

A new diverse charophyte flora and biozonation of the Eocene bauxite cover-sequence at Gánt (Vértes Hills, Hungary)

Journal:	Journal of Systematic Palaeontology
Manuscript ID	TJSP-2020-0103.R1
Manuscript Type:	Original Article
Keywords:	Characeae, Raskyellaceae, phylozone, gradualistic evolution, Paleogene, Central Europe



2		
3 4	1	A new diverse charophyte flora and biozonation of the Eocene bauxite
5 6 7	2	cover-sequence at Gánt (Vértes Hills, Hungary)
8 9 10 11	3	
12 13 14	4	Khaled Trabelsi ^{a,b,c*} , Benjamin Sames ^{c,d} , Michael Wagreich ^c , Miklós Kázmér ^e , Andrea
15 16 17	5	Mindszenty ^f , Carles Martín-Closas ^g
18 19	6	
20 21 22	7	^a Université de Sfax, Faculté des Sciences de Sfax, CP 3000, Sfax, Tunisie; ^b Université de
22 23 24	8	Tunis El Manar II, Faculté des Sciences de Tunis, LR18 ES07, C.P. 2092, Tunis, Tunisie;
25 26	9	^c Department of Geology, University of Vienna, UZA 2, Althanstrasse 14, 1090 Vienna,
27 28 20	10	Austria; ^d Sam Noble Museum, University of Oklahoma, 2401 Chautauqua Ave, Norman, OK
29 30 31	11	73072, USA; ^e Department of Palaeontology & MTA-ELTE Geological, Geophysical and
32 33	12	Space Science Research Group, Eötvös Loránd University, Pázmány Péter sétány 1/C, 1117
34 35 26	13	Budapest, Hungary; ^f Department of Physical and Applied Geology, Eötvös Loránd
30 37 38	14	University, Pázmány Péter sétány 1/C, 1117 Budapest, Hungary; ^g Departament de Dinàmica
39 40	15	de la Terra i de l'Oceà, Institut de Recerca de la Biodiversitat (IRBio), Facultat de Ciències
41 42 42	16	de la Terra, Universitat de Barcelona-UB, 08028 Barcelona, Catalonia, Spain.
45 44 45	17	
46 47 48	18	
49 50 51 52	19	
53 54 55	20	
57 58 59 60	21	* Corresponding author, e-mail: <u>trabkhalfss@yahoo.fr (Khaled Trabelsi</u>)

A diverse Eocene charophyte flora from a section at Gánt (Vértes Hills), Transdanubian Central Range, north-western Hungary, provides significant new information to previous studies only based on subsurface data published from the mid-20th century. This newly acquired material facilitates the taxonomic revision and emendation of the species *Raskyella peckii* and thereby defines a new anagenetic lineage based on three successive varieties which were formerly considered as separate species or subspecies: *Raskvella peckii* var. *peckii* (early Lutetian–early Bartonian), Raskyella peckii var. caliciformis (early Bartonian), and Raskyella peckii var. *vadaszii* (late Bartonian). Based on this lineage, we propose a new local charophyte biozonation that consists of a 'Raskyella peckii Superzone' (Lutetian-Bartonian), subdivided into three successive charophyte partial range zones: The 'Raskyella peckii peckii Zone' (Lutetian-lowermost Bartonian) characterized by an assemblage of R. peckii peckii, Gyrogona caelata forma caelata, G. caelata forma monolifera and Nitellopsis (Tectochara) aff. palaeohungarica, the 'Raskyella peckii caliciformis Zone' (lower Bartonian) characterized by the assemblage of R. peckii var. caliciformis, G. caelata forma caelata, G. caelata forma monolifera, G. caelata forma baccata, Nitellopsis (Tectochara) aff. palaeohungarica and Chara media, and the 'Raskyella peckii vadaszii Zone' (upper Bartonian) characterized by *R. peckii* var. vadaszii, *G.* caelata forma bicincta, G. caelata forma baccata, G. caelata forma fasciata, G. tuberosa, *Psilochara polita*, *Psilochara* sp., *Chara media* and *Chara subcylindrica*. Future research may show the new local biozonation as applicable to the whole of Europe and complementary to the current European charophyte biozonation. Our results show that the sequences from Gánt, which were previously regarded as upper mid-Eocene (upper Lutetian-lower Bartonian) in age, appear to represent a longer chronostratigraphic interval: lower Lutetian to upper Bartonian. Our chronostratigraphic results imply a longer and more stepwise Eocene major transgression in the Transdanubian Central Range than previously thought.

 Keywords: Characeae, Raskyellaceae, phylozone, gradualistic evolution, Paleogene, Central
Europe.

Introduction

Charophytes represent one of the most useful tools in the biostratigraphic analysis of Cenozoic non-marine deposits worldwide. During the Eocene, charophytes have been the object of significant taxonomic, biostratigraphic, palaeoecological and palaeobiogeographic interest, particularly in South European basins from France and Spain (Grambast 1958, 1962a, 1972; Feist-Castel 1970, 1972, 1975, 1977a; Feist & Ringeade 1977; Anadón & Feist 1981; Riveline 1986; Anadón et al. 1992; Sanjuan & Martín-Closas 2012). As a result, a European Charophyte Biozonation based largely on these basins was proposed by Riveline et al. (1996). For the Eocene, up to 11 charophyte biozones were defined based mainly on data from Western Europe. This biozonation has been updated since then, e.g., by Sanjuan et al. (2014) for the upper Eocene. In contrast, the Eocene charophyte flora from Central and Eastern Europe is relatively poorly known; and in the case of Hungary, the pioneer study by Rásky (1945) based on subsurface data is practically the only one available. This author described for the first time a species-rich charophyte flora from Hungary, at a time when charophyte taxonomy was still at an early stage. She had already assigned the flora studied in this area to the middle Eocene, and her work was the basis for future studies in charophyte taxonomy, including the definition of the new family Raskyellaceae by Grambast & Grambast (1954). Later, Bignot et al. (1985), based on an exhaustive palaeontological study of the Gánt section including molluscs,

foraminifers, ostracods, palynomorphs and charophytes, assigned the bauxite cover-sequence to the Upper Lutetian or Bartonian, respectively.

This study aims to update the compendium of knowledge on Eocene Hungarian and Central European charophytes, providing a taxonomic revision of the respective flora that is based on surface samples of sections at the Gánt locality, as well as a discussion regarding its biostratigraphic significance and utility.

Geological setting

Within the Transdanubian Central Range (TCR), several bauxite deposits that developed across the Cretaceous-Early Tertiary boundary interval are well known for their industrial use as sources of aluminium ore; and among these, the famous karst bauxite of the Vértes Hills from the Gánt locality (north-western Hungary) is a prominent example. Strata overlying the bauxite represent the sedimentary record of progressive subcrustal erosion along the East Alpine-West Carpathian forearc basin (Kázmér et al. 2003). The bauxite represents the base of the Eocene charophyte-bearing strata investigated in this study (Fig. 1). The Eocene succession sampled shows a remarkable lateral and vertical change of facies (Pálfalvi et al. 2006; Pálfalvi 2007) that has been attributed to tectonic forces acting on the sedimentary body (Fodor 2007). The development of the post-bauxite deposits was lain down during oscillation of the groundwater table and eustatic sea level variations (Carannante et al. 1994; Mindszenty 2010), occurring before the region was flooded by a marine incursion during the late Bartonian (Bignot et al. 1985).

At the Gánt section, the bauxite cover-sequence shows five stratigraphic units in a vertical orientation dating from the middle Eocene (Fig. 2), called 'Packets' in the sense of Bignot et al. (1985). 'Packet 1', about 1.5 m thick, corresponds to the bauxite itself, which unconformably overlies Triassic dolomites. 'Packet 2', 1.5-2 m in thickness, forms the 'blue-hole' freshwater limestone facies (Carannante et al. 1994; Pálfalvi 2007) alternating with clays, rich in charophytes, ostracods and gastropods. 'Packet 3', ca. 6.5 m thick, includes alternating sandy clay, coal and fresh- to brackish water limestone, rich in charophytes, ostracods, molluscs and large benthic foraminifera (Bignot et al. 1985). 'Packet 4' is ca. 12 m thick and is mainly dominated by shallow marine limestone, rich in nummulites, miliolids, molluscs and ostracods. Finally, 'Packet 5' is *ca*. 6m thick and displays an alternation of shallow marine marl and limestone, rich in Nummulites and Orbitolites. 'Packets' 2 and 3 were sampled for charophytes and are studied here.

-----Figures 1, 2 near here-----

Material and methods

Intensive sampling for charophytes during two consecutive field work sessions in 2018 and 2019 was carried out on the cover sequence of the bauxite at Gánt (Vértes Hills, Hungary). Moderately- to well-preserved gyrogonites were recovered from marly limestone to hard limestone using acetolysis. This method, first applied by Nötzold (1965) to the study of charophytes, has been recently improved by Trabelsi et al. (2010, 2016) and shown to be very effective in recovering charophyte fructifications and thalli from consolidated carbonate rocks. It consists of soaking the sample of hard calcareous rock, perfectly dried and mechanically comminuted into fragments of about 1–3 mm across, in equal amounts of anhydrous acetic acid

and anhydrous copper sulphate (acid reacts exothermically). After neutralization by ammonia, the residue is treated with ultra-sound, then washed and rinsed. Gyrogonites were measured using the software Motic Images Plus 2.0 ML with a Motic BA310 stereomicroscope in the Departament de Dinàmica de la Terra i de l'Oceà (University of Barcelona, Catalonia, Spain). Scanning electron microscopy on gold-sputtered selected specimens was conducted with a JEOL JSM-6400 at the Faculty of Earth Sciences, Geography and Astronomy, University of Vienna (Austria) and with a Quanta 200 device at the Centres Científics i Tecnològics of the University of Barcelona (CCiTUB). The studied materials are housed in the Hungarian Natural History Museum (Budapest, Hungary), Botanical Department, Palaeobotanical Collection. The figured specimens are deposited under the inventory numbers: HNHM-PBO 1501–1591.

127 Systematic palaeontology

The charophyte flora from the bauxite cover sequence at Gánt (Vértes Hills, Hungary) studied here yields gyrogonites from two families: Raskyellaceae and Characeae. The different charophyte species described below are stratigraphically distributed in the section as shown in the Fig. 2.

Revie

 43
 132

 44
 133

 45
 133

 46
 133

 47
 133

 49
 134

 50
 134

 51
 135

 52
 135

 53
 135

 54
 55

 55
 136

 56
 136

 59
 137

 59
 137

 50
 137

1 2 3 4	138	Genus <i>Raskyella</i> (Grambast et Grambast, 1954) emend. Grambast, 1962b
5 6 7	139	
8 9 10 11	140	Type species. Raskyella peckii Grambast et Grambast, 1954
12 13 14	141	
15 16 17	142	Remarks. This species is understood as including several traditional taxa belonging to the
18 19 20 21	143	genus Raskyella, which form a gradualistic lineage during the Eocene. These traditional taxa
	144	have been newly combined here to anagenetic varieties within a single evolutionary lineage or
22 23 24	145	an evolutionary species, following the recommendations of Wiley (1981) and Ax (1978).
25 26 27	146	
28 29 30	147	Raskyella peckii var. peckii Grambast et Grambast, 1954
31 32 33	148	(Fig. 3A–H)
34 35 36	149	
37 38 39	150	1954 Raskyella pecki sp. nov. L. & N. Grambast: p. 670, text-figs 1a-c.
40 41 42 43	151	1957 Raskyella pecki Grambast: p. 358, pl. 5, figs 7–9.
44 45 46	152	1958 Raskyella pecki Grambast: p. 190, figs 87, a-c; p. 191, text-fig. 88.
47 48 49	153	1959 Raskyella pecki Horn af Rantzien: pl. 19, figs 7–13.
50 51 52	154	1971 Raskyella peckii subsp. ganesensis Soulié-Märsche: pl. 2, 1–5.
53 54 55	155	1981 Raskyella pecki Anadón & Feist: pl. 1, figs 1–2; pl. 2, figs 3–4.
56 57 58	156	1986 Raskyella pecki Riveline: pl. 37, figs 7–9.
59 60	157	1999a <i>Raskyella pecki</i> Martín-Closas <i>et al.</i> : p. 11, figs 6, 1–3.
		/

Material. Up to 65 gyrogonites in sample G-2.4, and dozens in samples G-2.2 and G-2.3.
Collection numbers of figured specimens: HNHM-PBO 1501–1508.

Description. Gyrogonites are ovoidal to ellipsoidal in shape, spherical to subprolate (ISI 100-120) and of large size, 800–1050 µm in height and 750–1050 µm in width, showing laterally 7-10 (usually 9) convolutions (Fig. 4). Spiral cells often flat (Fig. 3A) to slightly convex (Fig. 3C), or concave (Fig. 3B), but regularly without any kind of ornamentation. Apex broadly rounded, truncated and flattened, with the spiral cells abruptly discontinue and ending acutely in the apical periphery to be replaced by the development of five opercular cells, each obliquely disposed at the end of a spiral cell (Fig. 3F). Germinated specimens (Fig. 3G) show a rose-shaped apical opening. Internal casts of these gyrogonites were also found (Fig. 3E).

169 Remarks. The contemporaneous unornamented gyrogonites of *Raskyella peckii* subsp. 170 ganesensis Soulié-Märsche, 1971 from the Aquitaine basin (France) appear to represent a 171 gyrogonite population of relatively smaller size within *R. peckii* var. *peckii* and both are here 172 considered synonymous. However, supplementary research on the type material is needed to 173 verify this synonymy, since the morphotype ganesesis is only known from the type locality.

Additionally, the subspecies *Raskyella peckii* subsp. *meridionale* Grambast, 1960, is kept
within the rank of subspecies due to its palaeogeographic restriction. The extremely large
gyrogonites of this subspecies are limited to the southernmost biogeographic range of *Raskyella peckii*, i.e. Algeria (Grambast 1960; Mebrouk *et al.* 1997), and the Betic Domain of the Balearic
Islands (Martín-Closas & Ramos 2005).

Distribution. This is the first record of *R. peckii* var. *peckii* in Hungary and Central Europe.
 This variety is widely distributed in the Lutetian and lower Bartonian of southern Europe,
 mainly in France (L. & N. Grambast 1954; Grambast 1958; Soulié-Märsche 1971, 1974;

182	Riveline 1984, 1986) and Spain (Anadón & Feist 1981; Anadón et al. 1992; Martín-Closas et
183	al. 1999a; Martín-Closas & Ramos 2005). The total range of this variety (early Lutetian-late
184	Bartonian) has been characterized in the Eastern Ebro basin (Catalonia) by Martín-Closas et al.
185	(1999) based on correlation with larger foraminifera (mainly Nummulites). Furthermore, R.
186	peckii var. peckii has been also reported from North Africa, i.e. in the lower Eocene of Algeria
187	(Gevin et al. 1974; Mebrouk et al. 1997; Vianey-Liaud 1994), in the late lower Eocene-early
188	middle Eocene of Tunisia (Abdeljaoued et al. 1984) and in the Lutetian of Libya (Megerisi &
189	Mamgain 1980).
190	
191	Figures 3, 4 near here
192	
193	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín-
193 194	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín- Closas
193 194 195	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín- Closas (Fig. 3I–P)
193 194 195 196	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín- Closas (Fig. 3I–P)
193 194 195 196 197	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín- Closas (Fig. 3I–P) Basionym. Raskyella caliciformis Soulié-Märsche, 1974, Compte Rendu 96 ^{ème} Congrès
193 194 195 196 197 198	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín-Closas (Fig. 3I–P) Basionym. Raskyella caliciformis Soulié-Märsche, 1974, Compte Rendu 96 ^{ème} Congrès National des Sociétés Savantes, Toulouse, 1971, Section Science, 2, p. 114, text-figure 2 (pl.
193 194 195 196 197 198 199	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín- Closas (Fig. 3I–P) Basionym. Raskyella caliciformis Soulié-Märsche, 1974, Compte Rendu 96 ^{ème} Congrès National des Sociétés Savantes, Toulouse, 1971, Section Science, 2, p. 114, text-figure 2 (pl. I), 1–5.
193 194 195 196 197 198 199 200	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín- Closas (Fig. 3I–P) Basionym. Raskyella caliciformis Soulié-Märsche, 1974, Compte Rendu 96 ^{ème} Congrès National des Sociétés Savantes, Toulouse, 1971, Section Science, 2, p. 114, text-figure 2 (pl. I), 1–5.
193 194 195 196 197 198 199 200 201	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi et Martín-Closas (Fig. 3I–P) Basionym. Raskyella caliciformis Soulié-Märsche, 1974, Compte Rendu 96 ^{eme} Congrès National des Sociétés Savantes, Toulouse, 1971, Section Science, 2, p. 114, text-figure 2 (pl. I), 1–5. 1974 Raskyella caliciformis sp. nov., Soulié-Märsche: p. 112, pl. 1, figs 1–5.

Material. 63 gyrogonites in sample G-2.5. Collection numbers of figured specimens: HNHM-PBO 1509–1516.

Description. Large sized gyrogonites (650–1000 µm high and 750–1050 µm wide) of globular to oblate shape (ISI 80–105), showing laterally 6–9 (usually 8) convolutions (Fig. 5). Spiral cells flat to slightly concave ornamented with stout, vertical to slightly inclined, well individualized tubercles (Fig. 3I-K), which are the main diagnostic character of this variety. Apex broadly rounded to truncated showing five opercular cells placed at the end of the spiral cells. The opercular cells are sometimes convex and somewhat polygonal in shape (Fig. 3N), rather than rounded, which is the reason why the dehiscence opening appears sometimes irregularly star-shaped (Fig. 3O), rather than rose-shaped (Fig. 3L), as already noted by Soulié-Märsche (1974) in the type material. The internal cast of the gyrogonite (Fig. 3M) shows straight ridges perpendicular to the well-marked, undulated spiral cells. This wavy surface is uncommon in the inside of other raskyellacean gyrogonites and is thought to correspond internally to the external tubercle ornamentation.

Distribution. This is the first record of *R. peckii* var. *caliciformis* in Hungary. It was previously
described from the Bartonian of South France (Soulié-Märsche 1974; Riveline 1986), and from
the lower Bartonian (Auversian local stage) of the Ebro Basin, Catalonia, Spain (Anadón &
Feist 1981; Anadón *et al.* 1992).

----- Figure 5 near here------

2		
3 4	225	Raskyella peckii var. vadaszii (Grambast et Grambast 1954) comb. nov. Trabelsi et Martín-
5 6	226	Closas
7		
8 9	227	(Fig. 6A–S)
10		
11		
12	228	
13		
14 15	220	Basionym <i>Bashvalla vadaszi</i> (Báshv) I & N. Gramhast (1954) Bayue Gánárale de
15 16	229	Dasionyin. <i>Ruskyetta vaduszi</i> (Rasky) L. & N. Otambast (1994), Revue Genetate de
17	220	Botanique (61) n 670
18	250	Botanique (01), p. 070.
19		
20	231	
21		
22 73		
23 24	232	1945 Aclistochara vadaszi, sp. nov. Rásky: p. 45, pl. II, figs 22–24.
25		
26	• • • •	
27	233	1954 Raskyella vadaszi, comb. nov. L. and N. Grambast: p. 670.
28		
29	231	1957 Rashvella vadaszi, Gramhast: n. 358 nl. 5 figs 1-6
30 31	234	1957 Ruskyellu vuluszi, Grunoust. p. 556, pl. 5, figs 1 °C.
32		
33	235	1959 Raskyella vadaszi, Horn af Rantzien: pl. 20, figs 1–3.
34		
35		
36	236	1981 <i>Raskyella vadaszi</i> , Anadón & Feist: pl. 1, fig. 5; pl. 2, fig. 5.
3/ 20		
39	227	1981 Rashvalla aff vadaszi Anadón & Feist: nl. 1. figs 3-1: nl. 2. figs 1-2. 6
40	257	1961 <i>Ruskyellu</i> all. <i>valuuszi, Miladoli & Felst. pl. 1, ligs 5 4, pl. 2, ligs 1 2, 0.</i>
41		
42	238	1985 Raskyella vadaszi, Bignot et al.: p. 36, pl. 3, figs 8–11.
43		
44 15		
45 46	239	1986 Raskyella vadaszi, Riveline: pl. 37, figs 1–6.
47		
48	240	
49	240	
50		
51 52	241	Material. Hundreds of gyrogonites in samples G-6a and G-6b. Collection numbers of figured
52 53		
55 54	242	specimens: HNHM-PBO 1517–1534
55		1
56		
57	243	Description. Large sized gyrogonites (800–1150 µm high and 800–1150 µm wide) of oblate to
58 50		
59 60	244	ovoidal shape (ISI 80-120), showing laterally 7-10 (usually 9) convolutions (Fig. 7). Spiral

cells often convex and ornamented with stout tubercles of different shapes and sizes, in most cases oriented parallel to the intercellular sutures (Fig. 6A–C), or more rarely tilted 20–30° but keeping parallelism between adjacent nodules (Fig. 6E–G), this being a diagnostic character of this morphotype. Three tubercle morphologies have been observed: (1) rounded tubercles, well individualized in the upper half of the gyrogonite, but fused to neighbouring tubercles in the lower half (Fig. 6I–K), (2) elongated tubercles more or less connected to each other and producing slightly wavy (undulated) sutures (Fig. 6H), (3) irregularly alternating round and elongated tubercles (Fig. 6D). Base of gyrogonite rounded (Fig. 6C, J) to slightly tapered (Fig. 6A, G) and showing a small, superficial and pentagonal basal pore, sometimes within a less-marked funnel (Fig. 6Q). Apex of gyrogonite truncated or broadly rounded and covered by five independent opercular cells at the end of the spiral cells. Opercular cells roughly prismatic, with its outer surface concave, flat or slightly convex (Fig. 6L-M). Germinated specimens show a rounded or rose-like opening (Fig. 6N–P). The inside of the gyrogonite allows observation of a characteristic crenulation of the intercellular sutures near their internal side (Fig. 6R–S), while, to the outside, sutures are flat. Besides, this crenulation occurs also between the opercule cells themselves and between spiral and opercule cells as already described by Feist in Anadón and Feist (1981).

Distribution. R. vadaszii has been first described by Rásky (1945) from middle Eocene borehole samples (60 m depth) at Gánt, in beds roughly equivalent laterally to the outcrop succession studied here. Therefore, the samples studied may be considered topotypes. Subsequently, the age of this variety was suggested to be upper Bartonian by Bignot et al. (1985) based on the associated microfossils (foraminifers, ostracods, and pollen) from the same beds of the bauxite cover-sequence at the Gánt section (Vértes Hills). R. peckii var. vadaszii has also been well documented in France, in the upper Bartonian of the Paris Basin (Grambast 1957, 1958, 1962a; Riveline 1986) and from several basins in southern France (Feist-Castel

1		
2 3 4	270	1976). Anadón & Feist (1981) and Anadón et al. (1992) documented this variety also in the
5 6 7	271	upper Bartonian of the Eastern Ebro Basin (Catalonia, Spain).
8 9 10	272	
11 12 13	273	Figures 6, 7 near here
14 15 16	274	
17 18 19 20	275	Family Characeae (Richard ex C.A. Agardh, 1824) emend. Martín-Closas et Schudack, 1991
20 21 22 23	276	Subfamily Charoideae Braun in Migula, 1897
23 24 25 26	277	Genus Gyrogona (Lamarck, 1804 ex Lamarck, 1822) emend. Grambast, 1956
27 28 29	278	
30 31 32	279	Gyrogona caelata (Reid et Groves, 1921) Grambast, 1956
33 34 35	280	(Fig. 8A–V)
36 37 38	281	
39 40 41	282	1921 Chara caelata sp. nov., Reid & Groves: p. 184, pl. 4, figs 4-6.
42 43 44 45	283	1927 Kosmogyra caelata, Pia: p. 90.
45 46 47 48	284	1954 Brachychara caelata, L. & N. Grambast: p. 667.
49 50 51	285	1956 Gyrogona caelata, Grambast: p. 280.
52 53 54	286	1977b Gyrogona caelata, Feist-Castel: p. 117.
55 56 57	287	1981 Gyrogona caelata, Grambast & Grambast-Fessard: p. 22, text-fig. 11, a-f; pl. 4, figs 1-9.
58 59 60	288	1981 Gyrogona cf. Caelata, Anadón & Feist: p. 163.

289 1986 *Gyrogona caelata*, Riveline: pl. 38, figs 1–5, 7–8.

290 1989 *Gyrogona caelata*, Choi: pl. 2, figs 1–11.

1991 Gyrogona caelata, Weidmann et al.: p. 900, fig. 3, C.

292 2014 Gyrogona caelata, Sanjuan & Martín-Closas: p. 403, fig. 7, A–C.

Material. 56 gyrogonites in sample G-2.2, 38 in sample G-2.3, 29 in sample G-2.4, 18 in sample
G-2.5, and 23 in sample G-6a. Collection numbers of figured specimens: HNHM-PBO 1535–
1553.

Description. Medium to large gyrogonites, 600–800 µm high and 700–1000 µm wide with generally oblate to suboblate spheroidal shape (ISI 80–100) and showing laterally 5–7 (usually 6) convolutions (Fig. 9). Apex and base broadly rounded to subtruncate. Apex showing a less-marked spiral cell periapical thinning (e.g., Fig. 8G) and, in some specimens, apical nodules of different shape, generally flat or slightly convex (e.g. Fig. 8S). Base showing a small pentagonal basal pore (Fig. 8F), sometimes flared by a shallow funnel (Fig. 8P). Basal plate unicellular and only visible from the gyrogonite interior (Fig. 8U–V). Spiral cells flat to slightly concave and ornamented with different patterns of tubercles, which allowed Grambast (1958) and Grambast & Grambast-Fessard (1981) to distinguish a number of morphotypes, ranking them as *formae* of the same species. The following five forms were recognized in the material studied: (1) G. caelata forma caelata characterized by small nodules well-spaced and irregularly ranged along the spiral cell median line (e.g. Fig. 8A, D), (2) G. caelata forma bicincta characterized by nodules irregularly ranged along two lines parallel to spiral cell sutures (Fig. 8H, I), (3) G. caelata forma monolifera showing medium-sized nodules close to each other, sometimes fused and forming a thin, irregular mid-cellular crest (e.g., Fig. 8L), (4) G. caelata forma baccata

3 4	312
5 6	313
7 8	314
9 10 11	315
12 13	316
14 15	217
16 17	210
18 19 20	510
20 21 22	319
22 23 24	320
25 26	321
27 28	322
29 30	323
31 32	324
33 34 35	325
36 37	
38 39	326
40 41 42	327
43 44	328
45 46	
47 48 40	329
49 50 51	330
52 53	221
54 55	551
56 57 58	332
59	

characterized by large nodules very closely ranged along the spiral-cell median line (e.g., Fig. 312 8N–Q), (5) G. caelata forma fasciata characterized by a broad median band of variable width 313 (e.g., Fig. 8R–T). 314

Distribution. The species *Gyrogona caelata* is reported here from Hungary for the first time. 315 According to Riveline (1986), this species was widely distributed in the upper Lutetian-316 Priabonian non-marine deposits of Western Europe. It was first recorded from the Isle of Wight, 317 England by Reid & Groves (1921). Thereafter, it was reported from the upper Lutetian to upper 318 319 Priabonian of France (Grambast 1958; Grambast & Grambast-Fessard 1981; Feist-Castel 1971; Feist & Ringeade 1977; Feist-Castel 1977a, b; Ollivier-Pierre et al. 1988), Spain (Anadón & 320 Feist 1981; Choi 1989; Anadón et al. 1992; Sanjuan & Martín-Closas 2014), Switzerland 321 (Weidmann et al. 1991), as well as from the middle Eocene of Romania (Iva 1987). In North 322 Africa, the species has also been documented from the central part of the Sahara, Algeria, by 323 Mebrouk et al. (1997). 324

----- Figures 8, 9, near here-----

328 Gyrogona tuberosa (Reid et Groves, 1921) Grambast in Grambast et Grambast-Fessard, 1981

- (Fig. 10A–J)
- 331 1921 Chara wrighti var. rhytidocarpa, Reid & Groves: p. 183, pl. 4, fig. 3.
- 1958 Gyrogona tuberosa, Grambast: p. 139, fig. 54. 332
- 1976 Gyrogona tuberosa, Feist-Castel: p. 26. 333 60

Gyrogona tuberosa, Grambast & Grambast-Fessard: p. 25, text-fig. 12, a–d; pl. 5, figs 1–
6.

1986 Gyrogona tuberosa, Riveline: pl. 14, figs 8–11.

Material. 35 gyrogonites in sample G-6b. Collection numbers of figured specimens: HNHM-PBO 1554–1563.

 Description. Large gyrogonites, 820–1000 μm high and 835–1050 μm wide with generally oblate to suboblate spheroidal shape (ISI 80–100) and showing laterally 6–7 convolutions (usually 6). Apex subtruncate (Fig. 10A, C) to somewhat pointed (Fig. 10D, G) with spiral cells protruding (Fig. 10H). Base broadly rounded (Fig. 10B, F) to slightly tapered (Fig. 10A, C, E) and showing a small, superficial and pentagonal basal pore, sometimes within a less-marked funnel (Fig. 10I–J). Spiral cells flat or concave, smooth or somewhat ornamented and separated by protruding narrow to weakly undulated intercellular ridges.

Distribution. *Gyrogona tuberosa* is reported here from Hungary and central Europe for the first time. According to Riveline (1986) and Riveline & Cavelier (1987), this species was widely distributed in upper Bartonian non-marine deposits of Western Europe. It was first recorded from the Isle of Wight and Hampshire (England) by Reid & Groves (1921). Subesequently, it was reported from the upper Bartonian of the Paris Basin (Grambast 1958; Grambast & Grambast-Fessard 1981; Riveline 1986; Riveline & Cavelier 1987), as well as from several basins of southern France (Feist-Castel 1976).

57 355

1		
2 3	356	Figure 10 near here
4 5		
6 7 8	357	
9 10 11	358	Genus Psilochara Grambast, 1959
12 13 14	359	
15 16 17	360	Psilochara polita (Reid et Groves, 1921) Grambast, 1959
18 19 20	361	(Fig. 11A–F)
21 22 23	362	
24 25 26	363	1921 Chara polita, sp. nov. Reid & Groves: p. 187, pl. 5, figs 9, 12.
27 28 29	364	1927 Gyrogona politus, Pia: p. 90.
30 31 32 33	365	1958 Ovochara polita, comb. nov. Grambast: p. 167.
34 35 36	366	1959 Peckichara polita, Horn af Rantzien: p. 116, pl. 13, figs 1–3.
37 38 39	367	1959 Psilochara polita, Grambast: p. 11.
40 41 42	368	1977b Psilochara polita, Feist-Castel: p. 153.
43 44 45	369	1986 Psilochara polita, Riveline: p. 59, pl. 22, figs 8-12.
46 47 48	370	
49 50 51	371	Material. 58 gyrogonites in sample G-6a. Collection numbers of figured specimens: HNHM-
52 53 54	372	PBO 1564–1569.
55 56 57	373	Description. Medium-sized gyrogonites (550–700 μ m in height 500–650 and μ m in width)
58 59	374	with ovoidal shape (ISI 100-120) and laterally showing 7-10 convolutions (Fig. 11A-C; Fig.
60	375	12). Apex round to pointed (Fig. 11A–C). Base tapering to prolonged into a stout basal column

376 (Fig 11A-C) and showing a small pentagonal basal pore (Fig. 11F). Spiral cells smooth,

377 concave or flat and separated by protruding narrow to weakly undulated intercellular ridges.

Distribution. *Psilochara polita* is described here from the middle Eocene of Hungary for the first time. It has previously been described from the upper Bartonian of the Isle of Wight, England, (Reid & Groves 1921; Feist-Castel 1977b; Riveline 1986) and of the Paris Basin (Grambast 1958; Riveline 1986).

----- Figures 11, 12 near here------Psilochara sp. (Fig. 11G–I) Material. 17 gyrogonites in sample G-6a. Collection numbers of figured specimens: HNHM-PBO 1570-1571. **Description.** Medium to large-sized gyrogonites (780–905 µm wide and 670–775 µm high) with elongated ovoidal (subprolate) shape (ISI 110-125) and laterally showing 8-10 convolutions (Fig. 11G). Apex truncated. Apical end of spiral cells enlarged and pointing upwards (Fig. 11H). Base truncated to somewhat tapering, bearing a small pentagonal basal pore. Basal plate unipartite and visible from the gyrogonite interior (Fig. 11I). Spiral cells concave or flat and smooth, except at the periapical area, where they are irregularly ornamented

1 2		
- 3 4	397	Remark. The low number of gyrogonites hinders a more precise taxonomic attribution of this
5 6	398	population. However, it is reported here since it differs in size and shape from the other species
7 8 9	399	of <i>Psilochara</i> found at Gánt.
10 11 12	400	
13 14 15 16	401	Genus <i>Nitellopsis</i> Hy, 1889
17 18 19	402	
20 21	403	Sub-genus Tectochara L. et N. Grambast, 1954
22 23 24 25	404	
26 27	405	Nitellopsis (Tectochara) aff. palaeohungarica (Rásky, 1945) Grambast et Soulié-Märsche,
28 29 30	406	1972
31 32 33	407	(Fig. 11J–N)
34 35 36	408	
37 38 39 40	409	1945 Chara palaeohungarica, sp. nov. Rásky: p. 38, pl. 1, figs 16–18.
41 42 43	410	1955 Tectochara palaeohungarica, comb. nov. Mädler: p. 298.
44 45 46	411	1959 Tectochara palaeohungarica, Horn af Rantzien: p. 90, pl. 8, figs 4-7.
47 48	412	1972 Nitellopsis (Tectochara) palaeohungarica, nov. comb. Grambast & Soulié-Märsche: p.
49 50 51	413	4.
52 53 54	414	
55 56 57	415	Material. 28 gyrogonites in sample G-2.3 and 33 in sample G-2.5. Collection numbers of
58 59 60	416	figured specimens: HNHM-PBO 1572–1576.

Description. Gyrogonites very large (900–1200 μm high and 800–1050 μm wide), oval, prolate
spheroidal (ISI 100–120) in shape, showing 8–11 (often 9) convolutions in lateral view (Fig.
11J–L; Fig. 13). Spiral cells concave to flat. Apex prominent with spiral cells protruding to
form a central rosette. Spiral cells show both narrowing and thinning in the periapical area (Fig.
11M). Base rounded to almost conical, occasionally lengthened in a short broad column. A
large (155–230 μm across) pentagonal basal pore occurs within a wide basal funnel (Fig. 11N).

Distribution. The species '*Chara' palaeohungarica* was first described from subsurface beds
attributed to the Paleocene in Dorog, Hungary, by Rásky (1945). Here this species is described
from beds cropping out at Gánt, which are probably time-equivalent to those of the type locality.
The present study supports reassignment of this species to the middle Eocene rather than to the
Paleocene.

Remarks. The gyrogonites studied here are diagenetically deformed, which hinders a more definitive taxonomic attribution. A re-study of the type material (HNHM 55.1458–55.1460) by one of the authors of this study (CMC) showed that, besides the holotype, which is a subspherical gyrogonite as illustrated by Rásky (1945, pl. I, Fig. 16), there were more oval to elongated gyrogonites present in the collection similar to those described here, which Rásky (1945, p. 38) termed 'cylindrical'.

Genus Chara Vaillant, 1719

----- Figure13 near here-----

1		
2		
4	439	Chara media Grambast, 1958
5		
6	440	(Fig. 14A–I)
7 8		
9		
10	441	
11		
12	442	1958 Chara media, Grambast: p. 178, fig. 81b.
14		
15	112	1086 Chara madia Riveline: n 68 nl 20 figs 6 12
16 17	445	1980. Chara meaia, Rivenne. p. 68, pl. 29, figs 0–12.
18		
19	444	
20		
21 22	115	Material Up to 80 gyrogonites in both samples G-2.5 and G-6a. Collection numbers of figured
23	445	Waterial. Op to ob gylogonites in both samples of 2.5 and of ou. Concetion numbers of figured
24	446	specimens: HNHM-PBO 1577–1585.
25 26		
20 27		
28	447	Description. Gyrogonites of medium size (400–650 µm high and 300–500 µm wide) ellipsoidal
29	118	subprolate (ISI 110-145) laterally showing $8-11$ (usually $9-10$) convolutions (Fig. 14A-G
30 31	440	subprotate (151 110 145), faterally showing 6 11 (usually 5 10) convolutions (11g. 1471 O,
32	449	Fig. 15). Maximum width nearly at the half to $2/3$ of height. Apex rounded to slightly conical.
33		
34 35	450	with distinctly widening of the spiral cell endings (Fig. 14H). Spiral cells concave, smooth and
36		
37	451	without any periapical modification. Base tapering showing a superficial pentagonal basal pore
38	452	(Fig. 14I)
39 40	452	(F1g. 141).
41		
42	453	Distribution. Chara media is described here from the upper Eocene (upper Bartonian) of
43 44		
45	454	Hungary for the first time. Grambast (1958) and Riveline (1986) documented this species from
46	455	the unper Partenian lower Oligogone of several basing in France Palgium and Cormany
47 49	455	the upper Bartoman-lower Ongocene of several basins in France, Bergium, and Germany.
40 49		
50	456	
51		
52 53	/157	Figures 14, 15 near here
55 54	457	rigues 14, 15 neur noie
55		
56 57	458	
57 58		
59	459	Chara subcvlindrica Reid et Groves. 1921
60		•

URL: http://mc.manuscriptcentral.com/tjsp

1 ว		
2 3	460	(Fig 14J–O)
4 5		
6	461	
7	401	
8 9	460	
10	462	1921 Chara subcyunarica, sp. nov. Reid & Groves: p. 187, pl. 5, fig. 4–5.
11 12		
13	463	1959 Grambastichara subcylindrica, Horn af Rantzien: p. 76, pl. 3, figs 5–7.
14 15		
16	464	1986 Chara cf. subcylindrica, Riveline: p. 67, pl. 30, figs 5-8.
17 18		
19	465	
20		
21	466	Material. Up to 250 gyrogonites in sample G-6b. Collection numbers of figured specimens:
23		
24 25	467	НNHM-PBO 1586–1591.
26		
27 28	468	Description. Medium-sized gyrogonites (500–750 µm high and 200–400 µm wide) ellipsoidal
29	460	
30 31	469	prolate to perprolate (ISI 130–200) in shape, laterally showing 8–11 (usually 9–10)
32	470	convolutions (Fig. 14J-M; Fig. 16). Maximum width at the equator. Apex rounded with
33 34		4
35	471	widening of the spiral cell endings (Fig. 14N). Spiral cells often slightly concave to flat,
36 27	472	separated by narrow intercellular ridges, cells non-ornamented and without any periapical
37 38		
39	473	modification. Base regularly tapering to round, showing a superficial pentagonal basal pore
40 41	171	(Fig. 140)
42	4/4	(Fig. 140).
43 44		
45	475	Distribution. This is the first report of <i>Chara subcylindrica</i> in Hungary. According to Reid &
46 47	476	Groves (1921) and Riveline (1986), this species occurs in the upper Bartonian-lower Oligocene
48		
49 50	477	of England, France, Belgium and Germany.
50 51		
52	478	
53 54		
55	479	Figure 16 near here
56 57		-
58	480	
59 60		
50		

Discussion

483 Definition of the *Raskyella peckii* anagenetic lineage

Evolutionary lineages formed by a succession of charophyte fructifications changing gradually in time were first described in the family Clavatoraceae by Grambast (1974). Later, similar lineages were found as well in the family Characeae (e.g. lineage Harrisichara vasiformis-tuberculata described by Feist-Castel 1977b; or lineage Peckichara pectinata by Vicente et al. 2018). Here we describe the first of such lineages in the family Raskyellaceae. In the Lutetian and Bartonian of Gánt (Hungary), three former species of the genus Raskyella - R. peckii, R. *caliciformis*, and *R. vadaszii* – have been found to form a continuous succession of gyrogonite morphologies, connected by intermediate morphotypes. This gradualistic lineage is interpreted as an evolutionary species in the sense of Wiley (1981) and Ax (1987), and the original taxa have been newly combined as an genetic varieties of the species with nomenclatural priority, which is *R. peckii*.

The first evolutionary stage of the *R. peckii* lineage (Fig. 17) is represented by *R. peckii* var. *peckii*, and includes, as well, the smallest gyrogonite morphotype initially described as Raskyella peckii ganesensis Soulié-Märsche, 1971. R. peckii var. peckii is characterized by gyrogonites which are very variable in size, but consistently unornamented. This stage has a long duration, since it was documented from the lower Lutetian to the upper Bartonian of the Ebro Basin, Catalonia, by Martín-Closas et al. (1999a) and can thus be superimposed onto some of the ulterior morphotypes of the lineage, this being quite a common situation in charophyte lineages (e.g., Grambast, 1974). The coeval Raskvella peckii subsp. meridionale Grambast, 1960, was not found in the section studied and corresponds to a southern geographic subspecies

of this lineage, thriving in North Africa and the Prebetic Domain in the Balearic Islands in Spain
(Grambast 1960; Martín-Closas & Ramos 2005).

The second evolutionary stage in the lineage of R. peckii is represented by R. peckii var. caliciformis. Intermediate morphotypes between R. peckii var. peckii and R. peckii var. *caliciformis* display a progressive increase in the gyrogonite size (up to 1000 µm in height), and a change in shape from elongated to rounded, between samples G-2.2 and G-2.4 of the Gánt section. Furthermore, there is a progressive development of the ornamentation corresponding to R. peckii var. caliciformis in the same sequence, with for instance 100% of gyrogonites corresponding to *R. peckii* var. *peckii* in sample G-2.4, while in sample G-2.5, there is only 10% of R. peckii var. peckii, resulting in 90% of R. peckii var. caliciformis (Fig. 17).

The third stage of the lineage is represented by *R. vadaszii* from the upper Bartonian, which shows an additional increase of the gyrogonite size of about 150–250 µm in height and 100– 125 µm in width, and especially the development of progressively more complex ornamentation patterns in comparison to the previous evolutionary step (Fig. 17). This gradual change can be observed between samples G-2.5 and G-6b of the Gánt section. Thus, sample G-5 displays a homogeneous population with 100% of gyrogonites corresponding to R. peckii var. caliciformis, while in the overlying sample G-6a, the gyrogonite population of R. peckii contains only 15% of specimens of R. peckii var. caliciformis and 85% of R. peckii var. vadaszii. Finally, in sample G-6b there is a homogeneous population of gyrogonites corresponding to R. peckii var. vadaszii.

524 Overall, the *R. peckii* lineage follows the general evolutionary trend in the evolutionary lineages 525 of other charophyte families, characterized by an increase in size and sphericity (Clavatoraceae 526 and Raskyellaceae), and a progressive development of ornamentation (Characeae), as shown

2		
3 4	527	by Feist-Castel (1977b), Martín-Closas et al. (1999b), Sille et al. (2004) and Vicente & Martín-
5 6 7	528	Closas (2018).
8 9 10	529	
11 12 13	530	Figure 17 near here
14 15 16	531	
17 18 19	532	New local charophyte biozonation
20 21 22	533	Among the charophyte species described from the bauxite cover-sequence at the Gánt section
23 24	534	(Vértes Hills, Hungary), Raskyella peckii represents the most significant species for use in
25 26 27	535	biostratigrpahy within the non-marine Lutetian and Bartonian, as previously suggested by
28 29	536	Riveline et al. (1996) and Martín-Closas et al. (1999a). The Raskyella peckii biozone was
30 31	537	defined by Riveline et al. (1996) as a 'partial range zone comprising the interval from the first
32 33	538	appearance of Raskyella peckii L. and N. Grambast, 1954, to the first appearance of Chara
34 35 36	539	friteli Grambast, 1958, lower Lutetian to lower Bartonian in age. This study proposes to extend
37 38	540	this biozone to cover also the upper Bartonian, and to redesignate it as a superzone subdivided
39 40 41	541	into the following three successive biozones (Fig. 18):
42 43	542	- Raskyella peckii peckii Zone: partial range zone defined from the first occurrence of the
44 45 46	543	morphotype peckii to the first occurrence of the morphotype caliciformis, Lutetian-lower
47 48	544	Bartonian in age. The local charophyte assemblage characterizing this zone in Gánt occurs in
49 50	545	the basal part of the studied section ('Packet 2', 'blue-hole' freshwater limestone facies,
51 52	546	samples G-2.2, G-2.3 and G-2.4), and is composed of R. peckii var. peckii, G. caelata forma
55 54 55	547	caelata, G. caelata forma monolifera, and Nitellopsis (Tectochara) aff. palaeohungarica, some
56 57	548	of which are well known to occur in several European basins (Rásky 1945; Grambast 1958;
58 59 60	549	Riveline 1986; Martín-Closas et al. 1999a) during the same time interval.
-		

- Raskyella peckii caliciformis Zone: partial range zone defined from the first occurrence of the morphotype *caliciformis* to the first occurrence of the morphotype *vadaszii*, lower Bartonian in age. This zone includes in Gánt the assemblage found in the lower part of 'Packet 3' (samples G-2.5) and composed of *R. peckii* var. *caliciformis*, *G. caelata* forma *caelata*, *G. caelata* forma *monolifera*, *G. caelata* forma *baccata*, *Nitellopsis (Tectochara)* aff. *palaeohungarica*, and *Chara media*.

- Raskyella pecki vadaszii Zone: partial range zone defined from the first occurrence of the
morphotype *vadaszii* to the first occurrence of the next zone defined in the Paris Basin, which
is *Psilochara repanda*. This zone would be upper Bartonian in age. The assemblage occurring
in the middle part of the Gánt section ('Packet 3', samples G-6a and G-6b), composed of *R*. *peckii* var. *vadaszii*, *G. caelata* forma *bicincta*, *G. caelata* forma *baccata*, *G. caelata* forma *fasciata*, *Psilochara polita*, *Psilochara* sp., *Chara media*, and *Chara subcylindrica*characterizes locally this biozone.

564 Implications on the age of the bauxite cover-sequence

The biostratigraphic analysis carried out suggests a relative age of Lutetian– Bartonian (Fig. 18) of the bauxite cover-sequence at the Gánt section (Vértes Hills, Hungary), rather than constraining it to the Bartonian as previously suggested by Bignot (1985). A Lutetian age, deduced from the charophytes and attributed to the lower part of the studied series, has been already suggested in several works preceding that of Bignot (1985), notably those of Szőts (1938), Kopek (1980), and Dudich & Kopek (1982), on the basis of mollusc and palynomorph biostratigraphy. The data presented herein support the idea that the Eocene succession in the studied area reflects a stepwise marine transgression upon the bauxite deposits, beginning in the Lutetian. Our new chronostratigraphic framework sheds new light on the timing of the long-

lasting subaerial exposure and alteration process generating the bauxite strata. Consequently, coeval strata from surrounding localities within the Transdanubian Central Range should be restudied and analysed from the viewpoint of charophyte biostratigraphy, in order to correlate the post-bauxite depositional event(s) on a regional scale and to improve the understanding of its tectono-eustatic control.

- ----- Figure 18, near here-----to pe
- Conclusions

Eocene (Lutetian-Bartonian) charophyte assemblages are taxonomically described for the first time from an outcrop of the bauxite cover-sequence at Gánt (Vértes Hills), Hungary's Transdanubian Central Range. The sections show for the first time that the raskyellacean charophytes also evolved in gradualistic lineages, similarly to what is already known for other charophyte families. The *Raskvella peckii* lineage is formed by three successive stages and is interpreted here in terms of the anagenesis of the evolutionary species Raskyella peckii, including its gradual change to the morphotype previously known as *R. vadaszi*.

From a biostratigraphic viewpoint, the assemblages studied belong to the Raskyella peckii biozone of Martín-Closas et al. (1999a), which is here reinterpreted as a superzone extending to cover the Raskyella vadaszii Zone of Riveline et al. (1996) and attributed to the Lutetian-Bartonian interval. In this study, this superzone is subdivided into three successive local partial range biozones, defined by each of the successive varieties of the evolutionary species R. peckii: (1) the Raskyella peckii peckii partial range zone is characterized by R. peckii peckii, G. caelata

forma caelata, G. caelata forma monolifera and Nitellopsis (Tectochara) aff. palaeohungarica, Lutetian-lowermost Bartonian in age; (2) the *Raskvella peckii caliciformis* partial range zone is characterized by R. peckii caliciformis, G. caelata forma caelata, G. caelata forma monolifera, G. caelata forma baccata, Nitellopsis (Tectochara) aff. palaeohungarica and Chara media, lower Bartonian in age; and (3) the Raskyella peckii vadaszii partial range zone is characterized by *R. peckii vadaszii*, *G. caelata* forma *bicincta*, *G. caelata* forma *baccata*, *G.* caelata forma fasciata, G. tuberosa, Psilochara polita, Psilohara sp., Chara media and Chara subcylindrica, upper Bartonian in age.

In light of the new results presented here, the charophyte-bearing sequences studied in this work represent a longer time span than previously thought, running from the Lutetian to the Bartonian. This has direct implications on the understanding of the Eocene regional stratigraphic scheme of the Transdanubian Central Range, particularly in terms of synchronism/diachronism in the regional stratigraphic correlation, as well as the timing of the tectono-sedimentary control and palaeogeographic evolution.

612 Acknowledgements

613 This study is a contribution to UNESCO-IGCP 632 'Continental Crises of the Jurassic: Major 614 Extinction Events and Environmental Changes within Lacustrine Ecosystems', subproject 'Late 615 Mesozoic lacustrine systems in Tunisia and their global correlation' (BS), and UNESCO IGCP 616 Project 661 'The Critical Zone in Karst Systems', subproject: 'Evolution of fossil blue hole 617 limestones and the critical zone in a greenhouse world' (MW, KT) within the scope of the Earth 618 System Science (ESS) programme funded by the Austrian Academy of Sciences (BS), as well 619 as project CGL2015-69805-P from the Spanish Ministry of Innovation and Competitiveness

620	and to project SGR2017-841 of the AGAUR (Catalan Research Agency) (CM-C). It also
621	benefitted from Austrian Science Fund (FWF) project P 27687-N29 (BS), and from the results
622	of the SYNTHESYS project HU-TAF6533 accorded to CM-C to study the Klára Rásky
623	collection at the Hungarian Natural History Museum in Budapest. We sincerely acknowledge
624	Dr. Ingeborg Soulié-Märsche (Université de Montpellier II, France), Dr. Josep Sanjuan
625	(University of Barcelona, Spain), and Dr. Alba Vicente (Universidad Nacional Autónoma de
626	México-UNAM, México), as well as the editor Dr. Paul M. Barrett for their valuable and
627	constructive observations and criticism that greatly improved the manuscript during the peer-
628	review process. Kevin Kearney (University of Vienna, Austria) is acknowledged for English
629	language editing.
630	
631	References
632	
632 633	Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions
632 633 634	Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur «
632 633 634 635	Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–
632 633 634 635 636	Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur– Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i> , 22 , 73–77.
632 633 634 635 636 637	 Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i>, 22, 73–77. Agardh, C. A. 1824. <i>Systema Algarum</i>. Lundae Literis Berlingianis, Lundae, 312 pp.
632 633 634 635 636 637 638	 Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i>, 22, 73–77. Agardh, C. A. 1824. <i>Systema Algarum</i>. Lundae Literis Berlingianis, Lundae, 312 pp. Anadón, P. & Feist, M. 1981. Charophytes et biostratigraphie du Paléogène inférieur du bassin
632 633 634 635 636 637 638 639	 Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i>, 22, 73–77. Agardh, C. A. 1824. <i>Systema Algarum</i>. Lundae Literis Berlingianis, Lundae, 312 pp. Anadón, P. & Feist, M. 1981. Charophytes et biostratigraphie du Paléogène inférieur du bassin de l'Ebre oriental. <i>Palaeontographica</i>, 178 B(4–6), 143–168
632 633 634 635 636 637 638 639 640	 Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i>, 22, 73–77. Agardh, C. A. 1824. <i>Systema Algarum</i>. Lundae Literis Berlingianis, Lundae, 312 pp. Anadón, P. & Feist, M. 1981. Charophytes et biostratigraphie du Paléogène inférieur du bassin de l'Ebre oriental. <i>Palaeontographica</i>, 178 B(4–6), 143–168 Anadón, P., Cabrera, L., Choi, SJ., Colombo, F., Feist, M. & Sáez A. 1992. Biozonación
632 633 634 635 636 637 638 639 640 641	 Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i>, 22, 73–77. Agardh, C. A. 1824. <i>Systema Algarum</i>. Lundae Literis Berlingianis, Lundae, 312 pp. Anadón, P. & Feist, M. 1981. Charophytes et biostratigraphie du Paléogène inférieur du bassin de l'Ebre oriental. <i>Palaeontographica</i>, 178 B(4–6), 143–168 Anadón, P., Cabrera, L., Choi, SJ., Colombo, F., Feist, M. & Sáez A. 1992. Biozonación del Paleógeno continental de la zona oriental de la Cuenca del Ebro mediante carofitas:
 632 633 634 635 636 637 638 639 640 641 642 	 Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur « l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur–Ludien. <i>Nouvelles Archives du Museum d'Histoire Naturelle de Lyon</i>, 22, 73–77. Agardh, C. A. 1824. <i>Systema Algarum</i>. Lundae Literis Berlingianis, Lundae, 312 pp. Anadón, P. & Feist, M. 1981. Charophytes et biostratigraphie du Paléogène inférieur du bassin de l'Ebre oriental. <i>Palaeontographica</i>, 178 B(4–6), 143–168 Anadón, P., Cabrera, L., Choi, SJ., Colombo, F., Feist, M. & Sáez A. 1992. Biozonación del Paleógeno continental de la zona oriental de la Cuenca del Ebro mediante carofítas: implicaciones en la biozonación general de carofítas de Europa occidental. <i>Acta</i>

Bignot, G., Blondeau, A., Guenet, C., Perreau, M., Poignant, A., Renard, M., Riveline. J.,

2
3
4
5
6
7
8
0
10
10
11
12
13
14
15
16
17
18
19
20
21
י∠ רר
∠∠ วว
23
24
25
26
27
28
29
30
31
27
J∠ 22
22
34
35
36
37
38
39
40
41
42
42
7J 11
-1-1 1 F
45
46
47
48
49
50
51
52
53
54
55
55
50 57
5/
58
59
60

1

646

Ax, P. 1987. The Phylogenetic System. The Systematization of Organisms on the Basis of Their
 Phylogenesis. Wiley-Interscience, New York, 340 pp.

Gruas, C., Dudich, E., Kázmér, M. & Kopek, G. 1985. Age and characteristics of the Eocene transgression at Gánt (Vértes Mts. Transdanubia, Hungary). *Acta Geologica Hungarica*, 28(1–2), 29–48.

650 Carannante, G., Mindszenty, A., Neumann, A. C., Rasmussen, K. A., Simone, L. & Tóth,

K. 1994. Inland blue-hole-type ponds in the Mesozoic–Tertiary karst-filling sequences.

Abstracts, 15th IAS Regional Meeting, April, 1994, Ischia, Italy. 25–59.

- 653 Choi, S. J. 1989. Les Charophytes du Bassin Potassique Catalan (Nord-Est de l'Espagne) à la
 654 limite Eocène-Oligocène. *Paléobiologie continentale*, 26, 1–67.
- Dudich, E. & Kopek, G. 1982. Outlines of the Eocene paleogeography of the Bakony
 Mountains, Transdanubia, Hungary. *Földtani Közlöny (Bulletin of the Hungarian Geological Society*), 3–4, 417–431. [In Hungarian with English Abstract].

Feist-Castel, M. 1970. Distribution verticale des Charophytes dans l'Eocène du Minervois. *Bulletin de la Société Géologique de France*, 12, 926–931.

Feist-Castel, M. 1971. Sur les Charophytes fossiles du Bassin tertiaire d'Alès (Gard). *Geobios*,
4, 157–172.

Feist-Castel, M. 1972. Charophytes Eocène de la région montpelliéraine. *Paléobiologie Continentale*, 3, 1–22.

Feist-Castel, M. 1975. Répartition des Charophytes dans le Paléocène et l'Eocène du bassin d'Aix-en-Provence. *Bulletin de la Société Géologique de France*, 17(7), 88–97.

Feist-Castel, M. 1976. Les charophytes dans le Paléocène du Sud de la France (Provence,
Languedoc, Aquitaine). Etude Systématique et biostratigraphique. Unpublished PhD

Page 31 of 103

1 2		
2 3 4	668	thesis, University of Montpellier II, Science and Technology, Languedoc. Article de
5 6 7 8 9 10 11	669	synthèse, 82 p.
	670	Feist-Castel, M. 1977a. Étude floristique et biostratigraphique des Charophytes dans les séries
	671	du Paléogène de Provence. Géologie Méditerranéenne, 4, 109–138.
12 13	672	Feist-Castel, M. 1977b. Evolution of the charophyte floras in the Upper Eocene and Lower
14 15	673	Oligocene of the Isle of Wight. Palaeontology, 20, 143–157.
16 17 18	674	Feist, M. & Ringeade, M. 1977. Étude biostratigraphique et paléobotanique (Charophytes) des
19 20	675	formations continentales d'Aquitaine de l'Eocène supérieur au Miocène inférieur.
21 22	676	Bulletin de la Société géologique de France, 19, 341–354.
23 24 25	677	Fodor, L. 2007. Segment linkage and stress field in transtensional strike-slip fault array: Field
25 26 27	678	examples from the Pannonian Basin. In: Cunningham, D.F. and Mann, P. (eds): Tectonics
28 29 30 31	679	of strike-slip restraining and releasing bends. Geological Society, London, Special
	680	Publications, 290 , 482 pp.
32 33 34	681	Gevin, P., Feist, M. & Mongereau, N. 1974. Découverte de charophytes d'âge Eocène au Glib
35 36	682	Zegdou (Sahara algérien). Bulletin de la Société d'Histoire naturelle d'Afrique du Nord,
37 38	683	65 , 371–374.
39 40 41	684	Gradstein, F. M., Ogg, J. G. & Smith, A. G. 2004. A Geologic Time Scale 2004. Cambridge:
42 43	685	Cambridge University Press, 589 pp.
44 45	686	Grambast, L. 1956. Le genre Gyrogona Lamarck (Characeae). Compte Rendu Sommaire des
46 47 48	687	Séances de la Société Géologique de France, 14, 278–280.
48 49 50	688	Grambast, L. 1957. Ornementation de la gyrogonite et systématique chez les charophytes
51 52	689	fossiles. <i>Revue générale de Botanique</i> , 64 , 339–362.
53 54	690	Grambast, L. 1958. Etude sur les Charophytes tertiaires d'Europe Occidentale et leurs rapports
55 56 57 58 59 60	691	avec les formes actuelles. Unpublished PhD thesis, University of Paris. 286 p.

2		
2 3 4	692	Grambast, L. 1959. Tendances évolutives dans le phylum des Charophytes. Comptes Rendus
5 6	693	des Séances de L'Académie des Sciences Paris, 249, 557–559.
7 8 0	694	Grambast, L. 1960. Description et signification stratigraphique de deux charophytes d'origine
9 10 11	695	saharienne. Revue de Micropaléontologie, 4, 192–198.
12 13	696	Grambast, L. 1962a. Sur l'intérêt stratigraphique des Charophytes fossiles: exemples
14 15	697	d'application au Tertiaire parisien. Comptes Rendus de la Société géologique de France,
16 17 18	698	7, 207–209.
19 20	699	Grambast, L. 1962b. Aperçu sur les Charophytes tertiaires du Languedoc et leur signification
21 22	700	stratigraphique. Comptes Rendus Sommaires des Séances de la Société Géologique de
23 24 25	701	France, 10 , 313–314.
25 26 27	702	Grambast, L. 1972. Principes de l'utilisation stratigraphique des charophytes. Applications au
28 29	703	Paléogène d'Europe occidentale. Mémoire du Bureau de Recherches géologiques et
30 31 32	704	minières, 77 , 319–328.
32 33 34	705	Grambast, L. 1974. Phylogeny of the Charophyta. Taxon 23, 463–481.
35 36	706	Grambast, L. & Grambast, N. 1954. Sur la position systématique de quelques Charophytes
37 38	707	tertiaires. <i>Revue Générale de Botanique</i> , 61 , 665–671.
39 40 41	708	Grambast, L. & Grambast, N. 1955. Les Raskyelloïdeae, sous-famille fossile des Characeae.
42 43	709	Compte Rendu de l'Académie des Sciences, Paris, 240, 999–1001.
44 45	710	Grambast, L. & Soulié-Märsche, I. 1972. Sur l'ancienneté et la diversification des Nitellopsis
46 47 49	711	(Charophytes). Paléobiologie continentale, III(3), 1–14.
40 49 50	712	Grambast, L. & Grambast-Fessard, N. 1981. Etude sur les Charophytes tertiaires d'Europe
51 52	713	occidentale. III. Le genre Gyrogona. Paléobiologie continentale, 12(2), 1-35.
53 54 55	714	Horn af Rantzien, H. 1959. Morphological types and organ-genera of Tertiary Charophyte
55 56 57 58 59 60	715	fructifications. Stockholm Contributions in Geology, 4, 45–197.

716	
717	
718	
719	
720	
721	
722	
723	
724	
725	
726	
727	
728	
729	
730	
731	
732	
732	
755	
754	
735	
/36	
737	
738	
739	
740	
	 716 717 718 719 720 721 722 723 724 725 726 727 728 729 730 731 732 733 734 735 736 737 738 739 740

716 Hy, F. 1889. Sur les modes de ramifications et cortication dans la famille des Characées.
717 *Bulletin de la Société Botanique de France*, 36, 393–398.

- Iva, M. 1987. Quelques espèces de Charophytes lutétiens du Nord-Ouest de la Transylvanie.
 Pp. 43–48. *In* Petrescu, I., Ghergari, L., Mészáros, N. & Nicorici, E. (eds) *The Eocene from the Transylvanian Basin, Romania*. Babeş-Bolyai Univesity Cluj, Cluj-Napoca.
- Kázmér, M., Dunkl, I., Frisch, W., Ozsvárt, P. 2003. The Palaeogene forearc basin of the
 Eastern Alps and Western Carpathians: subduction erosion and basin evolution. *Journal of Geological Society, London*, 160, 413–428.
- Kopek, G. 1980. A Bakony hegység ÉK-i részének eocénje [L'Éocéne de la partie Nord orientale de la Montagne du Bakony (Transdanubie, Hongrie)]. *Magyar Állami Földtani Intézet Évkönyve [Annales Instituti Publici Geologiae Hungarici]*, 63(1), 7–132 [133–
 176]. [In Hungarian and French]
- 728 Lamarck, J. B. 1822. *Histoire Naturelle des Animaux Sans Vertèbres*. Paris, Verdière Editeur,
 729 711 pp.
- 730 Lindley, J. 1836. *A Natural System of Botany*, second ed. Longman, London, 526 pp.
- 731 Mädler, K. 1955. Zur Taxinomie der tertiären Charophyten. *Geologisches Jahrbuch*, 70, 265–
 732 328.
- 733 Martin, E. 1971. Standard Tertiary and Quaternary calcareous nannoplankton zonation. *In*:
 734 Farinaci, A., Ed., *Proceedings of the II Planktonic Conference*, v. 2, 739–785. Roma:
 735 Edizioni Tecnoscienza.
- ⁹⁹ 736 Martín-Closas, C. & Schudack, M.E. 1991. Phylogenetic analysis and systematization of
 ¹⁰ 737 post-paleozoic Charophytes. *Bulletin de la Société Botanique de France*, 138. Actualités
 ¹³ botaniques 1, 53–71.
- 739 Martín-Closas, C. & Ramos, E. 2005. Palaeogene charophytes of the Balearic Islands (Spain).
 740 *Geologica Acta* 3, 39–58.

2	
3	
4	
5	
6	
7	
, o	
0	
9	
10	
11	
12	
13	
14	
15	
16	
17	
18	
19	
20	
21	
22	
22	
23	
24	
25	
20	
27	
28	
29	
30	
31	
32	
33	
34	
35	
36	
37	
38	
39	
40	
11	
71 12	
-ד∠ גע	
د ب	
44	
45	
46	
47	
48	
49	
50	
51	
52	
53	
54	
55	
56	
57	
58	
50	
59	

Martín-Closas, C., Serra-Kiel, J., Busquets, P. & Ramos-Guerrero, E. 1999a. New
 correlation between charophyte and larger foraminifera biozones (Middle Eocene,
 southeastern Pyrene es). *Geobios*, 32, 5–18.

Martín-Closas, C., Bosch-Casadevall, R., Serra-Kiel, J. 1999b. Biomechanics and evolution
of spiralization in charophyte fructifications. *In*: Kurmann M.H. & Hemsley A.R. (eds.).
The evolution of plant architecture. London, Royal Botanic Gardens Kew. p. 399-421.

- 747 Mebrouk, F., Mahboubi, M., Bessedik, M. & Feist, M. 1997. L'apport des charophytes à la
 748 stratigraphie des formations continentales Paléogènes de l'Algérie. *Geobios*, 30, 171–
 749 177.
 - Megerisi, M. F. & Mamgain, V. D. 1980. The Upper Cretaceous-Tertiary Formations of
 northern Libya: a synthesis. *Department of Geological Researches and Mining Bulletin*,
 12, 1–85.

Migula, W. 1897. Die Characeen Deutschlands, Österreichs und der Schweiz. *In* Rabenhorst, L. (ed.), *Dr. L. Rabenhorst's Kryptogamen-Flora von Deutschland, Oesterreich und der Schweiz. Zweite Auflage, Fünfter Band, Part 12.* Eduard Kummer, Leipzig, 765 pp.

756 Mindszenty, A. 2010. Bauxite deposits of Gánt (Vértes Hills, Hungary). Acta Mineralogica 757 Petrographica, Field Guide Series, 11, 1–11.

Nötzold, T. 1965. Die Präparation von Gyrogoniten und kalkigen Charophyten-Oogonien aus
 festen Kalksteinen. *Monatsberichte der Deutschen Akademie der Wissenschaften zu Berlin*, 7, 216–221.

Ollivier-Pierre, M. F., Riveline, J., Lautridou, J. P. & Cavelier, C. 1988. Le fossé de Céaucé
(Orne) et les bassins ludiens (Eocène supérieur) de la partie orientale du Massif
armoricain: sédimentologie, paléontologie. Intérêt stratigraphique, paléogéographique et
tectonique. *Géologie de France*, 1, 51–60.

2 3	765	Pálfalvi, S., Fodor, L. Kercsmár, Z., Báldi-Beke, M., Kollánvi, K. & Less, G. Y. 2006.
4 5		
6	766	Sedimentation pattern, tectonic control, and basin evolution of the northern
/ 8 9	767	Transdanubian Eocene basins (Vértes Hills, central Hungary). Geophysical Research
10 11	768	Abstracts, 8 , 08384.
12 13	769	Pálfalvi, S. 2007. Reconstruction of Eocene depositional environments in the Vértes Hills,
14 15 16	770	based on microfacies analysis. Unpublished PhD thesis, University of Budapest,
17 18	771	Budapest, 150 pp.
19 20	772	Pia, J. 1927. Charophyta. In Hirmer, M. (ed.), Handbuch der Paläobotanik, 1. R. Oldenbourg
21 22	773	Druck und Verlag, München-Berlin, 708 pp.
23 24 25	774	Rásky, K. 1945. Fossile Charophyten-Früchte aus Ungarn. Budapest, Verlag des Ungarischen
26 27	775	Naturwissenschaftlichen Museums, 75 pp.
28 29	776	Reid, C. & Groves, J. 1921. The Charophyta of the Lower Headon Beds of Hordle (Hordwell)
30 31 32	777	Cliffs (South Hampshire). Quarterly Journal of the Geological Society of London, 77,
32 33 34	778	175–192.
35 36	779	Riveline, J. 1984. Les gisements à charophytes du Cénozoïque (Danien à Burdigalien)
37 38	780	d'Europe occidentale: Lithostratigraphie, Biostratigraphie, chronostratigraphie. Bulletin
39 40 41	781	d'Information des Géologues du Bassin de Paris, 4, 583 pp.
42 43	782	Riveline, J. 1986. Les charophytes du Paléogène et du Miocène inférieur d'Europe occidentale.
44 45	783	Cahiers de Paléontologie (édition du C.N.R.S), 227 pp.
46 47 48	784	Riveline, J. & Cavelier, C. 1987. Les charophytes du passage Eocene moyen-Eocene superieur
49 50	785	en Europe occidentale; implications stratigraphiques. Bulletin de la Société Géologique
51 52	786	<i>de France</i> , III (2), 307–315.
53 54	787	Riveline, J, Berger J. P., Bilan W, Feist, M., Martín-Closas, C., Schudack, M. E. & Soulié-
55 56 57	788	Märsche, I. 1996. European Mesozoic-Cenozoic Charophyte Biozonation. Bulletin de la
58 59 60	789	Société Géologique de France, 167, 453–468.
2		
-------------	--	
2		
1		
4		
5		
6		
7		
8		
9		
10		
11		
11		
12		
13		
14		
15		
16		
17		
10		
10		
19		
20		
21		
22		
23		
24		
25		
25		
20		
27		
28		
29		
30		
31		
32		
22		
22		
34		
35		
36		
37		
38		
39		
10		
-10 /1 1		
41		
42		
43		
44		
45		
46		
47		
48		
⊿0		
77 50		
50		
51		
52		
53		
54		
55		
56		
50		
5/		
58		
59		
60		

Sanjuan, J. & Martín-Closas, C. 2012. Charophyte palaeoecology in the Upper Eocene of the
 Eastern Ebro basin (Catalonia, Spain). Biostratigraphic implications. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 365–366, 247–262.

Sanjuan, J. & Martín-Closas, C. 2014. Taxonomy and palaeobiogeography of charophytes
from the Upper Eocene–Lower Oligocene of the Eastern Ebro Basin (Catalonia, NE
Spain). *Geodiversitas*, 36, 385–420.

- Sanjuan, J., Martín-Closas, C., Costa, E., Barberà, X. & Garcés, M. 2014. Calibration of
 Eocene-Oligocene charophyte biozones in the eastern Ebro Basin (Catalonia, Spain).
 Stratigraphy, 11, 61–81.
- Schudack, M. E. 1993. Die Charophyten im Oberjura und Unterkreide Westeuropas. Mit einer
 phylogenetischen Analyse der Gesamtgruppe. *Berliner Geowissenschaftliche Abhandlungen, Reihe A*, 8, 1–209.
- Sille, N. P., Collinson, M. E., Kucera, M. & Hooker, J. J. 2004. Evolution within the
 charophyte genus *Harrisichara*, late Paleogene, southern England; environmental and
 biostratigraphic implications. *Palaeogeography Palaeoclimatology Palaeoecolgy*, 208,
 153–173.
- 806 Smith, G. M. 1938. *Cryptogamic Botany Volume 1, Algae and Fungi*. McGraw Hill, New York,
 807 547 pp. [Class Charophyceae, p. 127].
- 808 Szőts, E. 1938. A móri Antalhegy óharmadkori képződményei [Early Tertiary formations of
 809 Antalhegy at Mór, Hungary]. PhD thesis, Supplement to Földtani Szemle, Budapest, 42
 810 pp. [Published thesis, in Hungarian].
- 811 Soulié-Märsche, I. 1971. Description de nouvelles Charophytes éocènes. *Bulletin de la Société* d'Histoire naturelle de Toulouse, 107(1–2), 18–27.

- 3 4	813
5 6	814
7 8	815
9 10 11	816
12 13	817
14 15	818
16 17 19	819
19 20	820
21 22	821
23 24	822
25 26 27	823
28 29	824
30 31	825
32 33 34	826
35 36	827
37 38	828
39 40 41	829
41 42 43	830
44 45	831
46 47 48	832
40 49 50	833
51 52	834
53 54 55	835
55 56 57	836
58 59 60	

Soulié-Märsche, I. 1974. Nouvelles espèces de Charophytes éocènes du Bassin d'Aquitaine. 813 814 Compte Rendu 96ème Congrès national Sociétés Savantes Toulouse, 1971, section *Science Paris*, **2**, 109–144. 815

Trabelsi, K., Touir, J., Soulié-Märsche, I., Martín-Closas, C., Soussi, M. & Colin, J. P. 816 2010. Découverte des charophytes de l'Albien dans la Formation Kebar (Tunisie 817 centrale): implications paléoécologiques et paléobiogéographiques. Annales de 818 Paléontologie, 96, 117–133. 819

Trabelsi, K., Soussi, M., Touir, J., Houla, Yassine, Abbes, C. & Martin-Closas, C. 2016. 820 Charophytes biostratigraphy of the nonmarine Lower Cretaceous in the Central Tunisian 821 822 Atlas (North Africa). Paleobiogeographic implications. Cretaceous Research, 67, 66–83. Vaillant, S. 1719. Charactères de quatorze genres de plantes. Mémoires de l'Académie royale 823

des Sciences de Paris pour 1719, 17–20. 824

Vianey-Liaud, M., Jaeger, J.-J., Hartenberger, J.-L. & Mahboubi, M. 1994. Les rongeurs 825 de l'Eocène d'Afrique nord-occidental [Glib Zegdou (Algérie) et Chambi (Tunisie)] et 826 l'origine des Anomaluridae. Palaeovertebrata, 23, 93-118. 827

Vicente, A. & Martín-Closas, C. 2018. Gradualistic characean lineages in the Upper 828 Cretaceous–Palaeocene of southern Europe. *Historical Biology*, **30**(5), 593–607. 829

Weidmann, M., Franzen, E. & Berger, J. P. 1991. Sur l'âge des Couches à Cérithes ou 830 Couches des Diablerets de l'Eocène alpin. Eclogae Geologicae Helvetiae, 84(3), 893-831 919. 832

Willey E. O. 1981. Phylogenetics. John Wiley and Sons, New York. 833

Figure Captions 835

1	
2	
3	
4	
5	
6	
7	
/ Q	
0	
9	
10	
11	
12	
13	
14	
15	
16	
17	
18	
19	
20	
21	
22	
23	
24	
25	
26	
27	
28	
29	
30	
31	
32	
33	
34	
35	
36	
20	
27 20	
20	
39	
40	
41	
42	
43	
44	
45	
46	
47	
48	
49	
50	
51	
52	
53	
54	
55	
56	
57	
58	
59	
60	

837	Figure 1. A, geographical and geological setting of the study area (after Fodor 2007). B,
838	panoramic view of the studied Gánt section at the Vértes Hills (north-western Hungary).
839	
840	Figure 2. Distribution of the charophytes species in the bauxite cover-sequence of the studied
841	Gánt section (Vértes Hills, Hungary), according to Bignot et al. (1985), updated for
842	charophyte content.
843	
844	Figure 3. Raskyella peckii gyrogonites from the Gánt bauxite cover-sequence. A–H,
845	Raskyella pecki var. peckii (samples G-2.2, G-2.3, G-2.4, and G-2.5), HNHM-PBO 1501–
846	1508. A-E, lateral view; F-G, apical view; H, basal view. I-P, Raskyella peckii var.
847	caliciformis (samples G-2.5 and G-6a), HNHM-PBO 1509–1516. I-K, lateral view; L, lateral
848	view of gyrogonite partially broken showing internal cast. M, internal cast with well-
849	developed undulations. N–O, apical view; P, basal view.
850	
851	Figure 4. Frequency distribution of the height (A), width (B), number of convolutions (C),
852	and height/width ratio (ISI) (D) of the Raskyella peckii var. peckii population (50 gyrogonites
853	measured), from samples G-2.2, G-2.3, and G-2.4 in the bauxite cover-sequence of the Gánt
854	section.
855	
856	Figure 5. Frequency distribution of the height (A), width (B), number of convolutions (C),
857	and height/width ratio (ISI) (D) of the Raskyella peckii var. caliciformis population (50
858	gyrogonites measured), from sample G-2.5 in the bauxite cover-sequence of the Gánt section.

859	
860	Figure 6. <i>Raskyella peckii vadaszii</i> gyrogonites from the Gánt bauxite cover-sequence (A–S,
861	samples G-6a and G-6b, HNHM-PBO 1517–1534). A–K, lateral view. L–P, apical view; Q,
862	basal view; R–S , inside wall of a gyrogonite showing the crenate undulation of the cellular
863	sutures in contact with the spiral cells and the apical cells.
864	
865	Figure 7. Frequency distribution of the height (A), width (B), number of convolutions (C),
866	and height/width ratio (ISI) (D) of the Raskyella peckii var. vadaszii population (50
867	gyrogonites measured), from samples G-6a and G-6b in the bauxite cover-sequence of the
868	Gánt section.
869	
870	Figure 8. Gyrogona caelata gyrogonites from the Gánt bauxite cover-sequence. A–G,
871	Gyrogona caelata forma caelata (samples G-2.2, G-2.3, G-2.4, and G-2.5), HNHM-PBO
872	1535–1553. A–B, D–E, basal view C, detail of the ornamentation pattern; F, basal view; G,
873	apical view. H-K, Gyrogona caelata forma bicincta (samples G-6a and G-6b). H, lateral
874	view; I, detail of the ornamentation pattern; J-K, apical view. L-M, Gyrogona caelata forma
875	monolifera (samples G-2.2, G-2.3, G-2.4, and G-2.5). L, lateral view; M, basal view. N-Q,
876	<i>Gyrogona caelata</i> forma <i>baccata</i> (samples G-2.5, G-6a, and G-6b). N–O, lateral view; P,
877	basal view; Q , apical view. R–V , <i>Gyrogona caelata</i> forma <i>fasciata</i> (samples G-6a and G-6b).
878	R , lateral view; S – T , apical view; U – V , detail of the simple (unipartite) basal plate.
879	
880	Figure 9. Frequency distribution of the height (A), width (B), number of convolutions (C),
881	and height/width ratio (ISI) (D) of the Gyrogona caelata population (50 gyrogonites
	 859 860 861 862 863 866 867 868 869 870 871 872 873 874 875 876 877 878 879 880 881

URL: http://mc.manuscriptcentral.com/tjsp

measured), from samples G-2.2, G-2.3, G-2.4, G-2.5, and G-6a in the bauxite cover-sequence

2 3 4	882
5 6	883
7 8 9 10	884
10 11 12	885
13 14 15	886
16 17 18	887
19 20 21	888
22 23	889
24 25 26	890
20 27 28	891
29 30	892
31 32 33	893
34 35 36	894
37 38 39	895
40 41	896
42 43	897
44 45 46 47	898
48 49	899
50 51 52	900
53 54	901
55 56	902
57 58 59 60	903

1

of the Gánt section.

884	
885	Figure 10. Gyrogona tuberosa gyrogonites from the Gánt bauxite cover-sequence (sample G-
886	6b), HNHM-PBO 1554–1563. A–G, lateral view; H, apical view; I–J, basal view.
887	
888	Figure 11. Gyrogonites of genera Psilochara and Nitellopsis from the Gánt bauxite cover-
889	sequence. A-F, Psilochara polita (sample G-6a), HNHM-PBO 1564–1569. A-C, lateral
890	view; D–E , apical view; F , basal view. G–I , <i>Psilochara</i> sp. (sample G-6a), HNHM-PBO
891	1570–1571. G, lateral view; H, apical view; I, internal view showing simple (unipartite) basal
892	plate (arrowed). J-N, Nitellopsis (Tectochara) aff. palaeohungarica (samples G-2.3 and G-
893	2.5), HNHM-PBO 1572–1576. J–L, lateral view; M, apical view; N, basal view.
894	
895	Figure 12. Frequency distribution of the height (A), width (B), number of convolutions (C),
896	and height /width ratio (ISI) (D) of the Psilochara polita population (50 gyrogonites
897	measured), from sample G-6a in the bauxite cover-sequence of the Gánt section.
898	
899	Figure 13. Frequency distribution of the height (A), width (B), number of convolutions (C),
900	and height /width ratio (ISI) (D) of the Nitellopsis (Tectochara) aff. palaeohungarica
901	population (50 gyrogonites measured), from samples G-2.3 and G-2.5 in the bauxite cover-
902	sequence of the Gánt section.
903	

2		
3 4	904	Figure 14. Chara gyrogonites from the Gánt bauxite cover-sequence. A–I, Chara media
5 6	905	(samples G-2.5 and G-6a), HNHM-PBO 1577–1585. A–G, lateral view; H, apical view; I,
/ 8 9	906	basal view. J-O, Chara subcylindrica (sample G-6b), HNHM-PBO 1586–1591. J-M, lateral
) 10 11	907	view; N, apical view; O, basal view.
12 13 14 15	908	
16 17	909	Figure 15. Frequency distribution of the height (A), width (B), number of convolutions (C),
18 19	910	and height /width ratio (ISI) (D) of the Chara media population (50 gyrogonites measured),
20 21 22	911	from samples G-2.5 and G-6a in the bauxite cover-sequence of the Gánt section.
23 24 25 26	912	
20 27 28	913	Figure 16. Frequency distribution of the height (A), width (B), number of convolutions (C),
29 30	914	and height/width ratio (ISI) (D) of the Chara subcylindrica population (50 gyrogonites
31 32 33	915	measured), from sample G-6b in the bauxite cover-sequence of the Gánt section.
34 35 36	916	
37 38 39	917	Figure 17. Stratigraphic distribution of variants of the anagenetic lineage of the species
40 41 42	918	Raskyella peckii.
43 44 45	919	
46 47	920	Figure 18. Charophyte Biostratigraphy, age and correlation of the bauxite cover-sequence of
48 49 50 51 52 53 54 55 56 57 58 59 60	921	the Gánt section.

2	
2	
3	
4	
5	
5	
6	
7	
/	
8	
0	
9	
10	
11	
12	
12	
13	
1/	
14	
15	
16	
10	
17	
10	
١ð	
19	
20	
20	
21	
~ `	
22	
23	
~)	
24	
25	
25	
26	
27	
27	
28	
20	
29	
30	
50	
31	
22	
52	
33	
24	
54	
35	
20	
30	
37	
20	
38	
30	
40	
∆ 1	
71	
42	
⊿२	
-J	
44	
15	
40	
46	
47	
4/	
48	
49	
50	
51	
52	
52	
53	
51	
54	
55	
56	
20	
57	
EO	
28	
59	
60	

1

2

3

4

5

6

7

8

9

10

11

17

18

19

20

21

22

A new diverse charophyte flora and biozonation of the Eocene bauxite cover-sequence at Gánt (Vértes Hills, Hungary)

Khaled Trabelsi^{1,2,3}Trabelsi^{a,b,c*}, Benjamin Sames^{3,4}Sames^{c,d}, Michael Wagreich³Wagreich^c,

Miklós Kázmér^sKázmér^e, Andrea Mindszenty⁶Mindszenty^f, Carles Martín-Closas⁷Closas⁸ ¹Université^aUniversité de Sfax, Faculté des Sciences de Sfax, CP <u>30383000</u>, Sfax, Tunisie; ²Université^bUniversité de Tunis El Manar II, Faculté des Sciences de Tunis, LR18 ES07, C.P. 2092, Tunis, Tunisie; ³Department^eDepartment of Geology, University of Vienna, UZA 2, Althanstrasse 14, 1090

Vienna, Austria; ⁴Sam^dSam Noble Museum, University of Oklahoma, 2401 Chautauqua Ave,

12 Norman, OK 73072, USA; ⁵Department^eDepartment of Palaeontology & MTA-ELTE

13 Geological, Geophysical and Space Science Research Group, Eötvös Loránd University,

14 Pázmány Péter sétány 1/C, 1117 Budapest, Hungary; ⁶Department Department of Physical

15 and Applied Geology, Eötvös Loránd University, Pázmány Péter sétány 1/C, 1117 Budapest,

16 *Hungary*; ⁷Departament^gDepartament de Dinàmica de la Terra i de l'Oceà, Institut de

Recerca de la Biodiversitat (IRBio), Facultat de Ciències de la Terra, Universitat de Barcelona-UB, 08028 Barcelona, Catalonia, Spain.

23 * Corresponding author, e-mail: <u>trabkhalfss@yahoo.fr</u>(Khaled Trabelsi)

A largely new and A diverse Eocene charophyte flora from a section at Gánt (Vértes Hills), Transdanubian Central Range, north-western Hungary, provides significant new information to previous studies only based on subsurface data published from the mid-20th Century. The century. This newly acquired material facilitates the taxonomic study of this flora allows revision and emendation of the species Raskyella peckii facilitates the definition of and thereby defines a new evolutionary anagenetic lineage based on three successive anagenetic varieties of this species which were formerly considered as separate species or subspecies: Raskyella peckii var. peckii (early Lutetian-early Bartonian), Raskvella peckii var. caliciformis (early Bartonian), and Raskvella peckii var. vadaszii (late Bartonian). Based on these this lineage, we propose a new local charophyte biozonation with the new that consists of a 'Raskyella peckii Superzone Superzone' (Lutetian-Bartonian), subdivided into three successive charophyte partial range zones: The 'Raskyella peckii peckii Zone' (Lutetian-lowermost Bartonian) is locally characterized by an assemblage of *R. peckii peckii*, *Gyrogona caelata* forma *caelata*, *G.* caelata forma monolifera and Nitellopsis (Tectochara) aff. palaeohungarica. The, the 'Raskyella peckii caliciformis Zone' (lower Bartonian) includes characterized by the local assemblage of R. peckii var. caliciformis, G. caelata forma caelata, G. caelata forma monolifera, G. caelata forma baccata, Nitellopsis (Tectochara) aff. palaeohungarica and Chara media. The, and the 'Raskyella peckii vadaszii Zone' (upper Bartonian) is composed of the local assemblage of characterized by R. peckii var. vadaszii, G. caelata forma bicincta, G. caelata forma baccata, G. caelata forma fasciata, G. tuberosa, Psilochara polita, Psilochara sp., Chara media and Chara subcylindrica. Future research may show the new local biozonation as applicable to the whole of Europe and complementing complementary to the current European charophyte biozonation. Our results show that the sequences from Gánt, which were previously regarded as upper mid-Eocene (upper Lutetian-lower Bartonian) in age,

appear to compriserepresent a longer timechronostratigraphic interval, i.e.: lower Lutetian tillto
upper Bartonian, with also has implications on the understanding of the regional stratigraphy

50 of. Our chronostratigraphic results imply a longer and more stepwise Eocene major

51 <u>transgression in</u> the Transdanubian Central Range <u>during the Eocenethan previously thought</u>.

53 Keywords: Characeae, Raskyellaceae, biozonation, evolutionary lineagephylozone,
 54 gradualistic evolution, Paleogene, Central Europe.

or peries

56 Introduction

Charophytes represent one of the most useful tools in the biostratigraphic analysis of Cenozoic non-marine deposits worldwide. During the Eocene, charophytes have been the object of significant taxonomic, biostratigraphic, palaeoecological and palaeobiogeographic interest, particularly in South European basins from France and Spain (Grambast 1958, 1962a, 1972a1972; Feist-Castel 1970, 1972, 1975, 1977a; Feist & Ringeade 1977; Anadón & Feist 1981; Riveline 1986; Anadón et al. 1992; Sanjuan & Martín-Closas 2012;). As a result, a European Charophyte Biozonation based largely on these basins was proposed by Riveline et al. (1996). For the Eocene, up to 11 charophyte biozones were defined based mainly on data from Western Europe. This biozonation has been updated since then, e.g., by Sanjuan et al. (2014).) for the upper Eocene. In contrast, the Eocene charophyte flora from Central and Eastern Europe is relatively less well-poorly known; and in the case of Hungary, the pioneer

study by Rásky (1945) based on subsurface data is practically the only known. With the aimone available. This author described for the first time a species-rich charophyte flora from Hungary, at a time when charophyte taxonomy was still at an early stage. She had already assigned the flora studied in this area to the middle Eocene, and her work was the basis for future studies in charophyte taxonomy, including the definition of updating thethe new family Raskyellaceae by Grambast & Grambast (1954). Later, Bignot et al. (1985), based on an exhaustive palaeontological study of the Gánt section including molluscs, foraminifers, ostracods, palynomorphs and charophytes, assigned the bauxite cover-sequence to the Upper Lutetian or Bartonian, respectively.

This study aims to update the compendium of knowledge on Eocene Hungarian and Central
 European charophytes, the present study providesproviding a taxonomic revision of this flora
 from the respective flora that is based on surface samples of sections at the Gánt locality-based
 on an outcropping section, as well as a discussion onregarding its significance for
 biostratigraphic purposes significance and utility.

84 Geological setting

Within the Transdanubian Central Range (TCR), several bauxite deposits that developed duringacross the Cretaceous–Early Tertiary boundary <u>interval</u> are well known <u>for their</u> industrial use as economically exploited ores<u>sources</u> of aluminium, from which ore; and among these, the famous karst bauxite of the Vértes Hills from the Gánt locality,– (north-western Hungary–stands–out. Above) is a prominent example. Strata overlying the bauxite, the succession represents represent the sedimentary record of progressive subcrustal erosion along

the East Alpine-West Carpathian forearc basin (Kázmér et al. 2003). The bauxite represents the base of the Eocene charophyte-bearing strata studied hereininvestigated in this study (Fig. 1). The Eocene succession sampled shows a remarkable lateral and vertical change of sedimentary facies (Pálfalvi et al. 2006; Pálfalvi 2007) that has been attributed to tectonic controlforces acting on the sedimentationsedimentary body (Fodor 2007). The development of the postbauxite deposits occurred under dualwas lain down during oscillation of the groundwater table and the marine relative custatic sea level variations (Carannante et al. 1994; Mindszenty 2010), occurring before the region was invaded flooded by an opena marine incursion during the late Bartonian (Bignot et al. 1985). At the Gánt section, the bauxite cover-sequence vertically shows five stratigraphic units of in a vertical orientation dating from the middle Eocene age (Fig. 2), called 'Packets' in the sense of Bignot et al. (1985), from). 'Packet 1', about 1.5 m thick, corresponds to the bauxite itself, which only units 2 and 3 are studied herein.unconformably overlies Triassic dolomites. 'Packet 2', <u>1.5–2 m in thickness</u>, forms the 'blue-hole' freshwater limestone facies (Carannante et al. 1994; Pálfalvi 2007), while) alternating with clays, rich in charophytes, ostracods and gastropods. 'Packet 3', ca. 6.5 m thick, includes alternating sandy clay, coal and fresh- to brackish water limestone, rich in charophytes, ostracods, molluscs and large benthic foraminifera (Bignot et al. 1985). In the latter facies charophytes 1985). 'Packet 4' is ca. 12 m thick and is mainly dominated by shallow marine limestone, rich in nummulites, miliolids, molluses and ostracods. Finally, 'Packet 5' is *ca*. 6m thick and displays an alternation of shallow

marine marl and limestone, rich in Nummulites and Orbitolites. 'Packets' 2 and 3 were sampled-

for charophytes and are studied here.

------Figures 1, 2 near here------

117 Material and methods

Intensive sampling for charophytes during two consecutive field workswork sessions in 2018 and 2019 has been was carried out on the cover sequence of the bauxite at Gánt (Vértes Hills, Hungary). Moderately-preserved to well-preserved gyrogonites were recovered from marly limestone to hard limestone using acetolysis. This method, first applied by Nötzold (1965) to the study of charophytes, has been recently improved by Trabelsi et al. (2010, 2016) and shown to be very effective in recovering well preserved charophyte fructifications and thalli from consolidated carbonate rocks. It consists in taking of soaking the sample of hard calcareous rock, perfectly dried and mechanically comminuted ininto fragments of about 1-3 mm across-and adding similar, in equal amounts of anhydrous acetic acid and anhydrous copper sulfatesulphate (acid attacks in an exothermic reaction reacts exothermically). After neutralization by ammonia, the residue is treated with ultra-sound, then washed and rinsed. Gyrogonites were measured using the software Motic Images Plus 2.0 ML with a Motic BA310 stereomicroscope in the Departament de Dinàmica de la Terra i de l'Oceà (University of Barcelona, Catalonia, Spain). Scanning electron microscopy on gold-sputtered selected specimens was conducted with a Jeol JEOL JSM-6400 device at the Faculty of Earth Sciences, Geography and Astronomy, University of Vienna (Austria) and with a Quanta 200 device at the Centres Científics i Tecnològics of the University of Barcelona (CCiTUB), Spain.). The studied materials are housed in the Hungarian Natural History Museum (Budapest, Hungary), Botanical Department, Palaeobotanical Collection. The figured specimens are deposited under the inventory numbers: HNHM-PBO xxxx-yyyyy1501–1591.

140 Systematic palaeontology

2	
3	ŀ
4	
5	
7	
8	
9	
10	-
12	
13	
14	-
15	
10	
18	
19	
20 21	-
22	
23	
24	
25 26	
27	
28	L
29 30	-
31	
32	
33	
34 35	
36	
37	
38	•
39 40	
41	
42	
43 11	
45	
46	
47 40	-
48 49	
50	-
51	ŀ
52	
53 54	
55	
56	
57 58	
59	
60	

141	The charophyte flora from the bauxite cover sequence at Gánt (Vértes Hills, Hungary) studied
142	here yields gyrogonites from two families: Raskyellaceae and Characeae. The different
143	charophyte species described below are stratigraphically distributed in the section as shown in
144	the Fig. 2.
145	
146	Division Charophyta Migula, 1897
147	Class Charophyceae G. M. Smith, 1938 emend. Schudack, 1993
148	Order Charales Lindley, 1836
149	Family Raskyellaceae Grambast, 1957
150	Sub-Family Raskyelloideae, Grambast et Grambast, 1955
151	Genus <i>Raskyella</i> (L. & N. Grambast et Grambast, 1954) emend. Grambast, 1962b
152	
153	Type species. Raskyella peckii L. & N.Grambast et Grambast, 1954
154	
155	Remarks. This species is understood as including several traditional taxa belonging to the
156	genus Raskyella L. & N. Grambast (1954), which form a gradualistic lineage during the
157	Eocene. These traditional taxa have been newly combined here to anagenetic varieties within a
158	single evolutionary lineage or an evolutionary species, following the recommendations of
159	Wiley (1981) and Ax (1978).
160	
161	Raskyella peckii var. peckii L. & N.Grambast et Grambast, 1954

1 2 3		
3 4 5	162	(Fig. 3A–H)
6 7 8	163	
9 10 11	164	1954 Raskyella pecki sp. nov. L. & N. Grambast: p. 670, text-figs 1a-c.
12 13 14	165	1957 Raskyella pecki Grambast: p. 358, pl. 5, figs 7–9.
15 16 17	166	1958 Raskyella pecki Grambast: p. 190, figs 87, a-c; p. 191, text-fig. 88.
18 19 20	167	1959 Raskyella pecki Horn af Rantzien: pl. 19, figs 7–13.
21 22 23	168	1971 Raskyella peckii subsp. ganesensis Soulié-Märsche: pl. 2, 1–5.
24 25 26 27 28 29 30 31 32 33	169	1981 Raskyella pecki Anadón & Feist: pl. 1, figs 1–2; pl. 2, figs 3–4.
	170	1986 Raskyella pecki Riveline: pl. 37, figs 7–9.
	171	1999b <u>1999a</u> Raskyella pecki Martín-Closas et al.: p. 11, figs 6, 1–3.
34 35 36	172	
37 38	173	Material. Up to 65 gyrogonites in sample G-2.4, and dozens in samples G-2.2 and G-2.3.
39 40 41	174	Collection numbers of figures figured specimens: HNHM-PBO xxxxx-yyyyy1501-1508.
42 43 44	175	Description. Gyrogonites of are ovoidal to ellipsoidal in shape, spherical to subprolate (ISI 100–
45 46	176	120) and of large size, 800–1050 μm in height and 750–1050 μm in width, showing laterally
47 48	177	7–10 (usually 9) convolutions (Fig. 4). Spiral cells often flat (Fig. $3, A3A$) to slightly convex
49 50 51 52 53	178	(Fig. $3, \underline{C3C}$), or concave (Fig. $3, \underline{B3B}$), but regularly without any kind of ornamentation. Apex
	179	broadly rounded, truncated and flattened, with the spiral cells abruptly discontinue and ending
54 55	180	acutely in the apical periphery to be replaced by the development of five deciduous opercular
56 57 58	181	cells, each obliquely disposed at the end of a spiral cell (Fig. 3, F). In case of germinated

specimen<u>3F</u>). Germinated specimens (Fig. 3, G),3G) show a rose-shaped apical pore
 appearsopening. Internal casts of these gyrogonites were also found (Figs 3, EFig. 3E).

Remarks. The contemporaneous unornamented gyrogonites of *Raskyella peckii* subsp. *ganesensis* Soulié-Märsche, 1971 from the Aquitaine basin (France) appear to represent a relatively smaller gyrogonite population of relatively smaller size within *R. peckii* var. *peckii* and both are here considered here synonymous. However, supplementary research on the type material is needed to verify this synonymy, since this morphotype was not found in the studied Gánt material, nor elsewhere to date the morphotype *ganesesis* is only known from the type locality.

Additionally, the subspecies *Raskyella peckii* subsp. *meridionale* Grambast, 1960, is kept within the rank of subspecies due to its palaeogeographic restriction. The extremely large gyrogonites of this subspecies are limited to the southernmost biogeographic range of *Raskyella peckii*, i.e. Algeria (Grambast 1960; Mebrouk *et al.* 1997), and the Betic Domain inof the Balearic Islands (Martín-Closas *et al.*& Ramos 2005)).

Distribution. This is the first record of *R. peckii* var. *peckii* in Hungary, and Central Europe. This variety is widely distributed in the Lutetian and lower Bartonian of southern Europe, mainly in France (L. and& N. Grambast 1954; Grambast 1958; Soulié-Märsche 1971, 1974; Riveline 1984, 1986) and Spain (Anadón & Feist 1981; Ramos-Guerrero et al. 1989; Anadón et al. 1992; Martín-Closas et al. 1999b).1999a; Martín-Closas & Ramos 2005). The total range of this variety (early Lutetian-lowerlate Bartonian) has been characterized in the Eastern Ebro basin (Northeast SpainCatalonia) by Martín-Closas et al. (1999b1999) based on correlation with larger foraminifera (mainly Nummulites). Furthermore, R. peckii var. peckii has been also reported from North Africa, i.e. in the lower Eocene of Algeria (Gevin et al. 1974; Mebrouk et

2 3	205	al. 1997; Vianey-Liaud 1994), in the late lower Eocene-early middle Eocene of Tunisia
4 5 6	206	(Abdeljaoued et al. 1984) and in the Lutetian of Libya (Megerisi & Mamgain 1980).
7 8 9	207	
10 11 12 13	208	Figures 3, 4 near here
14 15 16	209	
17 18 19	210	Raskyella peckii var. caliciformis (Soulié-Märsche, 1974) comb. nov. Trabelsi & et Martín-
20 21	211	Closas
22 23 24	212	(Fig. 3I–P)
25 26 27 28	213	
29 30	214	Basionym. Raskyella caliciformis Soulié-Märsche, 1974, Compte Rendu 96ème Congrès
31 32 33	215	National des Sociétés Savantes, Toulouse, 1971, Section Science, 2, p. 114, text-figure 2 (pl.
33 34 35	216	I), 1–5.
36 37 38 39	217	
40 41 42	218	1974 Raskyella caliciformis sp. nov Soulié-Märsche: p. 112, pl. 1, figs 1–5.
43 44 45	219	1981 Raskyella caliciformis, Anadón & Feist: pl. 1, figs 6–7; pl. 2, figs 7–8.
46 47 48	220	
49 50 51	221	Material. 63 gyrogonites in sample G-2.5. Collection numbers of figures figured specimens:
52 53	222	HNHM-PBO xxxxx-yyyyy<u>1509–1516</u>.
55 56	223	Description. Large sized gyrogonites (650–1000 μ m high and 750–1050 μ m wide) of globular
57 58	224	to oblate shape (ISI 80-105), showing laterally 6-9 (usually 8) convolutions (Fig. 5). Spiral
59 60	225	cells flat to slightly concave ornamented with stout, vertical to slightly inclined, well

2
2
5
4
5
6
7
/
8
9
10
11
11
12
13
14
15
15
16
17
18
10
20
20
21
22
22
∠J]4
24
25
26
27
27
28
29
30
31
20
32
33
34
35
55
36
37
38
20
29
40
41
42
42
44
45
46
17
47
48
49
50
51
51
52
53
54
55
55
56
57
58
50
55

individualized tuberculestubercles (Fig. 3, 131–K), which are the main diagnostic character of this variety. Apex broadly rounded to truncated showing five deciduous opercular cells placed at the end of the spiral cells. The opercular cells are sometimes convex and somewhat polygonal in shape (Fig. 3, N)3N), rather than rounded, which is the reason why the dehiscence poreopening appears sometimes irregularly star-shaped (Fig. 3, O)3O), rather than rose-shaped (Fig. 3, M3L), as already noted by Soulié-Märsche (1974) in the type material. The internal cast of the gyrogonite (Fig. 3, M3M) shows low and straight ridges delimitingperpendicular to the well–marked, undulated spiral cells. This wavy surface is uncommon in the inside of other raskyellacean gyrogonites and is thought to correspond internally to the external tubercle ornamentation.

Distribution. This is the first record of *R. peckii* var. *caliciformis* in Hungary. It was previously
described from the Bartonian of South France (Soulié-Märsche 1974; Riveline 1986), and from
the lower Bartonian (Auversian local stage) of the Ebro Basin, in-Catalonia, Spain (Anadón &
Feist, 1981; Anadón *et al.* 1992).

----- Figure 5 near here-----

Raskyella peckii var. *vadaszii* (L. & N.Grambast et Grambast 1954) comb. nov. Trabelsi & et Martín-Closas
 244 Martín-Closas

(Fig. 6A–S)

2		
3 4	247	Basionym. Raskyella vadaszi (Rásky) L. & N. Grambast (1954), Revue Générale de
5 6 7	248	Botanique (61), p. 670.
8 9 10	249	
11 12 13	250	1945 Aclistochara vadaszi, sp. nov. Rásky: p. 45, pl. II, figs 22–24.
14 15 16	251	1954 Raskyella vadaszi, comb. nov. L. and N. Grambast: p. 670.
17 18 19	252	1957 Raskyella vadaszi, Grambast: p. 358, pl. 5, figs 1–6.
20 21 22 23	253	1959 Raskyella vadaszi, Horn af Rantzien: pl. 20, figs 1–3.
24 25 26	254	1981 Raskyella vadaszi, Anadón & Feist: pl. 1, fig. 5; pl. 2, fig. 5.
27 28 29	255	1981 Raskyella aff. vadaszi, Anadón & Feist: pl. 1, figs 3–4; pl. 2, figs 1–2, 6.
30 31 32	256	1985 Raskyella vadaszi, Bignot et al.: p. 36, pl. 3, figs 8–11.
33 34 35	257	1986 Raskyella vadaszi, Riveline: pl. 37, figs 1–6.
36 37 38	258	
39 40 41	259	Material. Hundreds of gyrogonites in samples G-6a and G-6b. Collection numbers of
42 43 44	260	figuresfigured specimens: HNHM-PBO xxxxx-yyyyy1517–1534.
45 46	261	Description. Large sized gyrogonites (800–1150 μ m in widthhigh and 800–1150 μ m in
47 48 49	262	highwide) of oblate to ovoidal shape (ISI 80-120), showing laterally 7-10 (usually 9)
50 51	263	convolutions (Fig. 7). Spiral cells often convex and ornamented with stout tubercles of different
52 53	264	shapes and sizes, in most cases oriented parallel to the intercellular sutures (Fig. $\frac{6}{6}$, $A\underline{6A}$ -C), or
54 55 56	265	more rarely tilted 20–30° but keeping parallelism between adjacent nodules (Fig. $\frac{6}{6}$, E <u>6E</u> –G),
57 58	266	this being a diagnostic character of this morphotype. Three tubercle morphologies have been

267 observed: (1) rounded tuberculestubercles, well individualized in the upper half of the

gyrogonite, but fused to neighbouring tubercules tubercules in the lower half (Fig. 6, I6I–K), (2) elongated tubercules tubercles more or less connected to each other and producing slightly wavy (undulated) sutures (Fig. 6, H6H), (3) irregularly alternating round and elongated tubercles (Fig. 6, A–D6D). Base of gyrogonite rounded (Fig. 6, C6C, J) to slightly tapered (Fig. 6, A6A, G) and showing a small, superficial and pentagonal basal pore, sometimes within a less-marked funnel (Fig. 6, Q6Q). Apex of gyrogonite truncated or broadly rounded showing and covered by five deciduous independent opercular cells at the end of the spiral cells. Opercular cells roughly prismatic, with its outer surface concave, flat or slightly convex (Fig. 6, L6L-M), leaving a rose-shaped dehiscence pore in germinated). Germinated specimens show a rounded or rose-like opening (Fig. 6, N6N-P). The inside of the gyrogonite allows observation of a characteristic crenulation of the intercellular sutures near their internal side (Fig. 6, R6R-S), while, to the outside, sutures are flat. Besides, this crenulation occurs also between the opercule cells themselves and between spiral and opercule cells as already described by Feist in Anadón &and Feist (1981).

Distribution. R. vadaszii has been first described by Rásky (1945) from middle Eocene borehole samples (60m60 m depth) at Gánt, in beds roughly equivalent laterally to the outcrop succession studied here. Therefore, the samples studied may be considered as topotypes. Subsequently, the age of this variety was suggested to be upper Bartonian by Bignot et al. (1985) based on the basis of the associated microfossils (foraminifers, ostracods, and pollen) from the same beds of the bauxite cover-sequence at the Gánt section (Vértes Hills). R. peckii var. *vadaszii* has also been also well documented in France, in the upper Bartonian of the Paris Basin (Grambast 1957, 1958, 1962a; Riveline 1986) and infrom several basins from in southern France (Feist-Castel 1976). Anadón & Feist (1981) and Anadón et al. (1992) documented also this variety also in the upper Bartonian of the Eastern Ebro Basin (Catalonia, Spain).

1		
2	• • •	
4	293	Figures 6, 7 near here
5		
6	294	
7		
8	1	
9 10	295	Family Characeae (Richard ex C.A. Agardh, 1824) emend. Martín-Closas andet Schudack,
11		
12	296	1991
13		
14	207	Subfamily Charoidean Proun in Migula 1907
15 16	297	Sublaminy Charolucae Braun <i>in</i> Wigula, 1897
17		
18	298	Genus Gyrogona (Lamarck, 1804 ex Lamarck, 1822) emend. Grambast, 1956
19		
20		
21	299	
22		
24	200	Guragana caplata (Reid & Groves 1921) Grambast 1956
25	300	Gyrogona caetata (Reid eer Groves, 1921) Granibast, 1930
26		
2/	301	(Fig. 8A–V)
20 29		
30		
31	302	
32		
33	303	1921 Chara caelata sp. nov- Reid & Groves: p. 184 pl. 4 figs 4–6
35	505	
36		
37	304	1927 Kosmogyra caelata, Pia: p. 90.
38		
39	205	1054 Durschuck zur angleten L. & N. Crombastin (67
40 41	305	1954 Brachychara caelala, L. & NGrambast. p. 667.
42		
43	306	1956 Gvrogona caelata, Grambast: p. 280.
44		
45		
40 47	307	1977b Gyrogona caelata, Feist-Castel: p. 117.
48		
49	208	1981 Guragana caplata Grambast & Grambast-Fessard: p. 22 text-fig. 11. a-f: pl. 4. figs. 1-9
50	508	$1561 \text{ Gyrogona caetata_c} Oranibast & Oranibast-ressard. p. 22, text-lig. 11, a-1, pl. 4, ligs 1-9.$
51		
52 53	309	1981 Gyrogona cf. caelata <u>Caelata,</u> Anadón & Feist: p. 163.
54		
55		
56	310	1986 Gyrogona caelata, Riveline: pl. 38, figs 1–5, 7–8.
57		
28 20	311	1989 Gyrogona caelata, Choi: pl. 2 figs 1–11
60		

Gyrogona caelata, Weidmann *et al*.: p. 900, fig. 3, C.

2014 Gyrogona caelata, Sanjuan & Martín-Closas: p. 403, fig. 7, A–C.

Material. Up to 10056 gyrogonites in samplessample G-2.2, <u>38 in sample G-2.3</u>, <u>29 in sample G-2.4</u>, <u>18 in sample G-2.5</u>, and <u>23 in sample G-6a</u>. Collection numbers of figures figured specimens: HNHM-PBO <u>xxxxx-yyyyy</u>1535–1553.

Description. Medium to large gyrogonites, 600–800 µm high and 700–1000 µm wide with generally oblate to suboblate spheroidal shape (ISI 80–100) and showing laterally 5–7 (usually 6) convolutions (Fig. 9). Apex and base broadly rounded to subtruncate. Apex showing a less-marked spiral cell periapical thinning (e.g., Fig. 8, R),8G) and, in some specimens, apical nodules of different shape, generally flat or slightly convex (e.g. Fig. 8, <u>S8S</u>). Base showing a small pentagonal basal pore (Fig. 8, G8F), sometimes flared by a shallow funnel (Fig. 8, P8P). Basal plate unicellular and only visible from the gyrogonite interior (Fig. 8, U8U–V). Spiral cells flat to slightly concave and ornamented with different patterns of tuberculestubercles, which allow distinction of a number of morphotypes (allowed Grambast, (1958; Grambast) and Grambast & Grambast-Fessard, (1981), from which) to distinguish a number of morphotypes, ranking them as formae of the same species. The following five forms were recognized in the material studied: (1) G. caelata forma caelata characterized by small nodules well-spaced and irregularly ranged along the spiral cell median line (e.g. Fig. 8, A, D), (2) G. caelata forma monolifera showing medium sized nodules close to each other, sometimes fused forming a thin, irregular mid-cellular crest (e.g. Fig. 8, L), (3) G.8A, D), (2) G. caelata forma bicincta characterized by nodules irregularly ranged along two lines parallel to spiral cell sutures (Fig. 8, H, I), (4) G.8H, I), (3) G. caelata forma monolifera showing medium-sized nodules close to each other, sometimes fused and forming a thin, irregular mid-cellular crest (e.g., Fig. 8L), (4)

336	<u><i>G. caelata</i></u> forma <i>baccata</i> characterized by large nodules very closely ranged along the spiral-
337	cell median line (e.g., Fig. 8, N8N–Q), (5) <i>G. caelata</i> forma <i>fasciata</i> characterized by a broad
338	median band of variable width (e.g., Fig. $\frac{8}{7}$, R <u>8R</u> -T).
339	Distribution. The species Gyrogona caelata is first-reported here from Hungary for the first
340	time. According to Riveline (1986), this species was widely distributed in the upper Lutetian-
341	Priabonian non-marine deposits of Western Europe. It was first recorded from the Isle of Wight,
342	England by Reid & Groves (1921). Thereafter, it was reported from the upper Lutetian to upper
343	Priabonian of France (Grambast 1958; Grambast & Grambast-Fessard 1981; Feist-Castel 1971;
344	Feist & Ringeade 1977; Feist-Castel 1977a, b; Ollivier-Pierre et al. 1988), Spain (Anadón &
345	Feist 1981; Choi 1989; Anadón et al. 1992; Sanjuan & Martín-Closas 2014), Switzerland
346	(Weidmann et al. 1991), as well as from the middle Eocene of Romania (Iva 1987). In North
347	Africa, the species has also been documented from the central part of the Sahara, Algeria, by
348	Mebrouk <i>et al.</i> (1997).
349	
350	Figures 8, 9, near here
351	
352	<i>Gyrogona tuberosa</i> (Reid <u>&et</u> Groves, 1921) Grambast <i>in</i> Grambast <u>&et</u> Grambast-Fessard,
353	1981
354	(Fig. 10A–J)
355	
356	1921 Chara wrighti var. rhytidocarpa, Reid & Groves: p. 183, pl. 4, fig. 3.
357	1958 Gyrogona tuberosa, Grambast: p. 139, fig. 54.

358 1976 Gyrogona tuberosa, Feist-Castel: p. 26.

Gyrogona tuberosa, Grambast & Grambast-Fessard: p. 25, text-fig. 12, a–d; pl. 5, figs 1–
6.

Gyrogona tuberosa, Riveline: pl. 14, figs 8–11.

Material. 35 gyrogonites in samples G-6b. sample G-6b. Collection numbers of figured
 specimens: HNHM-PBO 1554–1563.

Description. Large gyrogonites, 820–1000 µm high and 835–1050 µm wide with generally oblate to suboblate spheroidal shape (ISI 80-100) and showing laterally 6-7 convolutions (usually 6). Apex subtruncate (Fig. 10, A10A, C) to somewhat prominent and pointed (Fig. 10, **D10D**, G) with spiral cells protruding (Fig. 10, H10H). Base broadly rounded (Fig. 10, B10B), F) to slightly tapered (Fig. 10, A10A, C, E) and showing a small, superficial and pentagonal basal pore, sometimes within a less-marked funnel (Fig. 10, 110I–J). Spiral cells flat or concave, smooth or somewhat ornamented and separated by protruding narrow to weakly undulated intercellular ridges.

Distribution. *Gyrogona tuberosa* is first-reported here from Hungary, and from central Europe for the first time. According to Riveline (1986) and Riveline & Cavelier (1987), this species was widely distributed in upper Bartonian non-marine deposits of Western Europe. It was first recorded from the Isle of Wight and Hampshire (England) by Reid & Groves (1921). ThereafterSubesequently, it was reported from the upper Bartonian of the Paris Basin (Grambast 1958; Grambast & Grambast-Fessard 1981; Riveline 1986; Riveline & Cavelier 1987), as well as infrom several basins from four france (Feist-Castel 1976).

1		
2		
3	381	
5		
6	202	Figure 10 pear here
7	302	Tigure 10 heat here
8		
9	383	
10		
12		
13	384	Genus <i>Psilochara</i> Grambast, 1959
14		
15 16	385	
17		
18		
19	386	Psilochara polita (Reid <u>&et</u> Groves, 1921) Grambast, 1959
20		
21 22	207	(Fig. 11A F)
23	307	(11g. 11A-1)
24		
25	388	
26 27		
27	200	1021 Chara polita an nov Roid & Groves: n 197 nl 5 figs 0 12
29	389	1921 <i>Chara polita</i> , sp. nov. Reid & Gloves. p. 187, pl. 3, figs 9, 12.
30		
31	390	1927 Gyrogona politus, Pia: p. 90.
5∠ 33		
34		
35	391	1958 Ovochara polita, comb. nov. Grambast: p. 167.
36		
3/	392	1959 <i>Peckichara polita</i> , Horn af Rantzien; p. 116, pl. 13, figs 1–3.
30 39		
40		
41	393	1959 Psilochara polita, Grambast: p. 11.
42		
43 11	201	1977h Psilochara polita Feist-Castel: n 153
45	554	1) // 0 I subenuru politu, 1 eist-Castel. p. 135.
46		
47	395	1986 Psilochara polita, Riveline: p. 59, pl. 22, figs 8-12.
48		
49 50	206	
51	390	
52		
53	397	Material. 58 gyrogonites in sample G-6a. Collection numbers of figures figured specimens:
54 55		
56	398	HNHM-PBO xxxxx-yyyyy<u>1564–1569</u>.
57		
58	I	
59 60		
00		

Description. Medium--sized gyrogonites (520-830 µm in width and 580-710550-700 µm in height 500–650 and µm in width) with ovoidal shape (ISI 105100–120) and laterally showing 7-810 convolutions (Fig. 11, A11A-C; Fig. 12). Apex round to pointed (Fig. 11, A11A-C). Base tapering to prolonged into a stout basal column (Fig 11, A11A-C) and showing a small pentagonal basal pore (Fig. 11, F11F). Spiral cells smooth, concave or flat and separated by protruding narrow to weakly undulated intercellular ridges. **Distribution.** *Psilochara polita* is first described here from the middle Eocene of Hungary- for the first time. It has previously been described previously from the upper Bartonian of the Isle of Wight, England, (Reid & Groves 1921; Feist-Castel 1977b; Riveline 1986) as well as and of the Paris Basin (Grambast 1958; Riveline 1986). ------ Figures 11, 12 near here------*Psilochara* sp. (Fig. 11G–I) Material. 17 gyrogonites in sample G-6a. Collection numbers of figures figured specimens: HNHM-PBO xxxx-yyyyy1570–1571. **Description.** Medium to large-sized gyrogonites (780–905 µm wide and 670–775 µm high) with elongated ovoidal (subprolate) shape (ISI 110-125) and laterally showing 8-10 convolutions (Fig. 11, G11G). Apex truncated. Apical end of spiral cells enlarged and pointing

420 upwards (Fig. 11, H<u>11H</u>). Base truncated to somewhat tapering, bearing a small pentagonal

3 4	421
5 6	422
7 8	423
9 10	I
11 12	424
13 14	425
15 16	426
17 18	427
19 20	427
21 22	428
23 24	429
25 26	
27 28	430
29 30	431
31 32	
33 34 25	432
35 36 27	433
37 38	100
39 40	434
41 42 42	435
43 44 45	
45 46 47	436
47 48 49	437
50 51	I
52 53	438
54 55	439
56 57	
58 59	440
60	

basal pore. Basal plate unipartite and visible from the gyrogonite interior (Fig. 11, 111). Spiral
cells concave or flat and smooth, except at the periapical area, where they are irregularly
ornamented with a broad mid-cellular crest.

424 Distribution. *Psilochara* sp. is described here first time from the upper Eocene (upper
425 Bartonian) of Hungary.

426 Remark. The low number of gyrogonites hinders a more precise taxonomic attribution of this
427 population. However, it is reported here since it differs in size and shape from the other species

428 of *Psilochara* found at Gánt.

Genus Nitellopsis Hy, 1889

Sous-genreSub-genus Tectochara L. andet N. Grambast, 1954

34 Nitellopsis (Tectochara) aff. palaeohungarica (Rásky, 1945) Grambast &et Soulié-Märsche,

435	1972
436	(Fig. 11J–N)
437	
438	1945 Chara palaeohungarica, sp. nov. Rásky: p. 38, pl. 1, figs 16–18.
439	1955 Tectochara palaeohungarica, comb. nov. Mädler: p. 298.

1972 Nitellopsis (Tectochara) palaeohungarica, nov. comb. Grambast & Soulié-Märsche: p.

<u>4.</u>

Material. 6128 gyrogonites in both samplessample G-2.3 and <u>33 in sample G-2.5</u>. Collection
 numbers of figuresfigured specimens: HNHM-PBO xxxxx-yyyyy1572–1576.

Description. VeryGyrogonites very large gyrogonites (900–1200 μm high and 800–1050 μm wide), oval, prolate spheroidal (ISI 100–120) pear-shaped andin shape, showing 8–11 (often 9) convolutions in lateral view (Fig. 11, J11J–L; Fig. 13). Spiral cells concave to flat. Apex prominent with spiral cells protruding to form a central rosette. Spiral cells show both shorteningnarrowing and thinning in the periapical area (Fig. 11, M11M). Base rounded to almost conical, occasionally lengthened in a short broad column. A large (155–230 μm across) pentagonal basal pore occurs within a wide basal funnel (Fig. 11, N11N).

Distribution. The species *Tectochara 'Chara' palaeohungarica* was first described from subsurface beds attributed to the Paleocene in Dorog, Hungary, by Rásky (1945). Here this species is described from beds cropping out <u>inat</u> Gánt, which are probably <u>time-</u>equivalent to those of the type locality. The present study <u>allows reassigningsupports reassignment of</u> this species to the middle Eocene rather than to the Paleocene.

Remarks. The gyrogonites studied here are diagenetically deformed, which hinders a more definitive taxonomic attribution. A re-study of the type material (HNHM 55.1458–55.1460) by one of the authors of this study (CMC) showed that, besides the holotype, which is a subspherical gyrogonite as illustrated by Rásky (1945, pl. I, Fig. 16), there were more oval to elongated gyrogonites present in the collection similar to those described here, which Rásky (1945, p. 38) termed 'cylindrical'.

1		
2 3	101	
4	464	
5		
6	465	Figure13 near here
/ 8		
9		
10	466	
11		
12	467	Genus <i>Chara</i> Vaillant 1719
13 14		
15		
16	468	
17		
18	160	Chang madia Grambast 1058
19 20	409	Churu meutu Grambast, 1958
21		
22	470	(Fig. 14A–I)
23		
24 25	471	
26	471	
27		
28	472	1958 Chara media, Grambast: p. 178, fig. 81b.
29		
30 31		
32	473	1986. <i>Chara media</i> , Riveline: p. 68, pl. 29, figs 6–12.
33		
34 25	474	
35 36		
37		
38	475	Material. Up to 80 gyrogonites in both samples G-2.5 and G-6a. Collection numbers of
39	470	General marine and UNUM DDO marine and 1577 1595
40 41	476	nguresngured specimens: HNHM-PBO xxxxx-yyyyy1577–1585 .
42		
43	477	Description. Gyrogonites of medium size (400–650 µm high and 300–500 µm wide) ellipsoidal
44		
45 46	478	subprolate (ISI 110–145), laterally showing lateraly 8–11 (usually 9–10) convolutions (Fig. 14,
46 47		
48	479	A <u>14A</u> –G, Fig. 15). Maximum width nearly at the half to 2/3 of height. Apex rounded to slightly
49		
50	480	<u>conical</u> , with distinctly widening of the spiral cell endings (Fig. <u>14, H14H</u>). Spiral cells concave,
51 52		
52 53	481	smooth and without any periapical modification. Base tapering showing a superficial
54	400	
55	482	pentagonal basal pore (Fig. $\frac{14}{141}$).
56		
57 58	I.	
59		

483	Distribution. <i>Chara media</i> is first time described here from the upper Eocene (upper Bartonian)
484	of Hungary for the first time. Grambast (1958) and Riveline (1986) documented this species
485	from the upper Bartonian-lower Oligocene of several basins in France, Belgium, and Germany.
486	
487	Figures14Figures 14, 15 near here
488	
489	Chara subcylindrica Reid &et Groves, 1921
490	(Fig. 14J–O)
491	
492	1921 Chara subcylindrica, sp. nov. Reid & Groves: p. 187, pl. 5, fig. 4–5.
493	1959 Grambastichara subcylindrica, Horn af Rantzien: p. 76, pl. 3, figs 5–7.
494	1986 Chara cf. subcylindrica, Riveline: p. 67, pl. 30, figs 5–8.
495	
496	Material. Up to 250 gyrogonites in sample G-6b. Collection numbers of figures figured
497	specimens: HNHM-PBO xxxxx-yyyyy1586–1591.
498	Description. Medium-sized gyrogonites (500–750 μ m high and 200–400 μ m wide) ellipsoidal
499	prolate to perprolate (ISI 130-200);) in shape, laterally showing 8-11 (usually 9-10)
500	convolutions (Fig. 14, J14J-M; Fig. 16). Maximum width at the equator. Apex rounded with
501	widening of the spiral cell endings (Fig. 14, N14N). Spiral cells often slightly concave to flat,
502	separated by narrow intercellular ridges, cells non-ornamented and without any periapical

modification. Base regularly tapering to round, showing a superficial pentagonal basal pore
(Fig. 14, 0140).

Distribution. This is the first report of *Chara subcylindrica* in the upper Eocene (upper Bartonian) of Hungary. According to Reid and& Groves (1921) and Riveline (1986), this species occurs in the upper Bartonian–lower Oligocene of England, France, Belgium and Germany.

----- Figure 16 near here -----

Discussion

514 Definition of the *Raskyella peckii* anagenetic lineage

Evolutionary lineages formed by a succession of charophyte fructifications changing gradually in time have beenwere first described in the family Clavatoraceae by Grambast (1974). Later, similar lineages were found as well in the family Characeae (e.g. lineage Harrisichara vasiformis-tuberculata described by Feist-Castel 1977b; or lineage Peckichara pectinata by Vicente *et al.* 2018). The Here we describe the first one of such lineage is described herelineages in the family Raskyellaceae. ThreeIn the Lutetian and Bartonian of Gánt (Hungary), three former species of the genus Raskyella, -R. peckii, R. caliciformis, and R. vadaszii, -have beenfound-in the Lutetian and Bartonian of Gánt (Hungary) to form a continuous succession of gyrogonite morphologies, connected by intermediate morphotypes. This gradualistic lineage is interpreted as an evolutionary species in the sense of Wiley (1981) and Ax (1987), and the

525 original taxa have been newly combined as anagenetic varieties of the species with 526 nomenclatural priority, which is *R. peckii*.

The first evolutionary stage of the R. peckii lineage (Fig. 17) is represented by R. peckii var. peckii, and includes, as well, the smallest gyrogonite morphotype initially described as Raskyella peckii ganesesisganesensis Soulié-Märsche, 1971. R. peckii var. peckii is characterized by gyrogonites which are very variable in size, but consistently consistently unornamented. This stage has a long duration, since it has been was documented from the lower Lutetian to the lowerupper Bartonian of the Ebro basinBasin, Catalonia, by Martín-Closas et al. (1999b1999a) and can thus be superimposed toonto some of the ulterior morphotypes of the lineage-, this being quite a common situation in charophyte lineages (e.g., Grambast, 1974). The coeval Raskyella peckii subsp. meridionale Grambast, 1960, was not found in the section studied and corresponds to a southern geographic subspecies of this lineage, thriving in North Africa and the Prebetic Domain in the Balearic Islands in Spain (Grambast 1960; Martín-Closas et al. 2004& Ramos 2005).

The second stepevolutionary stage in the lineage of *R. peckii* is represented by *R. peckii* var. caliciformis. Intermediate morphotypes between R. peckii var. peckii and R. peckii var. *caliciformis* display a progressive increase in the gyrogonite size (up to 1000 µm in highheight), and a change in shape from elongated to rounded, between samples G-2.2 and G-2.4 of the Gánt section. Furthermore, there is a progressive development of the ornamentation corresponding to R. peckii var. caliciformis in the same sequence, with for instance 100% of gyrogonites corresponding to R. peckii var. peckii in sample G-2.4, while in sample G-2.5, there is only 10% of *R. peckii* var. *peckii* for, resulting in 90% of *R. peckii* var. *caliciformis* (Fig. 17).

547 The third stage of the lineage is represented by *R. vadaszii*₅ from the upper Bartonian, which 548 shows an additional increase of the gyrogonite size of about 150–250 μ m in height and 100–

125 µm in width, and especially the development of progressively more complex ornamentation patterns in comparison to the previous evolutionary step (Fig. 17). This gradual change can be observed between samples G-2.5 and G-6b of the Gánt section. Thus, sample G-5 displays a homogeneous population with 100% of gyrogonites corresponding to R. peckii var. caliciformis, while in the overlying sample G-6a, the gyrogonite population of R. peckii contains only 15% of specimens of R. peckii var. caliciformis and 85% of R. peckii var. vadaszii. Finally, in sample G-6b there is a homogeneous population of gyrogonites corresponding to R. peckii var. vadaszii. Overall, the *R. peckii* lineage follows the general evolutionary trend in the evolutionary lineages from of other charophyte families, characterized by an increase in size and sphericity (Clavatoraceae and Raskyellaceae), and a progressive development of ornamentation (Characeae), as shown by Feist-Castel (19771977b), Martín-Closas et al. (1999b), Sille et al. (2004) and Vicente & Martín-Closas (2018). ----- Figure 17 near here -----New local charophyte biozonation Among the charophytescharophyte species described from the bauxite cover-sequence at the Gánt section (Vértes Hills, Hungary), Raskyella peckii represents the most significant species

568 in terms of biostratigraphy for use in biostratigrpahy within the non-marine Lutetian and

Bartonian, as previously suggested by Riveline *et al.* (1996) and Martín-Closas *et al.* (1999a).

570 The *Raskyella peckii* biozone was defined by the latter authors (p. XRiveline *et al.* (1996) as a

571 'partial range zone comprising the interval from the first appearance of *Raskyella peckii* L. and

N. Grambast, 1954, to the first appearance of *Chara friteli* Grambast, 1958-1958, lower Lutetian to lower Bartonian in age. This study proposes to extend this biozone to cover also the upper Bartonian, and to renameredesignate it as a superzone subdivided into the following three successive biozones (Fig. 18):

- Raskyella peckii zoneZone: partial range zone defined from the first occurrence of the morphotype *peckii* to the first occurrence of the morphotype *caliciformis*, Lutetian-lower Bartonian in age. The local charophyte assemblage characterizing this zone in Gánt occurs in the basal part of the studied section ('Packet 2', 'blue-hole' freshwater limestone facies, samples G-2.2, G-2.3 and G-2.4), and is composed of R. peckii var. peckii, G. caelata forma caelata, G. caelata forma monolifera, and Nitellopsis (Tectochara) aff. palaeohungarica, some of which are well known to occur in several European basins (Rásky 1945; Grambast 1958; Riveline 1986; Martín-Closas et al. 1999a) during the same time interval.

- Raskyella peckii caliciformis zoneZone: partial range zone defined from the first occurrence
of the morphotype *caliciformis* to the first occurrence of the morphotype *vadaszii*, lower
Bartonian in age. This zone includes in Gánt the assemblage found in the lower part of 'Packet
3' (samples G-2.5) and composed of *R. peckii* var. *caliciformis*, *G. caelata* forma *caelata*, *G. caelata* forma *monolifera*, *G. caelata* forma *baccata*, *Nitellopsis* (*Tectochara*) aff.
palaeohungarica, and Chara media.

- Raskyella pecki vadaszii zoneZone: partial range zone defined from the first occurrence of
the morphotype *vadaszii* to the first occurrence of the next zone defined in the Paris Basin,
which is *Psilochara repanda*. This zone would be upper Bartonian in age. The assemblage
occurring in the middle part of the Gánt section ('Packet 3', samples G-6a and G-6b), composed
of *R. peckii* var. *vadaszii*, *G. caelata* forma *bicincta*, *G. caelata* forma *baccata*, *G. caelata* forma

 fasciata, *Psilochara polita*, *Psilochara* sp., *Chara media*, and *Chara subcylindrica* 596 characterizes locally this biozone.

598 Implications on the age of the bauxite cover-sequence

The biostratigraphic analysis carried out allows suggesting suggests a relative age of Lutetian-Bartonian (Fig. 18) toof the studied bauxite cover-sequence at the Gánt section (Vértes Hills, Hungary), rather than uniquely uppermost middle Eocene (constraining it to the Bartonian) as previously suggested by Bignot (1985). Thus, theA Lutetian age, deduced from the charophytes and attributed to the lower part of the studied series herein from the TCR, has been already suggested in several works preceding that of Bignot (1985), notably those of Szőts (1938), Kopek (1980), and Dudich & Kopek (1982), on the basis of mollusc and palynomorph biostratigraphy. The data presented herein support the idea that the Eocene succession in the studied area reflects a stepwise marine transgression beginning since the Lutetian, early Eocene upon the bauxite deposits. Such a, beginning in the Lutetian. Our new chronostratigraphic framework sheds new light on the timing of the long-lasting subaerial exposure and alteration process generating the bauxite strata, which stratigraphically seems to occur at a major regional unconformity between late Triassic and early Eocene rather than in the middle Eocene as previously assumed by Mindszenty (2010). Consequently, it will be significant to review the coeval strata from the surrounding localities within the TCRTransdanubian Central Range should be re-studied and analysed from the viewpoint of charophyte-based biostratigraphy point of view, in order to establish a regional charophyte biozonation scheme allowing better understanding of the regional stratigraphic correlation of correlate the post-bauxite sedimentarydepositional event(s) on a regional scale and their associated to improve the understanding of its tectono-eustatic control.

----- Figure 18, near here------Conclusions Eocene (Lutetian–Bartonian) charophyte assemblages are taxonomically described for the first time from an outcrop of the bauxite cover-sequence at the Gánt section (Vértes Hills), Hungary's Transdanubian Central Range. This section shows The sections show for the first time that Raskyellaceae didthe raskyellacean charophytes also evolveevolved in gradualistic lineages, similarly as it wasto what is already known for other charophyte families. The *Raskyella peckii* lineage is formed by three successive stages and is interpreted here in terms of the anagenetic evolutionanagenesis of the evolutionary species *Raskyella peckii*, including its gradual change to the morphotype previously known formerly as R. vadaszi. From the biostartigraphica biostratigraphic viewpoint, the assemblages studied belong to the Raskyella peckii biozone of Martín-Closas et al. (1999b1999a), which is here reinterpreted here as a superzone extending to cover the *Raskyella vadaszii* zoneZone of Riveline *et al.* (1996) and therefore attributed to the Lutetian-Bartonian interval age. This. In this study, this superzone is subdivided herein into three successive local partial range biozones, defined by each of the successive varieties of the evolutionary species R. peckii: (1) the Raskyella peckii *peckii* partial range zone is formed characterized by *R. peckii peckii*, *G. caelata* forma *caelata*, G. caelata forma monolifera and Nitellopsis (Tectochara) aff. palaeohungarica, Lutetian-lowermost Bartonian in age; (2) the Raskyella peckii caliciformis partial range zone is formed characterized by R. peckii caliciformis, G. caelata forma caelata, G. caelata forma

 monolifera, *G. caelata* forma *baccata*, *Nitellopsis* (*Tectochara*) <u>aff</u>. *palaeohungarica* and *Chara media*, lower Bartonian in age; and (3) the *Raskyella peckii <u>vadaszivadaszii</u>* partial range
zone is <u>formedcharacterized</u> by *R. peckii vadaszii*, *G. caelata* forma *bicincta*, *G. caelata* forma *baccata*, *G. caelata* forma *fasciata*, *G. tuberosa*, *Psilochara polita*, *Psilohara* sp., *Chara media*and *Chara subcylindrica*, upper Bartonian in age.

In the light of the new results presented here, the studied charophyte-bearing sequences studied in this work represent a largerlonger time span than previously thought, lasting during running from the Lutetian andto the Bartonian₅₂. This has direct implications inon the understanding of the Eocene_regional stratigraphic scheme of the Transdanubian Central Range-during the Eocene, particularly in terms of synchronism/diachronism; in the regional/supraregional stratigraphic correlation, as well as the timing of the tectono-sedimentary control and palaeogeographic reconstitution. evolution.

655 Acknowledgements

This study is a contribution to UNESCO-IGCP 632 'Continental Crises of the Jurassic: Major Extinction Events and Environmental Changes within Lacustrine Ecosystems', subproject 'Late Mesozoic lacustrine systems in Tunisia and their global correlation' (BS), and UNESCO IGCP Project 661 'The Critical Zone in Karst Systems', subproject: 'Evolution of fossil blue hole limestones and the critical zone in a greenhouse world' (MW, KT) within the scope of the Earth System Science (ESS) programme funded by the Austrian Academy of Sciences (BS), as well as project CGL2015-69805-P from the Spanish Ministry of Innovation and Competitiveness and to project SGR2017-841 of the AGAUR (Catalan Research Agency) (MCCM-C). It also benefitted from Austrian Science Fund (FWF) project P 27687-N29 (BS), and from the results
665	of the SYNTHESYS project HU-TAF6533 accorded to CM-C to study the Klára Rásky
666	collection at the Hungarian Natural History Museum in Budapest. We sincerely acknowledge
667	Dr. Ingeborg Soulié-Märsche (Université de Montpellier II, France), Dr. Josep Sanjuan
668	(University of Barcelona, Spain), and Dr. Alba Vicente (Universidad Nacional Autónoma de
669	México-UNAM, México), as well as the editor Dr. Paul M. Barrett for their valuable and
670	constructive observations and criticism that greatly improved the manuscript during the peer-
671	review process. Kevin Kearney (University of Vienna, Austria) is acknowledged for English
672	language editing.
673	
674	References
6/5	
676	Abdeljaoued, A., Sassi, S., Triat, JM. & Truc, G. 1984. Nouvelles précisions
677	stratigraphiques et biostratigraphiques (mollusques terrestres et charophytes) sur «
678	l'Eocène continental » de Tunisie centrale et méridionale: intervalle Paléocène supérieur-
679	Ludien. Nouvelles Archives du Museum d'Histoire Naturelle de Lyon, 22, 73–77.
680	Agardh, C. A. 1824. Systema Algarum. Lundae Literis Berlingianis, Lundae, 312 pp.
681	Anadón, P & Feist, M. 1981. Charophytes et biostratigraphie du Paléogène inférieur du bassin
682	de l'Ebre oriental. Palaeontographica, 178 B-(4-6), 143-168
683	Anadón, P., Cabrera, L., Choi, SJ., Colombo, F., Feist, M. & <u>SaezSáez</u> A. 1992.
684	BiozonacionBiozonación del PaleogenoPaleógeno continental de la zona oriental de la
685	Cuenca del Ebro mediante carofitas: implicaciones en la biozonacionbiozonación general
686	de carofitas de Europa occidental. Acta Geologica Hispanica, 27-(1-2), 69-94.
687	Ax, P. 1987. The Phylogenetic System. The Systematization of Organisms on the Basis of Their
688	Phylogenesis. New York: Wiley-Interscience, New York, 340 pp.

3 4	689	Bignot, G., Blondeau, A., Guenet, C., Perreau, M., Poignant, A., Renard, M., Riveline. J.,
5 6 7	690	Gruas, C., Dudich, E., Kázmér, M. & Kopek, G. 1985. Age and characteristics of the
7 8 9	691	Eocene transgression at Gánt (Vértes Mts. Transdanubia, Hungary). Acta Geologica
10 11	692	<i>Hungarica</i> , 28 -(1–2), 29–48.
12 13	693	Carannante, G., Mindszenty, A., Neumann, A. C., Rasmussen, K. A., Simone, L. & Tóth,
14 15 16	694	K. 1994. Inland blue-hole-type ponds in the Mesozoic–Tertiary karst-filling sequences.
17 18	695	Abstracts, 15th IAS Regional Meeting, April, 1994, Ischia, Italy. 25-59.
19 20	696	Choi, S. J. 1989. Les Charophytes du Bassin Potassique Catalan (Nord-Est de l'Espagne) à la
21 22 23	697	limite Eocène-Oligocène. Paléobiologie continentale, 26, 1–67.
23 24 25	698	Dudich, E. & Kopek, G. 1982. Outlines of the Eocene paleogeography of the Bakony
26 27	699	Mountains, Transdanubia, Hungary. Földtani Közlöny (Bulletin of the Hungarian
28 29	700	<i>Geological Society)</i> , 3–4 , 417–431. [In Hungarian with English Abstract].
30 31 32	701	Feist-Castel, M. 1970. Distribution verticale des Charophytes dans l'Eocène du Minervois.
33 34	702	Bulletin de la Société Géologique de France, 12, 926–931.
35 36	703	Feist-Castel, M. 1971. Sur les Charophytes fossiles du Bassin tertiaire d'Alès (Gard). Geobios,
37 38 39	704	4, 157–172.
40 41	705	Feist-Castel, M. 1972. Charophytes Eocène de la région montpelliéraine. Paléobiologie
42 43	706	Continentale, 3 , 1–22.
44 45 46	707	Feist-Castel, M. 1975. Répartition des Charophytes dans le Paléocène et l'Eocène du bassin
47 48	708	d'Aix-en-Provence. Bulletin de la Société Géologique de France, 17-(7), 88-97.
49 50	709	Feist-Castel, M. 1976. Les charophytes dans le Paléocène du Sud de la France (Provence,
51 52 53	710	Languedoc, Aquitaine). Etude Systématique et biostratigraphique. Unpublished PhD
54 55	711	thesis, University of Montpellier II, Science and Technology, Languedoc. Article de
56 57 58 59 60	712	synthèse, 82 p.

graphique des Charophytes dans les séries
néenne, 4 , 109–138.
e floras in the Upper Eocene and Lower
20 , 143–157.
biostratigraphique et paléobotanique
d'Aquitaine de l'Eocène supérieur au
gique de France, 19 , 341–354.
ranstensional strike-slip fault array: Field
gham, D.F. and Mann, P. (eds): Tectonics
s. Geological Society, London, Special
verte de charophytes d'âge Eocène au Glib
é d'Histoire naturelle d'Afrique du Nord,
A Geologic Time Scale 2004. Cambridge:
Characeae). Compte Rendu Sommaire des
4, 278–280.
ite et systématique chez les charophytes
-362.
res d'Europe Occidentale et leurs rapports
esis, University of Paris. 286 p.
nylum des Charophytes. Comptes Rendus
, 249 , 557–559.

2	
3	737
4 5	
5 6	738
7	730
8 9	139
10	740
11 12	
13	741
14 15	742
16	
17	743
18 19	711
20	/44
21 22	745
23	
24	746
25 26	7/17
27	/ 4 /
28 29	748
30	
31	749
32 33	750
34	/50
35	751
30 37	
38	752
39 40	753
40 41	/55
42	754
43 44	
44 45	755
46	
47 48	/56
49	757
50	
51 52	758
53	- - ^
54 55	759
56	760
57	
58 59	
60	

Grambast, L. 1960. Description et signification stratigraphique de deux charophytes d'origine
saharienne. *-Revue de Micropaléontologie*, 4, 192–198.

- 739 Grambast, L. 1962a. Sur l'intérêt stratigraphique des Charophytes fossiles: exemples
 740 d'application au Tertiaire parisien. *Comptes Rendus de la Société géologique de France*,
 741 7, 207–209.
- Grambast, L. 1962b. Aperçu sur les Charophytes tertiaires du Languedoc et leur signification
 stratigraphique. *Comptes Rendus Sommaires des Séances de la Société Géologique de France*, 10, 313–314.
- Grambast, L. <u>1972a1972</u>. Principes de l'utilisation stratigraphique des charophytes.
 Applications au Paléogène d'Europe occidentale. *Mémoire du Bureau de Recherches géologiques et minières*, 77, 319–328.
- Grambast, L. 1972b. Etude sur les Charophytes tertiaires d'Europe Occidentale. I: Genre
 Tectochara. *Paléobiologie Continentale*, 3 (2), 1–301974. Phylogeny of the Charophyta.
 Taxon 23, 463–481.
 - Grambast, L. & Grambast, N. 1954. Sur la position systématique de quelques Charophytes
 tertiaires. *Revue Générale de Botanique*, 61, 665–671.
 - Grambast, L. & Grambast, N. 1955. Les Raskyelloïdeae, sous-famille fossile des Characeae.
 Compte Rendu de l'Académie des Sciences, Paris, 240, 999–1001.

Grambast, L. & Soulié-Märsche, I. 1972. Sur l'ancienneté et la diversification des *Nitellopsis* (Charophytes). *Paléobiologie continentale*, III-(3), 1–14.

Grambast, L. & Grambast-Fessard, N. 1981. Etude sur les Charophytes tertiaires d'Europe
 occidentale. III. Le genre *Gyrogona*. *Paléobiologie continentale*, 12-(2), 1–35.

Horn Afaf Rantzien, H. 1959. Morphological types and organ-genera of tertiaryTertiary
 Charophyte fructifications. *Stockholm ContributionContributions in Geology*, 4, 45–197.

3
4
5
6
7
/ 0
0
9
10
11
12
13
14
15
16
17
18
19
20
21
22
~~ 72
∠⊃ ⊃4
24
25
26
27
28
29
30
31
32
33
34
35
36
37
20
20
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
57
54
55
50
5/
58
59
60

Hy, F. 1889. Sur les modes de ramifications et cortication dans la famille des Characées. *Bulletin de la Société Botanique de France*, 36, 393–398.

- **Iva, M.** 1987. Quelques espèces de Charophytes lutétiens du Nord-Ouest de la Transylvanie.
- Pp. 43–48. *In* Petrescu, I., Ghergari, L., Mészáros, N. & Nicorici, E. (eds).) *The Eocene from the Transylvanian Basin, Romania*. Babeş-Bolyai Univesity Cluj, Cluj-Napoca.
- 766 Kázmér, M., Dunkl, I., Frisch, W., Ozsvárt, P. 2003. The Palaeogene forearc basin of the
 767 Eastern Alps and Western Carpathians: subduction erosion and basin evolution. *Journal* 768 of Geological Society, London, 160, 413–428.

Kopek, G. 1980. A Bakony hegység ÉK-i részének eocénje [L'Éocéne de la partie Nordorientale de la Montagne du Bakony (Transdanubie, Hongrie)]. *Magyar Állami Földtani Intézet Évkönyve [Annales Instituti Publici Geologiae Hungarici]*, 63-(1), 7–132 [133– 176]. [In Hungarian and French]

- Lamarck, J. B. 1822. *Histoire Naturelle des Animaux Sans Vertèbres*. Paris, Verdière Editeur,
 774 711 pp.
- **Lindley, J.** 1836. *A Natural System of Botany*, second ed. Longman, London, 526 pp.
- 776 Mädler, K. <u>1955b1955</u>. Zur Taxinomie der tertiären Charophyten. *Geologisches Jahrbuch*, 70,
 777 265–328.
 - Martin, E. 1971. Standard Tertiary and Quaternary calcareous nannoplankton zonation. *In*:
 Farinaci, A., Ed., *Proceedings of the II Planktonic Conference*, v. 2, 739–785. Roma:
 Edizioni Tecnoscienza.
- Martín-Closas, C. & Schudack, M.E. 1991. Phylogenetic analysis and systematization of
 post-paleozoic Charophytes. *Bulletin de la Société Botanique de France*, 138. Actualités
 botaniques 1, 53–71.
- Martín-Closas, C. & Ramos, E. 2005. Palaeogene charophytes of the Balearic Islands (Spain).
 Geologica Acta 3, 39–58.

2		
3 4	786	Martín-Closas, C., Bosch, R. & Serra-Kiel, J. 1999a. Biomechanics and evolution of
5 6	787	spiralization in charophyte fructifications. In Kurmann MH, Hemsley A.R. (eds). The
7 8	788	evolution of plant architectureLondon Royal Botanic Gardens Kew, 506 pp.
9 10 11	789	Martín-Closas, C., Serra–Kiel, J., Busquets, P. & Ramos-Guerrero, E. 1999b<u>1999a</u>. New
12 13	790	correlation between charophyte and larger foraminifera biozones (Middle Eocene,
14 15	791	southeastern Pyrene es). <i>Géobios<u>Geobios</u></i> , 32 , 5–18.
16 17 18	792	Martín-Closas, C., Bosch-Casadevall, R., Serra-Kiel, J. 1999b. Biomechanics and evolution
19 20	793	of spiralization in charophyte fructifications. In: Kurmann M.H. & Hemsley A.R. (eds.).
21 22	794	The evolution of plant architecture. Mebrouk, F. & Feist, M. 1999. Nouvelles
23 24 25	795	charophytes de l'Eocène continental de l'Algérie. Géologie méditerranéenne, 26, 29-45.
25 26 27	796	London, Royal Botanic Gardens Kew. p. 399-421.
28 29	797	Mebrouk, F., Mahboubi, M., Bessedik, M. & Feist, M. 1997. L'apport des charophytes à la
30 31 32	798	stratigraphie des formations continentales Paléogènes de l'Algérie. Geobios, 30, 171-
33 34	799	177.
35 36	800	Megerisi, M. F. & Mamgain, V. D. 1980. The Upper Cretaceous-Tertiary Formations of
37 38 30	801	northern Libya: a synthesis. Department of Geological Researches and Mining Bulletin,
39 40 41 42	802	12 , 1–85.
42 43 44	803	Mindszenty, A. 2010. Bauxite deposits of Gánt (Vértes Hills, Hungary). Acta Mineralogica-
45 46 47	804	Petrographica, Field Guide Series, 11, 1–11.
48 49	805	Migula, W. 1897. Die Characeen Deutschlands-, Österreichs und der Schweiz. In Rabenhorst,
50 51 52	806	L. (ed.), Dr. L. Rabenhorst's Kryptogamen-Flora von Deutschland, Oesterreich und der
52 53 54 55	807	Schweiz. Zweite Auflage, Fünfter Band, Part 12. Eduard Kummer, Leipzig, 765 pp.
56 57	808	Mindszenty, A. 2010. Bauxite deposits of Gánt (Vértes Hills, Hungary). Acta Mineralogica-
58 59 60	809	<u>Petrographica, Field Guide Series, 11, 1–11.</u>

Nötzold, T. 1965. Die Präparation von Gyrogoniten und kalkigen Charophyten-Oogonien aus festen Kalksteinen. Monatsberichte der Deutschen Akademie der Wissenschaften zu Berlin, 7, 216–221.

Ollivier-Pierre, M. F., Riveline, J., Lautridou, J. P. & Cavelier, C. 1988. Le fossé de Céaucé (Orne) et les bassins ludiens (Eocène supérieur) de la partie orientale du Massif armoricain: sédimen-tologiesédimentologie, paléontologie. Intérêt stratigraphique, paléogéographique et tectonique. Géologie de France, 1, 51-60.

Pálfalvi, S., Fodor, L. Kercsmár, Z., Báldi-Beke, M., Kollányi, K. & Less, G. Y. 2006. Sedimentation pattern, tectonic control, and basin evolution of the northern Transdanubian Eocene basins (Vértes Hills, central Hungary). Geophysical Research 08384. Abstracts, 8,

https://meetings.copernicus.org/www.cosis.net/abstracts/EGU06/08384/EGU06-J-

08384.pdf

Pálfalvi, S. 2007. Reconstruction of Eocene depositional environmets environments in the Vértes Hills, based on microfacies analysis. Unpublished PhD thesis, University of Budapest, Budapest, 150 pp.

Pia, J. 1927. Charophyta. In Hirmer, M. (ed.), Handbuch der Paläobotanik, 1. R. Oldenbourg Druck und Verlag, München-Berlin, 708 pp.

Rásky, K. 1945. - Fossile Charophyten-Früchte aus Ungarn. Budapest, Verlag des Ungarischen Naturwissenschaftlichen Museums, 75 pp.

- Ramos-Guerrero, E., Rodríguez-Perea, A., Sabat, F. & Serra-Kiel, J. 1989. Cenozoic tectosedimentary evolution of Mallorca island. Geodinamica Acta, 3, 53-72.
- Reid, C. & Groves, J. 1921. The Charophyta of the Lower Headon Beds of Hordle (Hordwell) Cliffs (South Hampshire). Quarterly Journal of the Geological Society of London, 77,
- 175-192.

- 3 4	83
5 6	83
7 8 0	83
9 10 11	83
12 13	83
14 15 16	84
17 18	84
19 20	84
21 22 22	84
23 24 25	84
26 27	84
28 29	84
30 31 32	84
33 34	84
35 36	84
37 38 39	85
40 41	85
42 43	85
44 45 46	85
40 47 48	85
49 50	85
51 52	85
53 54 55	85
56 57	85
58 59	85
00	

Riveline, J. 1984. Les gisements à charophytes du Cénozoïque (Danien à Burdigalien)
 d'Europe occidentale: Lithostratigraphie, Biostratigraphie, chronostratigraphie. *Bulletin d'Information des Géologues du Bassin de* Paris, 4, 583 pp.

Riveline, J. 1986. Les charophytes du Paléogène et du Miocène inférieur d'Europe occidentale.
 Cahiers de Paléontologie (édition du C.N.R.S), 227 ppp.

Riveline, J. & Cavelier, C. 1987. Les charophytes du passage Eocene moyen-Eocene superieur
 en Europe occidentale; implications stratigraphiques. *Bulletin de la Société Géologique de France*, III-(2), 307–315.

Riveline, J, Berger J. P., Bilan W, Feist, M., Martín-Closas, C., Schudack, M. E. & Soulié Märsche, I. 1996. European Mesozoic-Cenozoic Charophyte Biozonation. *Bulletin de la Société Géologique de France*, 167, 453–468.

846 Sanjuan, J. & Martín-Closas, C. 2012. Charophyte palaeoecology in the Upper Eocene of the
 847 Eastern Ebro basin (Catalonia, Spain). Biostratigraphic implications. *Palaeogeography,* 848 *Palaeoclimatology, Palaeoecology*, 365–366, 247–262.

Sanjuan, J. & Martín-Closas, C. 2014. Taxonomy and palaeobiogeography of charophytes from the Upper Eocene–Lower Oligocene of the Eastern Ebro Basin (Catalonia, NE Spain). *Geodiversitas*, 36, 385–420.

Sanjuan, J., Martín-Closas, C., Costa, E., Barberà, X. & Garcés, M. 2014. Calibration of
 Eocene-Oligocene charophyte biozones in the eastern Ebro Basin (Catalonia, Spain).
 Stratigraphy, 11, 61–81.

855 Schudack, M. E. 1993. Die Charophyten im Oberjura und Unterkreide Westeuropas. Mit einer
 856 phylogenetischen Analyse der Gesamtgruppe. *Berliner Geowissenschaftliche* 857 *Abhandlungen, Reihe A*, 8, 1–209.

858 Sille, N. P., Collinson, M. E., Kucera, M. & Hooker, J. J. 2004. Evolution within the
859 charophyte genus *Harrisichara*, late Paleogene, southern England; environmental and

1 2		
2 3 4 5 6 7 8 9 10 11	860	biostratigraphic implications. Palaeogeography Palaeoclimatology Palaeoecolgy, 208,
	861	153–173.
	862	Smith, G. M. 1938. Cryptogamic Botany Volume 1, Algae and Fungi. McGraw Hill, New York,
	863	547 pp. [Class Charophyceae, p. 127].
12 13	864	Szőts, E. 1938. A móri Antalhegy óharmadkori képződményei [Early Tertiary formations of
14 15 16 17 18 19 20 21 22 23 24 25 26 27 28 29 30 31 32 33 34 35 36 37 38 39 40 41 42	865	Antalhegy at Mór, Hungary]. PhD thesis, Supplement to Földtani Szemle, Budapest, 42
	866	pp. [Published thesis, in Hungarian]].
	867	Soulié-Märsche, I. 1971. Description de nouvelles Charophytes éocènes. Bulletin de la Société
	868	d'Histoire naturelle de Toulouse, 107-(1–2), 18–27.
	869	Soulié-Märsche, I. 1974. Nouvelles espèces de Charophytes éocènes du Bassin d'Aquitaine.
	870	Compte Rendu 96ème Congrès national Sociétés Savantes Toulouse, 1971, section
	871	<i>Science Paris</i> , 2 , 109–144.
	872	Trabelsi, K., Touir, J., Soulié-Märsche, I., Martín-Closas, C., Soussi, M. & Colin, J. P.
	873	2010. Découverte des charophytes de l'Albien dans la Formation Kebar (Tunisie
	874	centrale): implications paléoécologiques et paléobiogéographiques. Annales de
	875	Paléontologie, 96, 117–133.
	876	Trabelsi, K., Soussi, M., Touir, J., Houla, Yassine, Abbes, C. & Martin-Closas, C. 2016.
	877	Charophytes biostratigraphy of the nonmarine Lower Cretaceous in the Central Tunisian
44 45	878	Atlas (North Africa). Paleobiogeographic implications. Cretaceous Research, 67, 66-83.
46 47 48 49 50 51 52 53 54	879	Vaillant, S. 1719. Charactères de quatorze genres de plantes. Mémoires de l'Académie royale
	880	des Sciences de Paris pour 1719, 17–20.
	881	Vianey-Liaud, M., Jaeger, JJ., Hartenberger, JL. & Mahboubi, M. 1994. Les rongeurs
	882	de l'Eocène d'Afrique nord-occidental [Glib Zegdou (Algérie) et Chambi (Tunisie)] et
55 56	883	l'origine des Anomaluridae. Palaeovertebrata, 23, 93-118.
57 58 50		
60		

1	
2	
3	
4 5	
5	
7	
8	
9	
10	
11	
12	
14	
15	
16	
17	
18	
19	
20 21	
∠ı 22	
23	
24	
25	
26	
27	
28	
29	
31	
32	
33	
34	
35	
36	
3/	
30	
40	
41	
42	
43	
44	
45	
46	
47 78	
49	
50	
51	
52	
53	
54	
55	
50 57	
57	
59	
60	

884	Vicente, A. & Martín-Closas, C. 2018. Gradualistic characean lineages in the Upper
885	Cretaceous–Palaeocene of southern Europe. <i>Historical Biology</i> , 30 -(5), 593–607.
886	Weidmann, M., Franzen, E. & Berger, J. P. 1991. Sur l'âge des Couches à Cérithes ou
887	Couches des Diablerets de l'Eocène alpin. Eclogae Geologicae Helvetiae, 84-(3), 893-
888	919.
889	Wiley, R. H. 1981. Social structure and individual ontogenies: problems of description,
890	mechanism, and evolution. In Bateson, P. P. G. and Klopfer, P. H. (ed.). Perspecitves in
891	Ethology. Vol. 4: Advantages of diversity. Plenum Press, New York, 262 pp.
892	Willey E. O. 1981. Phylogenetics. John Wiley and Sons, New York.
893	
894	Figure Captions
895	
896	Figure 1. A, geographical and geological setting of the study area (after Fodor 2007). B,
897	panoramic view of the studied Gánt section at the Vértes Hills (north-western Hungary).
898	
050	
899	Figure 2. Distribution of the charophytes species in the bauxite cover-sequence from of the
900	studied Gánt section (Vértes Hills, Hungary), according to Bignot et al. (1985).), updated for
901	charophyte content.
902	
903	Figure 3. Gyrogonites of species-Raskyella peckii gyrogonites from the Gánt bauxite cover-
904	sequence. A–H, Raskyella pecki var. peckii (samples G-2.2, G-2.3, G-2.4, and G-2.5).
905	HNHM-PBO 1501–1508. A–E, lateral view; F–G, apical view; H, basal view. I–P, Raskyella
906	peckii var. caliciformis (samplesamples G-2.5, and G-6a).), HNHM-PBO 1509–1516. I–K,
907	lateral view; L, lateral view of gyrogonite partially devoided from external layerbroken

showing the oospore.internal cast. M, the oosporeinternal cast with well-developed
undulations. N–O, apical view; P, basal view. Specimen numbers HNHM-PBO xxxxx-yyyyy
will be added for each respective specimen.

Figure 4. Frequency distribution of the lengthheight (A), width (B), number of convolutions

(C), lengthand height/width ratio (ISI) (D) of the *Raskvella peckii* var. *peckii* population (50

gyrogonites measured), from samples G-2.2, G-2.3, and G-2.4 in the bauxite cover-sequence

915 of the Gánt section.

Figure 5. Frequency distribution of the lengthheight (A), width (B), number of convolutions
(C), lengthand height/width ratio (ISI) (D) of the *Raskyella peckii* var. *caliciformis* population
(50 gyrogonites measured), from sample G-2.5 in the bauxite cover-sequence of the Gánt
section.

Figure 6. Gyrogonites of species *Raskyella peckii <u>vadaszii gyrogonites</u>* from the Gánt bauxite
cover-sequence. (A–S, *Raskyella peckii var. vadaszii* (samples G-6a and G-6b, <u>HNHM-PBO</u>
<u>1517–1534</u>). A–K, lateral view. L–P, apical view; Q, basal view; R–S, inside wall of a
gyrogonite showing the crenate undulation of the cellular sutures in contact with the spiral
cells and the apical cells. Specimen numbers HNHM-PBO xxxxx-yyyyy will be added for
each respective specimen.

Figure 7. Frequency distribution of the lengthheight (A), width (B), number of convolutions
(C), lengthand height/width ratio (ISI) (D) of the *Raskyella peckii* var. *vadaszii* population (50
gyrogonites measured), from samples G-6a and G-6b in the bauxite cover-sequence of the
Gánt section.

Figure 8. Gyrogonites of species Gyrogona caelata gyrogonites from the Gánt bauxite cover-sequence. A–G, Gyrogona caelata forma caelata (samples G-2.2, G-2.3, G-2.4, and G-2.5+), HNHM-PBO 1535–1553. A–B, D–E, basal view C, detail of the ornamentation pattern; F, basal view; G, apical view. H-K, Gyrogona caelata forma bicincta (samples G-6a, and G-6b). H, lateral view; I, detail of the ornamentation pattern; J-K, apical view. L-M, Gyrogona caelata forma monolifera (samples G-2.2, G-2.3, G-2.4, and G-2.5). L, lateral view; M, basal view. N–Q, Gyrogona caelata forma baccata (samples G-2.5, G-6a, and G-6b). N–O, lateral view; P, basal view; Q, apical view. R-V, Gyrogona caelata forma fasciata (samples G-6a; and G-6b). R, lateral view; S-T, apical view; U-V, detail of the simple (unipartite) basal plate. Specimen numbers HNHM-PBO xxxxx-yyyyy will be added for each respective specimen.

Figure 9. Frequency distribution of the lengthheight (A), width (B), number of convolutions
(C), lengthand height/width ratio (ISI) (D) of the *Gyrogona caelata* population (50 gyrogonites measured), from samples G-2.2, G-2.3, G-2.4, G-2.5, and G-6a in the bauxite cover-sequence of the Gánt section.

Figure 10. Gyrogonites of species Gyrogona tuberosa gyrogonites from the Gánt bauxite cover-sequence (samples sample G-6b), HNHM-PBO 1554–1563. A-G, lateral view; H, apical view; I-J, basal view. Specimen numbers HNHM-PBO xxxxx-yyyyy will be added for each respective specimen. Figure 11. Gyrogonites of genusgenera Psilochara and Nitellopsis from the Gánt bauxite cover-sequence. A–F, Psilochara polita (samplessample G-6a), HNHM-PBO 1564–1569. A-C, lateral view; D-E, apical view; F, basal view. G-I, Psilochara sp. (samplessample G-6a).), HNHM-PBO 1570–1571. G, lateral view; H, apical view; I, internal view showing simple (unipartite) basal plate (arrowed). J–N, Nitellopsis (Tectochara) aff. palaeohungarica (samples G-2.3 and G-2.5), HNHM-PBO 1572–1576. J–L, lateral view; M, apical view; N, basal view. Specimen numbers HNHM-PBO xxxxx-yyyyy will be added for each respective specimen. Figure 12. Frequency distribution of the lengthheight (A), width (B), number of convolutions (C), lengthand height /width ratio (ISI) (D) of the *Psilochara polita* population (50 gyrogonites measured), from sample G-6a in the bauxite cover-sequence of the Gánt section. Figure 13. Frequency distribution of the lengthheight (A), width (B), number of convolutions (C), lengthand height /width ratio (ISI) (D) of the Nitellopsis (Tectochara) aff. palaeohungarica population (50 gyrogonites measured), from samples G-2.3 and G-2.5 in the bauxite cover-sequence of the Gánt section.

1	
2	
2	
1	
5	
5	
0	
/	
8	
9	
10	
11	
12	
13	
14	
15	
16	
17	
18	
10	
עו 20	
∠∪ 21	
21	
22	
23	
24	
25	
26	
27	
28	
29	
30	
21	
21	
32	
33	
34	
35	
36	
37	
38	
39	
40	
41	
42	
43	
11	
 15	
40	
46	
4/	
48	
49	
50	
51	
52	
53	
54	
55	
56	
57	
50	
58 50	
59	

974	Figure 14. Gyrogonites of genus Chara gyrogonites from the Gánt bauxite cover-sequence.
975	A–I, <i>Chara media</i> (samples G-2.5 and G-6a).), <u>HNHM-PBO 1577–1585.</u> A–G, lateral view;
976	H, apical view; I, basal view. J–O, Chara subcylindrica (samplessample G-6b).), HNHM-
977	PBO 1586–1591. J-M, lateral view; N, apical view; O, basal view. Specimen numbers
978	HNHM-PBO xxxxx-yyyyy will be added for each respective specimen.
979	
980	Figure 15. Frequency distribution of the lengthheight (A), width (B), number of convolutions
981	(C), lengthand height/width ratio (ISI) (D) of the Chara media population (50 gyrogonites
982	measured), from samples G-2.5 and G-6a in the bauxite cover-sequence of the Gánt section.
983	
984	Figure 16. Frequency distribution of the lengthheight (A), width (B), number of convolutions
985	(C), lengthand height/width ratio (ISI) (D) of the Chara subcylindrica population (50
986	gyrogonites measured), from samplessample G-6b in the bauxite cover-sequence of the Gánt
987	section.
988	
989	Figure 17. Stratigraphic distribution of variants of the anagenetic evolutionary-lineage of <u>the</u>
990	species Raskyella peckii.
991	
992	Figure 18. Charophyte Biostratigraphy, age and correlation of the bauxite cover-sequence
993	from <u>of</u> the Gánt section.





Figure 1. A, geographical and geological setting of the study area (after Fodor 2007). B, panoramic view of the studied Gánt section at the Vértes Hills (north-western Hungary).

480x291mm (300 x 300 DPI)





467x589mm (300 x 300 DPI)





Figure 4. Frequency distribution of the height (A), width (B), number of convolutions (C), and height/width ratio (ISI) (D) of the Raskyella peckii var. peckii population (50 gyrogonites measured), from samples G-2.2, G-2.3, and G-2.4 in the bauxite cover-sequence of the Gánt section.

380x246mm (300 x 300 DPI)



Figure 5. Frequency distribution of the height (A), width (B), number of convolutions (C), and height/width ratio (ISI) (D) of the Raskyella peckii var. caliciformis population (50 gyrogonites measured), from sample G-2.5 in the bauxite cover-sequence of the Gánt section.

384x248mm (300 x 300 DPI)



Figure 6. Raskyella peckii vadaszii gyrogonites from the Gánt bauxite cover-sequence (A–S, samples G-6a and G-6b, HNHM-PBO 1517–1534). A–K, lateral view. L–P, apical view; Q, basal view; R–S, inside wall of a gyrogonite showing the crenate undulation of the cellular sutures in contact with the spiral cells and the apical cells.

461x634mm (300 x 300 DPI)



Figure 7. Frequency distribution of the height (A), width (B), number of convolutions (C), and height/width ratio (ISI) (D) of the Raskyella peckii var. vadaszii population (50 gyrogonites measured), from samples G-6a and G-6b in the bauxite cover-sequence of the Gánt section.

384x247mm (300 x 300 DPI)



Figure 8. Gyrogona caelata gyrogonites from the Gánt bauxite cover-sequence. A–G, Gyrogona caelata forma caelata (samples G-2.2, G-2.3, G-2.4, and G-2.5), HNHM-PBO 1535–1553. A–B, D–E, basal view C, detail of the ornamentation pattern; F, basal view; G, apical view. H–K, Gyrogona caelata forma bicincta (samples G-6a and G-6b). H, lateral view; I, detail of the ornamentation pattern; J–K, apical view. L–M, Gyrogona caelata forma monolifera (samples G-2.2, G-2.3, G-2.4, and G-2.5). L, lateral view; M, basal view. N–Q, Gyrogona caelata forma baccata (samples G-2.5, G-6a, and G-6b). N–O, lateral view; P, basal view; Q, apical view. R–V, Gyrogona caelata forma fasciata (samples G-6a and G-6b). R, lateral view; S–T, apical view; U–V, detail of the simple (unipartite) basal plate.

412x681mm (300 x 300 DPI)



Figure 9. Frequency distribution of the height (A), width (B), number of convolutions (C), and height/width ratio (ISI) (D) of the Gyrogona caelata population (50 gyrogonites measured), from samples G-2.2, G-2.3, G-2.4, G-2.5, and G-6a in the bauxite cover-sequence of the Gánt section.

384x248mm (300 x 300 DPI)



Figure 10. Gyrogona tuberosa gyrogonites from the Gánt bauxite cover-sequence (sample G-6b), HNHM-PBO 1554–1563. A–G, lateral view; H, apical view; I–J, basal view.

643x441mm (300 x 300 DPI)





Figure 12. Frequency distribution of the height (A), width (B), number of convolutions (C), and height /width ratio (ISI) (D) of the Psilochara polita population (50 gyrogonites measured), from sample G-6a in the bauxite cover-sequence of the Gánt section.

381x244mm (300 x 300 DPI)



Figure 13. Frequency distribution of the height (A), width (B), number of convolutions (C), and height /width ratio (ISI) (D) of the Nitellopsis (Tectochara) aff. palaeohungarica population (50 gyrogonites measured), from samples G-2.3 and G-2.5 in the bauxite cover-sequence of the Gánt section.

387x252mm (300 x 300 DPI)





Figure 15. Frequency distribution of the height (A), width (B), number of convolutions (C), and height /width ratio (ISI) (D) of the Chara media population (50 gyrogonites measured), from samples G-2.5 and G-6a in the bauxite cover-sequence of the Gánt section.

384x248mm (300 x 300 DPI)



Figure 16. Frequency distribution of the height (A), width (B), number of convolutions (C), and height/width ratio (ISI) (D) of the Chara subcylindrica population (50 gyrogonites measured), from sample G-6b in the bauxite cover-sequence of the Gánt section.

385x248mm (300 x 300 DPI)







413x587mm (300 x 300 DPI)

A g e (Ma)	Epoch	Stage	alcareous nannofossil (Martin, 1971)	ity Geomagnetic polarity time scale ns (Gradstein et al., 2004)		European Charophyte biozonation according to Riveline et al. (1996), modified par Martin-Closas et al. (1999a) and Sanjuan et al. (2014)		n), New proposal of Eocene charophyte) biozonation in Hungary				
			U	Polar	Chro			I		Age		
		z	ND20		C13	Lychnothamnus vectensis						
35-	Ð	NIA	NP20	NP20	NP20		C15	Harrisichara tuberculata superzone	zone	Disease		
E	a t	IABC	NP19		C16	Harrisichara lineata	ber	Biozones				
	_	R	NP18			Gyrogona tuberosa	SL					
40 40 45 50 50	_	IAN			C17	Raskvella vadaszi		Raskyella peckii	P p vadaszii			
			NP17					vadaszii		မ ပ		
	ш	BARTON			C18	Chara friteli	c k i	Raskyella peckii caliciformis	R. p. caliciformis	d n e u		
	• ح_		NP16		C19		b e			r - s e		
	E ™ C	LUTETIAN	NP15		C20	Raskyella pecki	askyella	Raskyella peckii peckii		nt bauxite cove		
	0		NP14				R		R. p. peckii	Gá r		
	ш	A N	NP13		C22	Nitellopsis(Tectochara) thaleri						
	E a r	Y P R E S I	NP12		C23							
			NP11									

Figure 18. Charophyte Biostratigraphy, age and correlation of the bauxite cover-sequence from the Gánt section.

457x532mm (300 x 300 DPI)