinced. I am not sure that this issue is resolvable. I think it represents an ambiguity that must be considered when working with this part of the timescale in rapidly deposited sequences. I think the best way to deal with this is honestly by recognising that additional magnetostratigraphic interpretations are possible and that the CB location cannot be dated with the precision stated with the information presented.

We thank the reviewer for such detailed explanations, but we feel that, while doing so, he is overemphasizing the likelihood of alternate interpretations based on the tiny wiggles (which, after all, would equally imply a Vallesian age for Castell de Barberà!). We acknowledge that we might have been too categorical in some portions of the manuscript,

such as for example the Abstract, where due to space constraints we did not mention other possible correlations. The revised version has been modified accordingly to be more explicit in such places about other alternate correlations, but it needs to be stressed that in the original version we already discussed several alternatives. For example, in the Discussion: L457-462: “The magnetostratigraphic data reported in this paper further enable a more accurate dating of CB within the early Vallesian, even if the shortness of the sampled sections allow for several possible interpretations. The most parsimonious interpretation, based exclusively on stratigraphic distance between the two sections as computed from dip measurements, implies as much as six different magnetozones. Other interpretations imply four to five different magnetozones.” We continued to argue that even under alternate correlations, the sampled sections were unlikely to represent more than 400 kyr, so that it was possible to exclude a correlation with C5n.2n. The reviewer’s criticism above does not apply to the need to recognize different possible correlations between the two sampled sections (which we already did), but rather only refers to our dismissal of a correlation with C5n.2n, which the reviewer think is feasible due to the existence of various cryptochrons. In the revised version, we adhere to our former preferred interpretation . However, instead of dismissing an alternate correlation with C5n.2n without further consideration, we have rephrased the relevant paragraph and added a new paragraph below to explicitly discuss such possibility and explain the reasons why we consider it unlikely. This has required adding some references. The new paragraph introduced in the revised version is reproduced below in italics:

An alternate correlation of the reversed polarity magnetozones R2 and R3 with some cryptochrons or geomagnetic excursions within C5n.2n (Cande and Kent, 1992; Roberts and Lewin-Harris, 2000; Evans et al., 2007) is considered unlikely because of the short duration, possibly less than 10 kyr, of these events. While chances of recording such short geomagnetic features in fluvial sediments are typically low, to record as many reversed polarity directions as normal polarity directions within C5n.2n is considered implausible. The fact that all the reversed polarity magnetozones documented at Castell de Barberà are recorded by more than a single paleomagnetic sample strongly argues against any of them representing a short-lived excursion event within C5n.2n. The similar thickness of the documented normal and reversed polarity magnetozones at Castell de Barberà strongly argues against this possibility, and favor a correlation of most of the sequence with chron C5r.

Further details about ‘tiny wiggles’ have been added to the caption of former Fig. 6, which has been partly redrawn to add the ‘excursion chrons’ identified by Evans et al. (2007). We have further rewritten a bit several sentences of the abstract to be less categorical about the age while making it clear that the controversy that we consider settled is that of the Aragonian vs. Vallesian age of the site.

5. Various minor information additions are needed for the figure captions. For Figure 2, please state which section was excavated. For Figure 5, please add descriptions for the c, s, and l lithological abbreviations. Also, please spell out the meaning of the black and white circles for paleomagnetic data. Presumably the sideways triangles on the stratigraphic column indicate sampling horizons — please specify this in the caption.

We assume that the reviewer refers to panel B in Figure 2, which read as “Detail of CB during the excavation of layer D [...]”; earlier in the caption, we defined CB as “the classical site of CB”, so we thought it was clear (at least looking at panel A) that CB corresponds to Section 2. Nevertheless, following the reviewer’s request we specified the section for the sake of clarity.

With regard to Figure 5, we added the lithological abbreviations that were inadvertently omitted in the previous version (c = conglomerates; s = sandstones; l = lutites). We also specified the meaning of triangles and white circles in the caption of this figure.

Reviewer 2

There are a lot of comments here, but most are quite minor and I think can be addressed without too much difficulty. I enjoyed reading this paper. In my opinion the work is suitable in quality and scope for the journal, and makes an important contribution to our understanding of an interesting catarrhine-bearing fossil site.

We thank the reviewer for the positive overall evaluation and the constructive criticism, see below for more detailed responses.

Major Comments:

There are some issues with the chronostratigraphic terminology being used in this manuscript. The ‘Aragonian’ is not a commonly understood term outside of Iberian Peninsula paleontology, and absolutely needs some explanation and justification for this audience. For this journal, it really might be better to stick to the recognized continent-wide ELMA bins (so Astaracian instead of Aragonian). Even better than that would be to employ the ICS (International Commission on Stratigraphy) names, which are the most widely recognized. Part of my concern is that the Aragonian covers such a long span of time, so it ends up being a rather less specific term than other, contemporary units that are the same age as the ‘late Aragonian’. Of course, I understand that the Aragonian does have a particular meaning, and that you may prefer to use it here, at least some of the time. In that case you must properly cite its origin, the precise meaning, and its specific relationship to these other better-known terms. I found the explanation in Steininger 1999 (p. 11) quite helpful. A brief explanation (like that by Steininger) explaining that the term was coined by Daams et al., and is generally used only in southwest Europe, would help. It’s worth noting

that the legend in Figure 3 uses the ICS names, so you need to at minimum clarify the relationship between those terms and the terms you use in the text. Also, line 116 introduces the term Vindobonian – is this strictly necessary? If so, then again, it needs explanation.

We know that the Aragonian is not generally used outside the Iberian Peninsula, but this is because the former has been defined as a regional unit restricted to Spain. In contrast, the Astaracian is not used in Iberia and to our knowledge it has not been formally defined anywhere (unlike the Aragonian). We have been using the latter term for many years in many publications without encountering any opposition from reviewers, so when preparing the original submission we did not even consider the need to provide further explanations. Following the reviewer’s concerns in this regard, we have clarified the meaning of the Aragonian and our reasons to favor this term (as well as the use of land mammal ages more generally), since our views clearly differ from those exposed in the reference cited by the reviewer (Steininger, 1999), whose alternate proposal is not exempt of problems. Given that a large portion of the readership of the journal will not be familiar with the controversies surrounding the Aragonian and other European land mammal ages, in the revised version we have not restricted ourselves to a short clarifying statement and have provided a new subsection of the Introduction (1.2. Vallesian and Aragonian European land mammal ages), with most of the text from former subsection 1.2 (The controversial age of Castell de Barberà) renumbered as 1.3. The reviewer’s concerns in this regard have made us realize that insufficient background was provided in the original version regarding these terms, which are central to the discussion, since our manuscript aims to resolve the longstanding controversy of whether Castell de Barberà is Aragonian or Vallesian in age. We therefore think that this is a welcome addition with regard to the international readership of the journal, most of which will not be familiar enough with these terms and even with the controversies and problems associated with the definition of land mammal ages (as rightly pointed out by the reviewer). Nevertheless, we remain open to editorial suggestions regarding the need to shorten this new passage after the next round of review. The added fragment has been reproduced below in italics:

[...] has been hampered by a longstanding controversy about the age of this site— including whether it correlates to the Aragonian or with the Vallesian European land mammal ages (ELMAs). Some clarification is required with regard to the concept of ELMAs (for a historical review, see Lindsay and Tedford, 1990 and Van Dam, 2003), with particular emphasis on the use of ‘Aragonian’ instead of Orleanian + Astaracian, despite the latter are more widely used outside Spain. The use of ELMAs as geochronological units has been criticized (e.g., Steininger, 1999) on the grounds that they are regional and that, with few exceptions, they do not correspond to properly defined chronostratigraphic units (stages). ‘Mammal ages’ are generally conceptualized as “biochronologic units” (e.g.,

Woodburne, 2004a: xiv; see also Hilgen et al., 2012) of regional applicability due to divergent paleobiogeographic histories among regions (Lindsay and Tedford, 1990). However, an initial definition of land mammal ages as biochronologic units (i.e., biozones) is not mutually exclusive with their subsequent formal definition (based on bio- and magnetostratigraphic data) as chronostratigraphic units (i.e., stages, based on bodies of rock formed during a given time interval; e.g., Garcés, 1995; Woodburne, 2004), which automatically implies the definition of their corresponding geochronologic units (ages). The definition of regional chronostratigraphic units is not at odds with the International Stratigraphic Guide, because “It is better to refer strata to local or regional units with accuracy and precision rather than to strain beyond the current limits of time correlation in assigning these strata to units of a global scale” (Murphy and Salvador, 1999:267). Even if not directly correlated with marine stages, regional continental units can be dated based on radiometric and/or paleomagnetic methods (e.g., Krijgsman et al., 1994; Garcés et al., 1996). Steininger (1999) conversely advocated abandoning a chronostratigraphic/geochronologic concept of ELMAs altogether and proposed to replace them by entirely biostratigraphic units (‘European land mammal mega-zones’) based on MN (Mammal Neogene) biozones (Mein, 1975). This proposal is not exempt of problems given that MN zones were defined as informal biochronologic units (Mein, 1975) and that their utility at a continental-wide scale is restricted due to the significant diachrony of most mammal biochronologic events (van der Meulen et al., 2011, 2012). This has led some authors to contend that a formal European biozonation is not possible (van der Meulen et al., 2012), which would imply that the regional nature of formally defined ELMAs such as the Aragonian (see review in van der Meulen et al., 2012) would rather be an advantage with regard to providing accurate correlations.

The lack of formal (biostratigraphic and chronostratigraphic) definition of widely used biochronological units such as Orleanian and Astaracian is a problem that does not apply to the roughly time-equivalent Aragonian, since like the younger Vallesian (which enjoys a wider geographic applicability) it has been formally defined on the basis of a specific stratotype. Crusafont Pairó (1950, 1951, 1953, 1955; see also Crusafont Pairó and Truyols Santonja, 1954, 1959) first used the term Vallesian in a largely biochronological sense to designate the Vallès-Penedès deposits with Hipparion (currently Hippotherium) that, based on the fauna, appeared intermediate in age between the sites of La Grive in France (i.e., the Aragonian) and Pikermi in Greece (i.e., the Turolian). The Vallesian was finally more formally defined in reference to Vallès-Penedès mammal successions by Crusafont Pairó and Truyols Santonja (1960) based on the entry of Hippotherium as its main defining criterion. The Vallesian was rapidly accepted throughout Eurasia as a simple solution for the complex stratigraphic terminology in use, although it was not until decades later that it was formally defined as a ‘mammal stage’ based on a specific stratotype from the type area (the Vallès-Penedès Basin) within an accurate bio- and magnetostratigraphic

framework (Garcés, 1995; Garcés et al., 1996; Agustí et al., 1997). Following the proposal of MN zones by Mein (1975), MN9 and MN10 had already been equated with the early and late Vallesian, respectively, in turn subdivided into multiple local biozones based on rodents (Agustí, 1981, 1982; Agustí and Moyà-Solà, 1991; Agustí et al., 1997; Casanovas-Vilar et al., 2011a, 2016b).

The Aragonian was in turn originally conceptualized (Falhbusch, 1976) as a chronostratigraphic unit defined by the presence of the equid *Anchitherium* and the lack of the more derived equid *Hipparion* (currently *Hippotherium*), and subdivided into two subunits, the Orleanian and the Astaracian. Soon thereafter, the Aragonian was defined by Daams et al. (1977) as a new stage for continental middle Miocene deposits preceding the Vallesian, with its stratotype located within the Calatayud-Montalbán Basin in Spain (see also Daams et al., 1999). However, Daams et al. (1977) refrained from dividing the Aragonian into Orleanian and Astaracian because the latter had yet to be formally defined, ultimately leading to a tripartite subdivision (Daams and Freudenthal, 1981; Daams and Freudenthal, 1990). Moreover, the original criterion used to define the base of the Aragonian (the dispersal of *Anchitherium*) was soon questioned (Daams and Freudenthal, 1981) and eventually abandoned following the definition of the Ramblian stage (Daams et al., 1987; see also Daams and Freudenthal, 1990). Further stratigraphic refinements of the Aragonian were later provided by Daams et al. (1987, 1999) and van der Meulen et al. (2012), who further distinguished multiple Aragonian local biozones based on rodents. Based on these works, currently the Aragonian may be considered a regional mammal-based chronostratigraphic unit (stage), whose scope is limited to the continental record from Spain (see recent reviews in van der Meulen et al., 2012 and García-Paredes et al., 2016). Given the scope of this paper and the detailed local zonation of the Aragonian available for the Vallès-Penedès Basin (Casanovas-Vilar et al., 2016b), we refrain from using the alternative and more loosely-defined term 'Astaracian'.

Based on high-resolution magnetostratigraphic correlation to the geomagnetic polarity time scale (GPTS), according to Van der Meulen et al. (2012) the early Aragonian (ca. 17.2–15.9 Ma) corresponds to the late Burdigalian (early Miocene, MN4), while the middle Aragonian (ca. 15.9–13.8 Ma) comprises most of the Langhian (early to middle Miocene, MN5), and the late Aragonian (ca. 13.8–11.2 Ma, roughly equivalent to the 'Astaracian') covers the latest Langhian, the Serravallian, and the earliest Tortonian (middle to late Miocene, MN6 to MN7+8). The Vallesian, in turn, entirely corresponds to the Tortonian (late Miocene), being subdivided into early Vallesian (11.2–10.0 Ma, MN9) and late Vallesian (10.0–8.9 Ma, MN10; e.g., Hilgen et al., 2012; Casanovas-Vilar et al., 2016b).

With regard to other recommendations by the reviewer, as it is clear from the text above we still generally refrain from using the ICS stage names instead of the ELMAs, as the former are not customarily used in works dealing with continental sediments and fossil

sites. The only reason we used the latter terms in the legend of Figure 3 is because this figure is a geological map that further includes marine sediments. Nevertheless, as it can be shown in the fragment reproduced above, in the revised version we have specified the correlation of the Aragonian and Vallesian subdivisions with the ICS marine stage names. Finally, in the original manuscript the term Vindobonian was used only once (and within quotation marks) because we were referring to the works of previous authors that “advocated for a ‘Vindobonian’ (currently, late Aragonian) instead of Vallesian age for CB” (these authors did not employ the term Aragonian because it had not been defined yet). Following the reviewer’s concern in this regard, in the revised version, we have simplified this sentence (“advocated for a pre-Vallesian age for CB”) in order to avoid introducing unnecessary confusion because of this outdated term.

Throughout the manuscript, three different names are being used to refer to your sections – section 1, section ‘s.l’, and CB1; and section 2, section ‘s.s’ and CB2. This seems unnecessarily complicated. If the problem is that previous publications or abstracts have used one name, and now you prefer another, then you need to just explain that. Then pick one name and stick with it throughout this manuscript. Personally, I like CB1 and CB2 – very simple and easy to understand. If you’re going to use ‘section 1’ and ‘section 2’, then I would suggest treating these as proper nouns and capitalizing (Section 1 and Section 2). **Note that throughout the previous version of the manuscript we mostly used section 1 and section 2, and we prefer to keep these names for the sections, in order to avoid confusion with the nomenclature used for the corresponding fossil localities. Therefore, in the revised version we have followed the reviewer’s suggestion to capitalize ‘Section 1’ and ‘Section 2’. CB1 and CB2 were only used in former Figure 6, and this has been changed by ‘Section 1’ and ‘Section 2’ as well, since CB1 and CB2 resemble acronyms used for fossil localities in the Vallès-Penedès Basin. As for s.l. and s.s., these abbreviations were included to denote the equivalence of the sections with the fossil sites, and were only mentioned twice in the main text and also in Figures 5 (both image and caption) and 6 (only caption). We have omitted these terms from the figures, but we would like to keep them somewhere in the revised manuscript to clearly state the equivalences with the fossil sites (since these terms have been used in field reports and also in the collections of the institution where the material is curated). To be more clear-cut in this regard, we have added further explanation in the supplementary material of the revised manuscript.**

Your newly recovered *Hippotherium* humerus is said to come from level ‘D’, which is also said to be the same level that the material collected in the 1960s and 1970s comes from. But on line 197 you say that ‘the main fossiliferous level’ was exhausted and that the fossils collected came from somewhere else. Please clarify. Calling your fossil level ‘layer D’ also begs the question of what other fossiliferous layers you found. What were A, B, and C?! Are

they worth explaining? Please note that you also use both the terms 'layer D' and 'level D'; you should use one, and it would also probably be helpful to treat it as a proper noun as well – Level/Layer D.

The exact quote from the previous version of the manuscript alluded by the reviewer is “our works confirmed that the bone accumulation of main fossiliferous level of CB is exhausted”. Therefore, we did not assert that the main fossiliferous layer was exhausted, only that the bone accumulation of this layer was so. This might seem a trivial distinction but it explains the apparent incongruence noted by the reviewer. Isolated fossil finds occur throughout the section, including this very same layer. To be more clear-cut in this regard, to our former fragment “enabled the recovery of additional fossil remains” (now modified as “we recovered additional fossils remains”, following the edits by reviewer 1) we have added “from the same layer and other layers”. As asserted above, in the supplementary material we have clarified how the various fossiliferous layers (or localities) within each site were termed, and have further referred to a new supplementary figure in which we report the stratigraphy in further detail (including the name of the various layers). The uppercase letters alluded by the reviewer denote different fossil localities within each site. Besides layer D, in the footnote to Table 1 we explained that layer B was situated 2 m below D and that E was located 1 m above D. However, we agree that further explanation is required. The new supplementary information that we have added in this regard clarify all these questions. The terms 'level' and 'layer' were used interchangeably as synonyms in the previous version. In the revised version, we have favored the use of the latter term. We did not capitalize this term when followed by an uppercase letter denoting a particular layer (e.g., Layer D), as suggested by the reviewer, because this expression has been substituted by CB-D in the revised manuscript.

I think the Materials and Methods section needs some reorganization. I'd suggest having some introductory material that explains the site location, section locations, and basic stratigraphy first, then separate sections (2.1 and 2.2) to explain the methods for biochronology and paleomag. I would think that it would make more sense to explain the paleomag methods first, and then the biochronology (particularly because as written, line 205 sounds like it should come after the section on paleomag). The entire first paragraph of this section (lines 205-220) would be clearer if it were reorganized – location of site, then basic stratigraphy before trying to make the argument that you can know that your layer D is the same as the original layer.

We have reorganized the Materials and methods section following the suggestions by the reviewer. Note, however, that the paragraph of L205-220 includes the portion about biostratigraphy that the reviewer requests to move after the paleomag methods. Therefore, only the sentences from L205-209 were moved there, whereas the rest of the

paragraph (L209-220) was kept in the first subsection but moved after the paragraph indicating the geological background and depositional setting of the site as known from the previous literature. The new paragraph that explains the stratigraphy and the definition of fossil localities (see above) was added to the end of the first subsection. To sum up, in the revised version the Materials and methods are divided into subsections 2.1 (“Location and stratigraphy of Castell de Barberà”) and 2.2 (“Paleomagnetism and biostratigraphy”). Note that in this particular context we prefer the term ‘biostratigraphy’ rather than that of ‘biochronology’ (as suggested by the reviewer), given that we are using the fauna to discuss the dating of the sediments, instead of discussing the correlation between different bioevents.

Despite your assertions that ‘most’ of the original material came from a single layer (line 213), there really does seem to be legitimate reasons for uncertainty as to the provenance of the original *Hippotherium* material (e.g. the existence of ‘layer 1’ and ‘higher layer’). Assuming that we really cannot know where those two specific dental specimens came from, your argument really rests largely on your new humerus. Given that, I would suggest adding a comparative photo of an anchitheriine equid to Figure 4 so that the reader can clearly see the two anatomical differences you describe. Trochlea height should be fairly obvious to any informed reader, but the origin of the extensor digitorum communis muscle should perhaps be highlighted with an arrow or dashed line so that it is clear. Note that you refer to this feature as both a ‘groove’ (line 321) and a ‘crest’ (line 325), which should be clarified.

In the original version we described the structure alluded by the reviewer as (L321-322) “a narrow but deep groove, medial to the lateral epicondyle and separated by the lateral surface of the radial fossa”. In the revised version, we have provided a new version of former Figure 4 with better photographs of the Castell de Barberà specimen (after reintegrating various fragments). The use of ‘crest’ instead of ‘groove’ was an inadvertent mistake that has been corrected in the revised version, which by the way has been rephrased at this passage to be more clear. With regard to the addition of an anchitheriine humerus, we have added a new supplementary figure in which the distal articular morphology of Castell de Barberà distal is compared with that of *Anchitherium*. An arrow has been added in the former specimen to denote the localition of the discussed anatomical structure.

Even though we agree that the reviewer’s suggestions for this figure represent a welcome addition, with regard to the reviewer’s comment as a whole we would like to stress that, in the previous manuscript, we did explicitly recognize the uncertainties about the exact stratigraphic provenance of the *Hippotherium* material from the classical collections, but never intended to question their provenance from Castell de Barberà (as the reviewer seems to imply). We further disagree to some extent with the reviewer’s assertion that

“your argument really rests largely on your new humerus”. This new specimen certainly makes the discussion about the provenance of the previous finds rather irrelevant, but as we already highlighted in the original version (L385-388) “Even if both teeth came from a stratigraphically higher layer, these remains would indicate, at the very least (given the thickness of the outcropping section) that the main fossiliferous level would be situated less than 10 m below unambiguously Vallesian levels.” Given average sedimentation rates for the basin of ca. 20 cm/kyr (see references in the manuscript), that would represent around 50 kyr below contrasted Vallesian levels, thereby constraining to a great extent the possible magnetostratigraphic correlations of the sampled magnetozones.

Furthermore, your correlation of the original ‘layer 1’ from the 1965-1981 collections to your ‘layer D’ seems to be based solely on the personal communication of JV Santafé. That’s potentially not a problem, but I would be better to explicitly explain how s/he did it. Did s/he come to the field? Were there notes or old photos that were consulted? This would help to persuade readers that you really do have the same layer.

Yes, Santafé came to the field as already indicated in our previous version (L193-195): “was located thanks to the collaboration of the late Josep V. Santafé, who had repeatedly excavated the site during the 1970s with Crusafont, and indicated the exact location while visiting the site during the 2015 season”. We were unable to find contextual photographs of the site that could help to locate it and the published notes about its location, as explained in the manuscript, were too vague. It must be stressed that one of the authors of the present manuscript also visited and briefly excavated the site in the late 1970s, but was not able to remember the location of the spot until Santafé came to the field. The latter had intensively worked at the site, and when he arrived he went straight to the spot and told us where to excavate. Once we removed the overlying soil and plants, a layer was uncovered that displays color and other sedimentological features that are in agreement not only with the remembrances of both Santafé and the coauthor of the paper, but also with sediment fragments preserved in the collections attached to some fossils from this site. Santafé confirmed that they stopped excavating the site because the bone accumulation from the main layer had been exhausted, but this does not mean that further fossils cannot be found, as this had been always the case for the whole section. Given that we already explained that Santafé came to the field in 2015, we do not think it is necessary to provide further explanations, but could do so if the editor and/or the reviewer consider it necessary.

Your stratigraphic layers should be shown in Figure 5.

As explained above, we have provided a new supplementary figure to provide more details about the stratigraphy, given that Figure 5 is focused on the paleomag results and the size of the profiles does not enable for additional details. Note that in the new figure

we have only named the layers that yielded fossil remains, as doing otherwise would represent an unnecessary complication.

If you reorganize M&M to go through the paleomag methods first, then I would suggest the same organization for the Results section.

Done. The subsections have been renumbered accordingly.

This may be my preference, and you should check with the AE, but I think that some material currently in 'Discussion' could be moved to 'Results'. For example, some of the basic chronostratigraphic interpretations of the rodents (starting on Line 357) could be moved.

We disagree. In our opinion, the paragraph beginning at L357 of the previous manuscript better fits the Discussion, as it provides the required context to understand the main implications of the rodent assemblage from the site. Only the passage beginning at L369 ("In the case of CB, the most significant datum provided by the rodent assemblage...") could be moved by the Results, but it would not make much sense without the rest of the paragraph and it should be restated again in the Discussion. So unless the editor has a different view, we prefer not to modify the Discussion in this regard.

Lines 414-417. If that's the case, then you should delete most of this paragraph. It's really not relevant to your paper to provide an explanation of taxonomy within *Hippotherium* if you aren't going to assign the humerus to a species. This is a good place to shorten up the rather lengthy discussion on this topic. I wonder if all of the following paragraph (lines 418-434) is necessary either.

The reviewer alludes to our sentence "we refrain from attributing the CB scanty specimens to this species [*Hippotherium (primigenium) catalaunicum*], because they do not allow to ascertain its purportedly derived distinguishing features from *Hippotherium primigenium* s.s." and suggests to delete most of the paragraph. We disagree from the reviewer, in the sense that s/he considers the discussion about the taxonomy irrelevant because we provide no assignment to species, whereas we consider it relevant precisely to justify our inability to provide an assignment to species. Furthermore, this paragraph is required to explain our assignment for other remains of the genus from the Vallesian of the Vallès-Penedès Basin, which are assigned to *H. catalaunicum* instead of *H. primigenium*. Given the controversies about hipparionin taxonomy, we prefer to be clear-cut in this regard. Similarly, we also would like to keep the paragraph that follows the sentence alluded by the reviewer. Given that our conclusions indicate that CB is probably roughly coeval with the earliest occurrence of *Hippotherium* in the Vallès-Penedès Basin at Creu de Conill 20, the fact that the latter represent one of the few well-dated earliest occurrences of this taxon in Western Europe, and the great biostratigraphic importance of

the taxon for defining the Vallesian, we feel that contextualizing our finds from these viewpoints is warranted. We nevertheless remain open to suggestions from the editor if he considers that the manuscript needs to be shortened.

Minor Comments:

The term 'classical' is heavily overused in the manuscript, and isn't a great translation into English. In many places the term 'original' could be substituted (as in 'original outcrop'). When it refers to the material excavated in the original phase of research from 1965-1981, you could explicitly explain that, or refer to it as "previously collected material" (e.g. line 44-45) or "museum collections" (line 64). Once you've explained that your 'level D/layer D' is the same as the 'classical fossiliferous level', then you can probably just call it Layer D.

We agree, and therefore changed all occurrences of 'classical' using the various alternatives suggested by the reviewer or even deleting the word when it was not necessary. As explained above, Layer D was most of the time substituted by CB-D.

Throughout – this may relate to how other previous publications have treated Spanish surnames, but the hyphen use for several names is inconsistent. For example, what I believe is a single person is sometimes referred to as just "Crusafont", sometimes as "Crusafont-Pairó", and sometimes as "Crusafont Pairó" (no hyphen). Other impacted names include Golpe (also Golpe-Posse and Golpe Posse) and Santafé (Santafé Llopis, Santafé i Llopis). To the extent possible, these should be standardized.

The reviewer is right, except that the use of these surnames is consistent with that employed by these authors in the original publications. People have two surnames in Spain, and sometimes use only one of them, sometimes both, sometimes unite them with a hyphen, and sometimes with an 'i' (in Catalonia). The original authors are to be blamed for not having consistently spelled their names, but we think it is preferable to be consistent with the published literature, and this is indeed the criterion that I (the first author) have used in all my publications for many years. Therefore, no changes were introduced in this regard.

Does it really help to abbreviate the site name to 'CB' in the text? It's done inconsistently (particularly in figure captions), and doesn't really make the paper easier to read. I might just use the full site name throughout the manuscript (except of course on figures themselves – e.g. using 'CB1' to refer to the section). If you don't abbreviate Castell de Barbarà, then it's probably not necessary to abbreviate CCN20 or ACM either [note that ACM is not explained in the text].

CB was spelled out in each figure caption because in principle figures should stand on their own grounds and it is better not to use undefined abbreviations (even if explained in

the text). With a few exceptions, the CB abbreviation was consistently used elsewhere in the text, like that of CCN, so that only ACM needed to be spelled out (this was an inadvertent mistake). However, following the reviewer's request we have mostly abandoned the use of site abbreviations in the revised manuscript. We have only specified once within parentheses the abbreviation for Creu de Conill 20 (for it is widely employed in the literature), and have also specified the abbreviations for the various Castell de Barberà localities that yielded fossil remains in 2014-2015 in the supplementary material. Therefore, we have not employed CB to refer to Castell de Barberà, but have used CB-D to refer to Layer D. Note that, as a result of this change, the Abstract would have been lengthened slightly above generally permitted extension of 300 words. However, this is not the case thanks to the edits provided for the abstract by reviewer 1.

Lines 129-136. I'm struggling to understand the logic here. First, who are the 'latter authors'? If it's Crusafont-Pairó and Golpe-Posse (1974), then I don't understand how they could have 'only contemplated two possible explanations', because they are the ones who thought the fossil came from a higher level (so that would be three explanations). If, instead, it was Santafé Llopis 1978 who suggested the fossil came from a higher level, then that isn't the 'latter author' (nor is it 'most authors' as reported on line 381). In this scenario, it might be worth clarifying that Crusafont-Pairó and Golpe-Posse (1974) accepted that the *Hippotherium* fossil came from the main fossiliferous level before outlining their interpretations. Furthermore, their second interpretation (that the boundary was located 'along the upper portion' of the section, doesn't seem to logically follow. I think these lines could use some clarification.

We have tried to clarify this passage, for example, by specifying that the "latter authors" are indeed Crusafont-Pairó and Golpe-Posse (1974). The reference by Santafé Llopis (1978) was only to back up the location of the main fossiliferous layer at about midheight of the ca. 20 m-thick stratigraphic section, but in the revised version we have omitted it, as this was also stated by the former authors. Note that Crusafont-Pairó and Golpe-Posse (1974) considered it likely that the *Hippotherium* fragment came from the upper layer, but they were not sure about it, which is the source of their two different interpretations. We nevertheless understand that the reviewer has problems understanding the logic of Crusafont-Pairó and Golpe-Posse (1974), because it is quite contorted and stems from Crusafont's a priori assumption that Castell de Barberà must undoubtedly predate the Vallesian (as explicitly stated in their paper). We reproduce below the relevant fragments from Crusafont-Pairó and Golpe-Posse (1974; our translation from the Spanish original), so the reviewer can judge by him/herself whether our summary of their interpretation is faithful to the original or not:

"...the finding of a somewhat rounded fragment (the only one among thousands of specimens) of *Hipparion*. [...] The bank of the Ripoll River, where the considered site

[Castell de Barberà] is located, has about 20 m of depth. The fossiliferous layer, undoubtedly Vindobonian, is located at about midheight. However, on the other hand, there is another layer closer to the top of the wall border, from where, probably, the aforementioned *Hipparion* fragment come from. [...] the upper layer, from which, on the other hand, no further *Hipparion* remains have been found, is still Vindobonian and represents the first appearance, very shy, of Christol's genus? Or, in contrast, the lower layer is the top of the Vindobonian and the upper one is already Vallesian? This is a very difficult problem that could not be solved except by means of better findings from the upper layer. It is not possible that the fragment comes from layers considered by us as Vallesian (Can Llobateres I, Can Llobateres II, Santiga, etc.), given the presence of a forested area in the upper portion of the riverbank, with dense vegetation that covers a relatively extensive area”.

We have expanded our sentences in this regard to be more clear-cut, as we acknowledge that our original fragment was too succinct and could led to misinterpretations. The revised version reads as follows: “Crusafont-Pairó and Golpe-Posse (1974) considered that the main fossiliferous layer was undoubtedly pre-Vallesian and only contemplated two possible explanations for the presence of *Hippotherium* at the site: either *Hippotherium* arrived in the Vallès-Penedès Basin during the latest Aragonian (irrespective of whether the aforementioned fragment came from the main or the upper fossiliferous layers); or the Aragonian/Vallesian boundary was located above the main fossiliferous level along the upper portion of the Castell de Barberà section (assuming the *Hippotherium* fragment came from the upper layer)”

Note that we talk about *Hippotherium* and Aragonian instead of *Hipparion* and Vindobonian, as in the original publication, to avoid introducing unnecessary confusion.

P. 6. Regarding the original *Hippotherium* material reported by Crusafont-Pairó and Golpe-Posse (1974), it would be worth clarifying what this specimen was. Was it one of the dental specimens mentioned later in the text?

As already asserted later in our original manuscript (L286-287, subsection 3.2., “Presence of *Hippotherium*”; currently renumbered as subsection 3.3.), “It is uncertain on the basis of what specimen Crusafont-Pairó and Golpe-Posse (1974) originally reported the presence of *Hipparion* (currently, *Hippotherium*) at CB”. We continued to explain that two dental specimens of this taxon (once already reported by Rotgers and Alba, 2011, the other never figured or described) are available from the collections of the site. Crusafont-Pairó and Golpe-Posse (1974) did not report a collection number, describe or figure the purported *Hippotherium* fossil from Castell de Barberà, and merely referred to it as “a fragment” without even specifying whereas it was dental or postcranial, but merely noting that it was “somewhat rounded”. The lower molar figured by Rotgers and Alba (2011) is well preserved and therefore unlikely to correspond to the original report by

Crusafont-Pairó and Golpe-Posse (1974). In contrast, it might apply to the deciduous premolar germ newly mentioned in our manuscript, although it is impossible to be sure. This has been noted, albeit with doubts, in the revised manuscript (although not in the section alluded by the reviewer, but after the report of the two hipparionin teeth from the site): “Crusafont-Pairó and Golpe-Posse (1974) did not figure or describe the specimen attributed by them to *Hippotherium*, but merely referred to it as a somewhat rounded fragment—which might apply to the above-mentioned deciduous premolar, although it is not possible to be certain.”

P. 7. The reader requires some explanation of the age of Can Llobateres in order to understand the significance of their close stratigraphic distance.

Done, we have specified that we referred to Can Llobateres 1 and provided an updated interpolated age for the locality with the corresponding reference.

Lines 184-188. You make it sound as if there was some huge delay, but if the fieldwork was planned for 2014, and some fieldwork really happened in 2014, then this can probably be omitted. Also, it would also probably suffice to say that the modifications of the river bank were to build ‘a road’.

The reviewer alludes to our sentence: “This fieldwork initiative was hampered by the vagueness of published indications about the site’s exact location (Crusafont-Pairó and Golpe, 1972; Golpe Posse, 1974), coupled with decades of vegetation growth and recent anthropic modifications of the river bank (to build a roadway in the framework of a riverine park).” Indeed there was a considerable delay, because the indicated reasons apparently precluded Miguel Garcés from sampling the site for his dissertation (1995) on the magnetostratigraphy of the Vallès Sector of the Vallès-Penedès Basin (see also Garcés et al., 1996, cited in the text). Nevertheless, in this sentence we did not want to imply a temporal delay, but merely refer to the difficulties to locate the site to begin with. In the revised version we have therefore rephrased the sentence as: “Finding the exact location of the site was hampered by the vagueness of published indications ...”. Following the reviewer’s suggestion, we have also omitted allusion to the riverine park, but have rephrase this portion of the sentence as “recent anthropic modifications (to enlarge the riverbank’s trackway)”, which in our opinion better describes the situation and is about as short as the alternative suggested by the reviewer “recent anthropic modifications of the river bank (to build a road)”

Lines 290-291. The presence of a second specimen in the collections is not a valid reason to accept the provenance of IPS57437. They are independent specimens and must be treated as such. Again, the presence of either of these specimens, even if they are really from

Castell de Barbarà, isn't particularly helpful if they could have come from the 'higher layer' of the 1965-1981 collections.

Rephrased not to give such impression. There is no reason to question the provenance of the specimens (as the reviewer seems to implicitly suggest), in the sense that the site was clearly recorded in the labels. We disagree that these specimens are not relevant, for the reasons already explained above. The two layers are quite close, implying a small difference in age and at the very least a correlation with the Vallesian for the uppermost portion of the section, which is of utmost significance for anchoring the paleomag results. This was already stressed in the Discussion section (L385-388: "Even if both teeth came from a stratigraphically higher layer, these remains would indicate, at the very least (given the thickness of the outcropping section), that the main fossiliferous layer would be situated less than 10 m below unambiguously Vallesian layers."). However, given the concerns raised by the reviewer, in the revised version we have also briefly stressed this fact in the alluded paragraph of the Results section.

Line 366. Why would the latest Aragonian subzone be called this? Why would *Hippotherium* be in the name if your whole argument is that *Hippotherium* is not present in the Aragonian?

The reviewer wonders why the latest Aragonian subzone is called "*Democricetodon crusafonti* - *Hippotherium* interval subzone" according to Casanovas-Vilar et al. (2016b). These authors first provided a formal diagnosis and description of each zone, and followed Murphy and Salvador (1999, International Stratigraphic Guide, An Abridged Version) in naming the zones after their diagnostic taxa, and distinguished between 'assemblage', 'concurrent' and 'interval' (sub)zones. The subzone alluded by the reviewer was defined by Casanovas-Vilar et al. (2016b, p. 205) as follows: "The base is defined by the FLO [first local occurrence] of *Democricetodon crusafonti*, the top is defined by the FLO of the equid *Hippotherium*". So the interval subzone alluded by the reviewer includes *Hippotherium* merely because it was defined following the relevant stratigraphic guidelines, which indicate that "the names given to interval zones may be derived from the names of the boundary horizons, the name of the basal boundary preceding that of the upper boundary; e.g. *Globigerina sicanus-Orbiculina suturalis* Interval Zone" (Murphy and Salvador, 1999, p. 263). This has been briefly explained in the revised version by means of a footnote to prevent misunderstandings: "Interval zones are defined by two specified biohorizons (in these cases, corresponding to the lowest occurrences of the specified taxa), and termed with the name of taxon defining the basal boundary preceding that of the taxon defining the top boundary (Murphy and Salvador, 1999)."

Lines 530-543 – discussion of paleoenvironments. I know that Sukselainen et al use the term 'humid' in their paper, but their work is really a better proxy for 'precipitation', which I think

would be the more appropriate term to use here. For many English readers, the term 'humid' is going to invoke 'humidity' and specific types of forest, but what you're really talking about here is a forest (as opposed to a non-forest), and not a necessarily a very 'humid' one. The presence of beavers doesn't necessarily indicate high humidity, just the presence of permanent water bodies. Your other rodents certainly do indicate 'forests' as opposed to non-forested conditions, and I think this would be the more important point to emphasize. Also, line 543 would make more sense if you explain what the predominant paleoenvironments are for most late Aragonian localities. Presumably, they are drier... but are they dry forests or something else? Finally, there's no reason that I can see that these have to be 'exceptionally' humid conditions (line 587). Consider removing 'exceptionally'.

We do not completely agree with the reviewer's interpretation. Sukselainen et al. (2015) relied on hypsodonty as "a proxy for vegetation structure on an axis from closed to open [...], which in turn has been used as a proxy for past humidity and precipitation" (p. 25; see also their p. 36: "Previous studies have shown that the mean ordinated hypsodonty of large herbivorous mammals can be used as a proxy for past humidity and precipitation"). Therefore, we acknowledge that the reviewer is right when s/he says that Sukselainen et al.'s results are more directly indicative of vegetation structure, but specifically disagree with the reviewer's contention that their results are more a proxy for precipitation than for humidity itself. We do not think these two aspects can be easily disentangled based on hypsodonty, since the latter is more directly related to vegetation structure, and only through the latter linked to both humidity and precipitation. We acknowledge that, following Sukselainen et al. (2015), we were using the term 'humid' too loosely to refer to both humidity s.s. and precipitation, so this has been corrected in the revised version by putting the term 'humid' (sensu Sukselainen et al.) between quotation marks and adding the following clarifying statement: "This study relied on the hypsodonty of larger herbivorous mammals as a proxy for vegetation structure, which in turn has been used as a proxy for past humidity and precipitation (Sukselainen et al., 2015, and references therein), so that their conclusions are more adequately interpreted as indicating more closely forested conditions for primate-bearing localities (particularly those in which the two groups co-occur) as a result of higher humidity (moisture), precipitation (rainfall), or both." Further clarification in this regard has been also added regarding the interpretation of the results for Castell de Barberà by Casanovas-Vilar et al. (2008) in the following sentence.

We further dissent to some extent with the reviewer's assertion that beavers do not indicate high humidity in the case of Castell de Barberà. As a general assertion, the reviewer is right, and in the previous version we already explicitly recognized that beaver abundance at this particular locality indicated "the presence of permanent water bodies" (L536), as noted by the reviewer. Nevertheless, in the Vallesian of the Vallès-Penedès Basin, sites located near water bodies, such as Can Llobateres, have been interpreted as

most representative of the habitats that enabled the presence hominoid primates (see Marmi et al., 2012, added to the revised version, as well as Alba et al., 2018c). This has been further clarified in the revised version by adding the following sentences: “These taxa indicate the presence of permanent water bodies nearby, although not necessarily of humid and densely forested habitats. Nevertheless, in the framework of Vallès-Penedès regional setting it is consistent with the fact that Vallesian hominoids from this basin have been linked to forested humid habitats providing a year-round fruit supply (Marmi et al., 2012; Alba et al., 2018c)—such as that from Can Llobateres 1, which on the basis of plant remains has been reconstructed as a very humid marshy area with nearby dense wetland forests including some (sub)tropical elements (Marmi et al., 2012). The record of the anuran *Latonia* is also indicative of locally humid and warm conditions (Villa et al., 2017), and besides the presence of primates and beavers, a closed forested environment at Castell de Barberà is further supported by the presence of certain forest-dwelling taxa [...] Finally, certain cricetids (*Eumyarion*, *Anomalomys*) may have also preferred densely forested habitats because they tend to be more abundant in fossil faunas rich in glirids and flying squirrels (Casanovas-Vilar and Agustí, 2007)”.

The sentence from L542-544 alluded by the reviewer (“The earliest Vallesian age supported here for CB indicates that this site must no longer be considered a paleoenvironmental oddity among late Aragonian localities, but rather the best known earliest Vallesian locality from the Vallès-Penedès Basin”) has been completely rewritten as “The apparently more humid and densely forested paleoenvironmental conditions of Castell de Barberà as compared to late Aragonian localities from the same basin could be explained to some extent by differences in age. However, even if the earliest Vallesian is not well represented in the Vallès-Penedès, Castell de Barberà should not be taken as representative of the dominant paleoenvironments in the basin as a whole during this time span—as suggested by differences in faunal composition from the roughly coeval Creu de Conill 20 site...”. Furthermore, additional considerations have been added at the end of this paragraph regarding the paleoenvironmental characterization of Creu de Conill 20 as compared to Castell de Barberà based on the rodent fauna: “The rodent fauna of Creu Conill 20 is overwhelmingly dominated by the cricetid *Megacricetodon ibericus*, which is very rare at Castell de Barberà and is considered to have been a generalist probably preferring more arid woodlands (Daams and Freudenthal, 1988; Casanovas-Vilar et al., 2006, 2008, 2010, Casanovas-Vilar and Agustí, 2007). At Castell de Barberà, *Eumyarion leemanni* is the most abundant cricetid, which is consistent with the occurrence of locally humid and forested environments.”

Finally, “exceptionally” was used in the sentence from L587 (and also L530) to mean “to a greater degree than normal; unusually” (definition from the New Oxford American Dictionary). Given the reviewer’s concern in this regard, we changed it into “particularly” (meaning “to a higher degree than is usual or average”), added “and/or densely forested”

after “humid”, and also added “as compared to most other sites from the same basin” at the end to be more clear-cut.

Line 576. You said on the previous page that no primates had been found at ‘CCN20’ and from your description there, it didn’t sound too promising. So why now say that this area is ‘most promising’? Please address this inconsistency.

We agree with the reviewer that our previous assertions regarding CCN20 were a bit contradictory. The site is promising with regard to the age, less so regarding the paleoenvironment. The case is that Miocene primates are very rare in the Vallès-Penedès Basin, and the currently available sample (amassed during the 2016-2017 campaigns) is insufficient to conclusively rule out the presence of primates at the site. This is what we wanted to highlight, although we were not clear-cut enough. This has been fixed in the revised version by slightly rephrasing to some extent the sentence alluded by the reviewer as well as by expanding a bit the explanations provided in this paragraph. The revised version reads as follows: “Two different areas from the Vallès-Penedès Basin have the potential to yield additional primate remains roughly contemporaneous with those of Castell de Barberà [...]. The currently available sample of ca. 2000 macrovertebrate remains from Creu de Conill 20 is insufficient to dismiss the ulterior find of primates after additional sampling efforts at this locality (given the rarity of these taxa among Vallès-Penedès fossil assemblages), although the paleoenvironmental hints provided by the recovered fauna are not particularly promising in this regard (see above). This contrasts with the situation in the fossiliferous area of els Hostalets de Pierola, where the presence of an indeterminate dryopithecine is documented close to the Aragonian/Vallesian boundary by...”.

Figure 2 caption. I don’t know what you mean by ‘the bottom of sections samples’. **Sorry, this was a typo (“samples” instead of “sampled”). It has been fixed in the revised version, which now reads as “bottom of sections sampled for magnetostratigraphy”.**

Figure 3 caption could benefit from using A and B to refer to the two images? Note also that the legend in this figure uses ICS terminology, which should be addressed.

This is the figure with the schematic geological map of the basin, which further includes an inset to show the location of the basin within the Iberian Peninsula as well as the legend for the depicted deposits. As we do not refer separately to the main map and to the inset in the main text, we think that adding lettering is unnecessary in this particular figure. As for the use of ICS terminology, we already explained our reasons above and added to the revised text the equivalence between Aragonian and Vallesian land mammal ages with official marine stages and the corresponding ages.

Figure 5 caption. The terms CB s.l and CB s.s are never really explained. Consider omitting, as pointed out above.

These terms have been omitted from this and other figures, but as stated above their meaning has been explained with additional details in the SOM.

Figure 6 caption. As far as I can determine, the 'Aragonian' is not an ELMA term in any standard continental treatment (it's not used throughout Europe). Consider revising? Does MN need its abbreviation explained?

This relates to the comments provided by reviewer 1 about the Aragonian, and as stated above additional explanations have been provided in this regard in the revised version. As for the caption of this particular figure, note that it is specified that European land mammal ages are after Casanovas-Vilar et al. (2016b). This paper is entitled "An updated biostratigraphy for the late Aragonian and Vallesian of the Vallès-Penedès Basin (Catalonia)" and throughout the text and figures of that paper the three ELMA employed are Aragonian, Vallesian and Turolian (e.g., Casanovas-Vilar et al., 2016b: Fig. 6). So it would be inconsistent not to use the term Aragonian. Authors dealing with the Iberian record consistently use 'Aragonian' and do not even mention the 'Astaracian' (e.g., Garcés et al., 1996; Agustí et al.; Casanovas-Vilar et al., 2016b), which is not surprising given that the former is a formally defined continental stage (even if of regional applicability), whereas the latter is still a loosely defined biochronological unit. As explained above in our response to reviewer 1, this issue has been discussed in greater detail in the revised version, and the fact that the Aragonian is only applicable to Spain does not make it less European than others that are not applied to Spain. The Spanish record has played a major role in the definition of widely accepted (and formally defined) European land mammal ages (Vallesian and Turolian), and the failure to recognize the Aragonian in nearby countries such as France is more attributable to the lack of tradition and will of French paleontologists than to the inability to adequately correlate the main bioevents. Nevertheless, one must acknowledge that the detailed biostratigraphical divisions in the Aragonian type area (which include subzones lasting just a few hundred kyr) are most likely local to regional, as it occurs with the detailed Vallesian zones in the type area (see Casanovas-vilar et al., 2016b for a discussion). Overall, we consider that the reviewer concerns are adequately addressed with the newly added introductory subsection on the Aragonian and the Vallesian. As for the meaning of MN ("Mammal Neogene zones"), in the revised version it has been specified when first used in the main text (subsection 1.2 of the Introduction).

Editorial Suggestions:

The English is very good, and my only real concern is the use (and overuse!) of the term 'classical', as outlined above. I have made a few suggestions on wording and caught some typos:

We thank the reviewer for the detailed corrections. 'Classical' has been corrected as explained above.

Line 15: Delete 'a' to modify to 'precluded more accurate dating'

Done.

Line 52 - clarify 'whose provenance' - does this refer to the hominoid or pliopithecoid remains? If it's the pliopithecoids, then you could simplify the sentence to read "remains, which are attributed to the main....".

It referred to the pliopithecoids, corrected as suggested by the reviewer.

Line 116 - delete the comma after 'currently.'

This no longer applies, following the simplification of the sentence to avoid mentioning 'Vindobonian', as explained above.

Line 146 - "real scarce" is too informal. Substitute something like "even if *Hippotherium* was scarce or even absent from the site."

Done.

Line 164-165 "remained open TO (refs), or even supported (refs), an alternate..."

Done.

Line 190. Do you need to say 'Miocene' here? Is there any other section represented? If so, that should probably get explained somewhere.

Deleted. There are Quaternary sediments on top of the sections, but this needs not be mentioned in the manuscript.

Also Line 190. 'situated less than 300 m FROM the....'?

Done.

Line 196 - 'although our work (no s) confirmed that bone accumulation of THE main fossiliferous level of CB is exhausted, they enabled the recovery of additional fossil remains...'. It would be important to say WHERE these additional remains come from.

The two typos noted by the reviewer have been fixed and the sentence has been modified as further explained above. More details about the provenance of new fossil remains has been provided in the SOM (see also above).

Line 235 – say ‘east’ instead of ‘E’.

Done.

Line 236 – what is ‘only minimally higher than 95m’? Perhaps give the actual measurement or say ‘approximately 95 m ‘.

We meant that both were at 95 m with differences in the first decimal, changed into “~95 m”.

Line 251. I’m not sure that ‘Natural Remanent Magnetization’ needs to be capitalized. Also, the abbreviation (NRM) should appear right after the term.

Agree, both capitalization and abbreviation fixed.

Line 252. Should this be ‘until complete demagnetization’? I’m not sure.

Reviewer 1 also suggested an alternate phrasing, but our expert in paleomagnetisms insists on the original expression. A search in Scholar Google reveals that “up to complete demagnetization” is a common expression in papers devoted to paleomagnetism published by other authors (e.g., Di Stefano et al., 2015 Newsletters on Stratigraphy 48(2), 135-152; Fernández-Lozano et al., 2016 Tectonophysics 681, 195-208; Veselovskiy et al., 2017 Izvestiya, Physics of the Solid Earth, 53(6), 898-907).

Line 295 – the second “Vallesian” in this line appears to be out of place?

Deleted.

Line 300 – for simplicity, you can probably delete ‘in several fragments’

We substituted “broken in several fragments” by “fragmentary”.

Line 308 – substitute ‘attribution’ for ‘assignment’

Done as per reviewer’s request, although we do not understand the rationale, because both terms are synonymous within a taxonomic context, and the reviewer did not request us to change other instances in which “assignments” was used with the same meaning throughout the manuscript.

Line 345 – modify for simplicity to ‘uncertainties remain, because it is impossible to completely dismiss the presence of one or more...’

Changed as suggested.

Line 361 – substitute ‘largely’ for ‘being mostly’.

Done.

Line 380 - change 'find' to 'presence'

Done.

Line 382 - delete comma after 'significance'.

Done.

Line 407 - delete 'the' to make it 'material from els Hostalets de Pierola'.

Done.

Line 521 - change 'such' to 'this'

Done.

Line 525 - change 'of' to 'from'

Done.

Line 528 - delete 'even'.

Done.

Line 558 - change 'between' to 'of'

Done.

Line 562 - simplify to "the results of this paper provide an accurate date for this taxon"

Done.

Line 571 - what is ACM? I don't think this abbreviation was defined.

The reviewer is right, we inadvertently failed to spell it out. It has been substituted by Abocador de Can Mata in the revised version.

Line 580 - I would just write out 'female lower canine'.

Done, even though dental abbreviations are considered standard and need not be spelled out in this journal.

Line 582 - is Can Mata 1 part of Hostalets de Pierola? (e.g. a subsite or a specific level?)

Els Hostalets de Pierola is a fossiliferous area that covers a time span of about 3 Myr and includes about two dozens of 'classical' sites (mostly, but not exclusively, located within the hominimous municipality), plus >200 late Aragonian fossil localities from the macrosite of Abocador de Can Mata (ca. 12.5-11.5 Ma) and several Vallesian additional ones from the nearby macrosite of Ecoparc de Can Mata. Can Mata 1 is one of the best

known 'classical' sites, and corresponds to a locality in a strict sense (i.e., equivalent to a single geological layer). This is explained in several of the cited references in this sentence (Alba, 2012; Alba et al., 2013), so we think it needs not be explained in greater detail in the manuscript. However, we have slightly modified the passage to be more clear-cut: addition of "fossiliferous area of" before "els Hostalets de Pierola" and of "fossil locality" before "of Can Mata 1".

Line 591 – change 'samplings' to 'sampling'

Done.

Line 596 – change 'dating' to 'date'

Done.

Lines 606-609. Consider "...provided in this paper is important to contextualize the faunal and paleoenvironmental changes that enabled the coexistence of pliopithecoids and hominoids at some sites in the Vallès-Penedès Basin during the late Miocene."

Thanks, rephrased as suggested by the reviewer.

Figure 4 caption. Do you mean '*H. catalaunicum*'?

Yes, spelled out as *Hippotherium catalaunicum* in the revised version.

Other changes:

We corrected a few additional typos and made some minor edits throughout the manuscript (see the revised version of the main text with tracked changes).

We added an affiliation that was inadvertently missing from the last author.

The reference corresponding to the citation Ogg (2012) was missing from our previous version and has been added. This reference was in a couple of occasions incorrectly cited as Ogg et al. (2012), and this has also been fixed in the revised version. Multiple additional references have been added due to the additions required to address the reviewer concerns.

In the Acknowledgments of the revised version (provided in the title page, which is not sent to reviewers), we have thanked the comments and suggestions by the editor and the reviewers (those of reviewer 1 by name, since he disclosed his identity).

1 Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of
2 Castell de Barberà to the earliest Vallesian: ~~End of the controversy~~

3

4 **Keywords:** Hominoidea; Pliopithecoidea; *Hippotherium*; Late Miocene;
5 Magnetostratigraphy; Paleomagnetism.

6

7 **Abstract**

8 Castell de Barberà (~~CB~~), located in the Vallès-Penedès Basin (NE Iberian
9 Peninsula), is one of the few European sites where pliopithecoids (*Barberapithecus*)
10 and hominoids (cf. *Dryopithecus*) co-occur. The dating of this Miocene site has
11 proven ~~very~~ controversial. A latest Aragonian (MN7+8, ca. 11.88–11.18 Ma) age was
12 long accepted by most authors, ~~in despite of~~ subsequent reports of hipparionin
13 remains that signaled a Vallesian age. On the latter basis, Castell de Barberà was
14 recently correlated to the early Vallesian (MN9, ca. 11.18–10.3 Ma) on tentative
15 grounds (~~ca. 11.18–10.3 Ma~~). Uncertainties about the provenance of the
16 *Hippotherium* material and the lack of magnetostratigraphic data precluded ~~a~~ more
17 accurate dating. After decades of inactivity, fieldwork was resumed in 2014–2015 at
18 ~~CB~~ Castell de Barberà, including the ~~classical original horizon~~ (~~‘layer D’~~ CB-D) that
19 in the past delivered most of the fossils. Here we report ~~the results of~~
20 magnetostratigraphic ~~results analyses performed in for~~ the ~~classical original~~ outcrop
21 and ~~in~~ another nearby section. Our results indicate that CB-layer D is located in a
22 normal polarity magnetozone at about midheight of ~~the a~~ short (~20 m-thick) ~~classical~~
23 stratigraphic section. They ~~also imply that the~~ composite magnetostratigraphic section
24 (~50 m) has as many as four to six ~~different~~ magnetozones. These multiple reversals,
25 coupled with the in situ recovery of a *Hippotherium* humerus from ~~layer~~ CB-D in

26 2015, ~~rule out~~make it very unlikely the correlation of any of the sampled normal
27 polarity magnetozones with the long normal polarity subchron C5n.2n (11.056–9.984
28 Ma), which is characteristic of the early Vallesian. Our results support instead a
29 correlation of ~~CB-layer~~D with C5r.1n (11.188–11.146 Ma), where the
30 Aragonian/Vallesian boundary is situated, and therefore ~~unambiguously~~ indicate an
31 earliest Vallesian age of ~11.2 Ma for ~~CB~~Castell de Barberà. ~~This~~Our results
32 ~~conclusively~~ settles the longstanding debate about the Aragonian vs. Vallesian age of
33 this site, ~~being~~which appears roughly coeval with the ~~site of~~ Creu de Conill 20
34 locality (11.18 Ma), ~~which represents the first appearance datum of where~~
35 are first recorded in the Vallès-Penedès Basin.

36

37 1. Introduction

38 1.1. ~~The P~~Primates from Castell de Barberà

39 The Miocene primate-bearing site of Castell de Barberà (~~CB~~) is situated along the
40 left bank of the Ripoll River, near the old farmhouse of Ca n'Altimira, in the
41 municipality of Barberà del Vallès (Catalonia, Spain). The site was named after a
42 former medieval castle (11th century) that was subsequently reconstructed as a country
43 house. According to museum labels, the site was mostly excavated from 1965 to 1981
44 by personnel from the former Instituto Provincial de Paleontología in Sabadell,
45 currently the Institut Català de Paleontologia Miquel Crusafont (ICP), although the
46 site was already mentioned in the catalog published by Crusafont and Truyols (1954).
47 Crusafont Pairó and Hürzeler (1969) first reported the find of primate remains, and
48 Crusafont-Pairó and Golpe (1972) later provided a summary account of the site,
49 noting that methodical excavations had been carried out there. Although publications
50 based on the ~~classical~~previously collected material from ~~CB~~Castell de Barberà have

51 continued until recently, the site was not excavated for many years, until it was
52 reopened in 2014 and 2015 under the direction of author#6 and author#7.

53 Among the Miocene primate-bearing sites of Europe, CB-Castell de Barberà is
54 ~~particularly~~ renowned ~~for~~ being one of the few in which pliopithecoids and
55 hominoids co-occur (Moyà-Solà et al., 1990; Andrews et al., 1996; Alba, 2012;
56 Sukselainen et al., 2015). ~~In fact, it~~ is unknown whether the hominoid remains come
57 from the ~~very~~ same level-layer as the pliopithecoid remains, ~~whose provenance~~
58 ~~is~~ which are attributed to the main fossiliferous layer that also yielded most of the
59 fossils from the site (Alba and Moyà-Solà, 2012). However, the shortness of the
60 stratigraphic section, coupled with the lack of ~~any~~ obvious hiatuses, implies that all of
61 the fossils from the site are roughly coeval.

62 Hominoids were first reported from CB-Castell de Barberà by Crusafont Pairó and
63 Hürzeler (1969), based on a purportedly female right upper canine (IPS1823 [IPS26
64 in old terminology]; Fig. 1L) that was subsequently attributed (either implicitly or
65 explicitly) to *Hispanopithecus laietanus* (sometimes included in *Dryopithecus*;
66 Crusafont-Pairó and Casanovas Cladellas, 1973; Golpe-Posse, 1974; Golpe Posse,
67 1982, 1993; Harrison, 1991). Even though other authors later assigned this specimen
68 to a male pliopithecoid (Begun, 2002a; Alba et al., 2011; Almécija et al., 2011, 2012;
69 Alba and Moyà-Solà, 2012), hominoids are also recorded at CB-Castell de Barberà by
70 several postcranial remains. These were found among the classical-museum
71 collections in the late 1980s, ~~as~~ being ~~preliminarily~~ reported preliminarily by Moyà-
72 Solà et al. (1990), and subsequently described in detail. The hominoid specimens
73 consist of: a distal humeral diaphysis (IPS4334; Fig. 1O; Alba et al., 2011; Almécija
74 et al., 2011), a proximal pollical phalanx (IPS4333; Fig. 1N; Moyà-Solà et al., 2005;
75 Almécija et al., 2012), and a distal pollical phalanx (IPS4335; Fig. 1M; Almécija et

76 al., 2012). Based on the large body mass inferred from the humeral shaft fragment, the
77 hominoid from Castell de Barberà was tentatively referred to cf. *Dryopithecus*
78 *fontani* (Alba et al., 2011; Almécija et al., 2011; Alba, 2012; Marigó et al., 2014),
79 although such a tentative attribution cannot be further substantiated due to the lack of
80 craniodental specimens (Almécija et al., 2012).

81 The pliopithecoid from Castell de Barberà, in turn, was first reported by
82 Crusafont-Pairó (1975), based on several upper and lower teeth of a single individual
83 (Fig. 1A–K) that was initially referred to *Pliopithecus* sp. Afterward, the lower
84 dentition was described by Crusafont-Pairó and Golpe-Posse (1981a, 1982a), who
85 similarly assigned the material to *Pliopithecus* sp., albeit noting similarities with
86 *Pliopithecus lockeri* (currently in *Plesiopliopithecus*, i.e., a crouzeliid; Begun, 2002a;
87 Alba and Berning, 2013). Some later authors advocated ~~for the~~ crouzeliid affinities
88 ~~for~~ the Castell de Barberà pliopithecoid (Ginsburg, 1986), but most ~~of them~~
89 considered it to be a pliopithecoid (Andrews et al., 1996; Begun, 2002a; Harrison et al.,
90 2002). Some authors even attributed the material (although usually ~~in a~~ tentatively
91 ~~fashion~~) to *Pliopithecus antiquus* (Moyà-Solà et al., 1990; Andrews et al., 1996;
92 Harrison et al., 2002), whereas others considered it to represent a new (unnamed)
93 species (Ginsburg, 1986; Begun, 2002a; Alba et al., 2010). Finally, a new genus and
94 species, *Barberapithecus huerzeleri*, was ~~erected~~ described by Alba and Moyà-Solà
95 (2012) to accommodate the currently available dental remains, which belong to three
96 different individuals: the holotype (IPS1724; Fig. 1A–K), including 15 upper and
97 lower teeth from a single individual (see also Urciuoli et al., 2018); the paratype
98 (IPS34548), consisting of a P₃ from another individual; and IPS1823, the
99 aforementioned male upper canine (Fig. 1L), which was included in the hypodigm but
100 not designated as a paratype. Finally, Moyà-Solà et al. (2013) preliminarily reported a

101 proximal fragment of right radius (IPS66267) assigned to *Barberapithecus*, found by
102 Lars van den Hoek Ostende while revising the collections, although this specimen has
103 yet to be described in detail.

104 During the Miocene, hominoids and pliopithecoids did not co-occur at many sites,
105 which constitutesing an apparent, although poorly-understood, pattern that has
106 received ~~some~~ attention in recent publications (Sukselainen et al., 2015; Alba et al.,
107 2017, and references therein). Both groups are represented in late Aragonian and
108 Vallesian sites of the Vallès-Penedès Basin (see subsection 1.2. below for further
109 explanation on these terms), although they seldom come from the same locality (exact
110 stratigraphic horizon), even when their ranges are known to overlap as in Abocador de
111 Can Mata (Alba et al., 2017). This fact might be due to a sampling artifact related to
112 the rarity of these primates (Andrews et al., 1996), or it might be ~~alternatively~~-related
113 to different ecological preferences between these groups (e.g., Sukselainen et al.,
114 2015). As one of the few sites recording both groups and amenable to paleoecological
115 analyses based on the accompanying fauna, Castell de BarberàCB has the potential to
116 shed ~~some~~ light ~~onto~~ the particular conditions that enabled the infrequent coexistence
117 of hominoids and pliopithecoids throughout the European Miocene.

118

119 1.2. The controversial age of Castell de BarberàVallesian and Aragonian European 120 land mammal ages

121 The contribution of Castell de BarberàCB to a broader understanding of ~~fa~~ primate
122 evolution during the European Miocene in general, and the succession of primate taxa
123 at the Vallès-Penedès Basin in particular, has been ~~until now~~ hampered by a
124 longstanding controversy about the age of this site—whether it correlates to the
125 Aragonian or with the Vallesian European land mammal ages (ELMAs).

126 Some clarification is required with regard to the concept of ELMAs (for a
127 historical review, see Lindsay and Tedford, 1990 and Van Dam, 2003), with particular
128 emphasis on the use of ‘Aragonian’ instead of Orleanian + Astaracian, despite the
129 latter are more widely used outside Spain. The use of ELMAs as geochronological
130 units has been criticized (e.g., Steininger, 1999) on the grounds that they are regional
131 and that, with few exceptions, they do not correspond to properly defined
132 chronostratigraphic units (stages). ‘Mammal ages’ are generally conceptualized as
133 “biochronologic units” (e.g., Woodburne, 2004a: xiv; see also Hilgen et al., 2012) of
134 regional applicability due to divergent paleobiogeographic histories among regions
135 (Lindsay and Tedford, 1990). However, an initial definition of land mammal ages as
136 biochronologic units (i.e., biozones) is not mutually exclusive with their subsequent
137 formal definition (based on bio- and magnetostratigraphic data) as chronostratigraphic
138 units (i.e., stages based on bodies of rock formed during a given time interval; e.g.,
139 Garcés, 1995; Woodburne, 2004), which automatically implies the definition of their
140 corresponding geochronologic units (ages).

141 The definition of regional chronostratigraphic units is not at odds with the
142 International Stratigraphic Guide, because “It is better to refer strata to local or
143 regional units with accuracy and precision rather than to strain beyond the current
144 limits of time correlation in assigning these strata to units of a global scale” (Murphy
145 and Salvador, 1999: 267). Even if not directly correlated with marine stages, regional
146 continental units can be dated based on radiometric and/or paleomagnetic methods
147 (e.g., Krijgsman et al., 1994; Garcés et al., 1996). Steininger (1999) conversely
148 advocated abandoning a chronostratigraphic/geochronologic concept of ELMAs
149 altogether and proposed to replace them by entirely biostratigraphic units (‘European
150 land mammal mega-zones’) based on MN (Mammal Neogene) biozones (Mein,

151 1975). This proposal is not exempt of problems given that MN zones were defined as
152 informal biochronologic units (Mein, 1975) and that their utility at a continental-wide
153 scale is restricted due to the significant diachrony of most mammal biochronologic
154 events (van der Meulen et al., 2011, 2012). This has led some authors to contend that
155 a formal European biozonation is not possible (van der Meulen et al., 2012), which
156 would imply that the regional nature of formally defined ELMAs such as the
157 Aragonian (see review in van der Meulen et al., 2012) would rather be an advantage
158 with regard to providing accurate correlations.

159 The lack of formal (biostratigraphic and chronostratigraphic) definition of widely
160 used biochronological units such as Orleanian and Astaracian is a problem that does
161 not apply to the roughly time-equivalent Aragonian, since like the younger Vallesian
162 (which enjoys a wider geographic applicability) it has been formally defined on the
163 basis of a specific stratotype. Crusafont Pairó (1950, 1951, 1953, 1995; see also
164 Crusafont Pairó and Truyols Santonja, 1954, 1955) first used the term Vallesian in a
165 largely biochronological sense to designate the Vallès-Penedès deposits with
166 *Hipparion* (currently *Hippotherium*) that, based on the fauna, appeared intermediate
167 in age between the sites of La Grive in France (i.e., the Aragonian) and Pikermi in
168 Greece (i.e., the Turolian). The Vallesian was finally more formally defined in
169 reference to Vallès-Penedès mammal successions by Crusafont Pairó and Truyols
170 Santonja (1960) based on the entry of *Hippotherium* as its main defining criterion.
171 The Vallesian was rapidly accepted throughout Eurasia as a simple solution for the
172 complex stratigraphic terminology in use, although it was not until decades later that
173 it was formally defined as a ‘mammal stage’ based on a specific stratotype from the
174 type area (the Vallès-Penedès Basin) within an accurate bio- and magnetostratigraphic
175 framework (Garcés, 1995; Garcés et al., 1996; Agustí et al., 1997). Following the

176 proposal of MN zones by Mein (1975), MN9 and MN10 had already been equated
177 with the early and late Vallesian, respectively, in turn subdivided into multiple local
178 biozones based on rodents (Agustí, 1981, 1982; Agustí and Moyà-Solà, 1991; Agustí
179 et al., 1997; Casanovas-Vilar et al., 2011a, 2016b).

180 The Aragonian was in turn originally conceptualized (Falhbusch, 1976) as a
181 chronostratigraphic unit defined by the presence of the equid *Anchitherium* and the
182 lack of the more derived equid *Hipparion* (currently *Hippotherium*), and subdivided
183 into two subunits, the Orleanian and the Astaracian. Soon thereafter, the Aragonian
184 was defined by Daams et al. (1977) as a new stage for continental middle Miocene
185 deposits preceding the Vallesian, with its stratotype located within the Calatayud-
186 Montalbán Basin in Spain (see also Daams et al., 1999). However, Daams et al.
187 (1977) refrained from dividing the Aragonian into Orleanian and Astaracian because
188 the latter had yet to be formally defined, ultimately leading to a tripartite subdivision
189 (Daams and Freudenthal, 1981; Daams and Freudenthal, 1990). Moreover, the
190 original criterion used to define the base of the Aragonian (the dispersal of
191 *Anchitherium*) was soon questioned (Daams and Freudenthal, 1981) and eventually
192 abandoned following the definition of the Ramblian stage (Daams et al., 1987; see
193 also Daams and Freudenthal, 1990). Further stratigraphic refinements of the
194 Aragonian were later provided by Daams et al. (1987, 1999) and van der Meulen et al.
195 (2012), who further distinguished multiple Aragonian local biozones based on
196 rodents. Based on these works, currently the Aragonian may be considered a regional
197 mammal-based chronostratigraphic unit (stage), whose scope is limited to the
198 continental record from Spain (see recent reviews in van der Meulen et al., 2012 and
199 García-Paredes et al., 2016). Given the scope of this paper and the detailed local
200 zonation of the Aragonian available for the Vallès-Penedès Basin (Casanovas-Vilar et

201 al., 2016b), we refrain from using the alternative and more loosely-defined term
202 ‘Astaracian’.
203 Based on high-resolution magnetostratigraphic correlation to the geomagnetic
204 polarity time scale (GPTS), according to Van der Meulen et al. (2012) the early
205 Aragonian (ca. 17.2–15.9 Ma) corresponds to the late Burdigalian (early Miocene,
206 MN4), while the middle Aragonian (ca. 15.9–13.8 Ma) comprises most of the
207 Langhian (early to middle Miocene, MN5), and the late Aragonian (ca. 13.8–11.2 Ma,
208 roughly equivalent to the ‘Astaracian’) covers the latest Langhian, the Serravallian,
209 and the earliest Tortonian (middle to late Miocene, MN6 to MN7+8). The Vallesian,
210 in turn, entirely corresponds to the Tortonian (late Miocene), being subdivided into
211 early Vallesian (11.2–10.0 Ma, MN9) and late Vallesian (10.0–8.9 Ma, MN10; e.g.,
212 Hilgen et al., 2012; Casanovas-Vilar et al., 2016b).

213

214 1.3. The controversial age of Castell de Barberà

215 Originally, Crusafont and Truyols (1954), Crusafont-Pairó and Hürzeler (1969)
216 and Crusafont-Pairó and Golpe (1972) advocated for a ~~‘Vindobonian’ (currently, late~~
217 ~~Aragonian) instead of pre-~~Vallesian age for Castell de Barberà^{CB}, because no
218 hipparionin remains were originally found. Crusafont-Pairó and Golpe (1972),
219 Crusafont-Pairó (1972), and Golpe-Posse (1974) remarked that the fauna appeared
220 somewhat intermediate between those of the late Aragonian (MN7+8) and those of
221 the early Vallesian (MN9) localities of the Vallès-Penedès Basin (e.g., Trinxera del
222 Ferrocarril in Sant Quirze vs. Can Llobateres, respectively). However, given the lack
223 of hipparionins~~*Hipparion* (currently, *Hippotherium*)~~ and the presence of giraffids, the
224 above-mentioned authors correlated the site with the latest Aragonian. For many
225 years, such a correlation was further supported by Crusafont and coworkers

226 (Crusafont Pairó and Golpe Posse, 1972a; Golpe-Posse, 1972, 1974, 1975, 1977,
227 1978; Crusafont-Pairó and Golpe Posse, 1972b, 1974, 1981a,b, 1982a,b; Petter, 1976;
228 Santafé Llopis, 1978; Santafé i Llopis, 1978a,b), even in the face of a surface-
229 collected fragmentary fossil of *Hippotherium* already reported by Crusafont-Pairó and
230 Golpe-Posse (1974). ~~The latter authors~~ According to Crusafont-Pairó and Golpe-Posse
231 (1974), the *Hippotherium* argued that this specimen was somewhat rounded and
232 probably eroded from the upper-most fossiliferous level-layer of the Castell de
233 BarberàCB section, whereas most of the fossil assemblage (lacking *Hippotherium*)
234 came from the main fossiliferous level-layer, (situated at about the midheight of the
235 ~20 m-thick outcropping section (~~Santafé Llopis, 1978~~). Crusafont-Pairó and Golpe-
236 Posse (1974) considered that the main fossiliferous layer was undoubtedly pre-
237 Vallesian and only contemplated two possible explanations for the presence of: either
238 *Hippotherium* at the site: either *Hippotherium* arrived in the Vallès-Penedès Basin
239 during the latest Aragonian (irrespective of whether the aforementioned fragment
240 came from the main or the upper fossiliferous layers); or the Aragonian/Vallesian
241 boundary was located above the main fossiliferous level along the upper portion of
242 the Castell de Barberà ~~local~~-section (assuming the *Hippotherium* fragment came from
243 the upper layer).

244 Needless to say, the first possibility discussed by Crusafont-Pairó and Golpe-Posse
245 (1974)—namely, a pre-Vallesian appearance of *Hippotherium*—is ~~just~~ an ad hoc
246 explanation that does not hold upona closer scrutiny, because it is at odds with the
247 original very same-definition of the Vallesian proposed by Crusafont ~~himself~~ on the
248 basis of the first appearance datum of *Hippotherium* (formerly *Hipparion*; Crusafont
249 Pairó, 1950, 1951, 1953; Crusafont Pairó and Truyols Santonja, 1960). The second
250 explanation, in contrast, remained a reasonable hypothesis that would be falsified by

251 ~~the find of~~ *Hippotherium* remains from the main Castell de Barberà fossiliferous level
252 layer of CB. ~~Of course, t~~There is a third possible explanation, namely, that Castell de
253 Barberà CB could be Vallesian in age, even if *Hippotherium* was ~~real~~ scarce or ~~even~~
254 absent ~~from there~~.

255 For many years and up to the present, most subsequent authors favored a
256 correlation of Castell de Barberà CB with the late Aragonian (Moyà-Solà, 1983;
257 Agustí et al., 1985, 2001, 2013; Moyà-Solà et al., 1990; Begun, 2002a; van den Hoek
258 Ostende and Furió, 2005; Alba et al., 2006, 2010, 2011; Casanovas i Vilar, 2007;
259 Casanovas-Vilar et al., 2008, 2011a; Robles et al., 2010; Almécija et al., 2011;
260 Sukselainen et al., 2015; Agustí, 2018). The various alternate possibilities discussed
261 ~~in the paragraph~~ above were generally not discussed further, or ~~at most~~ were
262 dismissed on the grounds that the specimen ~~would had been~~ was purportedly washed
263 down from upper layer levels of the outcropping section (e.g., Santafé Llopis, 1978;
264 van den Hoek Ostende and Furió, 2005; Casanovas i Vilar, 2007; Alba et al., 2011).
265 However, it should be taken into account that the uppermost levels layers are very
266 close (<10 m) to the main fossiliferous level layer of Castell de Barberà CB
267 (Casanovas-Vilar et al., 2016b). The relatively short stratigraphic distance (110 m,
268 according to Crusafont Pairó and Truyols Santonja, 1951) between Castell de
269 Barberà CB and Can Llobateres 1 (with an interpolated age of 9.76 Ma; Casanovas-
270 Vilar et al., 2016b (110 m, according to Crusafont Pairó and Truyols Santonja, 1951)
271 also further adds difficulty to ~~contradicts~~ the considerable age difference estimated
272 between the two localities (Alba and Moyà-Solà, 2012; Casanovas-Vilar et al.,
273 2014)—which would require ~~hypothesizing~~ the existence of considerable major faults
274 (e.g., Agustí et al., 1985) that, ~~albeit conceivable~~, have yet to be documented. ~~Not~~
275 ~~surprisingly,~~ S some authors remained open to about (Aguilar et al., 1979; Ginsburg,

276 1986), or even supported (Gibert, 1975; de Bruijn et al., 1992; Andrews et al., 1996),
277 an alternate correlation of Castell de BarberàCB with MN9. However, it was not until
278 ~~the~~ publication of a Hippotherium lower molar ~~of Hippotherium~~ discovered among
279 the ~~classical-museum~~ collections of Castell de BarberàCB (Rotgers and Alba, 2011)
280 that most authors started to favor a Vallesian age for the site (Casanovas-Vilar et al.,
281 2011b, 2014, 2016a,b; Alba, 2012; Alba and Moyà-Solà, 2012; Almécija et al., 2012;
282 Robles et al., 2013a,b; Furió et al., 2015; Villa et al., 2017). Casanovas-Vilar et al.
283 (2014) even estimated an age of 10.55 Ma for the site ~~on the~~ basis of the
284 stratigraphic distance between Castell de BarberàCB and Can Llobateres reported by
285 Crusafont Pairó and Truyols Santonja (1951). ~~Later on,~~ Casanovas-Vilar et al.
286 (2016a,b) then correlated Castell de BarberàCB to the *Hippotherium – Cricetulodon*
287 *hartenbergeri* interval subzone (11.88–10.3 Ma) of the Vallès-Penedès Basin without
288 further specifying its age. However, these authors considered that such a correlation,
289 which has yet to gain universal acceptance (e.g., Agustí et al., 2013; Sukselainen et
290 al., 2015; Agustí, 2018), must be considered tentative—given ~~the~~ uncertainties about
291 the exact provenance of the *Hippotherium* remains from Castell de BarberàCB, which
292 would still allow ~~for~~ a latest Aragonian age ~~for~~ of the main fossiliferous site layer.

293 With the aim ~~of~~ performing magnetostratigraphic analyses ~~that could~~ to more
294 conclusively settle ~~this issue~~ the age of the Castell de Barberà sequence and hopefully
295 lead to the find of additional fossils, a team from the ICP led by author⁷ and author⁶
296 planned to reopen the ~~site of~~ Castell de BarberàCB site in 2014. ~~This fieldwork~~
297 initiative Finding the exact location of the site was hampered by the vagueness of
298 published indications ~~about the site's exact location~~ (Crusafont-Pairó and Golpe,
299 1972; Golpe Posse, 1974), coupled with decades of vegetation growth and recent
300 anthropic modifications ~~of the river bank~~ (to ~~build a roadway in~~ enlarge the framework

301 ~~of a riverbank's trackway in park~~). Finally, ~~T~~ two paleontological campaigns were
302 performed there in 2014 and 2015 with the aid of an excavator machine (Fig. 2).
303 Excavations in 2014 ~~were~~ focused on a ~~Miocene~~ section (~~S~~section 1, ~~or Castell de~~
304 ~~Barberà s.l.~~) situated less than 300 m ~~from~~ the ~~classical-original~~ outcrop (~~S~~section 2,
305 ~~or Castell de Barberà s.s.~~), which was excavated in 2015. The exact geographic and
306 stratigraphic situation of the ~~classical-original~~ main fossiliferous ~~level layer~~ in
307 ~~S~~section 2 was located thanks to ~~the~~ collaboration of the late Josep V. Santafé, who
308 had repeatedly excavated at the site during the 1970s with Crusafont, and indicated
309 the exact location while visiting the site during the 2015 season. Although our works
310 confirmed that the bone accumulation of ~~the~~ main ~~Castell de Barberà~~ fossiliferous
311 ~~level layer of CB~~ is exhausted, ~~they enabled the recovery of~~ ~~we recovered~~ additional
312 fossil remains ~~from the same layer and other layers, and took as well as the~~
313 paleomagnetic sampling ~~of from~~ the two excavated sections. ~~In t~~This paper, ~~aims~~
314 ~~to~~ ~~we~~ report the magnetostratigraphic data, and, with the aid of biostratigraphic
315 considerations, provide a more accurate dating of ~~Castell de Barberà~~ ~~CB~~ and discuss
316 its implications for ~~Miocene~~ hominoid and pliopithecoid evolution ~~during the~~
317 ~~Miocene~~ (for a preliminary report of these data, see Alba et al., 2018a,b).

318

319 **2. Materials and methods**

320 *2.1. Location and stratigraphy of Castell de Barberà*

321 From a geological viewpoint (Fig. 3), Castell de Barberà ~~CB~~ is located on the distal
322 facies of the Castellar del Vallès aAlluvial fFan sSystem, which belong to the middle
323 to late Miocene uUpper cContinental cComplexes of the Vallès-Penedès Basin in (NE
324 Iberian Peninsula (Agustí et al., 1985; Garcés, 1995; de Gibert and Casanovas-Vilar,
325 2011). These outcropping sediments mostly consist of mudstones (especially siltstones

326 and some finer claystones) with some intercalated coarse to fine sandstones (Santafé
327 Llopis, 1978). The repeated alternation of these lithologies along the stratigraphic
328 series may be interpreted as energy cycles—from higher energy at the base of each
329 cycle, where sandstones are more abundant, to lower energy toward the top, where
330 fine siltstones and claystones were deposited.

331 ~~To better interpret the paleomagnetic data, we also provide an updated faunal list~~
332 ~~of the rodent assemblage from CB, which includes the material recovered in 2014–~~
333 ~~2015. Given its biostratigraphic significance, we further review the few hipparionin~~
334 ~~remains from the site, including a distal humeral fragment (IPS87652) that was found~~
335 ~~in 2015. All the fossils are housed at the ICP.~~ Most of the specimens from the
336 ~~classical~~ collections of Castell de Barberà (mostly amassed between 1965–1981) lack
337 ~~any kind of~~ associated information. When preserved, some old field labels refer to
338 ‘layer 1’, and more sporadically to a so-called ‘higher layer’. ~~In~~ Despite ~~of~~ this, it is
339 possible to confidently assert that most of the ~~classical~~ material came from a single
340 layer (Santafé Llopis, 1978), which according to J.V. Santafé (pers. comm. to author⁶
341 and author⁷) was stratigraphically equivalent to layer CB-D excavated in 2015 (Fig.
342 2B). Santafé Llopis (1978) provided coordinates for the site (31N 428291 E –
343 4597272 N), indicating that it would be situated close to Ca n’Altimira. However, the
344 published coordinates are at odds with the description of the site’s location, and that
345 confirmed by Santafé ~~in~~ the field is ~~indeed~~ situated more than 400 m southward
346 (31N 428314 E – 4596862 N; Fig. 2A)¹.

¹ Geographic coordinates are given in the Universal Transverse Mercator – European Terrestrial Reference System 1989 (UTM ETRS89). They were verified ~~with the aid~~

347 ~~From a geological viewpoint (Fig. 3), CB is located on the distal facies of the~~
348 ~~Castellar del Vallès Alluvial Fan System, which belong to the middle to late Miocene~~
349 ~~Upper Continental Complexes of the Vallès-Penedès Basin (NE Iberian Peninsula).~~
350 ~~These outcropping sediments mostly consist of mudstones (especially siltstones and~~
351 ~~some finer claystones) with some intercalated coarse to fine sandstones (Santafé~~
352 ~~Llopis, 1978). The repeated alternation of these lithologies along the stratigraphic~~
353 ~~series may be interpreted as energy cycles —from higher energy at the base of each~~
354 ~~cycle, where sandstones are more abundant, to lower energy toward the top, where~~
355 ~~fine siltstones and claystones are deposited.~~

356 Two magnetostratigraphic sections, located along the left bank of the Ripoll River
357 and ~~generated~~ exposed with the aid of digging machinery, were sampled for this study
358 (Figs. 2A and 4): Ssection 1 was excavated in 2014 close to Ca n'Altimira (bottom:
359 31 N 428232 E – 4597224 N), ~~whereas and~~ Ssection 2 was excavated in 2015 ~~in the~~
360 ~~location of~~ at the ~~classical original~~ outcrop (bottom: 31N 428316 E – 4596888 N). The
361 horizontal distance between the two points is 346 m. Strata ~~gently~~ dip gently toward
362 the east, and altitude differences ~~in altitude~~ between the two points are negligible
363 (~~both are, both being~~ situated only minimally higher than 95 m above sea level).

364 Both sections (Supplementary Online Material [SOM] S1 and Fig. S1) were
365 correlated on lithostratigraphic grounds ~~and~~ based on the assumption that there is no
366 major fault located between the two sections. In particular, by taking into
367 ~~account~~ considering a bedding orientation ~~of beds~~ of 002/10E (strike/dip), measured
368 in ~~the~~ outcropping layers ~~of in~~ both sections, the bottom of magnetostratigraphic

~~of~~ using topographic maps and orthophotos from the web application VISSIR v3.26 of the Institut Cartogràfic i Geològic de Catalunya (ICGC, 2018).

369 Ssection 1 would be stratigraphically situated 16.6 m below the bottom of Ssection 2.
370 Given that Ssection 1 has a thickness of almost 50 m and Ssection 2 of about 20 m,
371 these results imply that, unless there is a fault between the two sections, they must
372 overlap to a large extent (SOM Fig. S1). Such correlation methodology is far from
373 ideal, not only because it has to assume the lack of major faulting, but also because it
374 does not take into account lateral changes in lithology and local accumulation rates.
375 However, such an approach was unavoidable given the dense vegetation cover
376 between the two sections and the impossibility to deforest the whole riverbank in
377 between.

378

379 2.2. Paleomagnetism and biostratigraphy

380 -A total of 25 samples from 14 magnetostratigraphic sites were collected along the
381 ~50 m-thick Ssection 1, although most ~~of the~~ samples ~~were~~ are concentrated ~~in~~ the
382 lower portion (~20 m in thickness) of the section. ~~In turn~~ Also, 22 samples from 11
383 sites were collected along the 20 m-thick Ssection 2. No further samples were taken ~~in~~
384 ~~the area~~ due to the presence of unsuitable coarse-grained lithologies and the
385 inaccessibility of ~~the remaining~~ other outcrops.

386 All samples were subjected to stepwise thermal demagnetization (Fig. 4) ~~aiming~~ to
387 isolate the ~~different~~ paleomagnetic components contributing to the nNatural
388 rRemanent mMagnetization (NRM). Temperature increments of 50° to 30 °C were
389 applied up to complete demagnetization ~~of the samples of the NRM or to~~
390 temperatures at which acquisition of spurious magnetization caused ~~observation of~~
391 unstable behavior ~~caused by neoformation of magnetite upon heating~~. Maximum
392 unblocking temperatures close to 600 °C indicate that the magnetic remanence is
393 carried by iron oxides. In earlier studies ~~in of~~ these ~~same~~ sedimentary sequences

394 (Garcés et al., 1996), detrital magnetite has been reported as the main remanence
395 carrier-~~of the remanence~~, together with pigmentary and detrital hematite.

396 Stepwise demagnetization reveal~~ed~~s the presence of a stable magnetic component
397 at temperatures above 250 °C, which-~~could~~ demagnetize~~s~~ at temperatures up to about
398 600 °C. A characteristic direction-of-the-magnetization direction was calculated by
399 means of least-~~squares~~ analysis for each sample (Kirschvink, 1980). Complete
400 demagnetization could rarelynot be achieved, as-when thermal treatment often lead to
401 the acquisition of a spurious remanence at temperatures above 400 °C-~~—~~. This was
402 likely related to chemical alteration of Fe-bearing minerals upon heating and growth
403 of new magnetite upon heating. In these cases, ifFortunately, the direction of the
404 magnetic remanence often remained stable between 250 and 400 °C-~~remained stable~~,
405 and a paleomagnetic average direction was could be calculated from this temperature
406 range. Anchoring the solutions to the origin was needed in most of cases as a clean
407 progressive decay of the remanence could not often be observed (Fig. 4). A total of 47
408 paleomagnetic directions were obtained (SOM Table S1) and were ranked into three
409 quality categories. Q1 directions (9 samples) were obtained from samples showing a
410 gradual and nearly complete decay towards the origin. Q2 directions (32 samples)
411 were obtained from samples that underwent unstable behavior at temperatures
412 typically above 400 °C. Q3 directions (6 samples) were obtained from samples that
413 underwent unstable at temperatures lower than 400 °C and did not show a clear decay
414 of its remanence upon heating, and were not considered for polarity interpretation.

415 A polarity interpretation of the paleomagnetic directions was done by calculating
416 the latitude of the vVirtual gGeomagnetic pPole (VGP) for eachat sample-~~level~~.
417 Positive and negative VGP latitudes correspond~~ed~~ to normal and reversed polarity,
418 respectively (Fig. 4).

419 ~~To aid better interpretation of the paleomagnetic data, we also provide an updated~~
420 ~~faunal list of the rodent assemblage from Castell de BarberàCB, which includes the~~
421 ~~material recovered in 2014–2015. Given its biostratigraphic significance, we further~~
422 ~~review the few hipparionin remains from the site, including a distal humeral fragment~~
423 ~~(IPS87652) that was found in 2015. All the fossils are housed at the ICP.~~

424

425 **3. Results**

426 ~~3.1. Rodent biostratigraphy~~

427 ~~The scanty reports of rodents in the initial faunal lists from CB (Crusafont Pairó~~
428 ~~and Golpe, 1972; Crusafont Pairó and Casanovas Cladellas, 1973; Crusafont Pairó~~
429 ~~and Golpe Posse, 1974; Golpe Posse, 1974) were soon followed by a detailed study~~
430 ~~(Aguilar et al., 1979) that recognized four ericetids and five glirids. After multiple~~
431 ~~additions and refinements (Agustí, 1981; Agustí et al., 1985; Aldana Carraseo,~~
432 ~~1992a,b,c; Casanovas i Vilar, 2007), the most recently published list is that of~~
433 ~~Casanovas Vilar et al. (2016b), with up to 16 species. According to our own revision~~
434 ~~of the material, the classical rodent assemblage from Castell de Barberà, composed of~~
435 ~~785 identifiable rodent remains, records up to 16 species, plus four additional ones~~
436 ~~recognized on the basis of the material recovered in 2015 (Table 1).~~

437

438 ~~3.2. Presence of Hippotherium~~

439 ~~It is uncertain on the basis of what specimen Crusafont Pairó and Golpe Posse~~
440 ~~(1974) originally reported the presence of *Hipparion* (currently, *Hippotherium*) at CB.~~
441 ~~However, Rotgers and Alba (2011) reported and figured an M₃ (IPS57437) of~~
442 ~~*Hippotherium* found among the CB collections. The possibility that the specimen's~~
443 ~~provenance was misrecorded is unlikely, given the fact that during a recent revision of~~

444 ~~the collections we were able to find a dp^2 -germ (IPS92389) also attributable to this~~
445 ~~taxon. While the exact stratigraphic provenance of these remains is not recorded,~~
446 ~~during the 2015 field season we recovered from level D a distal humeral fragment~~
447 ~~(IPS87652) of an equid (Fig. 4A) that is virtually undistinguishable in both size and~~
448 ~~shape from Vallesian specimens of *Hippotherium* from Vallesian the same basin (Fig.~~
449 ~~4B, C). The taxonomic attribution of IPS87652 to *Hippotherium* is complicated by the~~
450 ~~fact that both anchitheriine and hipparionin equids co-occur in some early Vallesian~~
451 ~~localities of the Iberian Peninsula and elsewhere in Europe (Salesa et al., 2004;~~
452 ~~Daxner-Höck and Bernor, 2009; Rotgers and Alba, 2011; Bernor et al., 2017).~~
453 ~~Although IPS87652 is broken in several fragments, the preserved portion of the distal~~
454 ~~diaphysis and epiphysis enable taxonomically meaningful morphological~~
455 ~~comparisons, which are provided below given the biostratigraphic relevance of~~
456 ~~*Hippotherium*.~~

457 ~~In terms of size, IPS87652 fits well with *Hippotherium (primigenium)* humeri from~~
458 ~~elsewhere in Europe, as measured by the mediolateral width of the distal epiphysis~~
459 ~~(70.0 mm), which is well within the range of the Höwenegg (MN9, Germany) sample~~
460 ~~of *Hippotherium primigenium* ($X^- = 70.5$ mm, range = 65.3–74.1 mm, $n = 13$; Bernor~~
461 ~~et al., 1997: Table 6.2). In contrast, the size of the specimen rules out an assignment to~~
462 ~~any species of *Anchitherium* from the Iberian Peninsula (and elsewhere in Europe),~~
463 ~~which are clearly smaller (Sánchez et al., 1998; Rotgers et al., 2011; Rotgers and~~
464 ~~Alba, 2011). On size grounds, IPS87652 would only be consistent with the much~~
465 ~~larger anchitheriine *Sinohippus sampelayoi* (formerly in *Anchitherium*), which is~~
466 ~~recorded at Nombrevilla 1 (MN9) in Spain (Salesa et al., 2004). Nonetheless, the~~
467 ~~latter species has not been recorded in the Vallès-Penedès Basin, where available~~
468 ~~anchitheriine remains from both the late Aragonian (Abocador de Can Mata) and~~

469 ~~early Vallesian (Can Ponce 1) indicate a much smaller size (Crusafont-Pairó and~~
470 ~~Golpe-Posse, 1974; Crusafont-Pairó, 1976; Sánchez et al., 1998; Rotgers et al., 2011;~~
471 ~~Rotgers and Alba, 2011).~~

472 ~~Furthermore, on qualitative morphological grounds IPS87652 also more closely~~
473 ~~resembles *Hippotherium* than anchitheriines. The most relevant features include~~
474 ~~(author6, pers. obs.): (i) the presence of a narrow but deep groove, medial to the~~
475 ~~lateral epicondyle and separated by the lateral surface of the radial fossa, that marks~~
476 ~~the extensive origin attachment of the extensor digitorum communis muscle (Bernor~~
477 ~~et al., 1997); and (ii) a larger (wider and higher) trochlea relative to the smaller~~
478 ~~(narrower and lower) capitulum. The aforementioned crest (i) is much deeper in~~
479 ~~*Hippotherium* than in anchitheriines — where it is absent or only weakly expressed,~~
480 ~~and only the lateral epicondylar crest can be clearly observed. In turn, the latter~~
481 ~~feature (ii) is characteristic of *Hippotherium*, but absent from *Anchitherium*, where the~~
482 ~~capitulum is larger than the trochlea. Overall, both size and morphology support a~~
483 ~~confident assignment of IPS87652 to *Hippotherium*, in further agreement with the~~
484 ~~scarce dental material of this taxon previously reported from CB (Rotgers and Alba,~~
485 ~~2011).~~

486

487 3.13. Magnetostratigraphy

488 ~~The results of our paleomagnetic analyses results~~ (Fig. 5) indicate that at least
489 four magnetozones are recorded in Ssection 1 (beginning with a reversed polarity
490 one), whereas two magnetozones are recorded in Ssection 2 (beginning with a normal
491 polarity magnetozone, in which ~~the classical fossiliferous level CB-D~~ is situated). The
492 stratigraphic distance of ca. 17 m between the bases of the two sections would suggest
493 that six ~~different~~ magnetozones have been sampled (Fig. 6), with the first normal

494 polarity magnetozone of Ssection 2 (i.e., that of the ~~classical-original~~ main
495 fossiliferous levelayer) being correlated with the second normal polarity
496 magnetozone of Ssection 1 (N2), and with the reversed polarity magnetozone of
497 Ssection 2 (R3) not having been sampled in Ssection 1. Given the close distance
498 between the two sections, local ~~differences in~~ accumulation rate differences are
499 negligible. However, some uncertainties remain, because it is due to the
500 impossibility to completely dismiss ~~discount~~ the potential presence of one or more
501 small faults between the two sections, thus allowing for other (even if less likely)
502 alternate correlations. ~~Thus, A~~ assuming that both sections completely overlap, it
503 might be conceivable to correlate the first normal polarity magnetozone of Ssection 2
504 with the first normal polarity one of Ssection 1, which results ing in only four different
505 magnetozones. Other possible correlations, involving only ~~a~~ partial overlap between
506 the two sections, would imply up to five different magnetozones, with the normal
507 polarity magnetozone of Ssection 2 being older than ~~the whole~~ all of Ssection 1 or,
508 alternatively, with the normal polarity magnetozone of Ssection 2 being correlated
509 with the uppermost portion of Ssection 1.

510

511 3.24. Rodent biostratigraphy

512 The scanty reports of rodents in the initial faunal lists from Castell de Barberà
513 (Crusafont-Pairó and Golpe, 1972; Crusafont-Pairó and Casanovas Cladellas, 1973;
514 Crusafont-Pairó and Golpe-Posse, 1974; Golpe-Posse, 1974) were soon followed by a
515 detailed study (Aguilar et al., 1979) that recognized four cricetids and five glirids.
516 After multiple additions and refinements (Agustí, 1981; Agustí et al., 1985; Aldana
517 Carrasco, 1992a,b,c; Casanovas i Vilar, 2007), the most recently published list is that
518 of Casanovas-Vilar et al. (2016b), with up to 16 species. According to our own

519 taxonomic revision of the material, the classical previously collected Castell de
520 Barberà rodent assemblage from Castell de Barberà, which is composed of 785
521 identifiable rodent remains, records up to 16 species, plus four additional ones
522 recognized on the basis of the material recovered in 2015 (Table 1).

523

524 3.32. Presence of *Hippotherium*

525 It is uncertain on the basis of what specimen Crusafont-Pairó and Golpe-Posse
526 (1974) originally reported the presence of identified as *Hipparion* (currently,
527 *Hippotherium*) at Castell de BarberàCB. However, Rotgers and Alba (2011) reported
528 and figured an M₃ (IPS57437) of *Hippotherium* M₃ (IPS57437) found among from the
529 Castell de BarberàCB collections. The possibility that the specimen's of a misrecorded
530 provenance was misrecorded is unlikely given that, the specimen was clearly labeled,
531 given the fact that dIn oururing a recent revisionre-examination of the large mammal
532 collections, we were able to findalso found a dP² germ (IPS92389) also attributable to
533 this taxon. Crusafont-Pairó and Golpe-Posse (1974) did not figure or describe the
534 specimen attributed by them to *Hippotherium*, but merely referred to it as a somewhat
535 rounded fragment—which might apply to the above-mentioned deciduous premolar,
536 although it is not possible to be certain. While the exact stratigraphic provenance of
537 these remains is not recorded, even if they both came from the upper layer they would
538 still be relevant for interpreting the magnetostratigraphic results, indicating a
539 Vallesian age for at least the topmost portion of the section.

540 Moreover, during the 2015 field season we recovered from levelCB--D during the
541 2015 field season a distal humeral fragment (IPS87652) of an equid (Fig. 74A) that is
542 virtually indistinguishable in both size and shape from Vallesian specimens of
543 *Hippotherium* specimens from Vallesianthe same basin (Fig. 74B, C). The taxonomic

544 attribution of IPS87652 to *Hippotherium* is complicated by the fact that both
545 anchitheriine and hipparionin equids co-occur in some Iberian early Vallesian
546 localities of the Iberian Peninsula and elsewhere in Europe (Salesa et al., 2004;
547 Daxner-Höck and Bernor, 2009; Rotgers and Alba, 2011; Bernor et al., 2017).
548 Although IPS87652 is broken in several fragmentarys, the preserved portion of the
549 distal diaphysis and epiphysis enable taxonomically meaningful morphological
550 comparisons, which are provided below given the biostratigraphic relevance of
551 *Hippotherium*.

552 In terms of size, IPS87652 fits well with *Hippotherium (primigenium)* humeri from
553 elsewhere in Europe, as measured by the mediolateral width of the distal epiphysis
554 (70.0 mm), which is well within the range of the Höwenegg (MN9, Germany) sample
555 of *Hippotherium primigenium* ($X^- = 70.5$ mm, range = 65.3–74.1 mm, $n = 13$; Bernor
556 et al., 1997: Table 6.2). In contrast, the size of the specimen rules out an
557 assignment attribution to any species of *Anchitherium* species from the Iberian
558 Peninsula (and elsewhere in Europe), which are clearly smaller (Sánchez et al., 1998;
559 Rotgers et al., 2011; Rotgers and Alba, 2011). On size grounds, IPS87652 would only
560 be consistent with the much larger anchitheriine *Sinohippus sampelayoi* (formerly in
561 *Anchitherium*), which is recorded at Nombrevilla 1 (MN9) in Spain (Salesa et al.,
562 2004). Nonetheless, the latter species has not been recorded in the Vallès-Penedès
563 Basin, where available anchitheriine remains from both the late Aragonian (Abocador
564 de Can Mata) and early Vallesian (Can Poncic 1) indicate a much smaller size
565 (Crusafont-Pairó and Golpe-Posse, 1974; Crusafont-Pairó, 1976; Sánchez et al., 1998;
566 Rotgers et al., 2011; Rotgers and Alba, 2011).

567 Furthermore, on qualitative morphological grounds IPS87652 also more closely
568 resembles *Hippotherium* than anchitheriines. The most relevant features- is the larger

569 trochlea relative to the smaller capitulum. In particular, IPS87652 resembles
570 *Hippotherium* in the mediolaterally wide and cylindrical trochlea and proximodistally
571 short capitulum, in contrast to the mediolaterally narrower and more globular trochlea
572 and longer capitulum of anchitheriines (MacFadden, 2001). Therefore, the distal
573 articular surface of IPS87652 is clearly *Hippotherium*-like and distinct from the more
574 primitive pattern displayed by anchitheriines (SOM Fig. S2). Moreover, include
575 (author6, pers. obs.): (i) the presence of a narrow but deep groove, medial to the
576 lateral epicondyle and (between separated by the lateral surface of the radial fossa and
577 the lateral epicondylar crest), IPS87652 possesses a narrow but deep groove that marks
578 the extensive origin attachment of the extensor digitorum communis muscle (Bernor
579 et al., 1997). This feature is consistently present in the *Hippotherium* sample from
580 Höwenegg (Bernor et al., 1997) but absent or weakly developed in ; and (ii) a larger
581 (wider and higher) trochlea relative to the smaller (narrower and lower) capitulum.
582 The aforementioned crest (i) is much deeper in *Hippotherium* than in anchitheriines
583 (SOM Fig. S2)—where it is absent or only weakly expressed, and only the lateral
584 epicondylar crest can be clearly observed. In turn, the latter feature (ii) is
585 characteristic of *Hippotherium*, but absent from *Anchitherium*, where the capitulum is
586 larger than the trochlea. Overall, both size and morphologyshape support a confident
587 assignment of IPS87652 to *Hippotherium*, in further agreement with the scarce dental
588 material of this taxon previously reported from Castell de BarberàCB (Rotgers and
589 Alba, 2011).

590

591 **4. Discussion**

592 *4.1. Biochronology*

593 Among Miocene terrestrial assemblages, rodents are generally ~~of utmost~~ greatly
594 significant ~~tee~~ for ~~providing an~~ accurate dating of fossil sites, even when
595 magnetostratigraphic data are available. This is because paleomagnetism only
596 provides a binary signal, whose interpretation critically relies on independent
597 constraints. In the Vallès-Penedès Basin, a revised scheme of local biozones and
598 subzones largely based on rodents was recently provided by Casanovas-Vilar et al.
599 (2016b), ~~being mostly based on rodents~~. The only (but most remarkable) exception in
600 this regard ~~refers to~~ concerns the Aragonian/Vallesian boundary, which is defined by
601 the first appearance datum of the equid *Hippotherium*. This is because the beginning
602 of the Vallesian was not accompanied by important changes in the rodent faunas
603 (Agustí et al., 1997, 2001). Therefore, according to Casanovas-Vilar et al. (2016b),
604 the latest Aragonian ~~would~~ corresponds to the *Democricetodon crusafonti* –
605 *Hippotherium* interval subzone (MN7+8, 11.88–11.18 Ma), whereas the earliest
606 Vallesian ~~would~~ corresponds to the *Hippotherium* – *Cricetulodon hartenbergeri*
607 interval local subzone (MN9, 11.18–10.3 Ma)². ~~In the case of~~ For Castell de
608 Barberà CB, the most significant datum provided by ~~the rodents~~ assemblage is the lack
609 of *Cricetulodon*, which precludes ~~a~~ correlation with the *Cricetulodon hartenbergeri*
610 range subzone (MN9, 10.3–9.98 Ma); and ~~thus conclusively~~ indicates conclusively an
611 age older than 10.3 Ma. Otherwise, however, the rodent assemblage is consistent with
612 both a late Aragonian or early Vallesian age. In contrast, the presence of

² Interval zones are defined by two specified biohorizons (in these cases,
corresponding to the lowest occurrences of the specified taxa), and termed with the
name of the taxon defining the basal boundary preceding that of the taxon defining the
top boundary (Murphy and Salvador, 1999).

613 *Hippotherium* at Castell de BarberàCB (Fig. 74A) implies—by definition—a post-
614 Aragonian age.

615 Given the restricted evidence provided by the rodent faunas ~~about at~~ Castell de
616 Barberà, Casanovas-Vilar et al. (2014, 2016a, b) favored a correlation of this site
617 with the earliest Vallesian mostly based on the presence of hipparionins, even if they
618 did so tentatively, given ~~the~~ uncertainties about the provenance of ~~the~~ scarce available
619 material. Although the record find of *Hippotherium* ~~remains at~~ Castell de BarberàCB
620 was already reported by Crusafont-Pairó and Golpe-Posse (1974), most later authors
621 dismissed its biochronological significance, by arguing that the specimen came from a
622 layer ~~situated~~ above the main ~~classical~~ fossiliferous ~~level~~ layer. ~~Indeed, O~~ our revision
623 of the ~~classical~~ museum collections indicates that *Hippotherium* is recorded at Castell
624 de BarberàCB by two dental specimens. Even if both teeth came from a
625 stratigraphically higher layer, these remains would indicate, at ~~the very~~ least (given
626 the thickness of the outcropping section), that the main fossiliferous ~~level~~ layer would
627 be situated less than 10 m below unambiguously Vallesian layer levels. ~~The~~
628 Longstanding discussion about the exact stratigraphic provenance of the material is
629 now irrelevant after the find of the distal humeral fragment of *Hippotherium* reported
630 in this paper, from a layer stratigraphically equivalent to the ~~classical~~ original main
631 fossiliferous ~~level~~ layer. This find settles this issue, by ~~unambiguously~~ indicating
632 unambiguously that Castell de BarberàCB is Vallesian in age, even if a large portion
633 of the rest of the fauna is still reminiscent of ~~that from~~ the late Aragonian.

634 The first appearance datum of *Hippotherium* in the Vallès-Penedès Basin is
635 currently dated to 11.18 Ma, based on magnetostratigraphic correlation of the section
636 in which Creu de Conill 20 (CCN20) is situated (Garcés et al., 1997; Agustí et al.,
637 1997; Casanovas-Vilar et al., 2016a,b). The taxonomic identity of the earliest

638 hipparionins from the Vallès-Penedès Basin is uncertain, ~~in~~partly due to the lack of
639 diagnostic material and also because the alpha-taxonomy of *Hippotherium* is yet to be
640 ~~conclusively~~ settled conclusively—particularly with regard to the earliest
641 representatives of the genus (e.g., Bernor et al., 2017). Some authors have considered
642 that the type species of the genus, *H. primigenium*, is a single polymorphic species
643 with many regional variants (Pesquero and Arribas, 2002; Zouhri and Bensalmia,
644 2005). However, ~~here~~ we follow Bernor et al. (1996) in recognizing that *H.*
645 *primigenium* is best considered a species complex, which may be formally designated
646 as *Hippotherium (primigenium) spp.*³ The nominal species *Hippotherium*
647 *catalaunicum*, erected by Pirlot (1956) based on material from ~~the~~ els Hostalets de
648 Pierola, is available for the Vallesian *Hippotherium* remains from the Vallès-Penedès
649 Basin. This species has been considered a junior subjective synonym of *Hippotherium*
650 *primigenium* by several authors (Forstén, 1968, 1978; Alberdi, 1972; Pesquero and
651 Arribas, 2002; Zouhri and Bensalmia, 2005), but here we follow Bernor and
652 colleagues (Bernor et al., 1996; Bernor and Armour-Chelu, 1999) in accepting a
653 distinct species status for this taxon within the *Hippotherium primigenium* complex—
654 i.e., *Hippotherium (primigenium) catalaunicum*. Nevertheless, we refrain from
655 attributing the ~~CB~~ scanty Castell de Barberà specimens to this species, because they
656 do not allow to ~~ascertain~~ assess its purportedly derived distinguishing features from
657 *Hippotherium primigenium* s.s. (namely, a very elongate and anteroposteriorly
658 oriented preorbital fossa; Bernor et al., 1996).

³ The parentheses around the epithet denote a supraspecific species-group taxon, such as a species group or superspecies (see ICZN, 1999: Art. 6.2)

659 Pending a more detailed study, caution in ~~the~~ taxonomic assignment of the
660 *Hippotherium* remains from Castell de Barberà CB is further advised by the
661 plesiomorphic dental morphology ~~displayed by~~ of the earliest *Hippotherium* samples
662 from elsewhere in Europe (see below). In the Vallès-Penedès Basin, the interpolated
663 age of CCN20-Creu de Conill 20 fits well with that of the earliest Vallesian
664 hipparionins from Pannonian C (ca. 11.4–11.0 Ma) localities (Atzelsdorf and
665 Gaiselberg) of the Vienna Basin, Austria (Woodburne, 2007, 2009; Bernor et al.,
666 2017), which are as yet unassigned to species and appear more plesiomorphic than *H.*
667 *primigenium* from the type locality (Eppelsheim) and other Central European
668 localities. These earliest hipparionins from Europe most closely resemble
669 *Cormohipparion* sp. from the earliest late Miocene (ca. 11.5 Ma) of North America
670 (Woodburne, 2005, 2007), and suggest that *Hippotherium* is the descendant of a
671 single dispersal event of *Cormohipparion* from North America into Eurasia sometime
672 after 11.5 Ma (Woodburne, 2009; Bernor et al., 2017). According to Harzhauser
673 (2009), Atzelsdorf would be ca. 11.2–11.1 Ma in age—i.e., roughly coeval to Creu de
674 CN20Conill 20—whereas elsewhere in the Iberian Peninsula ~~the dispersal of~~
675 *Hippotherium* dispersal is not documented until somewhat later (within C5n.2n, ca.
676 10.8–10.7 Ma) in the Calatayud-Daroca Basin (Garcés et al., 2003).

677 With regard to the biostratigraphic significance of the lack of *Hippotherium*, it
678 should be ~~taken into account~~ considered that, as illustrated by some Vallès-Penedès
679 sites, this taxon may be ~~quite~~ rare during the earliest Vallesian. This is best illustrated
680 by the local section of Ecoparc de Can Mata, where *Hippotherium* has yet to be
681 recorded ~~despite of~~ despite of being stratigraphically situated above the Aragonian/Vallesian
682 transition—as indicated by litho-, magneto-, and rodent biostratigraphic data (Alba et
683 al., 2012; Casanovas-Vilar et al., 2016b). Therefore, cCaution is ~~therefore~~ required

684 when inferring the age of other sites chronologically close to the Aragonian/Vallesian
685 boundary, because the absence of *Hippotherium* does not necessarily imply a pre-
686 Vallesian age (e.g., Can Missert; see discussion in Robles et al., 2011; contra Agustí
687 et al., 2005)—~~being and is~~ alternatively attributable to ecological conditions (e.g.,
688 unsuitable local habitat) and/or taphonomic factors (i.e., sampling biases). Equating
689 the lack of *Hippotherium* with a late Aragonian age ~~is the reasoning that~~ originally led
690 to the misdating of Castell de BarberàCB, ~~and this reasoning has been being~~ further
691 perpetuated ~~afterward in de~~-spite ~~of~~ clear (even if scarce) evidence that *Hippotherium*
692 was recorded at least somewhere within the short section where the site is located.

693 When both rodents and *Hippotherium* are considered together, biostratigraphic
694 data enable ~~an~~ unambiguous correlation of Castell de BarberàCB with the
695 *Hippotherium – Cricetulodon hartenbergeri* interval local subzone (11.18–10.3 Ma;
696 Fig. 6), as previously supported by Casanovas-Vilar et al. (2016a, b) on tentative
697 grounds.

698

699 4.2. Magnetostratigraphy

700 The magnetostratigraphic data reported in this paper ~~further~~ enable ~~a~~ more accurate
701 dating of Castell de BarberàCB within the early Vallesian, even if the shortness of the
702 sampled sections allows ~~for~~ several possible interpretations. The most parsimonious
703 interpretation, based exclusively on ~~the~~ stratigraphic distance between the two
704 sections as computed from dip measurements, implies as much as six ~~different~~
705 magnetozones. Other interpretations imply four to five ~~different~~ magnetozones.
706 Irrespective of the preferred interpretation, regarding ~~the~~ correlation of the ~~sampled~~
707 ~~identified~~ magnetozones with the GPTS (Ogg ~~et al.~~, 2012), ~~there are~~ two important
708 factors ~~should~~ be considered. First, the ~~classical-main~~ fossiliferous ~~level-layer~~ of

709 Castell de Barberà is situated in a normal polarity magnetozone. Second, no matter
710 what correlation between the two sections is preferred, there are at least four different
711 magnetozones within a short ~~sequence of~~ 50–80 m sequence in thickness. Taking into
712 account an average sedimentation rate of ~20 cm/kyr, computed for both the Vallesian
713 (Garcés et al., 1996) and the late Aragonian (Alba et al., 2017) of the Vallès-Penedès
714 Basin, the sampled interval would not represent more than 400 kyr. The frequent
715 alternation of normal and reversed polarity magnetozones in the composite sequence
716 of Castell de Barberà indicates that is excludes the correlation of any of the sampled
717 normal magnetozones with the long normal subchron C5n.2n that is characteristic of
718 the early Vallesian is not recorded in the sampled interval. Only the uppermost normal
719 polarity magnetozone N3 could correlate with the base of C5n.2n. The
720 magnetostratigraphic results rather —except, perhaps, for the uppermost part of the
721 sampled composite sequence—and suggests instead that most of the composite
722 sequence must be correlated with the latest Aragonian and earliest Vallesian upper
723 C5r, where multiple reversals are recorded over in a short time interval (Ogg ~~et al.~~,
724 2012).

725 An alternate correlation of the reversed polarity magnetozones R2 and R3 with
726 some cryptochrons or geomagnetic excursions within C5n.2n (Cande and Kent, 1992;
727 Roberts and Lewin-Harris, 2000; Evans et al., 2007) is considered unlikely because of
728 the short duration, possibly less than 10 kyr, of these events. While chances of
729 recording such short geomagnetic features in fluvial sediments are typically low, to
730 record as many reversed polarity directions as normal polarity directions within
731 C5n.2n is considered implausible. The fact that all the reversed polarity magnetozones
732 documented at Castell de Barberà are recorded by more than a single paleomagnetic
733 sample strongly argues against any of them representing a short-lived excursion event

734 within C5n.2n. The similar thickness of the documented normal and reversed polarity
735 magnetozone at Castell de Barberà strongly argues against this possibility, and favor
736 a correlation of most of the sequence with chron C5r.

737

738 In our preferred interpretation (Fig. 6), the first-lowermost normal polarity
739 magnetozone of Ssection 1 would be correlated with subchron C5r.2r-1 (11.308–
740 11.263 Ma), and the first-lowermost normal polarity magnetozone of Ssection 2 (i.e.,
741 the main ~~classical~~ fossiliferous levellayer), like the second normal polarity
742 magnetozone of Ssection 1, with C5r~~m~~.1n (11.188–11.146 Ma). According to this
743 interpretation, the classical-original outcrop of Castell de Barberà~~CB~~, including the
744 main fossiliferous levellayer at about the middle of the sequence as well as~~and~~ its
745 uppermost levellayers, would cover a time span approximately between 11.2 and –
746 11.1 Ma. Such a correlation would be compatible with the composite sequence
747 recording six ~~different~~ magnetozone (with the uppermost portion of the composite
748 sequence, recorded in Ssection 1, ~~being~~ correlated to the lowermost part of subchron
749 C5n.2n). This correlation would be also compatible with an alternate interpretation of
750 the composite sequence recording only four magnetozone (~~with the top of section 2,~~
751 ~~of reverse polarity, being somewhat younger than the upper half of section 1, of~~
752 ~~normal polarity, and which would be correlated to C5n.1n).~~ Alternatively, with the
753 first-lower normal polarity magnetozone of each section ~~might be both~~ correlated to
754 the same subchron, either C5r.2r-1 or C5r~~m~~.1n. However, ~~a~~ correlation with the
755 former subchron is much more unlikely, because it would imply a first appearance
756 datum of *Hippotherium* about 100 kyr older than previously documented in the
757 Vallès-Penedès Basin and elsewhere in Eurasia (Garcés et al., 1997; see also above).
758 An older age for the bottom of Ssection 2 (normal polarity) relative to that of Ssection

759 1 (reversed polarity), implying five magnetozones, would similarly also imply a
760 correlation of the first normal polarity magnetozone of Ssection 2 with C5r.2r-1, ~~so~~
761 ~~that this possibility~~which is similarly unlikely.

762 In summary, although there are several possible interpretations about the number
763 and correlation of ~~the~~ magnetozones recorded in the two sampled sections, the
764 combined magnetostratigraphic and biostratigraphic data ~~clearly~~ favor a correlation of
765 the ~~classical-main~~ fossiliferous level-layer of Castell de BarberàCB with subchron
766 C5r.1n. An alternative correlation with the long normal polarity chron C5n.2n,
767 which is characteristic of the early Vallesian, is precluded by the presence of a
768 reversed polarity magnetozone ~~at the~~ top of the ~~classical-main~~ fossiliferous
769 ~~level-layer—unless such a reversed polarity magnetozone is interpreted to represent a~~
770 cryptochron, which seems highly unlikely as discussed above. I turn, —, whereas a
771 correlation with the older and short normal polarity subchron ~~within~~ C5r.2r-1 is
772 possible, but is not favored here, because it would imply a first appearance datum of
773 *Hippotherium* of ca. 100 kyr older than previously documented. According to our
774 interpretation of the results, the ~~classical-main~~ fossiliferous level-layer of Castell de
775 BarberàCB would be roughly equivalent in age to Creu de Conill CN20 (11.2 Ma),
776 and both sites would representing the earliest Vallesian faunas of the Vallès-Penedès
777 Basin, with the uppermost ~~levels-layers~~ of the ~~classical-level~~original Castell de
778 Barberà outcrop, which also delivered some fossil remains, being correlated with
779 C5r.1r and with an estimated age closer to 11.1 Ma. Only ~~levels-layers~~
780 stratigraphically situated belowunderlying the ~~classical-main~~ fossiliferous level-layer
781 might be latest Aragonian instead of earliest Vallesian, although they would be
782 correlated with the same subchron and would ~~have display~~ a similar estimated age to
783 the nearest 0.1 Ma.

784

785 *4.3. Implications for primate evolution*

786 As one of the few ~~localities of the~~ European Miocene localities where
787 pliopithecoids and hominoids are recorded, Castell de Barberà CB is a key site for
788 better understanding why their co-occurrence is so rare. However, ~~the~~ lack of ~~an~~
789 accurate dating for Castell de Barberà CB represented a serious drawback for
790 adequately contextualizing ~~this~~ phenomenon within the framework of faunal and
791 paleoenvironmental changes ~~recorded at~~ in the Vallès-Penedès Basin. Sukselainen et
792 al. (2015) concluded that pliopithecoids inhabited generally more ‘humid’
793 environments than hominoids, and further suggested that the few localities hosting a
794 species ~~from~~ of each group might have ~~displayed~~ had even more ‘humid’ conditions
795 than other primate-bearing ~~localities from the~~ European Miocene localities. This study
796 relied on the hypsodonty of larger herbivorous mammals as a proxy for vegetation
797 structure, which in turn has been used as a proxy for past humidity and precipitation
798 (Sukselainen et al., 2015, and references therein), so that their conclusions are more
799 adequately interpreted as indicating more closely forested conditions for primate-
800 bearing localities (particularly those in which the two groups co-occur) as a result of
801 higher humidity (moisture), precipitation (rainfall), or both. This interpretation is
802 consistent ~~This agrees~~ with ~~the~~ previous paleoecological analyses by Casanovas-Vilar
803 et al. (2008), which found that Castell de Barberà CB would have been ~~even~~ more
804 ‘humid’ (i.e., higher moisture and/or precipitation) and displayed a more marked
805 evergreen vegetation component than late Aragonian hominoid-bearing localities
806 from the Vallès-Penedès Basin.

807 ~~As far as rodents are concerned,~~ Pexceptionally articularly ‘humid’ conditions at
808 Castell de Barberà CB are ~~indicated by~~ consistent with the high ~~the~~ abundance of

809 swimming castorids. They are represented by two ~~different~~ species: the small-sized
810 *Euroxenomys minutus*—~~indeed~~, more than half of the rodent remains recovered using
811 screen-washing techniques from Castell de BarberàCB belong to this taxon—and the
812 much rarer *Chalicomys jaegeri*, which was the size of the extant beaver (Casanovas-
813 Vilar and Agustí, 2007; Casanovas-Vilar et al., 2008, 2010). ~~These taxa, overall~~
814 ~~indicating~~ the presence of permanent water bodies ~~at the area of the site~~ nearby, ~~a-~~
815 ~~lthough not necessarily of humid and densely forested habitats. Nevertheless, in the~~
816 ~~framework of Vallès-Penedès regional setting it is consistent with the fact that~~
817 ~~Vallesian hominoids from this basin have been linked to forested humid habitats~~
818 ~~providing a year-round fruit supply (Marmi et al., 2012; Alba et al., 2018c)—such as~~
819 ~~that from Can Llobateres 1, which on the basis of plant remains has been~~
820 ~~reconstructed as a very humid marshy area with nearby dense wetland forests~~
821 ~~including some (sub)tropical elements (Marmi et al., 2012). The record of the anuran~~
822 ~~Latonia is also indicative of locally humid and warm conditions (Villa et al., 2017),~~
823 ~~and besides the presence of primates and beavers, a closed forested environment at~~
824 ~~Castell de Barberà is further supported by the presence of~~ ~~In addition,~~ certain forest-
825 dwelling taxa, such as arboreal dormice (*Bransatoglis*, *Glirudinus*, *Muscardinus*,
826 *Myoglis*, *Paraglrirulus*) and flying squirrels (*Miopetaurista* and probably *Albanensia*),
827 ~~which~~ are more diverse and abundant at Castell de Barberà than in roughly
828 contemporary sites (Casanovas-Vilar and Agustí, 2007; Casanovas-Vilar et al., 2008,
829 2010). ~~Finally, certain cricetids (Eumyarion, Anomalomys) may have also preferred~~
830 ~~densely forested habitats because they tend to be more abundant in fossil faunas rich~~
831 ~~in glirids and flying squirrels (Casanovas-Vilar and Agustí, 2007).~~
832 The apparently more humid and densely forested paleoenvironmental conditions of
833 ~~The earliest Vallesian age supported here for Castell de BarberàCB as compared to~~

834 ~~indicates that this site must no longer be considered a paleoenvironmental oddity~~
835 ~~among late Aragonian localities from the same basin could be explained to some~~
836 ~~extent by differences in age. However, even if the ,but rather the best known earliest~~
837 ~~Vallesian locality from the Vallès-Penedès Basin—a time span is not very well~~
838 ~~represented in this basin the Vallès-Penedès and where primates are only poorly~~
839 ~~known-, Castell de Barberà should not be taken as representative of Interestingly, the~~
840 ~~moist paleoenvironment of CB (see also Villa et al., 2017) the dominant~~
841 ~~paleoenvironments in the basin as a whole during this time span—as suggested by~~
842 ~~differences in faunal composition from might be at odds with that of the roughly~~
843 ~~coeval site of Creu de Conill CN20 site, where Hippotherium is apparently more~~
844 ~~abundant (authors' unpubl. data), the micromammal assemblage indicates drier~~
845 ~~conditions (Casanovas-Vilar et al., 2006, 2008; Casanovas-Vilar and Agustí, 2007),~~
846 ~~and primates have yet to not been recorded yet in spite of recent (2016–2017)~~
847 ~~fieldwork campaigns (authors' unpubl. data). Indeed, the rodent assemblage from~~
848 ~~Creu de Conill CN20 is far less diverse than that of Castell de Barberà CB~~
849 ~~(Casanovas-Vilar et al., 2006, 2016a, b; authors' unpublished data), and includes very~~
850 ~~few forest-dwelling taxa (*Muscardinus*, cf. *Paraglitirulus*), with beavers being only~~
851 ~~represented by scarce remains of *Euroxenomys minutus*. The rodent fauna of Creu~~
852 ~~Conill 20 is overwhelmingly dominated by the cricetid *Megacricetodon ibericus*,~~
853 ~~which is very rare at Castell de Barberà and is considered to have been a generalist~~
854 ~~probably preferring more arid woodlands (Daams and Freudenthal, 1988; Casanovas-~~
855 ~~Vilar et al., 2006, 2008, 2010, Casanovas-Vilar and Agustí, 2007). At Castell de~~
856 ~~Barberà, *Eumyarion leemanni* is the most abundant cricetid, which is consistent with~~
857 ~~the occurrence of locally humid and forested environments.~~

858 The dating provided in this paper for Castell de Barberà ~~clearly~~ indicates that
859 renewed efforts are required to better characterize ~~the~~ earliest Vallesian faunas from
860 the Vallès-Penedès Basin, not only to investigate the particular conditions that
861 enabled ~~the~~ co-existence ~~between-of~~ hominoids and pliopithecoids, but especially to
862 ~~further~~ clarify their taxonomic status and/or phylogenetic relationships. ~~In the case~~
863 ~~of~~For pliopithecoids, *Barberapithecus huerzeleri* is exclusively known from Castell
864 de Barberà (Alba and Moyà-Solà, 2012). ~~The Our~~ results ~~of this paper enable to~~
865 provide an accurate date~~ing~~ for this taxon, but additional and more complete material
866 ~~would be~~is required to more definitively settle its crouzeliid status ~~as well as~~and its
867 relationships with other pliopithecoid taxa. Regarding hominoids, ~~the~~ attribution of
868 the postcranial remains from Castell de Barberà to *Dryopithecus fontani* by Alba et
869 al. (2011) was only tentative. Given that this taxon is otherwise known only from the
870 late Aragonian of the Vallès-Penedès Basin, France and Austria (Begun, 2002b;
871 Moyà-Solà et al., 2009; Pérez de los Ríos et al., 2013), if such assignment was
872 confirmed then Castell de Barberà might represent the last appearance datum of
873 this genus. However, additional (especially craniodental) remains would be necessary
874 to confirm the taxonomic status of the hominoid taxon recorded at Castell de
875 Barberà, which is intermediate in time between the late Aragonian dryopithecins
876 recorded at Abocador de Can Mata and the more derived hispanopithecins
877 recorded at the basin later during the Vallesian (Alba, 2012; Alba et al., 2017, 2018c).

878 Two ~~different~~ areas ~~appear most promising with regard to the prospect of finding of~~
879 ~~the Vallès-Penedès Basin have the potential to yield~~ additional primate remains
880 roughly contemporaneous with those of Castell de Barberà: the earliest Vallesian
881 site of Creu de Conill ~~CN~~20 in Terrassa; and the Vallesian levels of els Hostalets de
882 Pierola (located westward from Abocador de Can Mata; Moyà-Solà et al., 2009; Alba

883 et al., 2017). The currently available sample of ca. 2000 macrovertebrate remains
884 from Creu de Conill 20 is insufficient to dismiss the ulterior find of primates after
885 additional sampling efforts at this locality (given the rarity of these taxa among
886 Vallès-Penedès fossil assemblages), although the paleoenvironmental hints provided
887 by the recovered fauna are not particularly promising in this regard (see above). This
888 contrasts with the situation in the fossiliferous area of els Hostalets de Pierola, where
889 ~~While primates have yet to be found at the Creu de Conill area,~~ the presence of an
890 indeterminate dryopithecine is documented close to the Aragonian/Vallesian
891 boundary ~~is documented at els Hostalets de Pierola, being recorded~~ by a female lower
892 canine_{C₁} from the latest Aragonian ~~site~~ fossil locality of Can Mata 1 (Crusafont-Pairó
893 and Golpe-Posse, 1973; ~~+~~ Golpe Posse, 1982, 1993; Alba, 2012) ~~as well as~~ and by an
894 isolated upper molar of uncertain provenance from the same area (van der Made and
895 Ribot, 1999; Alba, 2012; Alba et al., 2013). ~~Hopefully,~~ F-future fieldwork in these and
896 other sites ~~will~~ should help to determine the taxonomic identity of earliest Vallesian
897 hominoids from the Vallès-Penedès Basin ~~as well as~~ and to clarify whether their rare
898 co-occurrence with pliopithecoids at Castell de Barberà_{CB} and other localities was
899 related to ~~exceptionally particularly~~ humid and/or densely forested environmental
900 conditions as compared to most other sites from the same basin.

901

902 **5. Summary and conclusions**

903 Here we report ~~the results of~~ new magnetostratigraphic ~~sampling results~~ for the
904 primate-bearing ~~site of~~ Castell de Barberà_{CB} ~~site~~ collected in 2014–2015, together
905 with ~~the a recent find of~~ *Hippotherium* find collected in 2015 from a stratigraphic
906 horizon (CB-D_{layer D}) equivalent to the original main ~~classical~~ fossiliferous layer at
907 Castell de Barberà. We interpret these data in ~~the~~ light of ~~the~~ previously published

908 literature and our own review of the rodent assemblage from Castell de Barberà~~CB~~,
909 with the aim of providing an unambiguous and more accurate date~~ing~~ for the site.

910 Our paleomagnetic results, coupled with ~~the~~-in situ recovery of a *Hippotherium*
911 humerus, ~~rule-out~~make it unlikely the ~~the~~-correlation of any of the various sampled
912 normal polarity magnetozones with the long normal polarity subchron C5n.2n
913 (11.056–9.984 Ma) that is characteristic of the early Vallesian, and support instead ~~a~~
914 correlation of ~~CB-layer~~-D with C5r.1n (11.188–11.146 Ma), where the
915 Aragonian/Vallesian boundary is situated. ~~All-in~~Overall, our results unambiguously
916 indicate an earliest-early Vallesian age ~~of ~11.2 Ma~~ for Castell de Barberà~~CB~~, thereby
917 settling ~~the~~-longstanding debate about the Aragonian vs. Vallesian age of this site.
918 Our results further support an earliest Vallesian (~11.2 Ma) age for ~~Therefore,~~ Castell
919 de Barberà, which~~CB~~ would be roughly coeval with the ~~site of CCN20~~Creu de Conill
920 20 site, which represents~~where~~ the first appearance datum of hipparionins in the
921 Vallès-Penedès Basin has been recorded.

922 ~~The~~ Aaccurate dating of the Castell de Barberà site~~CB~~ provided in this paper will
923 allow~~is~~ important to ~~better~~-contextualize ~~at a basin scale~~-the faunal and
924 paleoenvironmental changes that enabled ~~the~~-coexistence of pliopithecoids and
925 hominoids at ~~some~~-particular sites in the Vallès-Penedès Basin during the late
926 Miocene. It also provides useful hints ~~so as~~-to redirect future fieldwork efforts in the
927 Vallès-Penedès Basin, with the aim to clarify the taxonomic identity and/or
928 phylogenetic relationships of ~~the~~-catarrhine primates that lived there during the
929 earliest Vallesian.

930

931 **References**

932 Aguilar, J.-P., Agustí, J., Gibert, J., 1979. Rongeurs miocènes dans le Valles-Penedes.

- 933 2 - Les Rongeurs de Castell de Barbera. *Palaeovertebrata* 9, 17–32.
- 934 Agustí, J., 1981. Roedores miomorfos del Neógeno de Cataluña. PhD. dissertation,
935 Universitat de Barcelona.
- 936 Agustí, J., 1982. Biozonación del neogeno continental de Cataluña mediante roedores
937 (Mammalia). *Acta Geológica Hispánica* 17, 21–26.
- 938 Agustí, J., 2018. Las faunas de mamíferos del Mioceno continental de la Península
939 Ibérica. *Revista PH* 94, 182–205.
- 940 Agustí, J., Moyà-Solà, S., 1991. Spanish Neogene Mammal succession and its bearing
941 on continental biochronology. *Newsletters on Stratigraphy* 25, 91–114.
- 942 Agustí, J., Cabrera, L., Moyà-Solà, S., 1985. Sinopsis estratigráfica del Neógeno de la
943 fosa del Vallés-Penedés. *Paleontologia i Evolució* 18, 57–81.
- 944 Agustí, J., Cabrera, L., Garcés, M., Parés, J.M., 1997. The Vallesian mammal
945 succession in the Vallès-Penedès basin (northeast Spain): Paleomagnetic
946 calibration and correlation with global events. *Palaeogeography,*
947 *Palaeoclimatology, Palaeoecology* 133, 149–180.
- 948 Agustí, J., Cabrera, L., Garcés, M., Krijgsman, W., Oms, O., Parés, J.M., 2001. A
949 calibrated mammal scale for the Neogene of Western Europe. State of the art.
950 *Earth-Science Reviews* 52, 247–260.
- 951 Agustí, J., Casanovas-Vilar, I., Furió, M., 2005. Rodents, insectivores and
952 chiropterans (Mammalia) from the late Aragonian of Can Missert (Middle
953 Miocene, Vallès-Penedès basin, Spain). *Geobios* 38, 575–583.
- 954 Agustí, J., Cabrera, L., Garcés, M., 2013. The Vallesian Mammal Turnover: A late
955 Miocene record of decoupled land-ocean evolution. *Geobios* 46, 151–157.
- 956 Alba, D.M., 2012. Fossil apes from the Vallès-Penedès Basin. *Evolutionary*
957 *Anthropology* 21, 254–269.

958 Alba, D.M., Berning, B., 2013. On the holotype and original description of the
959 pliopithecoid *Plesiopliopithecus lockeri* (Zapfe, 1960). *Journal of Human Evolution*
960 65, 338–340.

961 Alba, D.M., Moyà-Solà, S., 2012. A new pliopithecoid genus (Primates:
962 Pliopithecoidae) from Castell de Barberà (Vallès-Penedès Basin, Catalonia, Spain).
963 *American Journal of Physical Anthropology* 147, 88–112.

964 Alba, D.M., Moyà-Solà, S., Casanovas-Vilar, I., Galindo, J., Robles, J.M., Rotgers,
965 C., Furió, M., Angelone, C., Köhler, M., Garcés, M., Cabrera, L., Almécija, S.,
966 Obradó, P., 2006. Los vertebrados fósiles del Abocador de Can Mata (els Hostalets
967 de Pierola, l’Anoia, Cataluña), una sucesión de localidades del Aragoniense
968 superior (MN6 y MN7+8) de la cuenca del Vallès-Penedès. Campañas 2002-2003,
969 2004 y 2005. *Estudios Geológicos* 62, 295–312.

970 Alba, D.M., Moyà-Solà, S., Malgosa, A., Casanovas-Vilar, I., Robles, J.M., Almécija,
971 S., Galindo, J., Rotgers, C., Bertó Mengual, J.V., 2010. A new species of
972 *Pliopithecus* Gervais, 1849 (Primates: Pliopithecidae) from the Middle Miocene
973 (MN8) of Abocador de Can Mata (els Hostalets de Pierola, Catalonia, Spain).
974 *American Journal of Physical Anthropology* 141, 52–75.

975 Alba, D.M., Moyà-Solà, S., Almécija, S., 2011. A partial hominoid humerus from the
976 middle Miocene of Castell de Barberà (Vallès-Penedès Basin, Catalonia, Spain).
977 *American Journal of Physical Anthropology* 144, 365–381.

978 Alba, D.M., Carmona, R., Bertó Mengual, J.V., Casanovas-Vilar, I., Furió, M.,
979 Garcés, M., Galindo, J., Luján, À.H., 2012. Intervenció paleontològica a l’Ecoparc
980 de Can Mata (els Hostalets de Pierola, conca del Vallès-Penedès). *Tribuna*
981 *d’Arqueologia* 2010–2011, 115–130.

982 Alba, D.M., Fortuny, J., Pérez de los Ríos, M., Zanolli, C., Almécija, S., Casanovas-

983 Vilar, I., Robles, J.M., Moyà-Solà, S., 2013. New dental remains of *Anoiapithecus*
984 and the first appearance datum of hominoids in the Iberian Peninsula. *Journal of*
985 *Human Evolution* 65, 573–584.

986 Alba, D.M., Casanovas-Vilar, I., Garcés, M., Robles, J.M., 2017. Ten years in the
987 dump: An updated review of the Miocene primate-bearing localities from
988 Abocador de Can Mata (NE Iberian Peninsula). *Journal of Human Evolution* 102,
989 12–20.

990 Alba, D.M., Moyà-Solà, S., DeMiguel, D., Casanovas-Vilar, I., Garcés, M., Robles, J.
991 M., Madurell-Malapeira, J., Almécija, S., 2018. Ape quest in the Vallès-Penedès
992 Basin (2014–2017): Fieldwork results and prospects for the future. *American*
993 *Journal of Physical Anthropology* 165 S66, 7.

994 Alba, D.M., Garcés, M., Pina, M., Casanovas-Vilar, I., Robles, J.M., Moyà-Solà, S.,
995 Almécija, S., 2018. Bio- and magnetostratigraphic correlation of the Miocene
996 primate-bearing site of Castell de Barberà: end of the controversy. *Proceedings of*
997 *the European Society for the study of Human Evolution* 7, 2.

998 Alba, D.M., Casanovas-Vilar, I., Furió, M., García-Paredes, I., Angelone, C., Jovells-
999 Vaqué, S., Luján, À.H., Almécija, S., Moyà-Solà, S., 2018c. Can Pallars i
1000 Llobateres: A new hominoid-bearing locality from the late Miocene of the Vallès-
1001 Penedès Basin (NE Iberian Peninsula). *Journal of Human Evolution* 121, 193–203.

1002 Alberdi, M.T., 1972. El género *Hipparion* en España, revisión e historia evolutiva.
1003 *Coloquios de la Cátedra de Paleontología* 21, 7–8.

1004 Aldana Carrasco, E.J., 1992a. Los Castoridae (Rodentia, Mammalia) del Neógeno de
1005 Cataluña (España). *Treballs del Museu de Geologia de Barcelona* 2, 99–141.

1006 Aldana Carrasco, E.J., 1992b. Los Esciurópteros del Mioceno de la cuenca del Vallès-
1007 Penedès (Cataluña, España). *Geogaceta* 11, 114–116.

- 1008 Aldana Carrasco, E.J., 1992c. Los Sciurinae (Rodentia, Mammalia) del Mioceno de la
1009 Cuenca del Vallès-Penedès (Cataluña, España). Treballs del Museu de Geologia de
1010 Barcelona 2, 69–97.
- 1011 Almécija, S., Alba, D.M., Moyà-Solà, S., 2011. Large-hominoid remains from the
1012 Middle Miocene locality of Castell de Barberà (Vallès-Penedès Basin, Catalonia,
1013 Spain). *American Journal of Physical Anthropology* 144 S52, 74.
- 1014 Almécija, S., Alba, D.M., Moyà-Solà, S., 2012. The thumb of Miocene apes: new
1015 insights from Castell de Barberà (Catalonia, Spain). *American Journal of Physical*
1016 *Anthropology* 148, 436–450.
- 1017 Álvarez Sierra, M.A., Calvo, J.P., Morales, J., Alonso-Zarza, A., Azanza, B., García
1018 Paredes, I., Hernández Fernández, M., van der Meulen, A.J., Peláez-Campomanes,
1019 P., Quiralte, V., Salesa, M.J., Sánchez, I.M., Soria, D., 2003. El tránsito
1020 Aragoniense-Vallesiense en el área de Daroca-Nombrevilla (Zaragoza, España).
1021 *Coloquios de Paleontología* Vol. Ext. 1, 25–33.
- 1022 Begun, D.R., 2002a. The Pliopithecoidea. In: Hartwig, W.C. (Ed.), *The Primate Fossil*
1023 *Record*. Cambridge University Press, Cambridge, pp. 221–240.
- 1024 Begun, D.R., 2002b. European hominoids. In: Hartwig, W.C. (Ed.), *The Primate*
1025 *Fossil Record*. Cambridge University Press, Cambridge, pp. 339–368.
- 1026 Bernor, R.L., Armour-Chelu, M., 1999. Family Equidae. In: Rössner, G.E., Heissig,
1027 K. (Eds.), *The Miocene Land Mammals of Europe*. Verlag Dr. Friedrich Pfeil,
1028 München, pp. 193–202.
- 1029 Bernor, R.L., Tobien, H., Hayek, L.-A.C., Mittmann, H.-W., 1997. *Hippotherium*
1030 *primigenium* (Equidae, Mammalia) from the late Miocene of Höwenegg (Hegau,
1031 Germany). *Andrias* 10, 5–230.

- 1032 Bernor, R.L., Göhlich, U., Harzhauser, M., Semprebon, G.M., 2017. The Pannonian C
1033 hipparions from the Vienna Basin. *Palaeogeography, Palaeoclimatology,*
1034 *Palaeoecology* 476, 28–41.
- 1035 [Cande, S.C., Kent, D.V., 1992. A new geomagnetic polarity time scale for the Late](#)
1036 [Cretaceous and Cenozoic. *Journal of Geophysical Research* 97, 13917–13951.](#)
- 1037 Casanovas i Vilar, I., 2007. The rodent assemblages from the Late Aragonian and the
1038 Vallesian (Middle to Late Miocene) of the Vallès-Penedès Basin (Catalonia,
1039 Spain). Ph.D. Dissertation, Universitat Autònoma de Barcelona.
- 1040 Casanovas-Vilar, I., Agustí, J., 2007. Ecogeographical stability and climate forcing in
1041 the Late Miocene (Vallesian) rodent record of Spain. *Palaeogeography,*
1042 *Palaeoclimatology, Palaeoecology* 248, 169–189.
- 1043 Casanovas-Vilar, I., Furió, M., Agustí, J., 2006. Rodents, insectivores and
1044 paleoenvironment associated to the first-appearing hipparionine horses in the
1045 Vallès-Penedès Basin (Northeastern Spain). *Beiträge zur Paläontologie* 30, 89–
1046 107.
- 1047 Casanovas-Vilar, I., Alba, D.M., Moyà-Solà, S., Galindo, J., Cabrera, L., Garcés, M.,
1048 Furió, M., Robles, J.M., Köhler, M., Angelone, C., 2008. Biochronological,
1049 taphonomical and paleoenvironmental background of the fossil great ape
1050 *Pierolapithecus catalaunicus* (Primates, Hominidae). *Journal of Human Evolution*
1051 55, 589–603.
- 1052 Casanovas-Vilar, I., Angelone, C., Alba, D.M., Moyà-Solà, S., Köhler, M., Galindo,
1053 J., 2010. Rodents and lagomorphs from the Middle Miocene hominoid-bearing site
1054 of Barranc de Can Vila 1 (els Hostalets de Pierola, Catalonia, Spain). *Neues*
1055 *Jahrbuch für Geologie und Paläontologie Abhandlungen* 257, 293–315.
- 1056 Casanovas-Vilar, I., Alba, D.M., Garcés, M., Robles, J.M., Moyà-Solà, S., 2011a.

1057 Updated chronology for the Miocene hominoid radiation in Western Eurasia.
1058 Proceedings of the National Academy of Sciences USA 108, 5554–5559.

1059 Casanovas-Vilar, I., Alba, D.M., Robles, J.M., Moyà-Solà, S., 2011b. Registro
1060 paleontológico continental del Mioceno de la cuenca del Vallès-Penedès.
1061 Paleontologia i Evolució memòria especial 6, 55–80.

1062 Casanovas-Vilar, I., Van den Hoek Ostende, L.W., Furió, M., Madern, A., 2014. The
1063 range and extent of the Vallesian Crisis (Late Miocene): new prospects based on
1064 the micromammal record from the Vallès-Penedès Basin (Catalonia, Spain).
1065 Journal of Iberian Geology 40, 29–48.

1066 Casanovas-Vilar, I., Madern, A., Alba, D.M., Cabrera, L., García-Paredes, I., Van den
1067 Hoek Ostende, L.W., DeMiguel, D., Robles, J.M., Furió, M., Van Dam, J., Garcés,
1068 M., Angelone, C., Moyà-Solà, S., 2016a. The Miocene mammal record of the
1069 Vallès-Penedès Basin (Catalonia). Comptes Rendus Palevol 15, 791–812.

1070 Casanovas-Vilar, I., Garcés, M., Van Dam, J., García-Paredes, I., Robles, J.M., Alba,
1071 D.M., 2016b. An updated biostratigraphy for the late Aragonian and the Vallesian
1072 of the Vallès-Penedès Basin (Catalonia). Geologica Acta 14, 195–217.

1073 Crusafont Pairó, M., 1950. La cuestión del llamado Meóico español. Arrahona 1950,
1074 41–48.

1075 Crusafont Pairó, M., 1951. El sistema miocénico en la depresión española del Vallés-
1076 Penedés. In: International Geological Congress "Report of the Eighteenth Session,
1077 Great Britain, 1948", Part XI, pp. 33–42.

1078 Crusafont Pairó, M., 1953. El sistema miocénico en la depresión española del Vallés-
1079 Penedés. Memorias y Comunicaciones del Instituto Geológico Provincial 10, 13–
1080 23.

1081 [Crusafont Pairó, M., 1955. Données biogéographiques relevées par la](#)

- 1082 [paléomammalogie du Miocène espagnol. In: Colloque International sur "Problèmes](#)
1083 [Actuels de Paléontologie" C. N. R. C. Paris, avril 1955, pp. 101–108.](#)
1084
- 1085 Crusafont-Pairó, M., 1972. Les *Ischyriactis* de la transition Vindobonien-Vallésien.
1086 *Palaeovertebrata* 5, 253–260.
- 1087 Crusafont-Pairó, M., 1975. El gibón fósil (*Pliopithecus*) del Vindoboniense terminal
1088 del Vallès. Boletín Informativo del Instituto de Paleontología de Sabadell 7, 36–38.
- 1089 Crusafont-Pairó, M., 1976. Corrigenda (A dos notas sobre la presencia del género
1090 *Anchitherium* en el Mioceno del Vallès. Boletín Informativo del Instituto de
1091 Paleontología de Sabadell 8, 27–28.
- 1092 Crusafont-Pairó, M., Casanovas Cladellas, L., 1973. Fossilium Catalogus. I:
1093 Animalia. Pars 121. Mammalia Tertiaria Hispaniae. Dr. W. Junk B.V., 's-
1094 Gravenhave.
- 1095 Crusafont-Pairó, M., Golpe, J.M., 1972. Dos nuevos yacimientos del vindoboniense
1096 en el Vallès. Acta Geológica Hispánica 7, 71–72.
- 1097 Crusafont Pairó, M., Golpe Posse, J.M., 1972a. Los yacimientos de mamíferos fósiles
1098 del Vallès. Boletín Informativo del Instituto Provincial de Paleontología Sabadell
1099 4, 20–24.
- 1100 Crusafont-Pairó, M., Golpe Posse, J.M., 1972b. Hallazgo del género *Anchitherium*
1101 Meyer, 1844 en el Vindoboniense terminal del Vallés-Penedés. Boletín de la Real
1102 Sociedad Española de Historia Natural 69, 297–298.
- 1103 Crusafont-Pairó, M., Golpe-Posse, J.M., 1973. New pongids from the Miocene of
1104 Vallès Penedès Basin (Catalonia, Spain). *Journal of Human Evolution* 2, 17–23.
- 1105 Crusafont-Pairó, M., Golpe-Posse, J.M., 1974. Asociación de *Anchitherium* Mey.,
1106 1834, con *Hipparion* Christ, 1832, en el Alto Mioceno del Vallès. Boletín de la

- 1107 Real Sociedad Española de Historia Natural 72, 75–93.
- 1108 Crusafont-Pairó, M., Golpe-Posse, J.M., 1981a. Estudio de la dentición inferior del
1109 primer pliopitécido hallado en España (Vindoboniense terminal de Castell de
1110 Barberà, Cataluña, España). Butlletí Informatiu de l’Institut de Paleontologia de
1111 Sabadell 13, 25–38.
- 1112 Crusafont-Pairó, M., Golpe-Posse, J.M., 1981b. Hallazgo de una nueva especie del
1113 género *Semigenetta* (*S. grandis*) del Vindoboniense terminal de Castell de Barberá
1114 (Depresión prelitoral catalana; España). Boletín de la Real Sociedad Española de
1115 Historia Natural 79, 67–76.
- 1116 Crusafont Pairó, M., Golpe Posse, J.M., 1982a. Los Pliopitécidos en España.
1117 Coloquios de Paleontología 37, 41–46.
- 1118 Crusafont Pairó, M., Golpe Posse, J.M., 1982b. Hallazgo de *Palaeomeles Pachecoi*
1119 Vill. et Crus., 1943, en Castell de Barberà (Vallés-Penedés). Acta Geológica
1120 Hispánica 17, 27–37.
- 1121 Crusafont Pairó, M., Hürzeler, J., 1969. Catálogo comentado de los póngidos fósiles
1122 de España. Acta Geológica Hispánica 4, 44–48.
- 1123 Crusafont Pairó, M., Truyols Santonja, J., 1951. Hallazgo del *plesiodimylus chantrei*
1124 Gaillard en el Meótico del Vallés. Notas y Comunicaciones del Instituto Geológico
1125 y Minero de España 22, 97–126.
- 1126 Crusafont, M., Truyols, J., 1954. Catálogo Paleomastológico del Mioceno del Vallés-
1127 Penedés y de Calatayud-Teruel. Segundo Cursillo Internacional de Paleontología.
1128 Museo de la Ciudad de Sabadell, Sabadell.
- 1129 Crusafont Pairó, M., Truyols Santonja, J., 1954. Sinopsis estratigráfico-paleontológica
1130 del Vallés-Penedés. Arrahona 1954, 1–15.
- 1131 Crusafont Pairó, M., Truyols Santonja, J., 1959. Sobre el nuevo proyecto de

- 1132 [estructuración y nomenclatura del Mioceno mediterráneo. Notas y Comunicaciones](#)
1133 [del Instituto Geológico y Minero de España 56, 33–53.](#)
1134
- 1135 Crusafont Pairó, M., Truyols Santonja, J., 1960. Sobre la caracterización del
1136 Vallesiense. *Notas y Comunicaciones del Instituto Geológico y Minero de España*
1137 60, 109–126.
- 1138 [Daams, R., Freudenthal, M., 1981. Aragonian: the stage concept versus Neogene](#)
1139 [Mammal zones. Scripta Geologica 62, 1–17.](#)
- 1140 [Daams, R., Freudenthal, M., 1988. Cricetidae \(Rodentia\) from the type-Aragonian;](#)
1141 [the genus *Megacricetodon*. Scripta Geologica Spec. Issue 1, 39–132.](#)
- 1142 [Daams, R., Freudenthal, M., 1990. The Ramblian and the Aragonian: Limits,](#)
1143 [subdivisions. In: Lindsay, E.H., Fahlbush, V., Mein, P. \(Eds.\), European Neogene](#)
1144 [Mammal Chronology. Plenum Press, New York, pp. 51–59.](#)
- 1145 [Daams, R., Freudenthal, M., van de Weerd, A., 1977. Aragonian, a new stage for](#)
1146 [continental deposits of Miocene age. Newsletters on Stratigraphy 6, 42–55.](#)
- 1147 [Daams, R., Van der Meulen, A.J., Álvarez Sierra, M.A., Peláez-Campomanes, P.,](#)
1148 [Calvo, J. P., Alonso Zarza, M.A., Krijgsman, W., 1999. Stratigraphy and](#)
1149 [sedimentology of the Aragonian \(Early to Middle Miocene\) in its type area \(North-](#)
1150 [Central Spain\). Newsletters on Stratigraphy 37, 103–139.](#)
- 1151 Daxner-Höck, G., Bernor, R.L., 2009. The early Vallesian vertebrates of Atzelsdorf
1152 (Late Miocene, Austria) 8. *Anchitherium*, Suidae and Castoridae (Mammalia).
1153 *Annalen des Naturhistorisches Museum in Wien* 111A, 557–584.
- 1154 de Bruijn, H., Daams, R., Daxner-Höck, G., Fahlbusch, V., Ginsburg, L., Mein, P.,
1155 Morales, J., Heinzmann, E., Mayhew, D.F., van der Meulen, A.J., Schmidt-Kittler,
1156 N., Telles Antunes, M., 1992. Report of the RCMNS working group on fossil

1157 mammals, Reinsenburg 1990. Newsletters on Stratigraphy 26, 65–118.

1158 [de Gibert, J.M., Casanovas-Vilar, I., 2011. Contexto geològic del Mioceno de la](#)
1159 [cuena del Vallès-Penedès. Paleontologia i Evolució memòria especial 6, 39–45.](#)

1160 [Evans, H.F., Westerhold, T., Paulsen, H., Channell, J.E.T., 2007. Astronomical ages](#)
1161 [for Miocene polarity chrons C4Ar–C5r \(9.3–11.2 Ma\), and for three excursion](#)
1162 [chrons within C5n.2n. Earth and Planetary Science Letters 256, 455–465.](#)

1163 [Fahlbusch, V., 1976. Report on the International Symposium on mammalian](#)
1164 [stratigraphy of the European Tertiary \(München, April 11-14, 1975\). Newsletters](#)
1165 [on Stratigraphy 5, 160–167.](#)

1166 Forstén, A.M., 1968. Revision of the Palearctic Hipparion. Acta Zoologica Fennica
1167 119, 1–130.

1168 Forstén, A., 1978. *Hipparion primigenium* (v. Meyer, 1829), an early three-toed
1169 horse. Annales Zoologici Fennici 15, 298–313.

1170 Furió, M., Prieto, J., Van den Hoek Ostende, L., 2015. Three million years of “Terror-
1171 Shrew” (*Dinosorex*, Eulipotyphla, Mammalia) in the Miocene of the Vallès-
1172 Penedès Basin (Barcelona, Spain). Comptes Rendus Palevol 14, 111–124.

1173 [Garcés, M., 1995. Magnetostratigrafía de las sucesiones del Mioceno Medio y](#)
1174 [Superior del Vallès Occidental \(Depresión del Vallès-Penedès, N.E. de España\):](#)
1175 [Implicaciones biocronológicas y cronoestratigráficas. Ph.D. Dissertation,](#)
1176 [Universitat de Barcelona.](#)

1177 Garcés, M., Agustí, J., Cabrera, L., Parés, J.M., 1996. Magnetostratigraphy of the
1178 Vallesian (late Miocene) in the Vallès-Penedès Basin (northeast Spain). Earth and
1179 Planetary Science Letters 142, 381–396.

1180 Garcés, M., Cabrera, L., Agustí, J., Parés, J.M., 1997. Old World first appearance
1181 datum of “Hipparion” horses: Late Miocene large-mammal dispersal and global

1182 events. *Geology* 25, 19–22.

1183 Garcés, M., Krijgsman, W., Peláez-Campomanes, P., Álvarez Sierra, M.A., Daams,
1184 R., 2003. Hipparion dispersal in Europe: magnetostratigraphic constraints from the
1185 Daroca area (Spain). *Coloquios de Paleontología* Vol. Ext. 1, 171–178.

1186 García-Paredes, I., Álvarez-Sierra, M.Á., Van den Hoek Ostende, L., Hernández-
1187 Ballarín, V., Hordijk, K., López-Guerrero, P., Oliver, A., Peláez-Campomanes, P.,
1188 2016. The Aragonian and Vallesian high-resolution micromammal succession from
1189 the Calatayud-Montalbán Basin (Aragón, Spain). *Comptes Rendus Palevol* 15,
1190 781–789.

1191 Gibert, J., 1975. New insectivores from the Miocene of Spain. I. Proceedings of the
1192 Koninklijke Nederlandse Akademie van Wetenschappen B 78, 108–123.

1193 Ginsburg, L., 1986. Chronology of the European pliopithecids. In: Else, J.C., Lee,
1194 P.C. (Eds.), *Primate Evolution*. Cambridge University Press, Cambridge, pp. 47–
1195 57.

1196 Golpe-Posse, J.M., 1972. Suiformes del Terciario español y sus yacimientos (Tesis
1197 doctoral-Resumen) (revisado y reimprimido en Diciembre de 1972). *Paleontología*
1198 *y Evolución* 2, 1–197.

1199 Golpe-Posse, J.M., 1974. Faunas de yacimientos con suiformes en el Terciario
1200 español. *Paleontología y Evolución* 8, 1–87.

1201 Golpe-Posse, J.M., 1975. Un nuevo tayasuido en el Vindoboniense terminal de Castell
1202 de Barberà (Cuenca del Vallès, España). *Boletín Informativo del Instituto de*
1203 *Paleontología de Sabadell* 7, 39–43.

1204 Golpe-Posse, J. M., 1977. *Barberahyus castellensis* n.g.n.sp., Tayasuido del
1205 Vindoboniense terminal de Castell de Barberà (Cuenca del Vallès, España).
1206 *Paleontología y Evolución* 12, 31–43.

- 1207 Golpe-Posse, J.M., 1978. Un nuevo tayasuido en el Vindoboniense terminal de Castell
1208 de Barberà (Cuenca del Vallès, España). Boletín Informativo del Instituto de
1209 Paleontología de Sabadell 13, 13–17.
- 1210 Golpe Posse, J.M., 1982. Los hispanopitecos (Primates, Pongidae) de los yacimientos
1211 del Vallès Penedès (Cataluña - España). I: Material ya descrito. Butlletí Informatiu
1212 de l'Institut de Paleontologia de Sabadell 14, 63–69.
- 1213 Golpe Posse, J.M., 1993. Los Hispanopitecos (Primates, Pongidae) de los yacimientos
1214 del Vallès-Penedès (Cataluña, España). II: Descripción del material existente en el
1215 Instituto de Paleontología de Sabadell. Paleontologia i Evolució 26–27, 151–224.
- 1216 Harrison, T., 1991. Some observations on the Miocene hominoids from Spain. Journal
1217 of Human Evolution 19, 515–520.
- 1218 Harzhauser, M., 2009. The early Vallesian vertebrates of Atzelsdorf (Late Miocene,
1219 Austria) 2. Geology. Annalen des Naturhistorischen Museums in Wien 111A, 479–
1220 488.
- 1221 ICGC (Institut Cartogràfic i Geològic de Catalunya), 2018. VISSIR v3.26 (VISor del
1222 Servidor d'Imatges Ràster). Generalitat de Catalunya, <http://www.icc.cat/vissir3/>
1223 (accessed July 1st 2018).
- 1224 ICZN (International Commission on Zoological Nomenclature), 1999. International
1225 Code of Zoological Nomenclature, 4th ed. The International Trust for Zoological
1226 Nomenclature, London.
- 1227 Kirschvink, J.L., 1980. The least-squares line and plane and the analysis of
1228 palaeomagnetic data. Geophysical Journal of the Royal Astronomical Society 62,
1229 699–718.
- 1230 Krijgsman, W., Langereis, C.G., Daams, R., van der Meulen, A.J., 1994.
1231 Magnetostratigraphic dating of the middle Miocene climate change in the

1232 continental deposits of the Aragonian type area in the Calatayud-Teruel basin
1233 (Central Spain). Earth and Planetary Science Letters 128, 513–526.

1234 Lindsay, E.H., Tedford, R.H., 1990. Development and application of land mammal
1235 ages in North American and Europe, a comparison. In: Lindsay, E.H., Fahlbusch,
1236 V., Mein, P. (Eds.), European Neogene Mammal Chronology. Plenum Press, New
1237 York, pp. 601–624.

1238 MacFadden, B. J. (2001). Three-toed browsing horse *Anchitherium clarencei* from the
1239 early Miocene (Hemingfordian) Thomas Farm, Florida. *Bulletin of the Florida*
1240 *Museum of Natural History*, 43, 79-109.

1241 Marigó, J., Susanna, I., Minwer-Barakat, R., Madurell-Malapeira, J., Moyà-Solà, S.,
1242 Casanovas-Vilar, I., Robles, J.M., Alba, D.M., 2014. The primate fossil record in
1243 the Iberian Peninsula. *Journal of Iberian Geology* 40, 179–211.

1244 Marmi, J., Casanovas-Vilar, I., Robles, J.M., Moyà-Solà, S., Alba, D.M., 2012. The
1245 paleoenvironment of *Hispanopithecus laietanus* as revealed by paleobotanical
1246 evidence from the Late Miocene of Can Llobateres 1 (Catalonia, Spain). *Journal of*
1247 *Human Evolution* 62, 412–423.

1248 Moyà-Solà, S., 1983. Los Boselaphini (Bovidae Mammalia) del Neógeno de la
1249 Península Ibérica. *Publicaciones de Geología*, Universitat Autònoma de Barcelona.

1250 Moyà Solà, S., Pons Moyà, J., Köhler, M., 1990. Primates catarrinos (Mammalia) del
1251 Neógeno de la península Ibérica. *Paleontologia i Eolució* 23, 41–45.

1252 Moyà-Solà, S., Köhler, M., Rook, L., 2005. The *Oreopithecus* thumb: a strange case
1253 in hominoid evolution. *Journal of Human Evolution* 49, 395–404.

1254 Moyà-Solà, S., Köhler, M., Alba, D.M., Casanovas-Vilar, I., Galindo, J., Robles,
1255 J.M., Cabrera, L., Garcés, M., Almécija, S., Beamud, E., 2009. First partial face
1256 and upper dentition of the Middle Miocene hominoid *Dryopithecus fontani* from

1257 Abocador de Can Mata (Vallès-Penedès Basin, Catalonia, NE Spain): taxonomic
1258 and phylogenetic implications. *American Journal of Physical Anthropology* 139,
1259 126–145.

1260 Moyà-Solà, S., Alba, D.M., Almécija, S., 2013. A proximal radius of *Barberapithecus*
1261 *huerzeleri* (Primates, Pliopithecidae) from the Miocene site of Castell de Barberà
1262 (NE Iberian Peninsula). *Journal of Vertebrate Paleontology* 33 S2, 182.

1263 [Murphy, M.A., Salvador, A. \(Eds.\), 1999. International Stratigraphic Guide — An](#)
1264 [abridged version. Episodes 22, 255–271.](#)

1265 [Ogg, J.G., 2012. Geomagnetic polarity time scale. In: Gradstein, F.M., Ogg, J.G.,](#)
1266 [Schmitz, M.D., Ogg, G.M. \(Eds.\), The Geologic Time Scale 2012, Volume 1.](#)
1267 [Elsevier, Amsterdam, pp. 85–113.](#)

1268 Pérez de los Ríos, M., Alba, D.M., Moyà-Solà, S., 2013. Taxonomic attribution of the
1269 La Grive hominoid teeth. *American Journal of Physical Anthropology* 151, 558–
1270 565.

1271 Pesquero, M.D., Arribas, A., 2002. Los restos de *Hipparion* (Equidae, Mammalia) en
1272 las colecciones de vertebrados del Museo Geominero (IGME): aspectos históricos
1273 y actualización taxonómica. *Boletín Geológico y Minero* 113, 97–108.

1274 Petter, G., 1976. Étude d'un nouvel ensemble de petits carnivores du Miocène
1275 d'Espagne. *Géologie Méditerranéenne* 3, 135–154.

1276 Pirlot, P.L., 1956. Les formes européennes du genre *Hipparion*. *Memorias y*
1277 *Comunicaciones del Instituto Geológico* 14, 1–121.

1278 [Roberts, A.P., Lewin-Harris, J.C., 2000. Marine magnetic anomalies: evidence that](#)
1279 ['tiny wiggles' represent short-period geomagnetic polarity intervals. Earth and](#)
1280 [Planetary Science Letters 183, 375–388.](#)

- 1281 Robles, J. M., Alba, D.M., Moyà-Solà, S., Casanovas-Vilar, I., Galindo, J., Rotgers,
1282 C., Almécija, S., Carmona, R., 2010. New craniodental remains of *Trocharion*
1283 *albanense* Major, 1903 (Carnivora, Mustelidae), from the Vallès-Penedès Basin
1284 (Middle to Late Miocene, Barcelona, Spain). *Journal of Vertebrate Paleontology*
1285 30, 547–562.
- 1286 Robles, J.M., Alba, D.M., Casanovas-Vilar, I., Galindo, J., Cabrera, L., Carmona, R.,
1287 Moyà-Solà, S., 2011. On the age of the paleontological site of Can Missert
1288 (Terrassa, Vallès-Penedès Basin, NE Iberian Peninsula). In: Pérez-García, A.,
1289 Gascó, J., Gasulla, J.M., Escaso, F. (Eds.) *Viajando a Mundos Pretéritos*.
1290 Ayuntamiento de Morella, Morella, pp. 339–346.
- 1291 Robles, J.M., Alba, D.M., Fortuny, J., De Esteban-Trivigno, S., Rotgers, C., Balaguer,
1292 J., Carmona, R., Galindo, J., Almécija, S., Bertó, J.V., Moyà-Solà, S., 2013a. New
1293 craniodental remains of the barbourfelid *Albanosmilus jourdani* (Filhol, 1883)
1294 from the Miocene of the Vallès-Penedès (NE Iberian Peninsula) and the phylogeny
1295 of the Barbourfelini. *Journal of Systematic Palaeontology* 11, 993–1022.
- 1296 Robles, J.M., Madurell-Malapeira, J., Abella, J., Rotgers, C., Carmona, R., Almécija,
1297 S., Balaguer, J., Alba, D.M., 2013b. New *Pseudaelurus* and *Styriofelis* remains
1298 (Carnivora: Felidae) from the middle Miocene of Abocador de Can Mata (Vallès-
1299 Penedès Basin). *Comptes Rendus Palevol* 12, 101–113.
- 1300 Rotgers, C., Alba, D.M., 2011. The genus *Anchitherium* (Equidae: Anchitheriinae) in
1301 the Vallès-Penedès Basin (Catalonia, Spain). In: Pérez-García, A., Gascó, J.,
1302 Gasulla, J.M., Escaso, F. (Eds.), *Viajando a Mundos Pretéritos*. Ayuntamiento de
1303 Morella, Morella, pp. 347–354.
- 1304 Rotgers, C., Alba, D.M., Robles, J.M., Casanovas-Vilar, I., Galindo, J., Bertó, J.V.,
1305 Moyà-Solà, S., 2011. A new species of *Anchitherium* (Equidae: Anchitheriinae)

1306 from the Middle Miocene of Abocador de Can Mata (Vallès-Penedès Basin, NE
1307 Iberian Peninsula). *Comptes Rendus Palevol* 10, 567–576.

1308 Salesa, M.J., Sánchez, I.M., Morales, J., 2004. Presence of the Asian horse *Sinohippus*
1309 in the Miocene of Europe. *Acta Palaeontologica Polonica* 49, 189–196.

1310 Sánchez, I.M., Salesa, M.J., Morales, J., 1998. Revisión sistemática del género
1311 *Anchitherium* Meyer 1834 (Equidae; Perissodactyla) en España. *Estudios*
1312 *Geológicos* 54, 39–63.

1313 Santafé Llopis, V., 1978. Rinocerótidos fósiles de España. PhD. dissertation,
1314 Universidad de Barcelona.

1315 Santafé i Llopis, J.V., 1978a. Síntesi de la distribució dels rinoceròtids fòssils a les
1316 conques Vallès-Penedès. *Butlletí Informatiu de l'Institut de Paleontologia de*
1317 *Sabadell* 10, 34–40.

1318 Santafé Llopis, J.V., 1978b. Revisión de los Rinocerótidos miocénicos del Vallès-
1319 Penedès. *Acta Geológica Hispánica* 13, 43–45.

1320 Steininger, F.F., 1999. Chronostratigraphy, geochronology and biochronology of the
1321 Miocene "European Land Mammal Mega-Zones" (ELMMZ) and the Miocene
1322 "Mammal-Zones (MN-Zones)". In: Rössner, G.E., Heissig, K. (Eds.), *The Miocene*
1323 *Land Mammals of Europe*. Verlag Fritz Pfeil, München, pp. 9–24.

1324 Sukselainen, L., Fortelius, M., Harrison, T., 2015. Co-occurrence of pliopithecoid and
1325 hominoid primates in the fossil record: An ecometric analysis. *Journal of Human*
1326 *Evolution* 84, 25–41.

1327 Urciuoli, A., Zanolli, C., Fortuny, J., Almécija, S., Schillinger, B., Moyà-Solà, S.,
1328 Alba, D.M., 2018. Neutron-based computed microtomography: *Pliobates*
1329 *cataloniae* and *Barberapithecus huerzeleri* as a test-case study. *American Journal*
1330 *of Physical Anthropology* 166, 987–993.

- 1331 van Dam, J., 2003. European Neogene mammal chronology: past, present and future.
1332 Deinsea 10, 85–95.
- 1333 van der Made, J., Ribot, F., 1999. Additional hominoid material from the Miocene of
1334 Spain and remarks on hominoid dispersals into Europe. Contributions to Tertiary
1335 and Quaternary Geology 36, 25–39.
- 1336 van der Meulen, A.J., García-Paredes, I., Álvarez-Sierra, M.Á., van den Hoek
1337 Ostende, L. W., Hordijk, K., Oliver, A., López-Guerrero, P., Hernández-Ballarín,
1338 V., Peláez-Campomanes, P., 2011. Biostratigraphy or biochronology? Lessons
1339 from the Early and Middle Miocene small Mammal Events in Europe. Geobios 44,
1340 309–321.
- 1341 van der Meulen, A.J., García-Paredes, I., Álvarez-Sierra, M.Á., van den Hoek
1342 Ostende, L. W., Hordijk, K., Oliver, A., Peláez-Campomanes, P., 2012. Updated
1343 Aragonian biostratigraphy: Small mammal distribution and its implications for the
1344 Miocene European chronology. Geologica Acta 10, 159–179.
- 1345 Villa, A., Delfino, M., Luján, À.H., Almécija, S., Alba, D.M., 2017. First record of
1346 *Latonia gigantea* (Anura, Alytidae) from the Iberian Peninsula. Historical Biology.
1347 <https://doi.org/10.1080/08912963.2017.1371712>
- 1348 Woodburne, M.O. (Ed.), 2004a. Late Cretaceous and Cenozoic Mammals of North
1349 America. Biostratigraphy and Geochronology. Columbia University Press, New
1350 York.
- 1351 Woodburne, M.O., 2004b. Principles and procedures. In: Woodburne, M.O. (Ed.),
1352 Late Cretaceous and Cenozoic Mammals of North America. Biostratigraphy and
1353 Geochronology. Columbia University Press, New York, pp. 1–20.

- 1354 Woodburne, M.O., 2005. A new occurrence of *Cormohipparion*, with implications for
1355 the Old World *Hippotherium* datum. *Journal of Vertebrate Paleontology* 25, 256–
1356 257.
- 1357 Woodburne, M.O., 2007. Phyletic diversification of the *Cormohipparion occidentale*
1358 complex (Mammalia; Perissodactyla, Equidae), Late Miocene, North America, and
1359 the origin of the Old World *Hippotherium* datum. *Bulletin of the American*
1360 *Museum of Natural History* 306, 1–138.
- 1361 Woodburne, M.O., 2009. The early Vallesian vertebrates of Atzelsdorf (Late
1362 Miocene, Austria) 9. *Hippotherium* (Mammalia, Equidae). *Annalen des*
1363 *Naturhistorischen Museums in Wien* 111A, 585–604.
- 1364 Zouhri, S., Bensalmia, A., 2005. Révision systématique des *Hipparion* sensu lato
1365 (Perissodactyla, Equidae) de l'ancien monde. *Estudios Geológicos* 61, 61–99.

1366

1367 **Figure captions**

1368

- 1369 **Figure 1.** Fossil primates from Castell de Barberà: the pliopithecoid *Barberapithecus*
1370 *huerzeleri* (A–L) and the large-bodied hominoid (M–O). A–K) IPS1724 (holotype),
1371 selected associated teeth from a single female individual, including the right I¹ (A),
1372 the left C¹ (B), and the left C₁ (C), in lingual (left) and labial (right) views, as well as
1373 the right M¹ (D), the right M² (E), the right M³ (F), the right P₃ (G), the left P₄ (H), the
1374 left M₁ (I), the right M₂ (J), and the right M₃ (K), in occlusal views. L) IPS1823, male
1375 right C¹ in lingual (left) and labial (right) views. M) IPS4335, partial distal pollical
1376 phalanx, in palmar (left) and dorsal (right) views. N) IPS4333, right proximal pollical
1377 phalanx, from left to right in palmar, radial, ulnar, and distal views. O) IPS4334, distal
1378 fragment of right humeral diaphysis assigned to cf. *Dryopithecus fontani* by Alba et

1379 al. (2012), from left to right in anterior, medial, posterior, and lateral views. Scale bars
1380 equal 1 cm, except for O (5 cm). Images are reproduced from previous papers by the
1381 authors: L) Alba and Moyà-Solà (2012: Fig. 2); M–N) Almécija et al. (2012: Fig. 1);
1382 O) Alba et al. (2011: Fig. 1).

1383

1384 **Figure 2.** A) Aerial photograph ~~showing the location~~ of the ~~classical~~ site of Castell de
1385 Barberà (CB), as well as that of the bottom of sections sampled ~~s~~ for
1386 magnetostratigraphy in this study. Modified from base orthophotos downloaded from
1387 VISSIR v3.26 (ICGC, 2018: sheets 288-119 and 288-120, scale 1:5000), ©Institut
1388 Cartogràfic i Geològic de Catalunya, with permission ~~allowed by from~~ licence
1389 Creative Commons (CC) – Attribution 4.0 International (CC BY 4.0; see
1390 <http://www.icgc.cat/Ajuda/Avis-legal> for ~~the~~ reuse policies allowed for ICGC web
1391 contents). B) Detail of Castell de Barberà CB (Section 2) during the excavation of CB-
1392 layer-D (equivalent to the original main ~~classical~~ fossiliferous level layer) in June
1393 2015.

1394

1395 **Figure 3.** Simplified geological map of the Vallès-Penedès Basin. The location of
1396 Castell de Barberà is denoted by a black star. ~~The T~~ top left inset: ~~shows the~~
1397 ~~situation~~ location of the Vallès-Penedès Basin within the Iberian Peninsula. Modified
1398 from Casanovas-Vilar et al. (2016a: Fig. 1).

1399

1400 **Figure 4.** Examples of NRM stepwise thermal demagnetization of samples of the
1401 Castell de Barberà composite section. Stratigraphic location of samples (uppercase
1402 letters) is indicated in Figure 5. Diagrams represent orthogonal projection of vector
1403 endpoint demagnetization data; black and white dots represent the projection on the

1404 horizontal and vertical planes, respectively. Red lines represent the least-square fit
1405 defining the paleomagnetic direction. Q1 and Q2 indicates ranked quality (see SOM
1406 Table S1).

1407

1408 **Figure 4.** Right distal humeral fragment of *Hippotherium* sp. IPS87652 from Castell
1409 de Barberà (layer D, equivalent to the classical fossiliferous level) recovered in 2015
1410 (A), as compared to those of *Heatalaunicum* IPS11117 from Can Llobateres (B) and
1411 IPS32449 from Polinyà (C, reversed). All specimens are depicted in anterior (left) and
1412 posterior (right) views.

1413

1414 **Figure 5.** Magnetostratigraphic results for the two sampled sections at Castell de
1415 Barberà (CB): A) Ssection 1 (CB s.l.); B) Ssection 2, where the site of Castell de
1416 Barberà (CB s.s.) is located. Sideways triangles denote the stratigraphic position of
1417 paleomagnetic samples. Red letters and triangles indicate the location of samples from
1418 Figure 4. Black circles indicate directions of higher quality (Q1 and Q2), while white
1419 circles indicate directions of low quality (Q3)—see SOM Table S1. Abbreviations: c
1420 = conglomerates; l = lutites; s = sandstones; VGP = vVirtual gGeomagnetic pPole.

1421

1422 **Figure 6.** Composite local magnetostratigraphic section ~~for~~ Castell de Barberà (CB)
1423 and its preferred correlation with the ~~Global geomagnetic pPolarity tTime sScale~~
1424 (~~GPTS~~; Ogg, 2012), as well as European land mammal ages (ELMA), MN (~~Mammal~~
1425 ~~Neogene~~) units, and local biozones of the Vallès-Penedès Basin (after Casanovas-
1426 Vilar et al., 2016b). See Figure 5 for ~~the~~ detailed magnetostratigraphic results of the
1427 two sampled sections upon which this composite magnetostratigraphic section is
1428 based. Abbreviations: ~~CB1 = CB s.l. (section 1); CB2 = CB s.s. (section 2); N =~~

1429 normal polarity magnetozones; R = reversed polarity magnetozones. The arrow next
1430 to N2 denotes the stratigraphic position of the main ~~classical~~-fossiliferous level-layer
1431 of Castell de Barberà (CB-D). Half-width polarity intervals in the GPTS represent
1432 short geomagnetic excursions after Evans et al. (2007). Short horizontal lines to the
1433 left of the polarity column represent ‘tinny wiggles’ from the sea floor magnetic
1434 anomaly stacks (Cande and Kent 1992), later interpreted as true geomagnetic polarity
1435 reversals (Roberts and Lewin-Harris, 2000).

1436

1437

1438 **Figure 7.** Right distal humeral fragment of *Hippotherium* sp. IPS87652 from Castell
1439 de Barberà CB-D (equivalent to the main fossiliferous layer) recovered in 2015 (A),
1440 compared to those of *Hippotherium catalaunicum* IPS11117 from Can Llobateres (B)
1441 and IPS32449 from Polinyà (C, reversed). All specimens are depicted in anterior (left)
1442 and posterior (right) views.

1443

1 Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of
2 Castell de Barberà to the earliest Vallesian: ~~End of the controversy~~

3

4 David M. Alba^{a,*}, Miguel Garcés^{b,c}, Isaac Casanovas-Vilar^a, Josep M. Robles^a, Marta
5 Pina^{d,a}, Salvador Moyà-Solà^{a,e,f}, Sergio Almécija^{g,h,a,*}

6

7 ^a *Institut Català de Paleontologia Miquel Crusafont, Universitat Autònoma de*
8 *Barcelona, c/ Columnes s/n, Campus de la UAB, 08193 Cerdanyola del Vallès,*
9 *Barcelona, Spain*

10 ^b *Departament de Dinàmica de la Terra i de l'Oceà, Facultat de Ciències de La*
11 *Terra, Universitat de Barcelona, c/ Martí i Franqués s/n, 08028, Barcelona, Spain*

12 ^c *Institut Geomodels, Grup de Recerca Consolidat de Geodinàmica i Anàlisi de*
13 *Conques, Universitat de Barcelona, c/ Martí i Franqués s/n, 08028, Barcelona, Spain*

14 ^d *Department of Zoology, Graduate School of Science, Kyoto University, Sakyo,*
15 *Kyoto, 606-8502, Japan*

16 ^e *Institució Catalana de Recerca i Estudis Avançats (ICREA), Pg. Lluís Companys 23,*
17 *08010, Barcelona, Spain*

18 ^f *Unitat d'Antropologia Biològica, Departament de Biologia Animal, Biologia Vegetal*
19 *i Ecologia, Universitat Autònoma de Barcelona, 08193 Cerdanyola del Vallès,*
20 *Barcelona, Spain*

21 ^g *Division of Anthropology, American Museum of Natural History, Central Park West*
22 *at 79th Street, New York, NY 10024, USA*

23 ^h *New York Consortium in Evolutionary Primatology, New York, USA*

24

25 *Corresponding authors.

26 *E-mail address:* david.alba@icp.cat (D.M. Alba); salmecija@amnh.org (S. Almécija).

27

28 **Acknowledgments**

29 We are especially grateful to the rest of the field team members for their continued
30 efforts (J. Manel Méndez, Manel Llenas, Àngel H. Luján, Ashley Hammond and
31 Victor Vinuesa). We also thank M. March for assistance during the study of the ICP
32 collections from Castell de Barberà, the staff of the Preparation Area of the ICP for
33 the preparation of the fossil specimens, [J. Morales for sending us pictures of](#)
34 [Anchitherium](#), and the late J.V. Santafé for sharing with us his remembrances about
35 the location of the site. This work has been funded by the Leakey Foundation, the
36 Agencia Estatal de Investigación (CGL2016-76431-P and CGL2017-82654-P,
37 MINECO AEI/FEDER EU; and RYC-2013-12470 to I.C.V.), and the Generalitat de
38 Catalunya (fieldwork grant 2014/100609 and CERCA Programme). Some of the
39 authors of this paper are members of the consolidated research groups 2017 SGR 116
40 GRC (D.M.A., I.C.V., J.M.R.) and 2017 SGR 86 GRC (S.M.S.) of the Generalitat de
41 Catalunya. We further acknowledge the support and cooperation of the Ajuntament de
42 Barberà del Vallès and the Servei d'Arqueologia i Paleontologia of the Generalitat de
43 Catalunya. [Finally, we thank Mike Plavcan \(editor\), Andrews P. Roberts \(reviewer\),](#)
44 [and an anonymous reviewer for helpful comments and suggestions that helped us to](#)
45 [improve a previous version of this paper.](#)

46

1 Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of
2 Castell de Barberà to the earliest Vallesian

3

4 **Keywords:** Hominoidea; Pliopithecoidea; *Hippotherium*; Late Miocene;
5 Magnetostratigraphy; Paleomagnetism.

6

7 **Abstract**

8 Castell de Barberà, located in the Vallès-Penedès Basin (NE Iberian Peninsula), is
9 one of the few European sites where pliopithecoids (*Barberapithecus*) and hominoids
10 (cf. *Dryopithecus*) co-occur. The dating of this Miocene site has proven controversial.
11 A latest Aragonian (MN7+8, ca. 11.88–11.18 Ma) age was long accepted by most
12 authors, despite subsequent reports of hipparionin remains that signaled a Vallesian
13 age. On the latter basis, Castell de Barberà was recently correlated to the early
14 Vallesian (MN9, ca. 11.18–10.3 Ma) on tentative grounds. Uncertainties about the
15 provenance of the *Hippotherium* material and the lack of magnetostratigraphic data
16 precluded more accurate dating. After decades of inactivity, fieldwork was resumed in
17 2014–2015 at Castell de Barberà, including the original layer (CB-D) that in the past
18 delivered most of the fossils. Here we report magnetostratigraphic results for the
19 original outcrop and another nearby section. Our results indicate that CB-D is located
20 in a normal polarity magnetozone at about midheight of a short (~20 m-thick)
21 stratigraphic section. The composite magnetostratigraphic section (~50 m) has as
22 many as four to six magnetozones. These multiple reversals, coupled with the in situ
23 recovery of a *Hippotherium* humerus from CB-D in 2015, make it very unlikely the
24 correlation of any of the sampled normal polarity magnetozones with the long normal
25 polarity subchron C5n.2n (11.056–9.984 Ma), which is characteristic of the early

26 Vallesian. Our results support instead a correlation of CB-D with C5r.1n (11.188–
27 11.146 Ma), where the Aragonian/Vallesian boundary is situated, and therefore
28 indicate an earliest Vallesian age of ~11.2 Ma for Castell de Barberà. Our results
29 settle the longstanding debate about the Aragonian vs. Vallesian age of this site,
30 which appears roughly coeval with the Creu de Conill 20 locality (11.18 Ma), where
31 hipparionins are first recorded in the Vallès-Penedès Basin.

32

33 **1. Introduction**

34 *1.1. Primates from Castell de Barberà*

35 The Miocene primate-bearing site of Castell de Barberà is situated along the left
36 bank of the Ripoll River, near the old farmhouse of Ca n'Altimira, in the municipality
37 of Barberà del Vallès (Catalonia, Spain). The site was named after a former medieval
38 castle (11th century) that was subsequently reconstructed as a country house.
39 According to museum labels, the site was mostly excavated from 1965 to 1981 by
40 personnel from the former Instituto Provincial de Paleontología in Sabadell, currently
41 the Institut Català de Paleontologia Miquel Crusafont (ICP), although the site was
42 already mentioned in the catalog published by Crusafont and Truyols (1954).
43 Crusafont Pairó and Hürzeler (1969) first reported the find of primate remains, and
44 Crusafont-Pairó and Golpe (1972) later provided a summary account of the site,
45 noting that methodical excavations had been carried out there. Although publications
46 based on the previously collected material from Castell de Barberà have continued
47 until recently, the site was not excavated for many years, until it was reopened in
48 2014 and 2015 under the direction of author#6 and author#7.

49 Among the Miocene primate-bearing sites of Europe, Castell de Barberà is
50 renowned for being one of the few in which pliopithecoids and hominoids co-occur

51 (Moyà-Solà et al., 1990; Andrews et al., 1996; Alba, 2012; Sukselainen et al., 2015).
52 It is unknown whether the hominoid remains come from the same layer as the
53 pliopithecoid remains, which are attributed to the main fossiliferous layer that also
54 yielded most of the fossils from the site (Alba and Moyà-Solà, 2012). However, the
55 shortness of the stratigraphic section, coupled with the lack of obvious hiatuses,
56 implies that all of the fossils from the site are roughly coeval.

57 Hominoids were first reported from Castell de Barberà by Crusafont Pairó and
58 Hürzeler (1969), based on a purportedly female right upper canine (IPS1823 [IPS26
59 in old terminology]; Fig. 1L) that was subsequently attributed (either implicitly or
60 explicitly) to *Hispanopithecus laietanus* (sometimes included in *Dryopithecus*;
61 Crusafont-Pairó and Casanovas Cladellas, 1973; Golpe-Posse, 1974; Golpe Posse,
62 1982, 1993; Harrison, 1991). Even though other authors later assigned this specimen
63 to a male pliopithecoid (Begun, 2002a; Alba et al., 2011; Almécija et al., 2011, 2012;
64 Alba and Moyà-Solà, 2012), hominoids are also recorded at Castell de Barberà by
65 several postcranial remains. These were found among the museum collections in the
66 late 1980s, as reported preliminarily by Moyà-Solà et al. (1990), and subsequently
67 described in detail. The hominoid specimens consist of: a distal humeral diaphysis
68 (IPS4334; Fig. 1O; Alba et al., 2011; Almécija et al., 2011), a proximal pollical
69 phalanx (IPS4333; Fig. 1N; Moyà-Solà et al., 2005; Almécija et al., 2012), and a
70 distal pollical phalanx (IPS4335; Fig. 1M; Almécija et al., 2012). Based on the large
71 body mass inferred from the humeral shaft fragment, the hominoid from Castell de
72 Barberà was tentatively referred to cf. *Dryopithecus fontani* (Alba et al., 2011;
73 Almécija et al., 2011; Alba, 2012; Marigó et al., 2014), although such a tentative
74 attribution cannot be further substantiated due to the lack of craniodental specimens
75 (Almécija et al., 2012).

76 The pliopithecoid from Castell de Barberà, in turn, was first reported by Crusafont-
77 Pairó (1975), based on several upper and lower teeth of a single individual (Fig. 1A–
78 K) that was initially referred to *Pliopithecus* sp. Afterward, the lower dentition was
79 described by Crusafont-Pairó and Golpe-Posse (1981a, 1982a), who similarly
80 assigned the material to *Pliopithecus* sp., albeit noting similarities with *Pliopithecus*
81 *lockeri* (currently in *Plesiopliopithecus*, i.e., a crouzeliid; Begun, 2002a; Alba and
82 Berning, 2013). Some later authors advocated crouzeliid affinities for the Castell de
83 Barberà pliopithecoid (Ginsburg, 1986), but most considered it to be a pliopithecoid
84 (Andrews et al., 1996; Begun, 2002a; Harrison et al., 2002). Some authors even
85 attributed the material (although usually tentatively) to *Pliopithecus antiquus* (Moyà-
86 Solà et al., 1990; Andrews et al., 1996; Harrison et al., 2002), whereas others
87 considered it to represent a new (unnamed) species (Ginsburg, 1986; Begun, 2002a;
88 Alba et al., 2010). Finally, a new genus and species, *Barberapithecus huerzeleri*, was
89 described by Alba and Moyà-Solà (2012) to accommodate the currently available
90 dental remains, which belong to three different individuals: the holotype (IPS1724;
91 Fig. 1A–K), including 15 upper and lower teeth from a single individual (see also
92 Urciuoli et al., 2018); the paratype (IPS34548), consisting of a P₃ from another
93 individual; and IPS1823, the aforementioned male upper canine (Fig. 1L), which was
94 included in the hypodigm but not designated as a paratype. Finally, Moyà-Solà et al.
95 (2013) preliminarily reported a proximal fragment of right radius (IPS66267) assigned
96 to *Barberapithecus*, found by Lars van den Hoek Ostende while revising the
97 collections, although this specimen has yet to be described in detail.

98 During the Miocene, hominoids and pliopithecoids did not co-occur at many sites,
99 which constitutes an apparent, although poorly-understood, pattern that has received
100 attention in recent publications (Sukselainen et al., 2015; Alba et al., 2017, and

101 references therein). Both groups are represented in late Aragonian and Vallesian sites
102 of the Vallès-Penedès Basin (see subsection 1.2. below for further explanation on
103 these terms), although they seldom come from the same locality (exact stratigraphic
104 horizon), even when their ranges are known to overlap as in Abocador de Can Mata
105 (Alba et al., 2017). This fact might be due to a sampling artifact related to the rarity of
106 these primates (Andrews et al., 1996), or it might be related to different ecological
107 preferences between these groups (e.g., Sukselainen et al., 2015). As one of the few
108 sites recording both groups and amenable to paleoecological analyses based on the
109 accompanying fauna, Castell de Barberà has the potential to shed light on the
110 particular conditions that enabled the infrequent coexistence of hominoids and
111 pliopithecoids throughout the European Miocene.

112

113 *1.2. Vallesian and Aragonian European land mammal ages*

114 The contribution of Castell de Barberà to a broader understanding of primate
115 evolution during the European Miocene in general, and the succession of primate taxa
116 at the Vallès-Penedès Basin in particular, has been hampered by a longstanding
117 controversy about the age of this site—whether it correlates to the Aragonian or with
118 the Vallesian European land mammal ages (ELMAs). Some clarification is required
119 with regard to the concept of ELMAs (for a historical review, see Lindsay and
120 Tedford, 1990 and Van Dam, 2003), with particular emphasis on the use of
121 ‘Aragonian’ instead of Orleanian + Astaracian, despite the latter are more widely used
122 outside Spain. The use of ELMAs as geochronological units has been criticized (e.g.,
123 Steininger, 1999) on the grounds that they are regional and that, with few exceptions,
124 they do not correspond to properly defined chronostratigraphic units (stages).
125 ‘Mammal ages’ are generally conceptualized as “biochronologic units” (e.g.,

126 Woodburne, 2004a: xiv; see also Hilgen et al., 2012) of regional applicability due to
127 divergent paleobiogeographic histories among regions (Lindsay and Tedford, 1990).
128 However, an initial definition of land mammal ages as biochronologic units (i.e.,
129 biozones) is not mutually exclusive with their subsequent formal definition (based on
130 bio- and magnetostratigraphic data) as chronostratigraphic units (i.e., stages based on
131 bodies of rock formed during a given time interval; e.g., Garcés, 1995; Woodburne,
132 2004), which automatically implies the definition of their corresponding
133 geochronologic units (ages).

134 The definition of regional chronostratigraphic units is not at odds with the
135 International Stratigraphic Guide, because “It is better to refer strata to local or
136 regional units with accuracy and precision rather than to strain beyond the current
137 limits of time correlation in assigning these strata to units of a global scale” (Murphy
138 and Salvador, 1999: 267). Even if not directly correlated with marine stages, regional
139 continental units can be dated based on radiometric and/or paleomagnetic methods
140 (e.g., Krijgsman et al., 1994; Garcés et al., 1996). Steininger (1999) conversely
141 advocated abandoning a chronostratigraphic/geochronologic concept of ELMAs
142 altogether and proposed to replace them by entirely biostratigraphic units (‘European
143 land mammal mega-zones’) based on MN (Mammal Neogene) biozones (Mein,
144 1975). This proposal is not exempt of problems given that MN zones were defined as
145 informal biochronologic units (Mein, 1975) and that their utility at a continental-wide
146 scale is restricted due to the significant diachrony of most mammal biochronologic
147 events (van der Meulen et al., 2011, 2012). This has led some authors to contend that
148 a formal European biozonation is not possible (van der Meulen et al., 2012), which
149 would imply that the regional nature of formally defined ELMAs such as the

150 Aragonian (see review in van der Meulen et al., 2012) would rather be an advantage
151 with regard to providing accurate correlations.

152 The lack of formal (biostratigraphic and chronostratigraphic) definition of widely
153 used biochronological units such as Orleanian and Astaracian is a problem that does
154 not apply to the roughly time-equivalent Aragonian, since like the younger Vallesian
155 (which enjoys a wider geographic applicability) it has been formally defined on the
156 basis of a specific stratotype. Crusafont Pairó (1950, 1951, 1953, 1995; see also
157 Crusafont Pairó and Truyols Santonja, 1954, 1955) first used the term Vallesian in a
158 largely biochronological sense to designate the Vallès-Penedès deposits with
159 *Hipparion* (currently *Hippotherium*) that, based on the fauna, appeared intermediate
160 in age between the sites of La Grive in France (i.e., the Aragonian) and Pikermi in
161 Greece (i.e., the Turolian). The Vallesian was finally more formally defined in
162 reference to Vallès-Penedès mammal successions by Crusafont Pairó and Truyols
163 Santonja (1960) based on the entry of *Hippotherium* as its main defining criterion.
164 The Vallesian was rapidly accepted throughout Eurasia as a simple solution for the
165 complex stratigraphic terminology in use, although it was not until decades later that
166 it was formally defined as a ‘mammal stage’ based on a specific stratotype from the
167 type area (the Vallès-Penedès Basin) within an accurate bio- and magnetostratigraphic
168 framework (Garcés, 1995; Garcés et al., 1996; Agustí et al., 1997). Following the
169 proposal of MN zones by Mein (1975), MN9 and MN10 had already been equated
170 with the early and late Vallesian, respectively, in turn subdivided into multiple local
171 biozones based on rodents (Agustí, 1981, 1982; Agustí and Moyà-Solà, 1991; Agustí
172 et al., 1997; Casanovas-Vilar et al., 2011a, 2016b).

173 The Aragonian was in turn originally conceptualized (Falhbusch, 1976) as a
174 chronostratigraphic unit defined by the presence of the equid *Anchitherium* and the

175 lack of the more derived equid *Hipparion* (currently *Hippotherium*), and subdivided
176 into two subunits, the Orleanian and the Astaracian. Soon thereafter, the Aragonian
177 was defined by Daams et al. (1977) as a new stage for continental middle Miocene
178 deposits preceding the Vallesian, with its stratotype located within the Calatayud-
179 Montalbán Basin in Spain (see also Daams et al., 1999). However, Daams et al.
180 (1977) refrained from dividing the Aragonian into Orleanian and Astaracian because
181 the latter had yet to be formally defined, ultimately leading to a tripartite subdivision
182 (Daams and Freudenthal, 1981; Daams and Freudenthal, 1990). Moreover, the
183 original criterion used to define the base of the Aragonian (the dispersal of
184 *Anchitherium*) was soon questioned (Daams and Freudenthal, 1981) and eventually
185 abandoned following the definition of the Ramblian stage (Daams et al., 1987; see
186 also Daams and Freudenthal, 1990). Further stratigraphic refinements of the
187 Aragonian were later provided by Daams et al. (1987, 1999) and van der Meulen et al.
188 (2012), who further distinguished multiple Aragonian local biozones based on
189 rodents. Based on these works, currently the Aragonian may be considered a regional
190 mammal-based chronostratigraphic unit (stage), whose scope is limited to the
191 continental record from Spain (see recent reviews in van der Meulen et al., 2012 and
192 García-Paredes et al., 2016). Given the scope of this paper and the detailed local
193 zonation of the Aragonian available for the Vallès-Penedès Basin (Casanovas-Vilar et
194 al., 2016b), we refrain from using the alternative and more loosely-defined term
195 ‘Astaracian’.

196 Based on high-resolution magnetostratigraphic correlation to the geomagnetic
197 polarity time scale (GPTS), according to Van der Meulen et al. (2012) the early
198 Aragonian (ca. 17.2–15.9 Ma) corresponds to the late Burdigalian (early Miocene,
199 MN4), while the middle Aragonian (ca. 15.9–13.8 Ma) comprises most of the

200 Langhian (early to middle Miocene, MN5), and the late Aragonian (ca. 13.8–11.2 Ma,
201 roughly equivalent to the ‘Astaracian’) covers the latest Langhian, the Serravallian,
202 and the earliest Tortonian (middle to late Miocene, MN6 to MN7+8). The Vallesian,
203 in turn, entirely corresponds to the Tortonian (late Miocene), being subdivided into
204 early Vallesian (11.2–10.0 Ma, MN9) and late Vallesian (10.0–8.9 Ma, MN10; e.g.,
205 Hilgen et al., 2012; Casanovas-Vilar et al., 2016b).

206

207 *1.3. The controversial age of Castell de Barberà*

208 Originally, Crusafont and Truyols (1954), Crusafont-Pairó and Hürzeler (1969)
209 and Crusafont-Pairó and Golpe (1972) advocated for a pre-Vallesian age for Castell
210 de Barberà because no hipparionin remains were originally found. Crusafont-Pairó
211 and Golpe (1972), Crusafont-Pairó (1972), and Golpe-Posse (1974) remarked that the
212 fauna appeared somewhat intermediate between those of the late Aragonian (MN7+8)
213 and those of the early Vallesian (MN9) localities of the Vallès-Penedès Basin (e.g.,
214 Trinxera del Ferrocarril in Sant Quirze vs. Can Llobateres, respectively). However,
215 given the lack of hipparionins and the presence of giraffids, the above-mentioned
216 authors correlated the site with the latest Aragonian. For many years, such a
217 correlation was further supported by Crusafont and coworkers (Crusafont Pairó and
218 Golpe Posse, 1972a; Golpe-Posse, 1972, 1974, 1975, 1977, 1978; Crusafont-Pairó and
219 Golpe Posse, 1972b, 1974, 1981a,b, 1982a,b; Petter, 1976; Santafé Llopis, 1978;
220 Santafé i Llopis, 1978a,b), even in the face of a surface-collected fragmentary fossil
221 of *Hippotherium* already reported by Crusafont-Pairó and Golpe-Posse (1974).
222 According to Crusafont-Pairó and Golpe-Posse (1974), the *Hippotherium* specimen
223 was somewhat rounded and probably eroded from the uppermost fossiliferous layer of
224 the Castell de Barberà section, whereas most of the fossil assemblage (lacking

225 *Hippotherium*) came from the main fossiliferous layer (situated at about the midheight
226 of the ~20 m-thick outcrop section). Crusafont-Pairó and Golpe-Posse (1974)
227 considered that the main fossiliferous layer was undoubtedly pre-Vallesian and only
228 contemplated two possible explanations for the presence of *Hippotherium* at the site:
229 either *Hippotherium* arrived in the Vallès-Penedès Basin during the latest Aragonian
230 (irrespective of whether the aforementioned fragment came from the main or the
231 upper fossiliferous layers); or the Aragonian/Vallesian boundary was located above
232 the main fossiliferous level along the upper portion of the Castell de Barberà section
233 (assuming the *Hippotherium* fragment came from the upper layer).

234 Needless to say, the first possibility discussed by Crusafont-Pairó and Golpe-Posse
235 (1974)—namely, a pre-Vallesian appearance of *Hippotherium*—is an ad hoc
236 explanation that does not hold upon closer scrutiny because it is at odds with the
237 original definition of the Vallesian proposed by Crusafont on the basis of the first
238 appearance datum of *Hippotherium* (formerly *Hipparion*; Crusafont Pairó, 1950,
239 1951, 1953; Crusafont Pairó and Truyols Santonja, 1960). The second explanation, in
240 contrast, remained a reasonable hypothesis that would be falsified by *Hippotherium*
241 remains from the main Castell de Barberà fossiliferous layer. There is a third possible
242 explanation, namely that Castell de Barberà could be Vallesian in age, even if
243 *Hippotherium* was scarce or absent.

244 For many years and up to the present, most subsequent authors favored a
245 correlation of Castell de Barberà with the late Aragonian (Moyà-Solà, 1983; Agustí et
246 al., 1985, 2001, 2013; Moyà-Solà et al., 1990; Begun, 2002a; van den Hoek Ostende
247 and Furió, 2005; Alba et al., 2006, 2010, 2011; Casanovas i Vilar, 2007; Casanovas-
248 Vilar et al., 2008, 2011a; Robles et al., 2010; Almécija et al., 2011; Sukselainen et al.,
249 2015; Agustí, 2018). The various alternate possibilities discussed above were

250 generally not discussed further, or were dismissed on the grounds that the specimen
251 was purportedly washed down from upper layers of the outcrop section (e.g., Santafé
252 Llopis, 1978; van den Hoek Ostende and Furió, 2005; Casanovas i Vilar, 2007; Alba
253 et al., 2011). However, it should be taken into account that the uppermost layers are
254 close (<10 m) to the main fossiliferous layer of Castell de Barberà (Casanovas-Vilar
255 et al., 2016b). The relatively short stratigraphic distance (110 m, according to
256 Crusafont Pairó and Truyols Santonja, 1951) between Castell de Barberà and Can
257 Llobateres 1 (with an interpolated age of 9.76 Ma; Casanovas-Vilar et al., 2016b) also
258 adds difficulty to the considerable age difference estimated between the two localities
259 (Alba and Moyà-Solà, 2012; Casanovas-Vilar et al., 2014)—which would require the
260 existence of major faults (e.g., Agustí et al., 1985) that have yet to be documented.
261 Some authors remained open to (Aguilar et al., 1979; Ginsburg, 1986), or even
262 supported (Gibert, 1975; de Bruijn et al., 1992; Andrews et al., 1996), an alternate
263 correlation of Castell de Barberà with MN9. However, it was not until publication of a
264 *Hippotherium* lower molar discovered among the museum collections of Castell de
265 Barberà (Rotgers and Alba, 2011) that most authors started to favor a Vallesian age
266 for the site (Casanovas-Vilar et al., 2011b, 2014, 2016a,b; Alba, 2012; Alba and
267 Moyà-Solà, 2012; Almécija et al., 2012; Robles et al., 2013a,b; Furió et al., 2015;
268 Villa et al., 2017). Casanovas-Vilar et al. (2014) even estimated an age of 10.55 Ma
269 for the site based of the stratigraphic distance between Castell de Barberà and Can
270 Llobateres reported by Crusafont Pairó and Truyols Santonja (1951). Casanovas-Vilar
271 et al. (2016a,b) then correlated Castell de Barberà to the *Hippotherium – Cricetulodon*
272 *hartenbergeri* interval subzone (11.88–10.3 Ma) of the Vallès-Penedès Basin without
273 further specifying its age. However, these authors considered that such a correlation,
274 which has yet to gain universal acceptance (e.g., Agustí et al., 2013; Sukselainen et

275 al., 2015; Agustí, 2018), must be considered tentative—given uncertainties about the
276 exact provenance of the *Hippotherium* remains from Castell de Barberà, which would
277 still allow a latest Aragonian age for the main fossiliferous layer.

278 With the aim of performing magnetostratigraphic analyses to more conclusively
279 settle the age of the Castell de Barberà sequence and hopefully lead to the find of
280 additional fossils, a team from the ICP led by author7 and author6 planned to reopen
281 the Castell de Barberà site in 2014. Finding the exact location of the site was
282 hampered by the vagueness of published indications (Crusafont-Pairó and Golpe,
283 1972; Golpe Posse, 1974), coupled with decades of vegetation growth and recent
284 anthropic modifications (to enlarge the riverbank's trackway). Two paleontological
285 campaigns were performed there in 2014 and 2015 with the aid of an excavator
286 machine (Fig. 2). Excavations in 2014 focused on a section (Section 1) situated less
287 than 300 m from the original outcrop (Section 2), which was excavated in 2015. The
288 exact geographic and stratigraphic situation of the original main fossiliferous layer in
289 Section 2 was located thanks to collaboration of the late Josep V. Santafé, who had
290 repeatedly excavated at the site during the 1970s with Crusafont, and indicated the
291 exact location while visiting the site during the 2015 season. Although our work
292 confirms that the bone accumulation of the main Castell de Barberà fossiliferous layer
293 is exhausted, we recovered additional fossil remains from the same layer and other
294 layers, and took paleomagnetic samples from the two excavated sections. In this
295 paper, we report the magnetostratigraphic data, and, with the aid of biostratigraphic
296 considerations, provide a more accurate dating of Castell de Barberà and discuss its
297 implications for Miocene hominoid and pliopithecoid evolution (for a preliminary
298 report of these data, see Alba et al., 2018a,b).

299

300 **2. Materials and methods**

301 *2.1. Location and stratigraphy of Castell de Barberà*

302 From a geological viewpoint (Fig. 3), Castell de Barberà is located on the distal
303 facies of the Castellar del Vallès alluvial fan system, which belong to the middle to
304 late Miocene upper continental complexes of the Vallès-Penedès Basin in NE Iberian
305 Peninsula (Agustí et al., 1985; Garcés, 1995; de Gibert and Casanovas-Vilar, 2011).
306 These outcrop sediments mostly consist of mudstones (especially siltstones and some
307 finer claystones) with intercalated coarse to fine sandstones (Santafé Llopis, 1978).
308 The repeated alternation of these lithologies along the stratigraphic series may be
309 interpreted as energy cycles—from higher energy at the base of each cycle, where
310 sandstones are more abundant, to lower energy toward the top, where fine siltstones
311 and claystones were deposited.

312 Most of the specimens from the collections of Castell de Barberà (mostly amassed
313 between 1965–1981) lack associated information. When preserved, some old field
314 labels refer to ‘layer 1’, and more sporadically to a so-called ‘higher layer’. Despite
315 this, it is possible to confidently assert that most of the material came from a single
316 layer (Santafé Llopis, 1978), which according to J.V. Santafé (pers. comm. to author⁶
317 and author⁷) was stratigraphically equivalent to layer CB-D excavated in 2015 (Fig.
318 2B). Santafé Llopis (1978) provided coordinates for the site (31N 428291 E –
319 4597272 N), indicating that it would be situated close to Ca n’Altimira. However, the
320 published coordinates are at odds with the description of the site location, and that

321 confirmed by Santafé in the field is situated more than 400 m southward (31N 428314
322 E – 4596862 N; Fig. 2A)¹.

323 Two magnetostratigraphic sections, located along the left bank of the Ripoll River
324 and exposed with the aid of digging machinery, were sampled for this study (Figs. 2A
325 and 4): Section 1 was excavated in 2014 close to Ca n'Altimira (bottom: 31 N 428232
326 E – 4597224 N), and Section 2 was excavated in 2015 at the original outcrop (bottom:
327 31N 428316 E – 4596888 N). The horizontal distance between the two points is 346
328 m. Strata dip gently toward the east, and altitude differences between the two points
329 are negligible (both are situated ~95 m above sea level). Both sections
330 (Supplementary Online Material [SOM] S1 and Fig. S1) were correlated on
331 lithostratigraphic grounds based on the assumption that there is no major fault located
332 between the two sections. In particular, considering a bedding orientation of 002/10E
333 (strike/dip), measured in outcropping layers in both sections, the bottom of
334 magnetostratigraphic Section 1 would be stratigraphically situated 16.6 m below the
335 bottom of Section 2. Given that Section 1 has a thickness of almost 50 m and Section
336 2 of about 20 m, these results imply that, unless there is a fault between the two
337 sections, they must overlap to a large extent (SOM Fig. S1). Such correlation
338 methodology is far from ideal, not only because it has to assume the lack of major
339 faulting, but also because it does not take into account lateral changes in lithology and
340 local accumulation rates. However, such an approach was unavoidable given the

¹ Geographic coordinates are given in the Universal Transverse Mercator – European Terrestrial Reference System 1989 (UTM ETRS89). They were verified using topographic maps and orthophotos from the web application VISSIR v3.26 of the Institut Cartogràfic i Geològic de Catalunya (ICGC, 2018).

341 dense vegetation cover between the two sections and the impossibility to deforest the
342 whole riverbank in between.

343

344 *2.2. Paleomagnetism and biostratigraphy*

345 A total of 25 samples from 14 magnetostratigraphic sites were collected along the
346 ~50 m-thick Section 1, although most samples are concentrated in the lower portion
347 (~20 m in thickness) of the section. Also, 22 samples from 11 sites were collected
348 along the 20 m-thick Section 2. No further samples were taken due to the presence of
349 unsuitable coarse-grained lithologies and the inaccessibility of other outcrops.

350 All samples were subjected to stepwise thermal demagnetization (Fig. 4) to isolate
351 the paleomagnetic components contributing to the natural remanent magnetization
352 (NRM). Temperature increments of 50 to 30 °C were applied up to complete
353 demagnetization of the samples or to temperatures at which acquisition of spurious
354 magnetization caused unstable behavior. Maximum unblocking temperatures close to
355 600 °C indicate that the magnetic remanence is carried by iron oxides. In earlier
356 studies of these sedimentary sequences (Garcés et al., 1996), detrital magnetite has
357 been reported as the main remanence carrier, together with pigmentary and detrital
358 hematite.

359 Stepwise demagnetization reveals the presence of a stable magnetic component at
360 temperatures above 250 °C, which demagnetizes at temperatures up to about 600 °C.
361 A characteristic magnetization direction was calculated by means of least-squares
362 analysis for each sample (Kirschvink, 1980). Complete demagnetization could rarely
363 be achieved, as thermal treatment often led to the acquisition of a spurious remanence
364 above 400 °C—likely related to chemical alteration of Fe-bearing minerals and
365 growth of new magnetite upon heating. Fortunately, the direction of the magnetic

366 remanence often remained stable between 250 and 400 °C, and a paleomagnetic
367 direction could be calculated from this temperature range. Anchoring the solutions to
368 the origin was needed in most of cases as a clean progressive decay of the remanence
369 could not often be observed (Fig. 4). A total of 47 paleomagnetic directions were
370 obtained (SOM Table S1) and were ranked into three quality categories. Q1 directions
371 (9 samples) were obtained from samples showing a gradual and nearly complete
372 decay towards the origin. Q2 directions (32 samples) were obtained from samples that
373 underwent unstable behavior at temperatures typically above 400 °C. Q3 directions (6
374 samples) were obtained from samples that underwent unstable at temperatures lower
375 than 400 °C and did not show a clear decay of its remanence upon heating, and were
376 not considered for polarity interpretation. A polarity interpretation of the
377 paleomagnetic directions was done by calculating the latitude of the virtual
378 geomagnetic pole (VGP) for each sample. Positive and negative VGP latitudes
379 correspond to normal and reversed polarity, respectively (Fig. 4).

380 To aid interpretation of the paleomagnetic data, we provide an updated faunal list
381 of the rodent assemblage from Castell de Barberà, which includes the material
382 recovered in 2014–2015. Given its biostratigraphic significance, we further review the
383 few hipparionin remains from the site, including a distal humeral fragment
384 (IPS87652) that was found in 2015. All the fossils are housed at the ICP.

385

386 **3. Results**

387 *3.1. Magnetostratigraphy*

388 Our paleomagnetic results (Fig. 5) indicate that at least four magnetozones are
389 recorded in Section 1 (beginning with a reversed polarity one), whereas two
390 magnetozones are recorded in Section 2 (beginning with a normal polarity

391 magnetozone, in which CB-D is situated). The stratigraphic distance of ca. 17 m
392 between the bases of the two sections would suggest that six magnetozones have been
393 sampled (Fig. 6), with the first normal polarity magnetozone of Section 2 (i.e., that of
394 the original main fossiliferous layer) being correlated with the second normal polarity
395 magnetozone of Section 1 (N2), and with the reversed polarity magnetozone of
396 Section 2 (R3) not having been sampled in Section 1. Given the close distance
397 between the two sections, local accumulation rate differences are negligible.
398 However, some uncertainties remain, because it is impossible to completely dismiss
399 the presence of one or more small faults between the two sections, thus allowing for
400 other (even if less likely) alternate correlations. Assuming that both sections
401 completely overlap, it might be conceivable to correlate the first normal polarity
402 magnetozone of Section 2 with the first normal polarity one of Section 1, which
403 results in only four magnetozones. Other possible correlations, involving only partial
404 overlap between the two sections, would imply up to five magnetozones, with the
405 normal polarity magnetozone of Section 2 being older than all of Section 1 or,
406 alternatively, with the normal polarity magnetozone of Section 2 being correlated with
407 the uppermost portion of Section 1.

408

409 *3.2. Rodent biostratigraphy*

410 Scanty reports of rodents in the initial faunal lists from Castell de Barberà
411 (Crusafont-Pairó and Golpe, 1972; Crusafont-Pairó and Casanovas Cladellas, 1973;
412 Crusafont-Pairó and Golpe-Posse, 1974; Golpe-Posse, 1974) were soon followed by a
413 detailed study (Aguilar et al., 1979) that recognized four cricetids and five glirids.
414 After multiple additions and refinements (Agustí, 1981; Agustí et al., 1985; Aldana
415 Carrasco, 1992a,b,c; Casanovas i Vilar, 2007), the most recently published list is that

416 of Casanovas-Vilar et al. (2016b), with up to 16 species. According to our own
417 taxonomic revision of the material, the previously collected Castell de Barberà rodent
418 assemblage, which is composed of 785 identifiable rodent remains, records up to 16
419 species, plus four additional ones recognized on the basis of the material recovered in
420 2015 (Table 1).

421

422 3.3. *Presence of Hippotherium*

423 It is uncertain what specimen Crusafont-Pairó and Golpe-Posse (1974) originally
424 identified as *Hipparion* (currently, *Hippotherium*) at Castell de Barberà. However,
425 Rotgers and Alba (2011) reported and figured a *Hippotherium* M₃ (IPS57437) from
426 the Castell de Barberà collections. The possibility of a misrecorded provenance is
427 unlikely given that the specimen was clearly labeled. In our recent re-examination of
428 the large mammal collections, we also found a dP² germ (IPS92389) attributable to
429 this taxon. Crusafont-Pairó and Golpe-Posse (1974) did not figure or describe the
430 specimen attributed by them to *Hippotherium*, but merely referred to it as a somewhat
431 rounded fragment—which might apply to the above-mentioned deciduous premolar,
432 although it is not possible to be certain. While the exact stratigraphic provenance of
433 these remains is not recorded, even if they both came from the upper layer they would
434 still be relevant for interpreting the magnetostratigraphic results, indicating a
435 Vallesian age for at least the topmost portion of the section.

436 Moreover, we recovered from CB-D during the 2015 field season a distal humeral
437 fragment (IPS87652) of an equid (Fig. 7A) that is virtually indistinguishable in both
438 size and shape from Vallesian *Hippotherium* specimens from the same basin (Fig. 7B,
439 C). The taxonomic attribution of IPS87652 to *Hippotherium* is complicated by the
440 fact that both anchitheriine and hipparionin equids co-occur in some Iberian early

441 Vallesian localities and elsewhere in Europe (Salesa et al., 2004; Daxner-Höck and
442 Bernor, 2009; Rotgers and Alba, 2011; Bernor et al., 2017). Although IPS87652 is
443 fragmentary, the preserved portion of the distal diaphysis and epiphysis enable
444 taxonomically meaningful morphological comparisons, which are provided below
445 given the biostratigraphic relevance of *Hippotherium*.

446 In terms of size, IPS87652 fits well with *Hippotherium (primigenium)* humeri from
447 elsewhere in Europe, as measured by the mediolateral width of the distal epiphysis
448 (70.0 mm), which is well within the range of the Höwenegg (MN9, Germany) sample
449 of *Hippotherium primigenium* ($X^- = 70.5$ mm, range = 65.3–74.1 mm, $n = 13$; Bernor
450 et al., 1997: Table 6.2). In contrast, the size of the specimen rules out an attribution to
451 any *Anchitherium* species from the Iberian Peninsula (and elsewhere in Europe),
452 which are smaller (Sánchez et al., 1998; Rotgers et al., 2011; Rotgers and Alba,
453 2011). On size grounds, IPS87652 would only be consistent with the much larger
454 anchitheriine *Sinohippus sampelayoi* (formerly in *Anchitherium*), which is recorded at
455 Nombrevilla 1 (MN9) in Spain (Salesa et al., 2004). Nonetheless, the latter species
456 has not been recorded in the Vallès-Penedès Basin, where available anchitheriine
457 remains from both the late Aragonian (Abocador de Can Mata) and early Vallesian
458 (Can Poncic 1) indicate a much smaller size (Crusafont-Pairó and Golpe-Posse, 1974;
459 Crusafont-Pairó, 1976; Sánchez et al., 1998; Rotgers et al., 2011; Rotgers and Alba,
460 2011). Furthermore, on qualitative morphological grounds IPS87652 also more
461 closely resembles *Hippotherium* than anchitheriines. The most relevant feature is the
462 larger trochlea relative to the smaller capitulum. In particular, IPS87652 resembles
463 *Hippotherium* in the mediolaterally wide and cylindrical trochlea and proximodistally
464 short capitulum, in contrast to the mediolaterally narrower and more globular trochlea
465 and longer capitulum of anchitheriines (MacFadden, 2001). Therefore, the distal

466 articular surface of IPS87652 is clearly *Hippotherium*-like and distinct from the more
467 primitive pattern displayed by anchitheriines (SOM Fig. S2). Moreover, medial to the
468 lateral epicondyle (between the lateral surface of the radial fossa and the lateral
469 epicondylar crest), IPS87652 possesses a narrow but deep groove that marks the
470 extensive origin attachment of the extensor digitorum communis muscle (Bernor et
471 al., 1997). This feature is consistently present in the *Hippotherium* sample from
472 Höwenegg (Bernor et al., 1997) but absent or weakly developed in anchitheriines
473 (SOM Fig. S2). Overall, both size and shape support a confident assignment of
474 IPS87652 to *Hippotherium*, in agreement with scarce dental material of this taxon
475 previously reported from Castell de Barberà (Rotgers and Alba, 2011).

476

477 **4. Discussion**

478 *4.1. Biochronology*

479 Among Miocene terrestrial assemblages, rodents are generally greatly significant
480 for accurate dating of fossil sites, even when magnetostratigraphic data are available.
481 This is because paleomagnetism only provides a binary signal, whose interpretation
482 critically relies on independent constraints. In the Vallès-Penedès Basin, a revised
483 scheme of local biozones and subzones largely based on rodents was recently
484 provided by Casanovas-Vilar et al. (2016b). The only (but most remarkable)
485 exception in this regard concerns the Aragonian/Vallesian boundary, which is defined
486 by the first appearance datum of the equid *Hippotherium*. This is because the
487 beginning of the Vallesian was not accompanied by important changes in the rodent
488 faunas (Agustí et al., 1997, 2001). Therefore, according to Casanovas-Vilar et al.
489 (2016b), the latest Aragonian corresponds to the *Democricetodon crusafonti* –
490 *Hippotherium* interval subzone (MN7+8, 11.88–11.18 Ma), whereas the earliest

491 Vallesian corresponds to the *Hippotherium* – *Cricetulodon hartenbergeri* interval
492 local subzone (MN9, 11.18–10.3 Ma)². For Castell de Barberà, the most significant
493 datum provided by rodents is the lack of *Cricetulodon*, which precludes correlation
494 with the *Cricetulodon hartenbergeri* range subzone (MN9, 10.3–9.98 Ma) and
495 indicates conclusively an age older than 10.3 Ma. Otherwise, however, the rodent
496 assemblage is consistent with both a late Aragonian or early Vallesian age. In
497 contrast, the presence of *Hippotherium* at Castell de Barberà (Fig. 7A) implies—by
498 definition—a post-Aragonian age.

499 Given the restricted evidence provided by the rodent faunas at Castell de Barberà,
500 Casanovas-Vilar et al. (2014, 2016a, b) favored a correlation of this site with the
501 earliest Vallesian mostly based on the presence of hipparionins, even if they did so
502 tentatively, given uncertainties about the provenance of scarce available material.
503 Although the record of *Hippotherium* at Castell de Barberà was already reported by
504 Crusafont-Pairó and Golpe-Posse (1974), most later authors dismissed its
505 biochronological significance by arguing that the specimen came from a layer above
506 the main fossiliferous layer. Our revision of the museum collections indicates that
507 *Hippotherium* is recorded at Castell de Barberà by two dental specimens. Even if both
508 teeth came from a stratigraphically higher layer, these remains would indicate, at least
509 (given the thickness of the outcrop section), that the main fossiliferous layer would be
510 situated less than 10 m below unambiguously Vallesian layers. Longstanding

² Interval zones are defined by two specified biohorizons (in these cases, corresponding to the lowest occurrences of the specified taxa), and termed with the name of the taxon defining the basal boundary preceding that of the taxon defining the top boundary (Murphy and Salvador, 1999).

511 discussion about the exact stratigraphic provenance of the material is now irrelevant
512 after the find of the distal humeral fragment of *Hippotherium* reported in this paper,
513 from a layer stratigraphically equivalent to the original main fossiliferous layer. This
514 find settles this issue by indicating unambiguously that Castell de Barberà is Vallesian
515 in age, even if a large portion of the rest of the fauna is still reminiscent of the late
516 Aragonian.

517 The first appearance datum of *Hippotherium* in the Vallès-Penedès Basin is
518 currently dated to 11.18 Ma, based on magnetostratigraphic correlation of the section
519 in which Creu de Conill 20 (CCN20) is situated (Garcés et al., 1997; Agustí et al.,
520 1997; Casanovas-Vilar et al., 2016a,b). The taxonomic identity of the earliest
521 hipparionins from the Vallès-Penedès Basin is uncertain, partly due to the lack of
522 diagnostic material and also because the alpha-taxonomy of *Hippotherium* is yet to be
523 settled conclusively—particularly with regard to the earliest representatives of the
524 genus (e.g., Bernor et al., 2017). Some authors have considered that the type species
525 of the genus, *H. primigenium*, is a single polymorphic species with many regional
526 variants (Pesquero and Arribas, 2002; Zouhri and Bensalmia, 2005). However, we
527 follow Bernor et al. (1996) in recognizing that *H. primigenium* is best considered a
528 species complex, which may be formally designated as *Hippotherium (primigenium)*
529 spp.³ The nominal species *Hippotherium catalaunicum*, erected by Pirlot (1956) based
530 on material from els Hostalets de Pierola, is available for the Vallesian *Hippotherium*
531 remains from the Vallès-Penedès Basin. This species has been considered a junior
532 subjective synonym of *Hippotherium primigenium* by several authors (Forstén, 1968,

³ The parentheses around the epithet denote a supraspecific species-group taxon, such as a species group or superspecies (see ICZN, 1999: Art. 6.2)

533 1978; Alberdi, 1972; Pesquero and Arribas, 2002; Zouhri and Bensalmia, 2005), but
534 here we follow Bernor and colleagues (Bernor et al., 1996; Bernor and Armour-
535 Chelu, 1999) in accepting a distinct species status for this taxon within the
536 *Hippotherium primigenium* complex—i.e., *Hippotherium (primigenium)*
537 *catalaunicum*. Nevertheless, we refrain from attributing the scanty Castell de Barberà
538 specimens to this species because they do not allow to assess its purportedly derived
539 distinguishing features from *Hippotherium primigenium* s.s. (namely, a very elongate
540 and anteroposteriorly oriented preorbital fossa; Bernor et al., 1996).

541 Pending a more detailed study, caution in taxonomic assignment of the
542 *Hippotherium* remains from Castell de Barberà is further advised by the
543 plesiomorphic dental morphology of the earliest *Hippotherium* samples from
544 elsewhere in Europe (see below). In the Vallès-Penedès Basin, the interpolated age of
545 Creu de Conill 20 fits well with that of the earliest Vallesian hipparionins from
546 Pannonian C (ca. 11.4–11.0 Ma) localities (Atzelsdorf and Gaiselberg) of the Vienna
547 Basin, Austria (Woodburne, 2007, 2009; Bernor et al., 2017), which are as yet
548 unassigned to species and appear more plesiomorphic than *H. primigenium* from the
549 type locality (Eppelsheim) and other Central European localities. These earliest
550 hipparionins from Europe most closely resemble *Cormohipparion* sp. from the
551 earliest late Miocene (ca. 11.5 Ma) of North America (Woodburne, 2005, 2007), and
552 suggest that *Hippotherium* is the descendant of a single dispersal event of
553 *Cormohipparion* from North America into Eurasia sometime after 11.5 Ma
554 (Woodburne, 2009; Bernor et al., 2017). According to Harzhauser (2009), Atzelsdorf
555 would be ca. 11.2–11.1 Ma in age—i.e., roughly coeval to Creu de Conill 20—
556 whereas elsewhere in the Iberian Peninsula *Hippotherium* dispersal is not documented

557 until somewhat later (within C5n.2n, ca. 10.8–10.7 Ma) in the Calatayud-Daroca
558 Basin (Garcés et al., 2003).

559 With regard to the biostratigraphic significance of the lack of *Hippotherium*, it
560 should be considered that, as illustrated by some Vallès-Penedès sites, this taxon may
561 be rare during the earliest Vallesian. This is best illustrated by the local section of
562 Ecoparc de Can Mata, where *Hippotherium* has yet to be recorded despite being
563 stratigraphically situated above the Aragonian/Vallesian transition—as indicated by
564 litho-, magneto-, and rodent biostratigraphic data (Alba et al., 2012; Casanovas-Vilar
565 et al., 2016b). Therefore, caution is required when inferring the age of other sites
566 chronologically close to the Aragonian/Vallesian boundary because the absence of
567 *Hippotherium* does not necessarily imply a pre-Vallesian age (e.g., Can Missert; see
568 discussion in Robles et al., 2011; contra Agustí et al., 2005)—and is alternatively
569 attributable to ecological conditions (e.g., unsuitable local habitat) and/or taphonomic
570 factors (i.e., sampling bias). Equating the lack of *Hippotherium* with a late Aragonian
571 age originally led to the misdating of Castell de Barberà, and this reasoning has been
572 further perpetuated despite clear (even if scarce) evidence that *Hippotherium* was
573 recorded at least somewhere within the short section where the site is located. When
574 both rodents and *Hippotherium* are considered together, biostratigraphic data enable
575 unambiguous correlation of Castell de Barberà with the *Hippotherium* – *Cricetulodon*
576 *hartenbergeri* interval local subzone (11.18–10.3 Ma; Fig. 6), as previously supported
577 by Casanovas-Vilar et al. (2016a, b) on tentative grounds.

578

579 4.2. Magnetostratigraphy

580 The magnetostratigraphic data reported in this paper enable more accurate dating
581 of Castell de Barberà within the early Vallesian, even if the shortness of the sampled

582 sections allows several possible interpretations. The most parsimonious interpretation,
583 based exclusively on the stratigraphic distance between the two sections as computed
584 from dip measurements, implies as much as six magnetozones. Other interpretations
585 imply four to five magnetozones. Irrespective of the preferred interpretation,
586 regarding correlation of the identified magnetozones with the GPTS (Ogg, 2012), two
587 important factors should be considered. First, the main fossiliferous layer of Castell de
588 Barberà is situated in a normal polarity magnetozone. Second, no matter what
589 correlation between the two sections is preferred, there are at least four magnetozones
590 within a short 50–80 m sequence. Taking into account an average sedimentation rate
591 of ~20 cm/kyr, computed for both the Vallesian (Garcés et al., 1996) and the late
592 Aragonian (Alba et al., 2017) of the Vallès-Penedès Basin, the sampled interval
593 would not represent more than 400 kyr. The frequent alternation of normal and
594 reversed polarity magnetozones in the composite sequence of Castell de Barberà
595 indicates that the long normal subchron C5n.2n that is characteristic of the early
596 Vallesian is not recorded in the sampled interval. Only the uppermost normal polarity
597 magnetozone N3 could correlate with the base of C5n.2n. The magnetostratigraphic
598 results rather suggest that most of the composite sequence correlates with the upper
599 C5r, where multiple reversals are recorded over a short time interval (Ogg, 2012).

600 An alternate correlation of the reversed polarity magnetozones R2 and R3 with
601 some cryptochrons or geomagnetic excursions within C5n.2n (Cande and Kent, 1992;
602 Roberts and Lewin-Harris, 2000; Evans et al., 2007) is considered unlikely because of
603 the short duration, possibly less than 10 kyr, of these events. While chances of
604 recording such short geomagnetic features in fluvial sediments are typically low, to
605 record as many reversed polarity directions as normal polarity directions within
606 C5n.2n is considered implausible. The fact that all the reversed polarity magnetozones

607 documented at Castell de Barberà are recorded by more than a single paleomagnetic
608 sample strongly argues against any of them representing a short-lived excursion event
609 within C5n.2n. The similar thickness of the documented normal and reversed polarity
610 magnetozone at Castell de Barberà strongly argues against this possibility, and favor
611 a correlation of most of the sequence with chron C5r.

612 In our preferred interpretation (Fig. 6), the lowermost normal polarity
613 magnetozone of Section 1 would be correlated with subchron C5r.2r-1 (11.308–
614 11.263 Ma), and the lowermost normal polarity magnetozone of Section 2 (i.e., the
615 main fossiliferous layer), like the second normal polarity magnetozone of Section 1,
616 with C5r.1n (11.188–11.146 Ma). According to this interpretation, the original
617 outcrop of Castell de Barberà, including the main fossiliferous layer at about the
618 middle of the sequence and its uppermost layers, would cover a time span
619 approximately between 11.2 and 11.1 Ma. Such a correlation would be compatible
620 with the composite sequence recording six magnetozone (with the uppermost portion
621 of the composite sequence, recorded in Section 1, correlated to the lowermost part of
622 subchron C5n.2n). This correlation would be also compatible with an alternate
623 interpretation of the composite sequence recording only four magnetozone, with the
624 lower normal polarity magnetozone of each section correlated to the same subchron,
625 either C5r.2r-1 or C5r.1n. However, correlation with the former subchron is much
626 more unlikely because it would imply a first appearance datum of *Hippotherium* about
627 100 kyr older than previously documented in the Vallès-Penedès Basin and elsewhere
628 in Eurasia (Garcés et al., 1997; see also above). An older age for the bottom of
629 Section 2 (normal polarity) relative to that of Section 1 (reversed polarity), implying
630 five magnetozone, would also imply a correlation of the first normal polarity
631 magnetozone of Section 2 with C5r.2r-1, which is similarly unlikely.

632 In summary, although there are several possible interpretations about the number
633 and correlation of magnetostratigraphic and biostratigraphic data favor correlation of the main
634 magnetostratigraphic and biostratigraphic data favor correlation of the main
635 fossiliferous layer of Castell de Barberà with subchron C5r.1n. An alternative
636 correlation with the long normal polarity chron C5n.2n, which is characteristic of the
637 early Vallesian, is precluded by the presence of a reversed polarity magnetozone at
638 the top of the main fossiliferous layer—unless such a reversed polarity magnetozone
639 is interpreted to represent a cryptochron, which seems highly unlikely as discussed
640 above. In turn, correlation with the older and short normal polarity subchron C5r.2r-1 is
641 possible, but is not favored here because it would imply a first appearance datum of
642 *Hippotherium* of ca. 100 kyr older than previously documented. According to our
643 interpretation of the results, the main fossiliferous layer of Castell de Barberà would
644 be roughly equivalent in age to Creu de Conill 20 (11.2 Ma), and both sites would
645 represent the earliest Vallesian faunas of the Vallès-Penedès Basin, with the
646 uppermost layers of the original Castell de Barberà outcrop, which also delivered
647 some fossil remains, being correlated with C5r.1r and with an estimated age closer to
648 11.1 Ma. Only layers underlying the main fossiliferous layer might be latest
649 Aragonian instead of earliest Vallesian, although they would be correlated with the
650 same subchron and would have a similar estimated age to the nearest 0.1 Ma.

651

652 *4.3. Implications for primate evolution*

653 As one of the few European Miocene localities where pliopithecoids and
654 hominoids are recorded, Castell de Barberà is a key site for better understanding why
655 their co-occurrence is so rare. However, lack of accurate dating for Castell de Barberà
656 represented a serious drawback for adequately contextualizing this phenomenon

657 within the framework of faunal and paleoenvironmental changes in the Vallès-
658 Penedès Basin. Sukselainen et al. (2015) concluded that pliopithecoids inhabited
659 generally more ‘humid’ environments than hominoids, and further suggested that the
660 few localities hosting a species from each group might have had even more ‘humid’
661 conditions than other primate-bearing European Miocene localities. This study relied
662 on the hypsodonty of larger herbivorous mammals as a proxy for vegetation structure,
663 which in turn has been used as a proxy for past humidity and precipitation
664 (Sukselainen et al., 2015, and references therein), so that their conclusions are more
665 adequately interpreted as indicating more closely forested conditions for primate-
666 bearing localities (particularly those in which the two groups co-occur) as a result of
667 higher humidity (moisture), precipitation (rainfall), or both. This interpretation is
668 consistent with previous paleoecological analyses by Casanovas-Vilar et al. (2008),
669 which found that Castell de Barberà would have been more ‘humid’ (i.e., higher
670 moisture and/or precipitation) and displayed a more marked evergreen vegetation
671 component than late Aragonian hominoid-bearing localities from the Vallès-Penedès
672 Basin.

673 Particularly ‘humid’ conditions at Castell de Barberà are consistent with the high
674 abundance of swimming castorids. They are represented by two species: the small
675 *Euroxenomys minutus*—more than half of the rodent remains recovered using screen-
676 washing techniques from Castell de Barberà belong to this taxon—and the much rarer
677 *Chalicomys jaegeri*, which was the size of the extant beaver (Casanovas-Vilar and
678 Agustí, 2007; Casanovas-Vilar et al., 2008, 2010). These taxa indicate the presence of
679 permanent water bodies nearby, although not necessarily of humid and densely
680 forested habitats. Nevertheless, in the framework of Vallès-Penedès regional setting it
681 is consistent with the fact that Vallesian hominoids from this basin have been linked

682 to forested humid habitats providing a year-round fruit supply (Marmi et al., 2012;
683 Alba et al., 2018c)—such as that from Can Llobateres 1, which on the basis of plant
684 remains has been reconstructed as a very humid marshy area with nearby dense
685 wetland forests including some (sub)tropical elements (Marmi et al., 2012). The
686 record of the anuran *Latonia* is also indicative of locally humid and warm conditions
687 (Villa et al., 2017), and besides the presence of primates and beavers, a closed
688 forested environment at Castell de Barberà is further supported by the presence of
689 certain forest-dwelling taxa, such as arboreal dormice (*Bransatoglis*, *Glirudinus*,
690 *Muscardinus*, *Myoglis*, *Paraglrirulus*) and flying squirrels (*Miopetaurista* and
691 probably *Albanensia*), which are more diverse and abundant at Castell de Barberà
692 than in roughly contemporary sites (Casanovas-Vilar and Agustí, 2007; Casanovas-
693 Vilar et al., 2008, 2010). Finally, certain cricetids (*Eumyarion*, *Anomalomys*) may
694 have also preferred densely forested habitats because they tend to be more abundant
695 in fossil faunas rich in glirids and flying squirrels (Casanovas-Vilar and Agustí,
696 2007).

697 The apparently more humid and densely forested paleoenvironmental conditions of
698 Castell de Barberà as compared to late Aragonian localities from the same basin could
699 be explained to some extent by differences in age. However, even if the earliest
700 Vallesian is not well represented in the Vallès-Penedès, Castell de Barberà should not
701 be taken as representative of the dominant paleoenvironments in the basin as a whole
702 during this time span—as suggested by differences in faunal composition from the
703 roughly coeval Creu de Conill 20 site, where *Hippotherium* is more abundant
704 (authors' unpubl. data), the micromammal assemblage indicates drier conditions
705 (Casanovas-Vilar et al., 2006, 2008; Casanovas-Vilar and Agustí, 2007), and primates
706 have not been recorded yet (authors' unpubl. data). The rodent assemblage from Creu

707 de Conill 20 is far less diverse than that of Castell de Barberà (Casanovas-Vilar et al.,
708 2006, 2016a, b; authors' unpublished data), and includes few forest-dwelling taxa
709 (*Muscardinus*, cf. *Paraglitirulus*), with beavers only represented by scarce remains of
710 *Euroxenomys minutus*. The rodent fauna of Creu Conill 20 is overwhelmingly
711 dominated by the cricetid *Megacricetodon ibericus*, which is very rare at Castell de
712 Barberà and is considered to have been a generalist probably preferring more arid
713 woodlands (Daams and Freudenthal, 1988; Casanovas-Vilar et al., 2006, 2008, 2010,
714 Casanovas-Vilar and Agustí, 2007). At Castell de Barberà, *Eumyarion leemanni* is the
715 most abundant cricetid, which is consistent with the occurrence of locally humid and
716 forested environments.

717 The dating provided in this paper for Castell de Barberà indicates that renewed
718 efforts are required to better characterize earliest Vallesian faunas from the Vallès-
719 Penedès Basin, not only to investigate the particular conditions that enabled co-
720 existence of hominoids and pliopithecoids, but especially to clarify their taxonomic
721 status and/or phylogenetic relationships. For pliopithecoids, *Barberapithecus*
722 *huerzeleri* is exclusively known from Castell de Barberà (Alba and Moyà-Solà, 2012).
723 Our results provide an accurate date for this taxon, but additional and more complete
724 material is required to more definitively settle its crouzeliid status and its relationships
725 with other pliopithecoid taxa. Regarding hominoids, attribution of the postcranial
726 remains from Castell de Barberà to *Dryopithecus fontani* by Alba et al. (2011) was
727 only tentative. Given that this taxon is otherwise known only from the late Aragonian
728 of the Vallès-Penedès Basin, France and Austria (Begun, 2002b; Moyà-Solà et al.,
729 2009; Pérez de los Ríos et al., 2013), if such assignment was confirmed then Castell
730 de Barberà might represent the last appearance datum of this genus. However,
731 additional (especially craniodental) remains would be necessary to confirm the

732 taxonomic status of the hominoid taxon recorded at Castell de Barberà, which is
733 intermediate in time between the late Aragonian dryopithecins recorded at Abocador
734 de Can Mata and the more derived hispanopithecins recorded at the basin later during
735 the Vallesian (Alba, 2012; Alba et al., 2017, 2018c).

736 Two areas of the Vallès-Penedès Basin have the potential to yield additional
737 primate remains roughly contemporaneous with those of Castell de Barberà: the
738 earliest Vallesian site of Creu de Conill 20 in Terrassa; and the Vallesian levels of els
739 Hostalets de Pierola (located westward from Abocador de Can Mata; Moyà-Solà et
740 al., 2009; Alba et al., 2017). The currently available sample of ca. 2000
741 macrovertebrate remains from Creu de Conill 20 is insufficient to dismiss the ulterior
742 find of primates after additional sampling efforts at this locality (given the rarity of
743 these taxa among Vallès-Penedès fossil assemblages), although the
744 paleoenvironmental hints provided by the recovered fauna are not particularly
745 promising in this regard (see above). This contrasts with the situation in the
746 fossiliferous area of els Hostalets de Pierola, where the presence of an indeterminate
747 dryopithecine is documented close to the Aragonian/Vallesian boundary by a female
748 lower canine from the latest Aragonian fossil locality of Can Mata 1 (Crusafont-Pairó
749 and Golpe-Posse, 1973; Golpe Posse, 1982, 1993; Alba, 2012) and by an isolated
750 upper molar of uncertain provenance from the same area (van der Made and Ribot,
751 1999; Alba, 2012; Alba et al., 2013). Future fieldwork in these and other sites should
752 help to determine the taxonomic identity of earliest Vallesian hominoids from the
753 Vallès-Penedès Basin and to clarify whether their rare co-occurrence with
754 pliopithecoids at Castell de Barberà and other localities was related to particularly
755 humid and/or densely forested environmental conditions as compared to most other
756 sites from the same basin.

757

758 **5. Summary and conclusions**

759 Here we report new magnetostratigraphic results for the primate-bearing Castell de
760 Barberà site, together with a *Hippotherium* find collected in 2015 from a stratigraphic
761 horizon (CB-D) equivalent to the original main fossiliferous layer at Castell de
762 Barberà. We interpret these data in light of previously published literature and our
763 own review of the rodent assemblage from Castell de Barberà, with the aim of
764 providing an unambiguous and more accurate date for the site.

765 Our paleomagnetic results, coupled with in situ recovery of a *Hippotherium*
766 humerus, make it unlikely the correlation of any of the various sampled normal
767 polarity magnetozones with the long normal polarity subchron C5n.2n (11.056–9.984
768 Ma) that is characteristic of the early Vallesian, and support instead correlation of CB-
769 D with C5r.1n (11.188–11.146 Ma), where the Aragonian/Vallesian boundary is
770 situated. Overall, our results unambiguously indicate an early Vallesian age for
771 Castell de Barberà, thereby settling longstanding debate about the Aragonian vs.
772 Vallesian age of this site. Our results further support an earliest Vallesian (~11.2 Ma)
773 age for Castell de Barberà, which would be roughly coeval with the Creu de Conill 20
774 site, where the first appearance datum of hipparionins in the Vallès-Penedès Basin has
775 been recorded.

776 Accurate dating of the Castell de Barberà site provided in this paper is important to
777 contextualize the faunal and paleoenvironmental changes that enabled coexistence of
778 pliopithecoids and hominoids at particular sites in the Vallès-Penedès Basin during
779 the late Miocene. It also provides useful hints to redirect future fieldwork efforts in
780 the Vallès-Penedès Basin, with the aim to clarify the taxonomic identity and/or

781 phylogenetic relationships of catarrhine primates that lived there during the earliest
782 Vallesian.

783

784 **References**

785 Aguilar, J.-P., Agustí, J., Gibert, J., 1979. Rongeurs miocènes dans le Valles-Penedes.

786 2 - Les Rongeurs de Castell de Barbera. *Palaeovertebrata* 9, 17–32.

787 Agustí, J., 1981. Roedores miomorfos del Neógeno de Cataluña. PhD. dissertation,

788 Universitat de Barcelona.

789 Agustí, J., 1982. Biozonación del neogeno continental de Cataluña mediante roedores

790 (Mammalia). *Acta Geológica Hispánica* 17, 21–26.

791 Agustí, J., 2018. Las faunas de mamíferos del Mioceno continental de la Península

792 Ibérica. *Revista PH* 94, 182–205.

793 Agustí, J., Moyà-Solà, S., 1991. Spanish Neogene Mammal succession and its bearing

794 on continental biochronology. *Newsletters on Stratigraphy* 25, 91–114.

795 Agustí, J., Cabrera, L., Moyà-Solà, S., 1985. Sinopsis estratigráfica del Neógeno de la

796 fosa del Vallès-Penedés. *Paleontologia i Evolució* 18, 57–81.

797 Agustí, J., Cabrera, L., Garcés, M., Parés, J.M., 1997. The Vallesian mammal

798 succession in the Vallès-Penedès basin (northeast Spain): Paleomagnetic

799 calibration and correlation with global events. *Palaeogeography,*

800 *Palaeoclimatology, Palaeoecology* 133, 149–180.

801 Agustí, J., Cabrera, L., Garcés, M., Krijgsman, W., Oms, O., Parés, J.M., 2001. A

802 calibrated mammal scale for the Neogene of Western Europe. State of the art.

803 *Earth-Science Reviews* 52, 247–260.

804 Agustí, J., Casanovas-Vilar, I., Furió, M., 2005. Rodents, insectivores and

805 chiropterans (Mammalia) from the late Aragonian of Can Missert (Middle

806 Miocene, Vallès-Penedès basin, Spain). *Geobios* 38, 575–583.

807 Agustí, J., Cabrera, L., Garcés, M., 2013. The Vallesian Mammal Turnover: A late
808 Miocene record of decoupled land-ocean evolution. *Geobios* 46, 151–157.

809 Alba, D.M., 2012. Fossil apes from the Vallès-Penedès Basin. *Evolutionary*
810 *Anthropology* 21, 254–269.

811 Alba, D.M., Berning, B., 2013. On the holotype and original description of the
812 pliopithecoid *Plesiopliopithecus lockeri* (Zapfe, 1960). *Journal of Human Evolution*
813 65, 338–340.

814 Alba, D.M., Moyà-Solà, S., 2012. A new pliopithecoid genus (Primates:
815 Pliopithecoidea) from Castell de Barberà (Vallès-Penedès Basin, Catalonia, Spain).
816 *American Journal of Physical Anthropology* 147, 88–112.

817 Alba, D.M., Moyà-Solà, S., Casanovas-Vilar, I., Galindo, J., Robles, J.M., Rotgers,
818 C., Furió, M., Angelone, C., Köhler, M., Garcés, M., Cabrera, L., Almécija, S.,
819 Obradó, P., 2006. Los vertebrados fósiles del Abocador de Can Mata (els Hostalets
820 de Pierola, l’Anoia, Catalunya), una sucesión de localidades del Aragoniense
821 superior (MN6 y MN7+8) de la cuenca del Vallès-Penedès. *Campañas 2002-2003,*
822 *2004 y 2005. Estudios Geológicos* 62, 295–312.

823 Alba, D.M., Moyà-Solà, S., Malgosa, A., Casanovas-Vilar, I., Robles, J.M., Almécija,
824 S., Galindo, J., Rotgers, C., Bertó Menguà, J.V., 2010. A new species of
825 *Pliopithecus* Gervais, 1849 (Primates: Pliopithecidae) from the Middle Miocene
826 (MN8) of Abocador de Can Mata (els Hostalets de Pierola, Catalonia, Spain).
827 *American Journal of Physical Anthropology* 141, 52–75.

828 Alba, D.M., Moyà-Solà, S., Almécija, S., 2011. A partial hominoid humerus from the
829 middle Miocene of Castell de Barberà (Vallès-Penedès Basin, Catalonia, Spain).
830 *American Journal of Physical Anthropology* 144, 365–381.

831 Alba, D.M., Carmona, R., Bertó Mengual, J.V., Casanovas-Vilar, I., Furió, M.,
832 Garcés, M., Galindo, J., Luján, À.H., 2012. Intervenció paleontològica a l'Ecoparc
833 de Can Mata (els Hostalets de Pierola, conca del Vallès-Penedès). Tribuna
834 d'Arqueologia 2010–2011, 115–130.

835 Alba, D.M., Fortuny, J., Pérez de los Ríos, M., Zanolli, C., Almécija, S., Casanovas-
836 Vilar, I., Robles, J.M., Moyà-Solà, S., 2013. New dental remains of *Anoiapithecus*
837 and the first appearance datum of hominoids in the Iberian Peninsula. *Journal of*
838 *Human Evolution* 65, 573–584.

839 Alba, D.M., Casanovas-Vilar, I., Garcés, M., Robles, J.M., 2017. Ten years in the
840 dump: An updated review of the Miocene primate-bearing localities from
841 Abocador de Can Mata (NE Iberian Peninsula). *Journal of Human Evolution* 102,
842 12–20.

843 Alba, D.M., Moyà-Solà, S., DeMiguel, D., Casanovas-Vilar, I., Garcés, M., Robles, J.
844 M., Madurell-Malapeira, J., Almécija, S., 2018. Ape quest in the Vallès-Penedès
845 Basin (2014–2017): Fieldwork results and prospects for the future. *American*
846 *Journal of Physical Anthropology* 165 S66, 7.

847 Alba, D.M., Garcés, M., Pina, M., Casanovas-Vilar, I., Robles, J.M., Moyà-Solà, S.,
848 Almécija, S., 2018. Bio- and magnetostratigraphic correlation of the Miocene
849 primate-bearing site of Castell de Barberà: end of the controversy. *Proceedings of*
850 *the European Society for the study of Human Evolution* 7, 2.

851 Alba, D.M., Casanovas-Vilar, I., Furió, M., García-Paredes, I., Angelone, C., Jovells-
852 Vaqué, S., Luján, À.H., Almécija, S., Moyà-Solà, S., 2018c. Can Pallars i
853 Llobateres: A new hominoid-bearing locality from the late Miocene of the Vallès-
854 Penedès Basin (NE Iberian Peninsula). *Journal of Human Evolution* 121, 193–203.

- 855 Alberdi, M.T., 1972. El género *Hipparion* en España, revisión e historia evolutiva.
856 Coloquios de la Cátedra de Paleontología 21, 7–8.
- 857 Aldana Carrasco, E.J., 1992a. Los Castoridae (Rodentia, Mammalia) del Neógeno de
858 Cataluña (España). Treballs del Museu de Geologia de Barcelona 2, 99–141.
- 859 Aldana Carrasco, E.J., 1992b. Los Esciurópteros del Mioceno de la cuenca del Vallès-
860 Penedès (Cataluña, España). Geogaceta 11, 114–116.
- 861 Aldana Carrasco, E.J., 1992c. Los Sciurinae (Rodentia, Mammalia) del Mioceno de la
862 Cuenca del Vallès-Penedès (Cataluña, España). Treballs del Museu de Geologia de
863 Barcelona 2, 69–97.
- 864 Almécija, S., Alba, D.M., Moyà-Solà, S., 2011. Large-hominoid remains from the
865 Middle Miocene locality of Castell de Barberà (Vallès-Penedès Basin, Catalonia,
866 Spain). American Journal of Physical Anthropology 144 S52, 74.
- 867 Almécija, S., Alba, D.M., Moyà-Solà, S., 2012. The thumb of Miocene apes: new
868 insights from Castell de Barberà (Catalonia, Spain). American Journal of Physical
869 Anthropology 148, 436–450.
- 870 Álvarez Sierra, M.A., Calvo, J.P., Morales, J., Alonso-Zarza, A., Azanza, B., García
871 Paredes, I., Hernández Fernández, M., van der Meulen, A.J., Peláez-Campomanes,
872 P., Quiralte, V., Salesa, M.J., Sánchez, I.M., Soria, D., 2003. El tránsito
873 Aragoniense-Vallesiense en el área de Daroca-Nombrevilla (Zaragoza, España).
874 Coloquios de Paleontología Vol. Ext. 1, 25–33.
- 875 Begun, D.R., 2002a. The Pliopithecoidea. In: Hartwig, W.C. (Ed.), The Primate Fossil
876 Record. Cambridge University Press, Cambridge, pp. 221–240.
- 877 Begun, D.R., 2002b. European hominoids. In: Hartwig, W.C. (Ed.), The Primate
878 Fossil Record. Cambridge University Press, Cambridge, pp. 339–368.

879 Bernor, R.L., Armour-Chelu, M., 1999. Family Equidae. In: Rössner, G.E., Heissig,
880 K. (Eds.), *The Miocene Land Mammals of Europe*. Verlag Dr. Friedrich Pfeil,
881 München, pp. 193–202.

882 Bernor, R.L., Tobien, H., Hayek, L.-A.C., Mittmann, H.-W., 1997. *Hippotherium*
883 *primigenium* (Equidae, Mammalia) from the late Miocene of Höwenegg (Hegau,
884 Germany). *Andrias* 10, 5–230.

885 Bernor, R.L., Göhlich, U., Harzhauser, M., Semprebon, G.M., 2017. The Pannonian C
886 hipparions from the Vienna Basin. *Palaeogeography, Palaeoclimatology,*
887 *Palaeoecology* 476, 28–41.

888 Cande, S.C., Kent, D.V., 1992. A new geomagnetic polarity time scale for the Late
889 Cretaceous and Cenozoic. *Journal of Geophysical Research* 97, 13917–13951.

890 Casanovas i Vilar, I., 2007. The rodent assemblages from the Late Aragonian and the
891 Vallesian (Middle to Late Miocene) of the Vallès-Penedès Basin (Catalonia,
892 Spain). Ph.D. Dissertation, Universitat Autònoma de Barcelona.

893 Casanovas-Vilar, I., Agustí, J., 2007. Ecogeographical stability and climate forcing in
894 the Late Miocene (Vallesian) rodent record of Spain. *Palaeogeography,*
895 *Palaeoclimatology, Palaeoecology* 248, 169–189.

896 Casanovas-Vilar, I., Furió, M., Agustí, J., 2006. Rodents, insectivores and
897 paleoenvironment associated to the first-appearing hipparionine horses in the
898 Vallès-Penedès Basin (Northeastern Spain). *Beiträge zur Paläontologie* 30, 89–
899 107.

900 Casanovas-Vilar, I., Alba, D.M., Moyà-Solà, S., Galindo, J., Cabrera, L., Garcés, M.,
901 Furió, M., Robles, J.M., Köhler, M., Angelone, C., 2008. Biochronological,
902 taphonomical and paleoenvironmental background of the fossil great ape
903 *Pierolapithecus catalaunicus* (Primates, Hominidae). *Journal of Human Evolution*

904 55, 589–603.

905 Casanovas-Vilar, I., Angelone, C., Alba, D.M., Moyà-Solà, S., Köhler, M., Galindo,
906 J., 2010. Rodents and lagomorphs from the Middle Miocene hominoid-bearing site
907 of Barranc de Can Vila 1 (els Hostalets de Pierola, Catalonia, Spain). *Neues*
908 *Jahrbuch für Geologie und Paläontologie Abhandlungen* 257, 293–315.

909 Casanovas-Vilar, I., Alba, D.M., Garcés, M., Robles, J.M., Moyà-Solà, S., 2011a.
910 Updated chronology for the Miocene hominoid radiation in Western Eurasia.
911 *Proceedings of the National Academy of Sciences USA* 108, 5554–5559.

912 Casanovas-Vilar, I., Alba, D.M., Robles, J.M., Moyà-Solà, S., 2011b. Registro
913 paleontológico continental del Mioceno de la cuenca del Vallès-Penedès.
914 *Paleontologia i Evolució memòria especial* 6, 55–80.

915 Casanovas-Vilar, I., Van den Hoek Ostende, L.W., Furió, M., Madern, A., 2014. The
916 range and extent of the Vallesian Crisis (Late Miocene): new prospects based on
917 the micromammal record from the Vallès-Penedès Basin (Catalonia, Spain).
918 *Journal of Iberian Geology* 40, 29–48.

919 Casanovas-Vilar, I., Madern, A., Alba, D.M., Cabrera, L., García-Paredes, I., Van den
920 Hoek Ostende, L.W., DeMiguel, D., Robles, J.M., Furió, M., Van Dam, J., Garcés,
921 M., Angelone, C., Moyà-Solà, S., 2016a. The Miocene mammal record of the
922 Vallès-Penedès Basin (Catalonia). *Comptes Rendus Palevol* 15, 791–812.

923 Casanovas-Vilar, I., Garcés, M., Van Dam, J., García-Paredes, I., Robles, J.M., Alba,
924 D.M., 2016b. An updated biostratigraphy for the late Aragonian and the Vallesian
925 of the Vallès-Penedès Basin (Catalonia). *Geologica Acta* 14, 195–217.

926 Crusafont Pairó, M., 1950. La cuestión del llamado Meóico español. *Arrahona* 1950,
927 41–48.

928 Crusafont Pairó, M., 1951. El sistema miocénico en la depresión española del Vallés-

- 929 Penedés. In: International Geological Congress "Report of the Eighteenth Session,
930 Great Britain, 1948", Part XI, pp. 33–42.
- 931 Crusafont Pairó, M., 1953. El sistema miocénico en la depresión española del Vallés-
932 Penedés. Memorias y Comunicaciones del Instituto Geológico Provincial 10, 13–
933 23.
- 934 Crusafont Pairó, M., 1955. Données biogéographiques relevées par la
935 paléomammalogie du Miocène espagnol. In: Colloque International sur "Problèmes
936 Actuels de Paléontologie" C. N. R. C. Paris, avril 1955, pp. 101–108.
- 937 Crusafont-Pairó, M., 1972. Les *Ischyriactis* de la transition Vindobonien-Vallésien.
938 *Palaeovertebrata* 5, 253–260.
- 939 Crusafont-Pairó, M., 1975. El gibón fósil (*Pliopithecus*) del Vindoboniense terminal
940 del Vallès. Boletín Informativo del Instituto de Paleontología de Sabadell 7, 36–38.
- 941 Crusafont-Pairó, M., 1976. Corrigenda (A dos notas sobre la presencia del género
942 *Anchitherium* en el Mioceno del Vallès. Boletín Informativo del Instituto de
943 Paleontología de Sabadell 8, 27–28.
- 944 Crusafont-Pairó, M., Casanovas Cladellas, L., 1973. Fossilium Catalogus. I:
945 Animalia. Pars 121. Mammalia Tertiaria Hispaniae. Dr. W. Junk B.V., 's-
946 Gravenhage.
- 947 Crusafont-Pairó, M., Golpe, J.M., 1972. Dos nuevos yacimientos del vindoboniense
948 en el Vallès. *Acta Geológica Hispánica* 7, 71–72.
- 949 Crusafont Pairó, M., Golpe Posse, J.M., 1972a. Los yacimientos de mamíferos fósiles
950 del Vallès. Boletín Informativo del Instituto Provincial de Paleontología Sabadell
951 4, 20–24.

- 952 Crusafont-Pairó, M., Golpe Posse, J.M., 1972b. Hallazgo del género *Anchitherium*
953 Meyer, 1844 en el Vindoboniense terminal del Vallés-Penedés. Boletín de la Real
954 Sociedad Española de Historia Natural 69, 297–298.
- 955 Crusafont-Pairó, M., Golpe-Posse, J.M., 1973. New pongids from the Miocene of
956 Vallès Penedès Basin (Catalonia, Spain). *Journal of Human Evolution* 2, 17–23.
- 957 Crusafont-Pairó, M., Golpe-Posse, J.M., 1974. Asociación de *Anchitherium* Mey.,
958 1834, con *Hipparion* Christ, 1832, en el Alto Mioceno del Vallés. Boletín de la
959 Real Sociedad Española de Historia Natural 72, 75–93.
- 960 Crusafont-Pairó, M., Golpe-Posse, J.M., 1981a. Estudio de la dentición inferior del
961 primer pliopitécido hallado en España (Vindoboniense terminal de Castell de
962 Barberà, Cataluña, España). *Butlletí Informatiu de l'Institut de Paleontologia de*
963 *Sabadell* 13, 25–38.
- 964 Crusafont-Pairó, M., Golpe-Posse, J.M., 1981b. Hallazgo de una nueva especie del
965 género *Semigenetta* (*S. grandis*) del Vindoboniense terminal de Castell de Barberà
966 (Depresión prelitoral catalana; España). Boletín de la Real Sociedad Española de
967 Historia Natural 79, 67–76.
- 968 Crusafont Pairó, M., Golpe Posse, J.M., 1982a. Los Pliopitécidos en España.
969 *Coloquios de Paleontología* 37, 41–46.
- 970 Crusafont Pairó, M., Golpe Posse, J.M., 1982b. Hallazgo de *Palaeomeles Pachecoi*
971 Vill. et Crus., 1943, en Castell de Barberà (Vallés-Penedés). *Acta Geológica*
972 *Hispánica* 17, 27–37.
- 973 Crusafont Pairó, M., Hürzeler, J., 1969. Catálogo comentado de los póngidos fósiles
974 de España. *Acta Geológica Hispánica* 4, 44–48.
- 975 Crusafont Pairó, M., Truyols Santonja, J., 1951. Hallazgo del *plesiodimylus chantrei*
976 Gaillard en el Meótico del Vallés. *Notas y Comunicaciones del Instituto Geológico*

- 977 y Minero de España 22, 97–126.
- 978 Crusafont, M., Truyols, J., 1954. Catálogo Paleomastológico del Mioceno del Vallés-
979 Penedés y de Calatayud-Teruel. Segundo Cursillo Internacional de Paleontología.
980 Museo de la Ciudad de Sabadell, Sabadell.
- 981 Crusafont Pairó, M., Truyols Santonja, J., 1954. Sinopsis estratigráfico-paleontológica
982 del Vallés-Penedés. Arrahona 1954, 1–15.
- 983 Crusafont Pairó, M., Truyols Santonja, J., 1959. Sobre el nuevo proyecto de
984 estructuración y nomenclatura del Mioceno mediterráneo. Notas y Comunicaciones
985 del Instituto Geológico y Minero de España 56, 33–53.
- 986 Crusafont Pairó, M., Truyols Santonja, J., 1960. Sobre la caracterización del
987 Vallesiense. Notas y Comunicaciones del Instituto Geológico y Minero de España
988 60, 109–126.
- 989 Daams, R., Freudenthal, M., 1981. Aragonian: the stage concept versus Neogene
990 Mammal zones. Scripta Geologica 62, 1–17.
- 991 Daams, R., Freudenthal, M., 1988. Cricetidae (Rodentia) from the type-Aragonian;
992 the genus *Megacricetodon*. Scripta Geologica Spec. Issue 1, 39–132.
- 993 Daams, R., Freudenthal, M., 1990. The Ramblian and the Aragonian: Limits,
994 subdivisions. In: Lindsay, E.H., Fahlbush, V., Mein, P. (Eds.), European Neogene
995 Mammal Chronology. Plenum Press, New York, pp. 51–59.
- 996 Daams, R., Freudenthal, M., van de Weerd, A., 1977. Aragonian, a new stage for
997 continental deposits of Miocene age. Newsletters on Stratigraphy 6, 42–55.
- 998 Daams, R., Van der Meulen, A.J., Álvarez Sierra, M.A., Peláez-Campomanes, P.,
999 Calvo, J. P., Alonso Zarza, M.A., Krijgsman, W., 1999. Stratigraphy and
1000 sedimentology of the Aragonian (Early to Middle Miocene) in its type area (North-
1001 Central Spain). Newsletters on Stratigraphy 37, 103–139.

1002 Daxner-Höck, G., Bernor, R.L., 2009. The early Vallesian vertebrates of Atzelsdorf
1003 (Late Miocene, Austria) 8. *Anchitherium*, Suidae and Castoridae (Mammalia).
1004 Annalen des Naturhistorisches Museum in Wien 111A, 557–584.

1005 de Bruijn, H., Daams, R., Daxner-Höck, G., Fahlbusch, V., Ginsburg, L., Mein, P.,
1006 Morales, J., Heinzmann, E., Mayhew, D.F., van der Meulen, A.J., Schmidt-Kittler,
1007 N., Telles Antunes, M., 1992. Report of the RCMNS working group on fossil
1008 mammals, Reinsenburg 1990. Newsletters on Stratigraphy 26, 65–118.

1009 de Gibert, J.M., Casanovas-Vilar, I., 2011. Contexto geológico del Mioceno de la
1010 cuenca del Vallès-Penedès. Paleontologia i Evolució memòria especial 6, 39–45.

1011 Evans, H.F., Westerhold, T., Paulsen, H., Channell, J.E.T., 2007. Astronomical ages
1012 for Miocene polarity chrons C4Ar–C5r (9.3–11.2 Ma), and for three excursion
1013 chrons within C5n.2n. Earth and Planetary Science Letters 256, 455–465.

1014 Fahlbusch, V., 1976. Report on the International Symposium on mammalian
1015 stratigraphy of the European Tertiary (München, April 11-14, 1975). Newsletters
1016 on Stratigraphy 5, 160–167.

1017 Forstén, A.M., 1968. Revision of the Palearctic Hipparion. Acta Zoologica Fennica
1018 119, 1–130.

1019 Forstén, A., 1978. *Hipparion primigenium* (v. Meyer, 1829), an early three-toed
1020 horse. Annales Zoologici Fennici 15, 298–313.

1021 Furió, M., Prieto, J., Van den Hoek Ostende, L., 2015. Three million years of “Terror-
1022 Shrew” (*Dinosorex*, Eulipotyphla, Mammalia) in the Miocene of the Vallès-
1023 Penedès Basin (Barcelona, Spain). Comptes Rendus Palevol 14, 111–124.

1024 Garcés, M., 1995. Magnetoestratigrafía de las sucesiones del Mioceno Medio y
1025 Superior del Vallès Occidental (Depresión del Vallès-Penedès, N.E. de España):
1026 Implicaciones biocronológicas y cronoestratigráficas. Ph.D. Dissertation,

- 1027 Universitat de Barcelona.
- 1028 Garcés, M., Agustí, J., Cabrera, L., Parés, J.M., 1996. Magnetostratigraphy of the
1029 Vallesian (late Miocene) in the Vallès-Penedès Basin (northeast Spain). *Earth and*
1030 *Planetary Science Letters* 142, 381–396.
- 1031 Garcés, M., Cabrera, L., Agustí, J., Parés, J.M., 1997. Old World first appearance
1032 datum of “Hipparion” horses: Late Miocene large-mammal dispersal and global
1033 events. *Geology* 25, 19–22.
- 1034 Garcés, M., Krijgsman, W., Peláez-Campomanes, P., Álvarez Sierra, M.A., Daams,
1035 R., 2003. Hipparion dispersal in Europe: magnetostratigraphic constraints from the
1036 Daroca area (Spain). *Coloquios de Paleontología Vol. Ext. 1*, 171–178.
- 1037 García-Paredes, I., Álvarez-Sierra, M.Á., Van den Hoek Ostende, L., Hernández-
1038 Ballarín, V., Hordijk, K., López-Guerrero, P., Oliver, A., Peláez-Campomanes, P.,
1039 2016. The Aragonian and Vallesian high-resolution micromammal succession from
1040 the Calatayud-Montalbán Basin (Aragón, Spain). *Comptes Rendus Palevol* 15,
1041 781–789.
- 1042 Gibert, J., 1975. New insectivores from the Miocene of Spain. I. *Proceedings of the*
1043 *Koninklijke Nederlandse Akademie van Wetenschappen B* 78, 108–123.
- 1044 Ginsburg, L., 1986. Chronology of the European pliopithecids. In: Else, J.C., Lee,
1045 P.C. (Eds.), *Primate Evolution*. Cambridge University Press, Cambridge, pp. 47–
1046 57.
- 1047 Golpe-Posse, J.M., 1972. *Suiformes del Terciario español y sus yacimientos (Tesis*
1048 *doctoral-Resumen) (revisado y reimprimido en Diciembre de 1972)*. *Paleontología*
1049 *y Evolución* 2, 1–197.
- 1050 Golpe-Posse, J.M., 1974. Faunas de yacimientos con suiformes en el Terciario
1051 español. *Paleontología y Evolución* 8, 1–87.

- 1052 Golpe-Posse, J.M., 1975. Un nuevo tayasuido en el Vindoboniense terminal de Castell
1053 de Barberà (Cuenca del Vallès, España). Boletín Informativo del Instituto de
1054 Paleontología de Sabadell 7, 39–43.
- 1055 Golpe-Posse, J. M., 1977. *Barberahyus castellensis* n.g.n.sp., Tayasuido del
1056 Vindoboniense terminal de Castell de Barberà (Cuenca del Vallès, España).
1057 Paleontología y Evolución 12, 31–43.
- 1058 Golpe-Posse, J.M., 1978. Un nuevo tayasuido en el Vindoboniense terminal de Castell
1059 de Barberà (Cuenca del Vallès, España). Boletín Informativo del Instituto de
1060 Paleontología de Sabadell 13, 13–17.
- 1061 Golpe Posse, J.M., 1982. Los hispanopitecos (Primates, Pongidae) de los yacimientos
1062 del Vallès Penedès (Cataluña - España). I: Material ya descrito. Butlletí Informatiu
1063 de l’Institut de Paleontologia de Sabadell 14, 63–69.
- 1064 Golpe Posse, J.M., 1993. Los Hispanopitecos (Primates, Pongidae) de los yacimientos
1065 del Vallès-Penedès (Cataluña, España). II: Descripción del material existente en el
1066 Instituto de Paleontología de Sabadell. Paleontologia i Evolució 26–27, 151–224.
- 1067 Harrison, T., 1991. Some observations on the Miocene hominoids from Spain. Journal
1068 of Human Evolution 19, 515–520.
- 1069 Harzhauser, M., 2009. The early Vallesian vertebrates of Atzelsdorf (Late Miocene,
1070 Austria) 2. Geology. Annalen des Naturhistorischen Museums in Wien 111A, 479–
1071 488.
- 1072 ICGC (Institut Cartogràfic i Geològic de Catalunya), 2018. VISSIR v3.26 (VISor del
1073 Servidor d’Imatges Ràster). Generalitat de Catalunya, <http://www.icc.cat/vissir3/>
1074 (accessed July 1st 2018).

1075 ICZN (International Commission on Zoological Nomenclature), 1999. International
1076 Code of Zoological Nomenclature, 4th ed. The International Trust for Zoological
1077 Nomenclature, London.

1078 Kirschvink, J.L., 1980. The least-squares line and plane and the analysis of
1079 palaeomagnetic data. *Geophysical Journal of the Royal Astronomical Society* 62,
1080 699–718.

1081 Krijgsman, W., Langereis, C.G., Daams, R., van der Meulen, A.J., 1994.
1082 Magnetostratigraphic dating of the middle Miocene climate change in the
1083 continental deposits of the Aragonian type area in the Calatayud-Teruel basin
1084 (Central Spain). *Earth and Planetary Science Letters* 128, 513–526.

1085 Lindsay, E.H., Tedford, R.H., 1990. Development and application of land mammal
1086 ages in North American and Europe, a comparison. In: Lindsay, E.H., Fahlbusch,
1087 V., Mein, P. (Eds.), *European Neogene Mammal Chronology*. Plenum Press, New
1088 York, pp. 601–624.

1089 MacFadden, B. J. (2001). Three-toed browsing horse *Anchitherium clarencei* from the
1090 early Miocene (Hemingfordian) Thomas Farm, Florida. *Bulletin of the Florida*
1091 *Museum of Natural History*, 43, 79-109.

1092 Marigó, J., Susanna, I., Minwer-Barakat, R., Madurell-Malapeira, J., Moyà-Solà, S.,
1093 Casanovas-Vilar, I., Robles, J.M., Alba, D.M., 2014. The primate fossil record in
1094 the Iberian Peninsula. *Journal of Iberian Geology* 40, 179–211.

1095 Marmi, J., Casanovas-Vilar, I., Robles, J.M., Moyà-Solà, S., Alba, D.M., 2012. The
1096 paleoenvironment of *Hispanopithecus laietanus* as revealed by paleobotanical
1097 evidence from the Late Miocene of Can Llobateres 1 (Catalonia, Spain). *Journal of*
1098 *Human Evolution* 62, 412–423.

- 1099 Moyà-Solà, S., 1983. Los Boselaphini (Bovidae Mammalia) del Neógeno de la
1100 Península Ibérica. Publicaciones de Geología, Universitat Autònoma de Barcelona.
- 1101 Moyà Solà, S., Pons Moyà, J., Köhler, M., 1990. Primates catarrinos (Mammalia) del
1102 Neógeno de la península Ibérica. *Paleontologia i Eolució* 23, 41–45.
- 1103 Moyà-Solà, S., Köhler, M., Rook, L., 2005. The *Oreopithecus* thumb: a strange case
1104 in hominoid evolution. *Journal of Human Evolution* 49, 395–404.
- 1105 Moyà-Solà, S., Köhler, M., Alba, D.M., Casanovas-Vilar, I., Galindo, J., Robles,
1106 J.M., Cabrera, L., Garcés, M., Almécija, S., Beamud, E., 2009. First partial face
1107 and upper dentition of the Middle Miocene hominoid *Dryopithecus fontani* from
1108 Abocador de Can Mata (Vallès-Penedès Basin, Catalonia, NE Spain): taxonomic
1109 and phylogenetic implications. *American Journal of Physical Anthropology* 139,
1110 126–145.
- 1111 Moyà-Solà, S., Alba, D.M., Almécija, S., 2013. A proximal radius of *Barberapithecus*
1112 *huerzeleri* (Primates, Pliopithecidae) from the Miocene site of Castell de Barberà
1113 (NE Iberian Peninsula). *Journal of Vertebrate Paleontology* 33 S2, 182.
- 1114 Murphy, M.A., Salvador, A. (Eds.), 1999. International Stratigraphic Guide — An
1115 abridged version. *Episodes* 22, 255–271.
- 1116 Ogg, J.G., 2012. Geomagnetic polarity time scale. In: Gradstein, F.M., Ogg, J.G.,
1117 Schmitz, M.D., Ogg, G.M. (Eds.), *The Geologic Time Scale 2012, Volume 1*.
1118 Elsevier, Amsterdam, pp. 85–113.
- 1119 Pérez de los Ríos, M., Alba, D.M., Moyà-Solà, S., 2013. Taxonomic attribution of the
1120 La Grive hominoid teeth. *American Journal of Physical Anthropology* 151, 558–
1121 565.

- 1122 Pesquero, M.D., Arribas, A., 2002. Los restos de *Hipparion* (Equidae, Mammalia) en
1123 las colecciones de vertebrados del Museo Geominero (IGME): aspectos históricos
1124 y actualización taxonómica. Boletín Geológico y Minero 113, 97–108.
- 1125 Petter, G., 1976. Étude d'un nouvel ensemble de petits carnivores du Miocène
1126 d'Espagne. Géologie Méditerranéenne 3, 135–154.
- 1127 Pirlot, P.L., 1956. Les formes européennes du genre *Hipparion*. Memorias y
1128 Comunicaciones del Instituto Geológico 14, 1–121.
- 1129 Roberts, A.P., Lewin-Harris, J.C., 2000. Marine magnetic anomalies: evidence that
1130 'tiny wiggles' represent short-period geomagnetic polarity intervals. Earth and
1131 Planetary Science Letters 183, 375–388.
- 1132 Robles, J. M., Alba, D.M., Moyà-Solà, S., Casanovas-Vilar, I., Galindo, J., Rotgers,
1133 C., Almécija, S., Carmona, R., 2010. New craniodental remains of *Trocharion*
1134 *albanense* Major, 1903 (Carnivora, Mustelidae), from the Vallès-Penedès Basin
1135 (Middle to Late Miocene, Barcelona, Spain). Journal of Vertebrate Paleontology
1136 30, 547–562.
- 1137 Robles, J.M., Alba, D.M., Casanovas-Vilar, I., Galindo, J., Cabrera, L., Carmona, R.,
1138 Moyà-Solà, S., 2011. On the age of the paleontological site of Can Missert
1139 (Terrassa, Vallès-Penedès Basin, NE Iberian Peninsula). In: Pérez-García, A.,
1140 Gascó, J., Gasulla, J.M., Escaso, F. (Eds.) Viajando a Mundos Pretéritos.
1141 Ayuntamiento de Morella, Morella, pp. 339–346.
- 1142 Robles, J.M., Alba, D.M., Fortuny, J., De Esteban-Trivigno, S., Rotgers, C., Balaguer,
1143 J., Carmona, R., Galindo, J., Almécija, S., Bertó, J.V., Moyà-Solà, S., 2013a. New
1144 craniodental remains of the barbourfelid *Albanosmilus jourdani* (Filhol, 1883)
1145 from the Miocene of the Vallès-Penedès (NE Iberian Peninsula) and the phylogeny
1146 of the Barbourfelini. Journal of Systematic Palaeontology 11, 993–1022.

- 1147 Robles, J.M., Madurell-Malapeira, J., Abella, J., Rotgers, C., Carmona, R., Alméjida,
1148 S., Balaguer, J., Alba, D.M., 2013b. New *Pseudaelurus* and *Styriofelis* remains
1149 (Carnivora: Felidae) from the middle Miocene of Abocador de Can Mata (Vallès-
1150 Penedès Basin). *Comptes Rendus Palevol* 12, 101–113.
- 1151 Rotgers, C., Alba, D.M., 2011. The genus *Anchitherium* (Equidae: Anchitheriinae) in
1152 the Vallès-Penedès Basin (Catalonia, Spain). In: Pérez-García, A., Gascó, J.,
1153 Gasulla, J.M., Escaso, F. (Eds.), *Viajando a Mundos Pretéritos*. Ayuntamiento de
1154 Morella, Morella, pp. 347–354.
- 1155 Rotgers, C., Alba, D.M., Robles, J.M., Casanovas-Vilar, I., Galindo, J., Bertó, J.V.,
1156 Moyà-Solà, S., 2011. A new species of *Anchitherium* (Equidae: Anchitheriinae)
1157 from the Middle Miocene of Abocador de Can Mata (Vallès-Penedès Basin, NE
1158 Iberian Peninsula). *Comptes Rendus Palevol* 10, 567–576.
- 1159 Salesa, M.J., Sánchez, I.M., Morales, J., 2004. Presence of the Asian horse *Sinohippus*
1160 in the Miocene of Europe. *Acta Palaeontologica Polonica* 49, 189–196.
- 1161 Sánchez, I.M., Salesa, M.J., Morales, J., 1998. Revisión sistemática del género
1162 *Anchitherium* Meyer 1834 (Equidae; Perissodactyla) en España. *Estudios*
1163 *Geológicos* 54, 39–63.
- 1164 Santafé Llopis, V., 1978. Rinocerótidos fósiles de España. PhD. dissertation,
1165 Universidad de Barcelona.
- 1166 Santafé i Llopis, J.V., 1978a. Síntesi de la distribució dels rinoceròtids fòssils a les
1167 conques Vallès-Penedès. *Butlletí Informatiu de l'Institut de Paleontologia de*
1168 *Sabadell* 10, 34–40.
- 1169 Santafé Llopis, J.V., 1978b. Revisión de los Rinocerótidos miocénicos del Vallès-
1170 Penedès. *Acta Geológica Hispánica* 13, 43–45.
- 1171 Steininger, F.F., 1999. Chronostratigraphy, geochronology and biochronology of the

1172 Miocene "European Land Mammal Mega-Zones" (ELMMZ) and the Miocene
1173 "Mammal-Zones (MN-Zones)". In: Rössner, G.E., Heissig, K. (Eds.), The Miocene
1174 Land Mammals of Europe. Verlag Fritz Pfeil, München, pp. 9–24.

1175 Sukselainen, L., Fortelius, M., Harrison, T., 2015. Co-occurrence of pliopithecoid and
1176 hominoid primates in the fossil record: An ecometric analysis. *Journal of Human*
1177 *Evolution* 84, 25–41.

1178 Urciuoli, A., Zanolli, C., Fortuny, J., Almécija, S., Schillinger, B., Moyà-Solà, S.,
1179 Alba, D.M., 2018. Neutron-based computed microtomography: *Pliobates*
1180 *cataloniae* and *Barberapithecus huerzeleri* as a test-case study. *American Journal*
1181 *of Physical Anthropology* 166, 987–993.

1182 van Dam, J., 2003. European Neogene mammal chronology: past, present and future.
1183 *Deinsea* 10, 85–95.

1184 van der Made, J., Ribot, F., 1999. Additional hominoid material from the Miocene of
1185 Spain and remarks on hominoid dispersals into Europe. *Contributions to Tertiary*
1186 *and Quaternary Geology* 36, 25–39.

1187 van der Meulen, A.J., García-Paredes, I., Álvarez-Sierra, M.Á., van den Hoek
1188 Ostende, L. W., Hordijk, K., Oliver, A., López-Guerrero, P., Hernández-Ballarín,
1189 V., Peláez-Campomanes, P., 2011. Biostratigraphy or biochronology? Lessons
1190 from the Early and Middle Miocene small Mammal Events in Europe. *Geobios* 44,
1191 309–321.

1192 van der Meulen, A.J., García-Paredes, I., Álvarez-Sierra, M.Á., van den Hoek
1193 Ostende, L. W., Hordijk, K., Oliver, A., Peláez-Campomanes, P., 2012. Updated
1194 Aragonian biostratigraphy: Small mammal distribution and its implications for the
1195 Miocene European chronology. *Geologica Acta* 10, 159–179.

1196 Villa, A., Delfino, M., Luján, À.H., Almécija, S., Alba, D.M., 2017. First record of

- 1197 *Latonia gigantea* (Anura, Alytidae) from the Iberian Peninsula. Historical Biology.
1198 <https://doi.org/10.1080/08912963.2017.1371712>
- 1199 Woodburne, M.O. (Ed.), 2004a. Late Cretaceous and Cenozoic Mammals of North
1200 America. Biostratigraphy and Geochronology. Columbia University Press, New
1201 York.
- 1202 Woodburne, M.O., 2004b. Principles and procedures. In: Woodburne, M.O. (Ed.),
1203 Late Cretaceous and Cenozoic Mammals of North America. Biostratigraphy and
1204 Geochronology. Columbia University Press, New York, pp. 1–20.
- 1205 Woodburne, M.O., 2005. A new occurrence of *Cormohipparion*, with implications for
1206 the Old World *Hippotherium* datum. Journal of Vertebrate Paleontology 25, 256–
1207 257.
- 1208 Woodburne, M.O., 2007. Phyletic diversification of the *Cormohipparion occidentale*
1209 complex (Mammalia; Perissodactyla, Equidae), Late Miocene, North America, and
1210 the origin of the Old World *Hippotherium* datum. Bulletin of the American
1211 Museum of Natural History 306, 1–138.
- 1212 Woodburne, M.O., 2009. The early Vallesian vertebrates of Atzelsdorf (Late
1213 Miocene, Austria) 9. *Hippotherium* (Mammalia, Equidae). Annalen des
1214 Naturhistorischen Museums in Wien 111A, 585–604.
- 1215 Zouhri, S., Bensalmia, A., 2005. Révision systématique des *Hipparion* sensu lato
1216 (Perissodactyla, Equidae) de l'ancien monde. Estudios Geológicos 61, 61–99.
1217
- 1218 **Figure captions**
- 1219
- 1220 **Figure 1.** Fossil primates from Castell de Barberà: the pliopithecoid *Barberapithecus*
1221 *huerzeleri* (A–L) and the large-bodied hominoid (M–O). A–K) IPS1724 (holotype),

1222 selected associated teeth from a single female individual, including the right I¹ (A),
1223 the left C¹ (B), and the left C₁ (C), in lingual (left) and labial (right) views, as well as
1224 the right M¹ (D), the right M² (E), the right M³ (F), the right P₃ (G), the left P₄ (H), the
1225 left M₁ (I), the right M₂ (J), and the right M₃ (K), in occlusal views. L) IPS1823, male
1226 right C¹ in lingual (left) and labial (right) views. M) IPS4335, partial distal pollical
1227 phalanx, in palmar (left) and dorsal (right) views. N) IPS4333, right proximal pollical
1228 phalanx, from left to right in palmar, radial, ulnar, and distal views. O) IPS4334, distal
1229 fragment of right humeral diaphysis assigned to cf. *Dryopithecus fontani* by Alba et
1230 al. (2012), from left to right in anterior, medial, posterior, and lateral views. Scale bars
1231 equal 1 cm, except for O (5 cm). Images are reproduced from previous papers by the
1232 authors: L) Alba and Moyà-Solà (2012: Fig. 2); M–N) Almécija et al. (2012: Fig. 1);
1233 O) Alba et al. (2011: Fig. 1).

1234

1235 **Figure 2.** A) Aerial photograph of the site of Castell de Barberà (CB), as well as that
1236 of the bottom of sections sampled for magnetostratigraphy in this study. Modified
1237 from base orthophotos downloaded from VISSIR v3.26 (ICGC, 2018: sheets 288-119
1238 and 288-120, scale 1:5000), ©Institut Cartogràfic i Geològic de Catalunya, with
1239 permission from licence Creative Commons (CC) – Attribution 4.0 International (CC
1240 BY 4.0; see <http://www.icgc.cat/Ajuda/Avis-legal> for reuse policies allowed for ICGC
1241 web content). B) Detail of Castell de Barberà (Section 2) during the excavation of
1242 CB-D (equivalent to the original main fossiliferous layer) in June 2015.

1243

1244 **Figure 3.** Simplified geological map of the Vallès-Penedès Basin. The location of
1245 Castell de Barberà is denoted by a black star. Top left inset: location of the Vallès-

1246 Penedès Basin within the Iberian Peninsula. Modified from Casanovas-Vilar et al.
1247 (2016a: Fig. 1).

1248

1249 **Figure 4.** Examples of NRM stepwise thermal demagnetization of samples of the
1250 Castell de Barberà composite section. Stratigraphic location of samples (uppercase
1251 letters) is indicated in Figure 5. Diagrams represent orthogonal projection of vector
1252 endpoint demagnetization data; black and white dots represent the projection on the
1253 horizontal and vertical planes, respectively. Red lines represent the least-square fit
1254 defining the paleomagnetic direction. Q1 and Q2 indicates ranked quality (see SOM
1255 Table S1).

1256

1257 **Figure 5.** Magnetostratigraphic results for the two sampled sections at Castell de
1258 Barberà: A) Section 1; B) Section 2, where the site of Castell de Barberà is located.
1259 Sideways triangles denote the stratigraphic position of paleomagnetic samples. Red
1260 letters and triangles indicate the location of samples from Figure 4. Black circles
1261 indicate directions of higher quality (Q1 and Q2), while white circles indicate
1262 directions of low quality (Q3)—see SOM Table S1. Abbreviations: c =
1263 conglomerates; l = lutites; s = sandstones; VGP = virtual geomagnetic pole.

1264

1265 **Figure 6.** Composite local magnetostratigraphic section for Castell de Barberà and its
1266 preferred correlation with the geomagnetic polarity time scale (GPTS; Ogg, 2012), as
1267 well as European land mammal ages (ELMA), MN (Mammal Neogene) units, and
1268 local biozones of the Vallès-Penedès Basin (after Casanovas-Vilar et al., 2016b). See
1269 Figure 5 for detailed magnetostratigraphic results of the two sampled sections upon
1270 which this composite magnetostratigraphic section is based. Abbreviations: N =

1271 normal polarity magnetozones; R = reversed polarity magnetozones. The arrow next
1272 to N2 denotes the stratigraphic position of the main fossiliferous layer of Castell de
1273 Barberà (CB-D). Half-width polarity intervals in the GPTS represent short
1274 geomagnetic excursions after Evans et al. (2007). Short horizontal lines to the left of
1275 the polarity column represent ‘tinny wiggles’ from the sea floor magnetic anomaly
1276 stacks (Cande and Kent 1992), later interpreted as true geomagnetic polarity reversals
1277 (Roberts and Lewin-Harris, 2000).

1278

1279 **Figure 7.** Right distal humeral fragment of *Hippotherium* sp. IPS87652 from Castell
1280 de Barberà CB-D (equivalent to the main fossiliferous layer) recovered in 2015 (A),
1281 compared to those of *Hippotherium catalaunicum* IPS11117 from Can Llobateres (B)
1282 and IPS32449 from Polinyà (C, reversed). All specimens are depicted in anterior (left)
1283 and posterior (right) views.

1284

Pliopithecoids

A



C



B



D



E



F



G



H



I



J



K



L



Hominoids

M



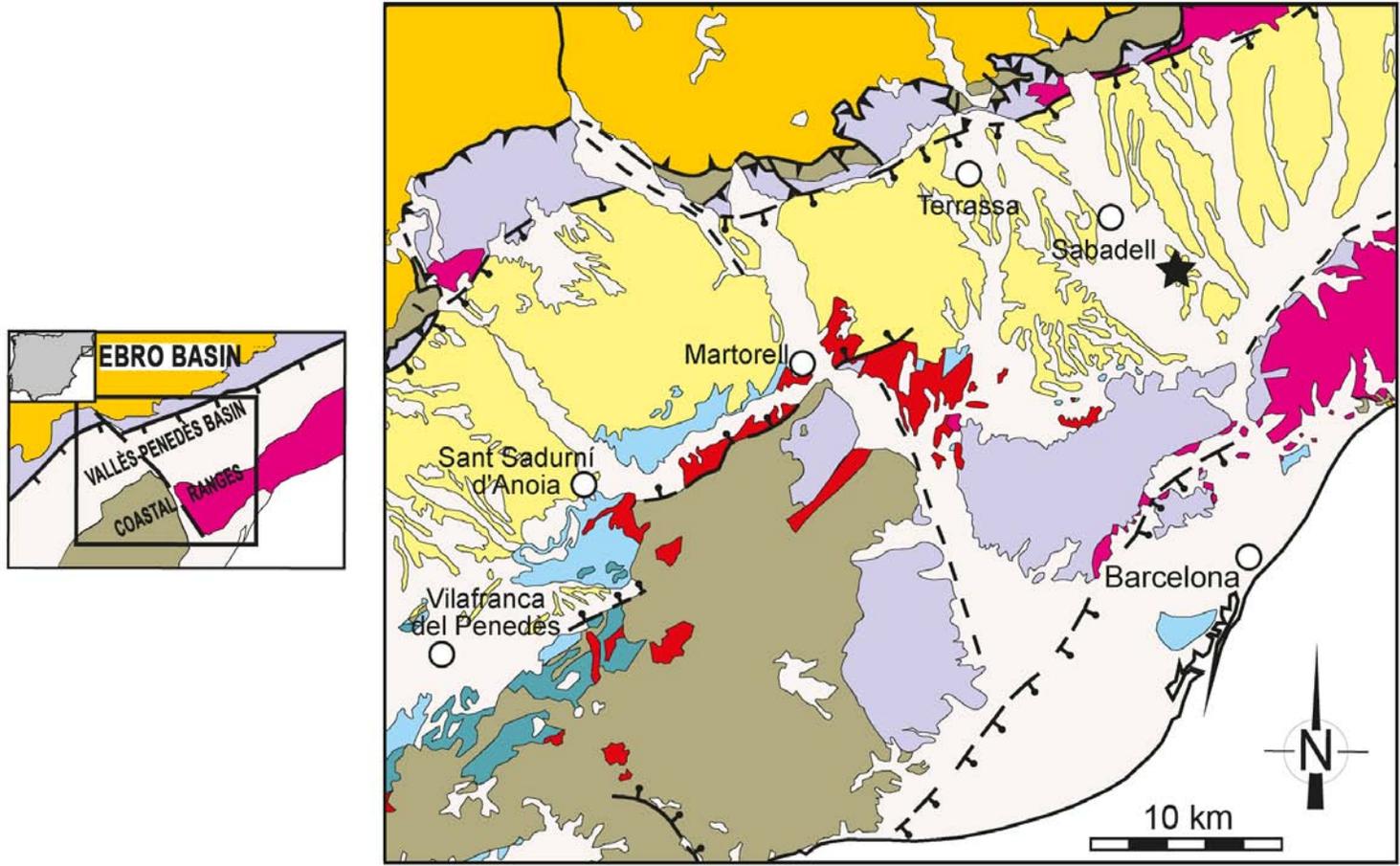
N



O







NEOGENE-QUATERNARY

- Pliocene and Quaternary sediments
- Middle to late Miocene continental units (Langhian-Tortonian): alluvial fan facies
- Miocene marine and transitional units (Late Burdigalian-Langhian): corallgal carbonate facies
- Miocene marine and transitional units (Late Burdigalian-Serravallian): marls, silts, bioclastic sands and sandstones
- Early Miocene continental units (Early and Late Burdigalian): alluvial fan and shallow lacustrine facies

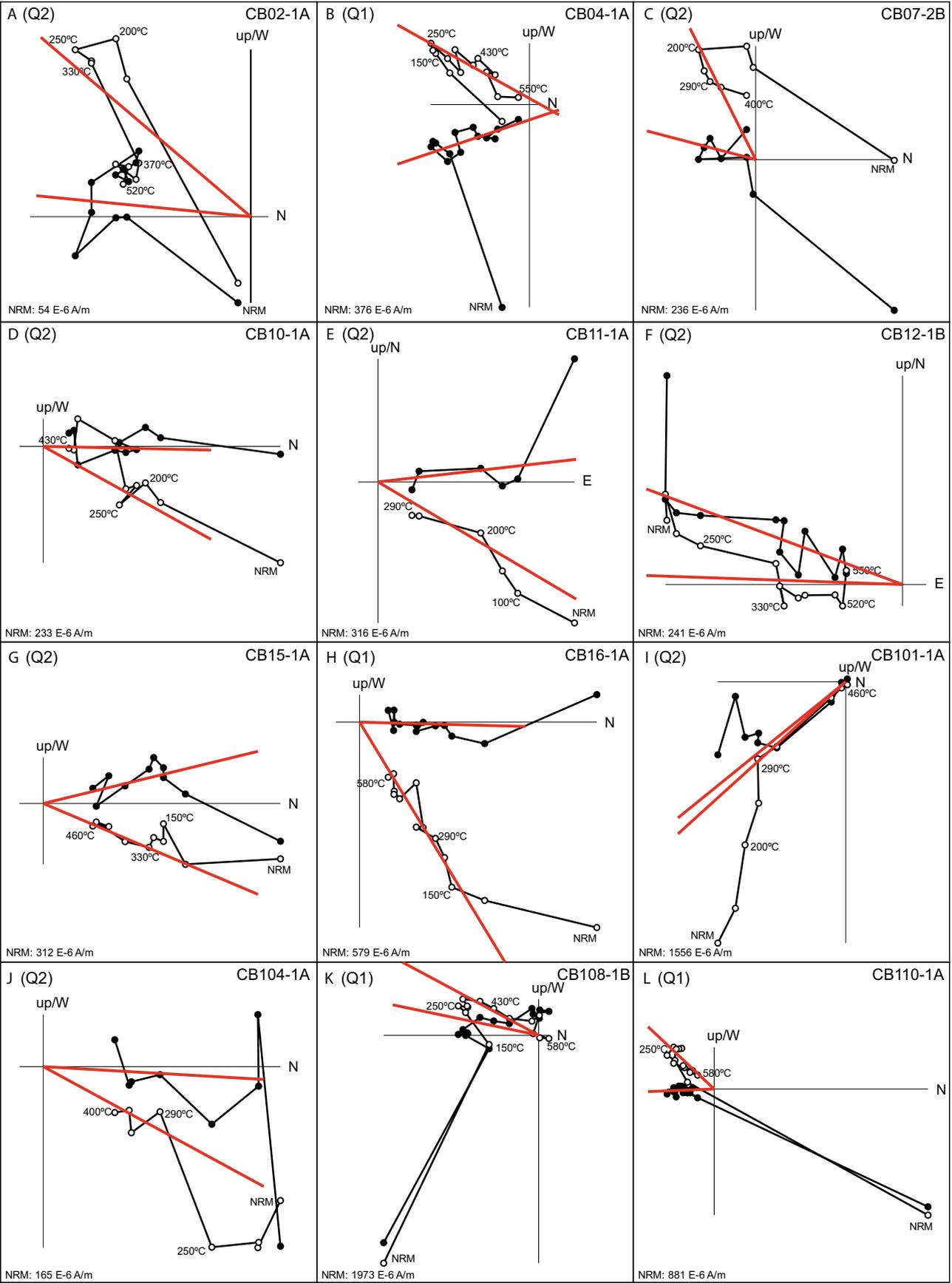
EBRO FORELAND BASIN INFILL

- Paleogene (conglomerates, sandstones, lutites)

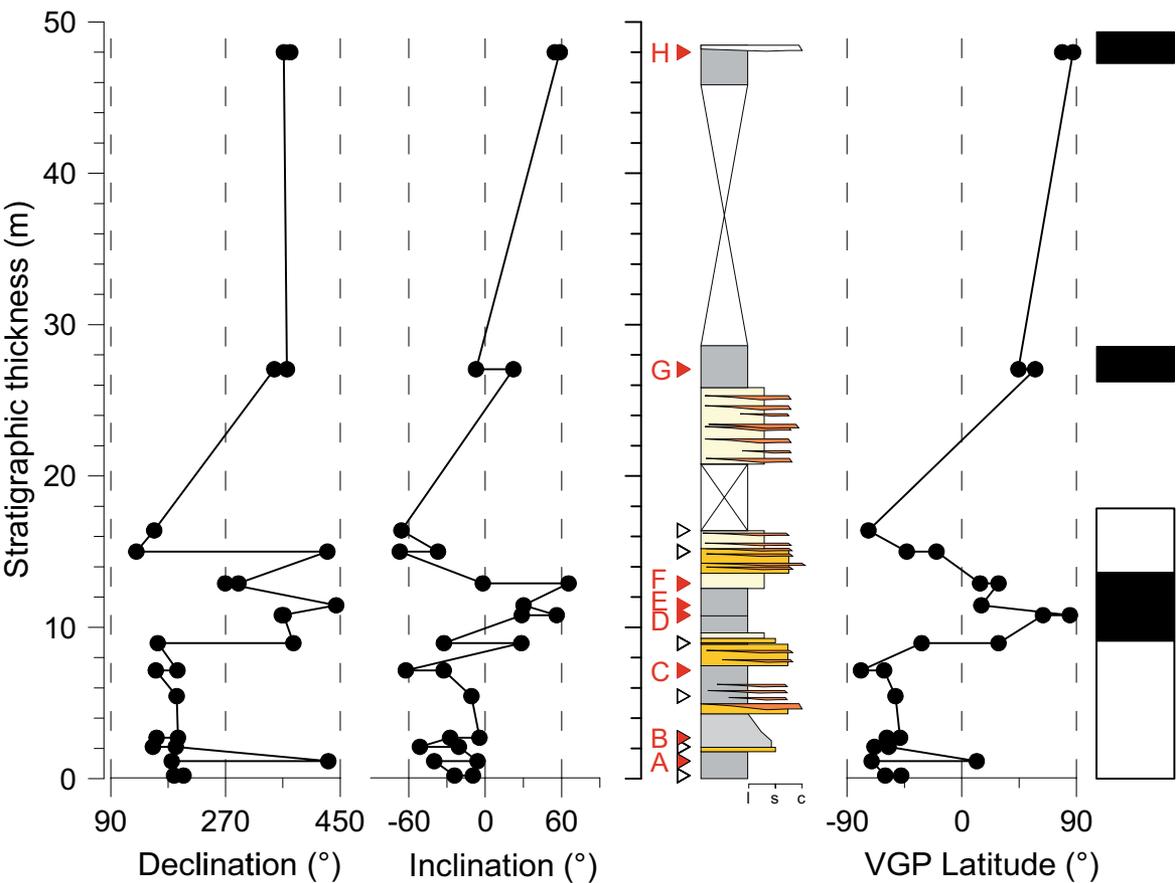
PALEOZOIC BASEMENT AND MESOZOIC COVER

- Mesozoic (carbonates, sandstones, lutites)
- Hercynian intrusive rocks (mostly granitoids)
- Paleozoic sedimentary and metamorphic rocks

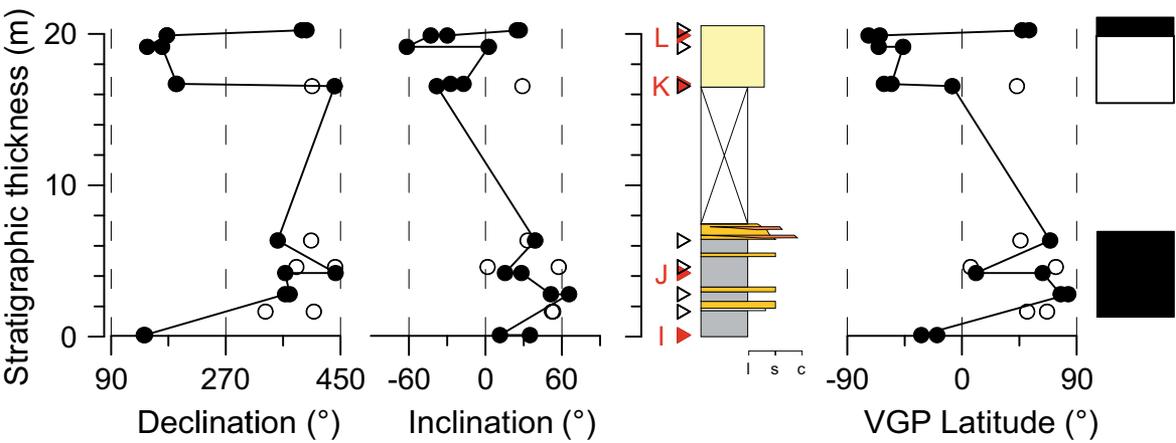
- Fault
- Inferred fault
- T Normal fault
- T Thrust fault

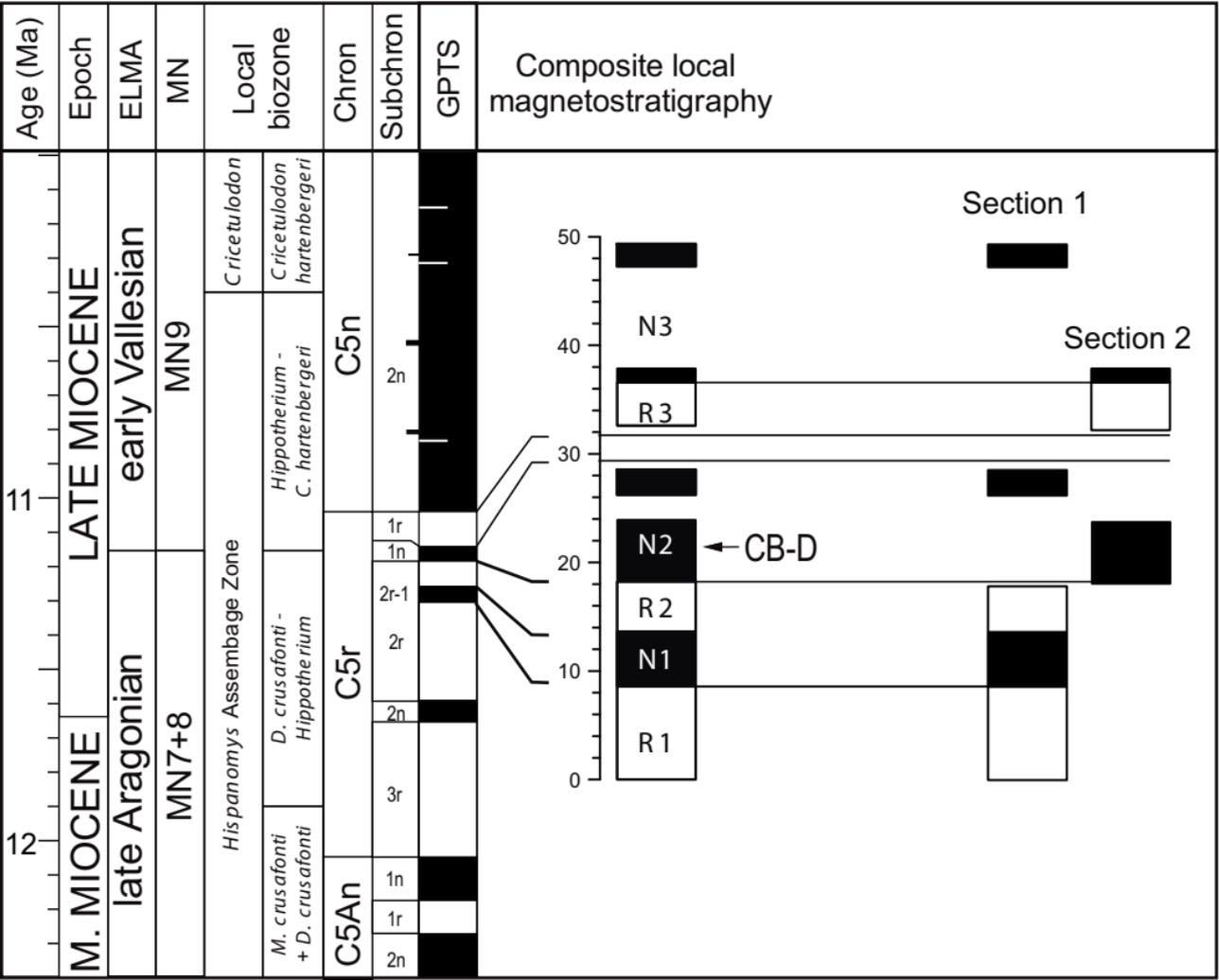


A Section 1



B Section 2





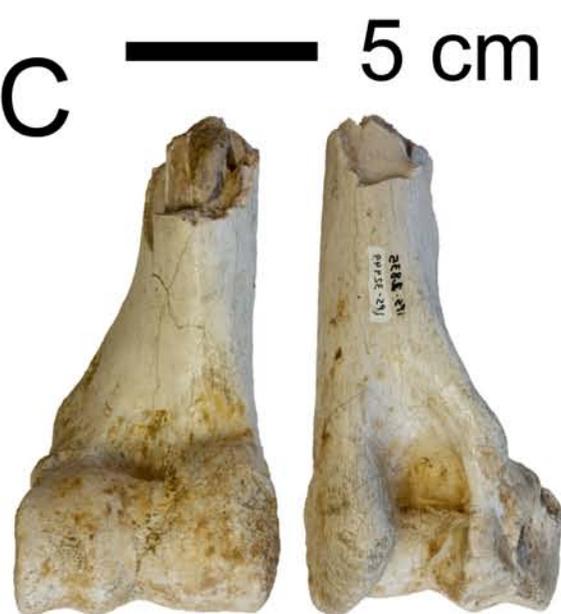
A



B



C



1 **Table 1**

2 Updated faunal list of the rodent assemblage from Castell de Barberà.

Family	Species
Castoridae	<i>Euroxenomys minutus</i>
	<i>Chalicomys jaegeri</i>
Sciuridae	<i>Spermophilinus bredai</i>
	<i>Albanensia albanensis quiricensis</i>
	<i>Miopetaurista neogrivensis</i>
Gliridae	<i>Bransatoglis astaracensis</i>
	<i>Paraglirulus werenfelsi</i>
	<i>Muscardinus</i> aff. <i>sansaniensis</i>
	<i>Muscardinus hispanicus</i>
	<i>Myomimus</i> sp. ^a
	<i>Glirudinus undosus</i>
Cricetidae	<i>Myoglis meini</i>
	<i>Anomalomys</i> sp. ^b
	<i>Eumyarion leemanni</i>
	<i>Hispanomys daamsi</i>
	<i>Megacricetodon minutus</i>
	<i>Megacricetodon</i> cf. <i>ibericus</i> ^c
	<i>Democricetodon brevis nemoralis</i>
<i>Democricetodon</i> cf. <i>crusafonti</i> ^b	

3 ^aNot reported by Casanovas-Vilar et al. (2016b).

4 ^b New taxa identified from CB, based on the scarce material recovered in 2015:

5 *Anomalomys* sp. ^cComes from CB-layer D (equivalent to the classical original main

6 fossiliferous level), whereas *Democricetodon* cf. *crusafonti* from ~~layer~~ CB-B (~2 m
7 below CB-D) and *Megacricetodon* cf. *ibericus* from ~~layer~~ CB-E (~1 m above CB-D).

8

Supplementary Online Material (SOM):

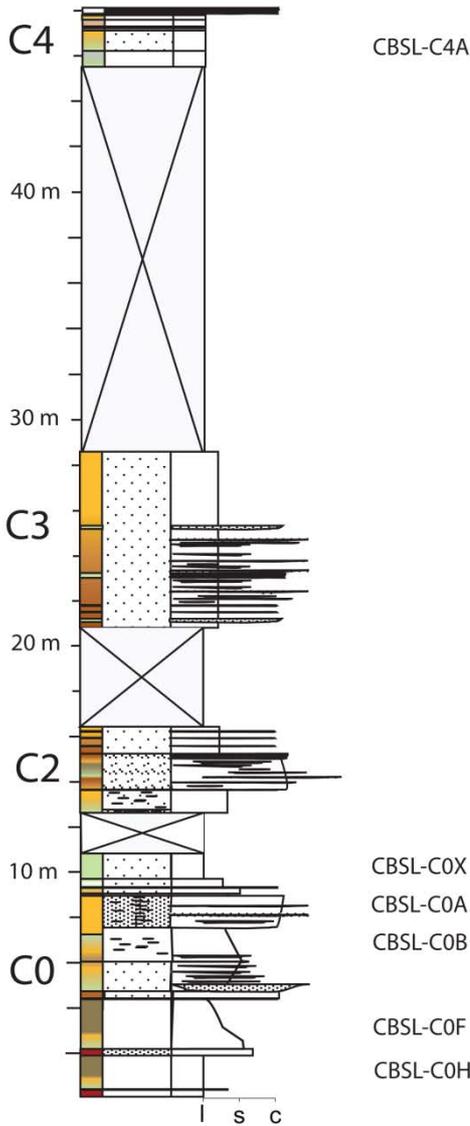
Bio- and magnetostratigraphic correlation of the Miocene primate-bearing site of Castell de Barberà to the earliest Vallesian: End of the controversy

SOM S1

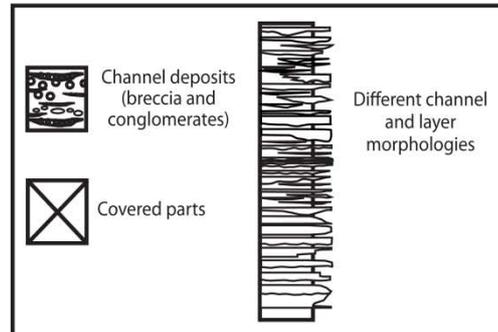
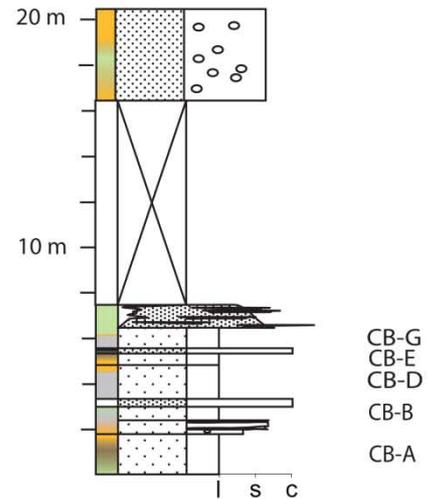
Castell de Barberà stratigraphic profiles

The stratigraphic profiles performed during the 2014 and 2015 campaigns and their correlation have been depicted in SOM Fig. S1. The various geological layers identified during fieldwork were termed with uppercase letters in each one of the outcrops generated by the digger machine, but these letters have not been indicated in SOM Figure 1, because they are not equivalent between the two sections or even among different outcrops within Section 1. Instead, only those layers that delivered fossils (being considered paleontological localities) have been depicted in SOM Figure S1. The fossil material recovered from Section 1 in 2014 and from Section 2 in 2015 was recorded as coming from the sites of Castell de Barberà s.l. (abbreviated CBSL) and Castell de Barberà (*sensu stricto*, abbreviated as CB), respectively. Section 1 is composed of four different outcrops (C0, C2, C3, C4) separated by covered portions (note that outcrop C1 is a short section that overlaps with outcrop 0 and yielded no fossils, so it has been omitted from SOM Fig. S1). Section 2, in turn, is composed of two different outcrops that were given no distinct numbers. The exact locality of provenance for each fossil was termed using the abbreviation of the site (see above), followed by a hyphen and an uppercase letter corresponding to the stratigraphic layer; in the case of Section 1, the outcrop number was intercalated before the letter denoting the layer. In Section 1, fossils were recovered from CBSL-C0H, C0F, C0B, C0A, C0X, and C4A (from bottom to top), whereas in Section 2 fossils were found at CB-A, B, D, E, and G (in the same stratigraphic order). CB-D is stratigraphically equivalent to the original layer that delivered most of the fossils from the site.

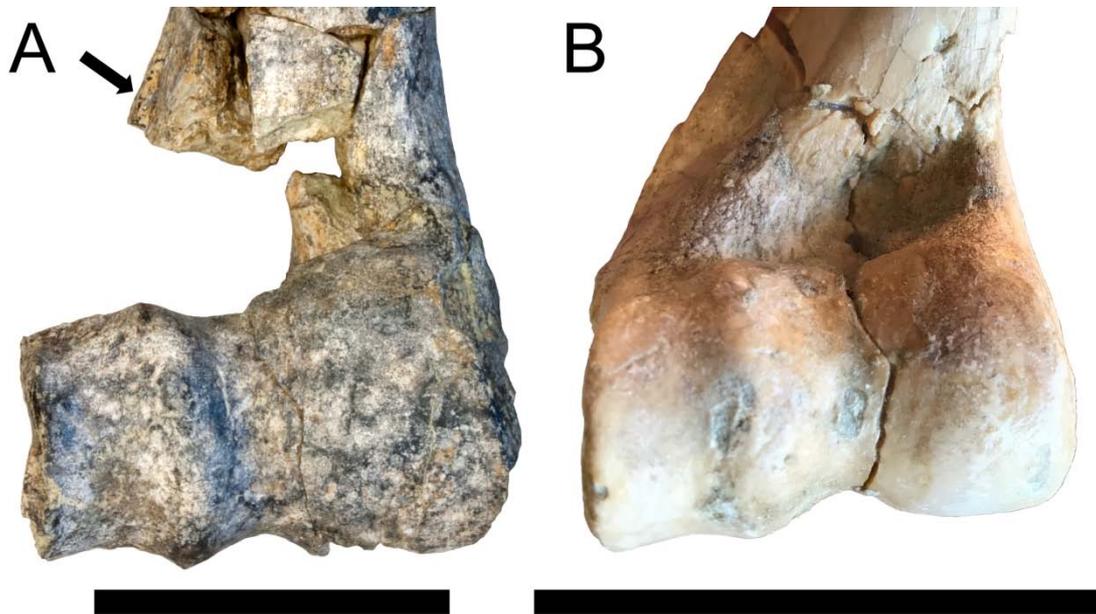
Section 1



Section 2



SOM Figure S1. Lithostratigraphic correlation of the two sections of Castell de Barberà excavated in 2014 and 2015. Only those layers that yielded some fossil remains are defined as fossiliferous localities and denoted in the figure (see SOM S1 for further details about their nomenclature)—including CB-D, equivalent to the main fossiliferous layer of Castell de Barberà. Abbreviations: c = conglomerates; l = lutites; s = sandstones.



SOM Figure S2. Anterior view of the humerus distal articular morphology of: A) *Hippotherium* sp. IPS87652 from Castell de Barberà CB-D; B) *Anchitherium* cf. *matritense* (without number) from Mahou site (MN5, Madrid Basin). Arrow in A denotes the position of the proximal end of the narrow but deep groove that marks the attachment of the extensor digitorum communis muscle. Specimen in B is from the left side but has been mirrored to ease comparison with IPS87652. Specimens scaled to approximately the same distal width; scale bars = 5 cm. Image B kindly provided by Jorge Morales.

SOM Table S1

Calculated paleomagnetic directions from Castell de Barberà sections.

Section 1 samples	Level (m)	Declination (°)	Inclination (°)	Intensity (E-6 A/m)	Temperature range (°C)	MAD (°)	VGP latitude (°)	Q
CB01-1A	0.2	204	-10	134	370–460	1.8	-47.5	2
CB01-1B	0.2	189	-24	84	400–520	5.8	-60.0	2
CB02-1A	1.2	186	-40	121	250–430	14.6	-70.7	2
CB02-1B	1.2	431	-6	266	290–490	7.8	11.8	2
CB03-1A	2.1	192	-21	100	200–370	19.4	-57.4	2
CB03-1B	2.1	156	-51	97	400–550	9.8	-68.5	2
CB04-1A	2.7	162	-28	178	250–580 ^a	19.1	-58.8	1
CB04-1B	2.7	195	-4	68	400–430	3.0	-48.4	2
CB06-2A	5.5	193	-11	245	250–370	9.9	-52.0	2
CB07-1A	7.2	160	-33	158	250–330	10.8	-60.9	2
CB07-2B	7.2	195	-62	142	200–400	10.0	-79.1	2
CB08-1A	9.0	164	28	83	250–370	20.6	-31.4	2
CB08-1C	9.0	376	-32	191	330–490	9.9	28.9	2
CB10-1A	10.8	361	29	84	250–330	6.7	63.8	2
CB10-2A	10.8	358	56	72	250–330	6.9	84.9	2
CB11-1A	11.5	444	30	135	200–290	9.0	15.4	2
CB12-1A	12.9	269	65	155	250–430	8.2	28.9	2
CB12-1B	12.9	290	-2	163	250–550	11.5	14.4	2

CB13-1A	15.0	430	-67	150	290–370	4.0	-19.8	2
CB13-2A	15.0	130	-37	68	370–400	8.6	-43.1	2
CB14-1A	16.4	158	-66	85	290–400	17.1	-73.2	2
CB15-1A	27.1	347	22	151	330–400	6.3	57.7	2
CB15-2A	27.1	366	-7	63	330–400	5.4	44.5	2
CB16-1A	48.0	361	58	293	200–580	5.4	87.4	1
CB16-1B	48.0	372	55	173	430–610	5.2	78.9	1

Section 2 samples	Level (m)	Declination (°)	Inclination (°)	Intensity (E-6 A/m)	Temperature range (°C)	MAD (°)	VGP latitude (°)	Q
CB101-1A	0.1	141	35	687	290–400	4.4	-19.5	2
CB101-1B	0.1	143	11	522	250–370	17.9	-31.9	2
CB102-1A	1.7	408	52	95	290–370	6.3	51.2	3
CB102-1B	1.7	332	53	125	290–330	9.2	66.8	3
CB103-1A	2.8	363	66	129	290–400	1.9	83.4	2
CB103-1B	2.8	370	51	161	330–430	5.9	77.5	2
CB104-1A	4.2	363	28	64	290–400	11.4	63.3	2
CB104-1B	4.2	442	15	148	250–330	5.2	11.0	2
CB105-1A	4.6	442	2	106	370–500	10.7	6.8	3
CB105-1B	4.6	381	57	125	200–400	16.1	73.7	3
CB106-1B	6.4	351	39	45	330–430	22.9	69.2	2
CB106-1A	6.4	404	33	80	330–450	5.2	45.9	3
CB107-2A	16.6	441	-38	137	330–540	11.7	-7.6	2

CB107-1A	16.6	405	29	35	290–330	18.1	43.1	3
CB108-1B	16.7	192	-27	485	370–460	6.3	-61.2	1
CB108-1A	16.7	193	-17	84	500–580	19.0	-55.3	2
CB109-1B	19.2	170	3	188	330–450	3.2	-46.2	1
CB109-1C	19.2	147	-62	261	250–430 ^a	18.0	-65.4	2
CB110-1A	19.9	178	-43	186	290–580	8.4	-73.3	1
CB110-1B	19.9	177	-30	639	250–450	6.5	-64.5	1
CB111-1A	20.3	389	27	301	250–610	9.1	52.7	1
CB111-1B	20.3	397	25	299	250–580	7.2	47.2	1

Abbreviations: MAD = maximum angular deviation; Q = Quality rank.

^a Directions not anchored to the origin.