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### ABSTRACT

Granitic rocks are ubiquitous worldwide in ancient and active tectonic settings, representing powerful sources of information about the Earth's past and present geodynamic behaviour. Numerous recent milestones fostering our knowledge of granites would have not been possible without a long-lasting, sometimes controversial, discussion on their origin and significance that has taken place over the last two centuries. Here we present a chronological review of how granites have been defined and interpreted in the context of the major theories that have successively governed the history of Earth Science. The main authors, scientific approaches, interpretations, and type-localities that have influenced knowledge about granitic rocks are summarized from the 18th and 19th centuries, when Earth Science was governed by the Neptunism, Plutonism and Uniformitarianism paradigms, to the acceptance of the Plate Tectonics theory and the very end of the magmatism vs. transformism debate in the late 20th century. Some of the most influential scientific advances in Earth Science, such as the invention of the polarizing microscope and the birth of geochemistry, as well as the role of schools of thought in these successive debates, are further discussed. Moreover, we review the recent and ongoing discussions on the mechanisms of magma generation, segregation, ascent and emplacement leading to the formation of granitic batholiths, as well as the observational, analytical, experimental, and numerical modelling approaches currently used for investigating granitic rocks. The history of granite science is classified into different periods of stasis or "normal" science, which were followed by scientific revolutions triggered by a growing number of inconsistencies. Our current understanding of granitic rocks is inevitably influenced by the preceding paradigms and disputes. Consequently, gathering and valuing the chronology, historical milestones, and overall evolution of ideas and theories on what granites are is crucial for the future directions of granite research.

### 1. Introduction

"The competition between paradigms is not the sort of battle that can be resolved by proofs" Thomas S. Khun, 1970. The Structure of Scientific Revolutions

Granitic rocks (Fig. 1) constitute a major proportion of the upper continental crust (~86–88 vol%; Wedepohl, 1969, 1995), and the average geochemical composition of the whole continental crust is often interpreted to be granodioritic (Clarke, 1992; Gao et al., 1998; Cobbing,

2000). Investigations of granitoid rocks have historically provided, and are still providing, important constraints on ancient and active geodynamic processes of the Earth. Our current knowledge of the global timing and secular changes of tectonic processes, the temporal variations in the strength of the lithosphere, and the creation and recycling of the continental crust, have greatly benefited from the study of granitic rocks (e.g., Pupin, 1980; Pearce et al., 1984; Pitcher, 1997a, 1997b; Bonin et al., 2002; Brown, 2013; Moyen et al., 2017; Palin et al., 2020; Gómez-Frutos et al., 2023). Moreover, one of the major mechanisms of heat and mass transfer within the Earth's crust is the transport of magma

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from zones of melting and melt segregation in the lower to middle crust to zones of emplacement in the upper crust (Sandiford et al., 1991; Petford et al., 2000; Bons et al., 2009; Cruden and Weinberg, 2018 and references therein). The origin and significance of granitic rocks have been debated for more than two centuries, and many milestones have been achieved after controversial discussions that divided different schools of thought and generations of scientists. However, despite their importance for laying the foundations and influencing current knowledge of granitic rocks are often addressed separately (Read, 1957; Rudwick, 1962; Porter, 1976; Pitcher, 1997a, 1997b), and have rarely been reviewed in one place (Clarke, 1996; Young, 2018). Furthermore, the logic through which knowledge of granitic rocks was progressively gained within the Neptunism, Plutonism, and Uniformitarianism paradigms has been often overlooked.

According to the Oxford English Dictionary, a 'concept' is "an idea or mental image which corresponds to some distinct entity or class of entities, or to its essential features, or determines the application of a term (especially a predicate), and thus plays a part in the use of reason or language". Here we therefore use the term "granite concept" to refer to how humans have seen and defined granitic rocks as a distinct and meaningful class and how they perceived, interpreted, and understood granitic rocks during different periods of recent history, either by the interpretation of facts, mental impressions, or opinions and beliefs. Considering the historical perspective of this review, as well as the bountiful terminology employed over the ca. 240 yr. time span that it covers, the broad terms "granitic rock", "granitoid rock", or "granite" are used in an interchangeable manner throughout this review to refer to coarse-grained quartz-bearing igneous or igneous-looking rocks (e.g., Marmo, 1967a, 1967b; Pearce et al., 1984; Cobbing, 2000; Chen and Grapes, 2007; Young, 2018). When referring to specific sub-classes of granitic rocks, the non-genetic normative terms and compositions of the subcommission on the Systematics of Igneous Rocks from the International Union of Geological Sciences (IUGS) are used (Streckeisen, 1974; Le Bas and Streckeisen, 1991; Le Maitre et al., 2002). Outdated, obsolete, or archaic terms are solely employed in the context of classical studies or when mentioning rock names in the language they were originally defined.

The term granite, as well as the rock type of that name (Fig. 1), has played a significant role in the history of modern society. The origin of the term granite, from the Latin word granum, is attributed to Andrea Celsapino in 1596 CE, an Italian physician, philosopher, and botanist at the Sapienza-Università di Roma (Italy). The influence of the granite concept on the modern history of human society becomes clear, for example, in the toponomy of different regions worldwide (e.g., New Hampshire, the Granite State in the United States of America (USA), or Aberdeen, the Granite City in Scotland), as well as in its historical and current usage as a building, paving, and sculpting stone (e.g., Langer, 2001). Among other famous monuments, the Amenhotep III statue, now on display at the British Museum (London, United Kingdom) but originally located at the Mut at Karnak temple (Egypt), as well as the Brihadeeswarar Temple building (India) or the Mount Rushmore National Memorial (USA), are made up of granite. The classical literature of the 19th and 20th centuries further provides evidence of the linkage between human cultures and the granite concept, as evidenced for example in Jules Verne's 1864 novel Journey to the Center of the Earth. Even wellknown historical characters, who were not geoscientists per se, speculated about the origin and significance of granitic rocks. The German poet, playwright, novelist, and scientist Johann Wolfgang von Goethe, for example, wrote extensively about the Brocken granite in the Harz Mountains (Germany) and its significance for understanding the formation of the Earth (Baldridge, 1984; Wolf et al., 1989). Other examples are Leonardo da Vinci, who made detailed drawings and observations of granitic formations in the mountains of northern Italy, and Benjamin Franklin, who wrote about the geology of the Appalachian Mountains in the USA and about the granitic rocks therein (e.g., Gortani, 1962; Dean,

2009).

Granitic rocks have drawn the attention of scientists, particularly geologists, for centuries because of their ubiquity in the Earth's crust (Fig. 1), the historical and ongoing debates on their origin and significance, and their association with orogenesis and ore deposit formation, among other processes. Granites are, however, difficult rocks to work with. In the field, geologists have typically encountered problems related to the common lack of discernible lower and upper contacts of granitic plutons, their internal structural and compositional complexity, and the wide variety of textures and accessory minerals that they contain. Granite research is equally challenging in the laboratory due to the wide range of geochemical and isotopic compositions that these rocks have, and the difficulties in experimentally replicating the physicochemical conditions of their formation. The history of the research on granitic rocks is marked by the fact that different investigations into their origin and significance have yielded numerous contradictory results and interpretations (e.g., Backlund, 1938; Miller et al., 1988; Klötzli et al., 2001, 2002; Finger and Clemens, 2002). Nevertheless, such challenges have not discouraged the scientific community from keeping momentum, as demonstrated by the constant development of new analytical techniques and instruments, and by the paradigm shifts that have represented consecutive scientific revolutions (Kuhn, 1970; Bak, 1996) during the long-lasting and controversial history of the granite concept.

The non-genetic German term Massige Gesteine (massive rocks) (Rosenbusch, 1877a, 1896) began to be used in the late 19th century to define crystalline or glassy rocks that, when observed in outcrop, show uniform textures and a massive character that allowed them to be distinguished from metamorphic or sedimentary rocks. Long before that time, however, investigations into and discussions about granitic rocks were already frequent. A primitive origin of granites supported by neptunists (i.e., being the oldest and hardest rocks of the Earth, formed under deep water conditions; Werner, 1787) contrasted the magmatic origin suggested by plutonists (e.g., Hutton, 1794), marking the debate on the origin and significance of granites during the 18th century. Subsequently, other influential works within the uniformitarianism paradigm (e.g., Hutton, 1795a, 1795b; Lyell, 1838; Keilhau, 1843; Darwin, 1845), as well as the invention of the petrographic microscope and the birth of geochemistry (Clarke, 1868; Rosenbusch, 1877a; Iddings, 1890), resulted in the diversification and specialization of research on granites during the 19th century. Ideas about transformism, granitization, and the metasomatic origin of granites later gained popularity due to a growing number of textures, compositions, and field features that could not be explained by the current knowledge of the epoch. With these new ideas, research on granitic rocks abruptly entered a long-lasting debate between the concepts of magmatism vs. transformism (e.g., von Bunsen, 1861; Read, 1957; Marmo, 1967a, 1967b; Pitcher, 1987; Pitcher, 1997a, 1997b; Young, 2018). At the end of the 20th century, after more than two centuries of the emergence, integration, and eventual obsolescence of terminology used to refer to different types of Massige Gesteine, the IUGS Commission on Systematics in Petrology (CSP, founded in 1970) officially proposed the term igneous rock (e.g., Streckeisen, 1974). This term, although semantically genetic (i.e., from the Latin ignis, meaning fire), had non-genetic implications and could be used for both igneous and igneous-looking rocks, irrespective of their genesis (Streckeisen, 1974; Le Bas and Streckeisen, 1991; Le Maitre et al., 2002). Since the proposal of this term, several genetic and non-genetic classifications for granitic rocks have been proposed, and many of them have already been abandoned. Presently, granites sensu stricto are understood to be coarsegrained, crystalline igneous rocks with a modal abundance of quartz between 20 vol.% and 60 vol.% relative to feldspar, and which may include one or more micas, hornblende and many other accessory minerals (Streckeisen, 1974; Le Maitre et al., 2002). As a sub-class of plutonic rocks, granites are presumed to have crystallized at depth from molten material (Hutton, 1794; Walton, 1960; Le Maitre et al., 2002), and references to a possible metasomatic origin of granitic rocks are



**Fig. 1.** Examples of type localities where granitic rocks have played a significant role in the history of Earth Sciences: (a) Outcrop of the Glen Tilt complex (granite dykes with pink colour) at its contact area with the metasedimentary rocks of the Dalradian Supergroup (grey-coloured rocks), Tilt river, Perthshire, Scottish Highlands (described in Hutton, 1794); photograph courtesy of Robert Butler (river width is ~3 m; see detailed outcrop sketch in Fig. 2b). (b) Syenite sill intruding parallel to bedding of Cambrian and lower Ordovician shales at Enerhaugen, Oslo, south Norway (described in Lyell, 1841; Keilhau, 1843); photograph reproduced from Hestmark (2011). (c) Contact between the Cape Granite Suite (light coloured) and the metasediments of the Malmesbury group (dark coloured) at the Sea Point locality, Cape Town, South Africa (described in Masson, 1776; Playfair and Hall, 1813; Darwin, 1845); photograph courtesy of Cape Town (buildings for scale). (d) Typical field aspect of the granites at Sierra Madre de Chiapas, southwestern Mexico (described in von Humboldt, 1822); photograph courtesy Gabriel Serrano López (compass for scale). (e) Granite gneiss outcrops at the southwestern sector of Geographe Bay, Western Australia (described in Baudin, 1974) (cars in a parking lot on the right side for scale). (f) Granitic gneisses at the Aston-Hospitalet Massif, eastern Pyrenees, SW Europe (described in Lacroix, 1896, 1898; Daly, 1898; Adams, 1901); photograph courtesy of Pilar Clariana (field of view is ~30 cm). See locations in Figs. 6, 7. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

almost non-existent in the modern literature. Present controversies regarding granitic rocks are, mostly, focused on specific issues such as melt generation and segregation, magma ascent and emplacement, and post-crystallization evolution. Here we present a chronological review of the mutual interactions between the granite concept and the history of Earth Sciences. The review is structured following the different currents of geological thought that have existed during the past two centuries, with special emphasis on the arguments, explanations, and controversies on the origin and significance of granitic rocks (Figs. 2-4). The history of the scientific knowledge of granites is divided into different paradigm-governed periods according to the ideas of Kuhn (1970) (Fig. 5), including the 18th, 19th, and late 19th to late 20th centuries. The current state of the art of the granite concept is further summarized by reviewing the different genetic and non-genetic classification schemes for granitoid rocks, as well as the observational, analytical, experimental, and numerical modelling approaches that are currently being employed. We also evaluate the role of FAIR (Findable, Accessible, Interoperable, Reusable) initiatives that have emerged in recent years, which have fostered the quantity, and quality, of granite-related data during the Digital Age (Figs. 6, 7).

### 2. The 18th century: Neptunism and Plutonism paradigms

### "I have been particularly anxious about this subject of granite" James Hutton. Theory of the Earth: With Proofs and Illustrations, in Four Parts: Volume 3

The term granite has been used without any quantitative mineralogical connotation for centuries. The Swiss naturalist Horace-Bénédict de Saussure recognized and described, in his Voyages dans les Alpes travel journals, different granitic rocks cropping out in the European Alps (Figs. 6, 7) (de Saussure, 1779, 1786). From Saussure's works, two main concepts that became central to the debate on granites during the following tens of years stand out: the "stratification" of granites, inferred from the fact that some minerals appear aligned in bands, and their nonprimitive origin (i.e., not a primordial part of the Earth), inferred from observations of granitic veins intruding surrounding schists (i.e., Granit Veiné). Five years later, Patrin (1791) used the term Pierre Graphique to refer to a quartz-feldspar rock that he found forming a mountain in the eastern part of Siberia (Figs. 6, 7), and highlighted its resemblance to rocks from other parts of the world. It was only three years later that the Scotsman James Hutton officially published his interpretations on the veined granites hosted in schists and marbles at the type-locality of Glen Tilt in the Central Highlands of Scotland (Fig. 1a, 2b, 6) (Hutton, 1794). Although the basic dynamics of terrestrial processes had already been discussed in his essay Theory of the Earth: or an investigation of the laws observable in the Composition, Dissolution, and Restoration of Land upon the Globe (Hutton, 1785), it was in 1794 when Hutton stated that he found "the most perfect evidence that the granite had been made to break the Alpine strata and invade that country in a fluid state". In this groundbreaking work, which had benefitted from discussions with John Clerk of Eldin and the Duke of Athol (landowner of the Glen Tilt area), and corroboration by the investigations of Sir James Hall in southwest Scotland (Hall, 1794), Hutton concluded that: "Granite, which has been hitherto considered by naturalists as being the original or primitive part of the Earth, is now found to be posterior to the Alpine schists" (Fig. 2b-d). This work, together with the first two parts of Hutton's treatise on the Theory of the Earth with proofs and illustrations in four parts (Hutton, 1795a, 1795b), exceptionally illustrated by Playfair (1802), is believed to have established the basic ideas of Plutonism (Fig. 2b-d). The third part of Hutton's Theory of the Earth (Hutton, 1899) remained lost for 60 yr. and was published after being released to the Geological Society by the Scottish merchant and geologist Leonard Horner in 1856 (e.g., Bonney, 1899). The fourth part is still missing.

Hutton's observations reinforced the ideas that had originally been proposed by the Italian abbot, geologist, and naturalist Anton Lazzaro Moro during his research on volcanic islands, while trying to explain the occurrence of crustaceans and other marine organisms on mountains (Moro, 1740a, 1740b) (Fig. 6). The breakthrough ideas of Moro and Hutton about the Earth's origin and rock-forming processes were incompatible with the dominant school of geology at the end of the 18th century in Freiberg (Saxony), where Abraham Gottlob Werner was the inspector of mines and professor of mining and mineralogy at the Mining Academy. Unlike Hutton, Werner never published the whole of his theory, and it was only by unpublished materials and the lecture notes of his students that his thoughts could be pieced together (Rudwick, 1962; Seddon, 1973; Hallam, 1983; Stone, 2020). Werner's theory was partially stated in his book Kurze Klassifikation und Beschreibung der verschiedenen Gesteinsarten (Werner, 1787), which since its publication constituted the basis of the Neptunism movement. Neptunists advocated, based on previous classifications from other authors (e.g., Lehmann, 1756; Füchsel, 1761; Desmarest, 1774) (Fig. 2a), that rocks were formed as a result of chemical precipitation from cosmic material in the early Earth's oceans (i.e., Allgemeines Gewaesser) in a sequence of five series: (1) Primeval (Urgebirge), (2) Transition (Übergangsgebirge), (3) Secondary or Stratified (Flötz), (4) Alluvial or Tertiary (Aufgeschwemmte) and (5) Volcanic Series. Crystalline rocks, such as granites, gneisses, and metasediments, were attributed by Werner and other Neptunism advocates to the Primitive series and were thus considered the oldest and hardest rocks of the Earth, formed under very deep, calm water conditions. Because no life was thought to have existed during the Primitive series, granite rocks were free of fossils.

The Primitive origin of granites as suggested in Werner's theory clashed with numerous field observations reported during the late 18th and early 19th centuries, as evidenced above with the Saussure's and Hutton's Granit Veiné. Another example were the observations of Marzari-Pencati (1806, 1820), who demonstrated that, in the Fassa Valley (Dolomites), igneous dykes crosscut and thus postdate banded limestones (Predazzite) that in turn record contact metamorphism. Another example is the work of the Prussian zoologist and botanist Peter Simon Pallas during his expeditions in Russia (e.g., Pallas, 1777, 1812; Parker, 1973). Close to the city of Chelyabinsk, Pallas described porphyritic granites belonging to Hercynian batholiths that intruded as large veinlike bodies into metamorphic schists and marbles of the eastern flank of the Urals (Fig. 6). Early in the 19th century, other intrusive granites than those reported by de Saussure (1786) and Patrin (1791) were described outside of Scotland, in the Western Cape Province (South Africa) (Figs. 1c, 6). As pointed out by Master (2009), the Cape Granites (Fig. 1c) played an important role in the Neptunism vs. Plutonism debate during the late 18th and early 19th centuries. Some descriptions of these rocks pointed out their similarity to the granites cropping out in the Alps (e.g., Hamilton, 1778), and many other naturalists and travellers also reported the textural and mineralogical features of the Cape Granites and discussed their intrusive nature and relationships with the surrounding rocks (e.g., Masson, 1776; Anderson, 1778; Sonnerat, 1782; Sparrman, 1785; Barrow, 1801; Degrandpré, 1801).

### 3. The 19th century: "normal" science

"This intrusion the schists all around it does roast,

and when it is absent it alters them most"

The Grizzly Bears Book of the Geological Survey of Scotland (c. 1875)

### 3.1. Uniformitarianism

During his lifetime (1726–1797), James Hutton did not witness the acceptance of his ideas by the scientific community, in part because his theories denied the influence of supernatural forces on the Earth system. However, at the beginning of the 19th century, currents of thought about the origin of granites, veins, and even mountain building started



**Fig. 2.** Representative illustrations from some of the most scientifically influential works on the geological knowledge during the 18th century. (a) Reproduction of an illustration from Lehmann (1756) representing one of the first known geological cross-sections, used by Werner to propose the idea of Neptunism (i.e., rocks belonging to Primitive, Transition, Secondary, Tertiary, and Volcanic series are continuously superimposed) (scale is unknown). (b) Outcrop sketch from Hutton (1794, 1899) of the crosscutting relationships between granite veins and pre-existing schists exposed at Glen Tilt (Scotland) (hammer for scale). (c) Redrawn sketch of a representative cross-section of the Island of Arran (which has a length of ~15 km) from Hutton (1899), showing the relationships of a granite intrusion and the pre-existing sediments and metasediments. (d) Redrawn illustration from James Hutton in the late 18th century and discovered in 1968 by a descendant of John Clerk of Eldin, representing a sequence of sedimentary rocks crosscut by an igneous intrusion in a trench for an artificial canal near Frederick Street in Edinburgh (Craig, 1978) (scale is unknown). See text for further references.

### to drift towards the basis of Plutonism (Hallam, 1983; Master, 2009; Leddra, 2010; Rapprich et al., 2019; Stone, 2020).

From 1800 to 1830, granitic rocks were reported around the globe in numerous studies, travel journeys, and correspondence among scientists with no consensus on their origin and significance (among others, Ramond, 1801; Herrgen, 1802; Breislak, 1811; Hall, 1815; Cortesi, 1819; Bakewell, 1823; Gemmellaro, 1823; Palassou, 1823; Crawford, 1824; Enys, 1833; Cormack and Bruton, 1928) (Figs. 3, 6). The progressive shift of thinking on the origin of granite from the 1830s is illustrated by the lectures given between 1809 and 1833 by Robert Jameson, Regius Professor of Natural History at Edinburgh University, who had previously studied with Werner at Freiberg (Bailey and Tait, 1921; Stone, 2020). It is known that, by the 1830s, Jameson had accepted some of the Huttonian ideas about intrusive granites, as revealed by his students' notes (McCormick, lecture 84, 28 March 1831; in Stone, 2020): "The granite sometimes forms veins in the surrounding mica-slate - and as veins are generally considered newer than the formation in which they occur - the mica slates, it would appear have existed prior to the Granite".

The main people responsible for promoting and extending Hutton's theory on Plutonism (which formerly was, in fact, a theoretical construct; e.g., Kuhn, 1970; Baker, 1998; Rossetter, 2018) were Charles Lyell (1797–1875) and Mary Elizabeth Horner (Rudwick, 1970, 1998; Dott, 1998; Virgili, 2003) (Fig. 3b, e). In the three volumes of the *Principles of Geology*, Lyell revisits the work of Hutton and settles an even more solid paradigm for the rock-forming processes across space and time: the theory of Uniformitarianism (Lyell, 1830, 1832, 1833) (Fig. 5). This theory postulated that the Earth's landscape had been sculpted over the course of an inconceivably long natural history by gradual processes that are visible at the surface and are governed by natural laws. For a summary of different aspects of the research, lectures, theories, letters, and travels of Lyell and Horner, which are out of the scope of this work, the reader is referred to Lyell (1849, 1853, 1881), Bailey and Hartley (1960), Greene (1973), Porter (1976), Dott (1998), Wool (2001) and

Virgili (2007). Concerning granites, their relevance and influence on the novel ideas is clearly recognisable in Lyell's work.

One of the best documented studies on granites carried out by Lyell and Horner is their 1837 trip to Norway and their observations, together with Prof. Baltazar M. Keilhau, in the vicinity of Christiania (present-day Oslo) (e.g., Holtedahl, 1963; Marmo, 1967a, 1967b; Hestmark, 2011) (Figs. 1b, 3b, 6). They documented contact zones between sedimentary rocks and large intrusive bodies of granite and syenite, revisiting the phenomenon of metamorphism that they had already coined in the Principles of Geology (Lyell, 1833) and citing examples from the Highlands of Scotland (Fig. 1a), the Alps, Cornwall, and Table Mountain at the Cape of Good Hope (Figs. 1c, 6). The observations and interpretations in Oslo confirmed Lyell's theories and were key for the development of the well-known book *Elements of Geology* (Lyell, 1838), which was presented to the research community three years later (Lyell, 1841) and became an influential work in Lyell's lectures in North America between 1841 and 1853 (Dott Jr., 1996; Dott, 1998). According to Lyell's view, rocks could be classified in four great classes (Lyell, 1838 - Part I, Chapter I): Aqueous (i.e., sedimentary), Volcanic, Plutonic, and Metamorphic (Fig. 3e). Regarding Plutonic rocks, Lyell wrote: "The granites have been formed at great depths in the earth and have cooled and crystallized slowly under enormous pressure where the contained gases could not expand. [...] large masses of granite are found to send forth dikes and veins into the contiguous strata, very much in the same way as lava and volcanic matter penetrate aqueous deposits" (Fig. 3e).

On the other hand, during the same trip to Christiania (Fig. 1b, 6), Baltazar M. Keilhau interpreted a gradual and complete passage of granitic rocks into stratified rocks with fossils, and described gradual changes from granite to a "primitive" gneiss (Hestmark, 2011). With his ideas, Keilhau was probably one of the pioneers of the transformism (i. e., granitization) theory that divided many geologists some tens of years later. In fact, widely used terms during the 19th century, such as *transmutation* and *granitification*, were already coined by Keilhau, who probably also was the first geologist noting the *room problem* (see below) in the emplacement of igneous rocks (Keilhau, 1828, 1838, 1843; Marmo, 1967a).

Another important source of information for Lyell during the development of his theories on the origin and significance of granite were the observations of Charles Robert Darwin (1809-1882), who in turn benefited from Lyell's theories during his investigations aboard the HMS Beagle (e.g., Burchfield, 1974; Rosen, 1982; Herbert, 1991, 2005; Secord, 1991; Master, 2012) (Fig. 6). Among Darwin's numerous geological observations, those made at Cape Town in 1836 and in South America in 1831-1836 stand out (Fig. 6). In Cape Town, Darwin revisited previous interpretations of the Green Point Contact (also known as the Sea Point Contact; Fig. 1c), (Masson, 1776; Anderson, 1778; Sonnerat, 1782; Playfair and Hall, 1813) and interpreted the schist enclaves lying within the granitic mass as "interconnected pendant slivers of infolded suprapositional schists that were intruded parallel to their schistosity by thin fingers of granite, and then eroded to reveal apparently detached schist fragments in the granite" (Master, 2009, 2012). In South America (Fig. 3c), Darwin also made abundant contributions to the understanding of granitic rocks as demonstrated by the number of times (172) that the word granite appears in his Geological Observations on South America (Darwin, 1846). During his visits along the eastern coast of Brazil (Fig. 6), Darwin wrote (Darwin, 1845, 1846): "Along the whole coast of Brazil, for a length of at least 2000 miles, and certainly for a considerable space inland, wherever solid rock occurs, it belongs to a granitic formation. The circumstance of this enormous area being constituted of materials which most geologists believe to have been crystallized when heated under pressure, gives rise to many curious reflections". With curious reflections Darwin was referring in part to the room or space problem, first noted by Keilhau (1838): "Was this effect produced beneath the depths of a profound ocean? or did a covering of strata formerly extend over it, which has since been removed?". Darwin's observations in Chile also raised some controversial points for scientific knowledge and for Darwin himself, e.g., about the structure and crosscutting relationships in the red granites from the Chilean Cordillera (Darwin, 1846): "There are gigantic mountain-like masses of red granite, which have been injected whilst liquefied, and which, nevertheless, display in parts a decidedly laminar structure" (Fig. 3c).

Darwin was not, however, a pioneer in the study of granites in Latin America, which always had been seen as a valuable building stone. During the Incan Empire, granitic rocks were exploited and worked from the Cañar mountains, in the present-day province of Cuenca (Ecuador) (Velasco, 1844), and granite was also used to build most of the city of Machu Picchu in the Vilcabamba range (south Peru) (Delgadillo, 2013). The first known geological reports on the granites of Latin America are those from the German explorer Alexander von Humboldt in Mexico (von Humboldt, 1822, 1825), where he described pics granitiques (i.e., granitic summits) in the Sierra Madre of western Mexico (Figs. 1d, 6). Humboldt also showed large granitic bodies and their crosscutting relationships with sedimentary and volcanic rocks in many representative cross-sections of the Earth's crust. The von Humboldt, 1841 crosssection (Fig. 3f) provides an excellent synthesis of the geological knowledge of that epoch, where granites are depicted as large intrusive masses with different ages relative to the surrounding strata. The French naturalist Alcide Dessalines d'Orbigny also provided very detailed descriptions of the geology of South America during his trip from 1826 to 1834 (Orbigny, 1834).

Other detailed maps and descriptions carried out before Darwin's publications were those from the Salesian Juan de Velasco and his expedition colleagues (e.g., Velasco, 1844; Garrido, 2020) (Fig. 6). Another important contribution to the knowledge of granitic rocks in South America during the 19th century (postdating the work of Darwin) is considered to be the travel chronicle of Jean Baptiste Boussingault, Joaquin Acosta, and Francois Desiree Roulin in Nueva Granada (a Spanish Viceroyalty that extended over present-day Colombia, Venezuela, Ecuador, Panama, and Guyana) (Boussingault, 1849). In this work, granite bodies from different areas are described in detail, including their structure, mineralogy, and relationship with the

surrounding rocks. In Ecuador, the Professor of Geology and Mineralogy at the National Polytechnic School of Quito, Franz Theodor Wolf, also made important contributions to the knowledge of the geology of the country, including investigations on granitic rocks (Wolf, 1892). Besides numerous descriptions of the field occurrence, mineralogy, and structure of granitic rocks across the country, it becomes clear that Wolf was familiar with the ongoing debate on the origin of granites: "geologists do not agree (...) it is one of the most difficult questions in geology (...) we do not claim that all plutonic rocks have the same origin of an igneous-fluid magma, or that they are in their primitive state, because I admit for many a very deep metamorphosis" (Wolf, 1892).

The Australian continent also played an important role on the origin and development of the knowledge of granites during the 19th century (Figs. 6, 7). Darwin visited Australia in 1836 and, although his investigations on granites there were limited, he reported the existence of granites in the Vale of Clwydd (New South Wales) and on the Vancouver Peninsula (Western Australia) (Darwin, 1845; Armstrong, 1985; Nicholas and Nicholas, 2008) (Fig. 6). The geology of Australia had been previously described by the Frenchman Nicolas Baudin and his team during expeditions in 1800-1803 (Mayer, 2009), and later on by other several exploration projects that focused on the biodiversity, geography and geology of the region (e.g., Péron et al., 1807; King, 1827; Lesson and Duperrey, 1830). The mineralogist of the Baudin expedition, Louis Depuch, was responsible for most of the geological observations that can be found within the travel chronicles (Baudin, 1974). Among the observations, similarities between the granites in the Baie du Géographe (now Geographe Bay; Western Australia; Fig. 1e) and those described by de Saussure (1779) in the European Alps are remarkable (Baudin, 1974; Mayer, 2009): "it is impossible not to be convinced here of the fact, which Mr. de Saussure was the first to recognize and which several naturalists still contest, that granites are apt to stratify". The Baudin expedition also reported and described granite outcrops in north Tasmania, the Blue Mountains in New South Wales, and the southern Australian Coast. According to their observations and the dominant, Wernerian, rock classification scheme of the epoch, they suggested a mostly primitive origin for this whole part of the continent (Werner, 1787; Baudin, 1974; Mayer, 2009).

The second half of the 19th century saw both a diversification and specialization in the study of granites. The Plutonism paradigm proposed by Moro (1740a, 1740b) and Hutton (1795a, 1795b), popularized and embedded within Uniformitarianism by Lyell (1833, 1838), and supported by the observations and descriptions of Darwin (1845, 1846), among others, gave rise to a period of paradigm-governed, "normal" science (Kuhn, 1970) (Fig. 5). This period, comparable to the states of stasis according to Bak's (1996) model of self-organised criticality, was characterized by the accumulation of inconsistencies and the emergence of new ideas leading to a collapse and to the occurrence of a scientific revolution (Kuhn, 1970), or an "avalanche" in the sandpile analogy (Bak, 1996) (Fig. 5). The study of granites moved from being purely observational and based on reports of their existence and mode of occurrence in different parts of the world, to become somehow reductionist. Thus, many scientists started to carry out more detailed investigations of granites focused on, for instance: the development of foliations and other structural features (Sharpe, 1852; Mackintosh, 1868; Ormerod, 1869; McMahon, 1887), their mineralogy and petrology (Haughton, 1858, 1859, 1864; Cooke, 1866; Hull, 1874) (Fig. 4); their formation mechanisms and effects on their host rocks (Ramsay, 1841; Von Buch, 1845; Duke, 1853; Bonney, 1888; Reade, 1889), their mafic and felsic enclaves (Phillips, 1880, 1882; McCormick, 1886), and their relationship with ore deposits and mining districts (Ansted, 1849; Hunt, 1849; Phillips, 1849; Pissis, 1850; Domeyko, 1855).

The knowledge of granites was probably further fostered by the emergence of several large-scale geological maps. Although the first known maps representing geology were drawn by the Count Luigi Ferdinando Marsili (Marsili, 1717; Romano et al., 2016), William Smith (Smith, 1815), and Carlos de Gimbernat (de Gimbernat, 1806, 1808; see a review in Vila, 2016), those produced during the mid-19th century allowed larger-scale comparisons between different regions of the world and provided a global overview of the occurrence of granitic rocks and their distribution. For example, the geological maps of Europe (Dumont, 1857) and the United States (Marcou, 1853) provided very detailed information about the classes of granitic rocks in these territories, in all

cases accompanied by detailed mineralogical information. The geological maps of the globe from Boué (1844) and Marcou (1861), described in detail by Oldroyd (2014), also represented the worldwide distribution of granites, which were in some cases lumped together with different metamorphic schist series, and in other cases distinguished as *granite* in the legend.



**Fig. 3.** Representative illustrations from some of the most scientifically influential works on scientific understanding of granitic rocks during the 19th century. (a) Emplacement dynamics of a granite according to Scrope (1825) (scale is unknown). (b) Granite veins intruding Silurian strata and gneiss in the vicinity of Christiania, Norway (Lyell, 1838) (scale is unknown). (c) Representative cross-section from the town of Copiapo to the western base of the Andean Main Cordillera (Darwin, 1846) (cross section length is ~225 km); pink and red colours represent granites and andesites; blues correspond to porphyries and conglomerates; yellows represent gypsum or gypsiferous formations; dark-green refers to feldspathic clays and slates. (d) Observations on the structural relationships between granite, schist, and mixture zones in the Torry Beach area (Aberdeen, Scotland) from Harris (1888) (mapped area is ~67 m<sup>2</sup>). (e) Idealized illustration of the origin and mode of occurrence of different rock types published as the frontispiece from Charles Lyell's *Principles of Geology* (second American editior; Lyell, 1857) (scale is unknown). (f) Representative cross-section summarizing the mid-19th century understanding of the structure Earth's crust and its major constituents, including granites (von Humboldt, 1841) (scale is unknown). See text for further references. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

### 3.2. The birth of geochemistry

The last decades of the 19th century saw an increasing interest in granites through the emerging discipline of Geochemistry. This term was coined by the chemist Christian Friedrich Schönbein in 1838 in Basel (Switzerland) (e.g., Fairbridge, 1998) based on the work of Lunn et al. (1835). Noteworthy is the work on granite geochemistry, and more specifically on fluid and melt inclusions, by the British geologist Henry Clifton Sorby, to whom the manufacture of the first rock thin sections is attributed. Through the comparison of the properties of experimentallygrown and natural crystals (the latter belonging to metamorphic rocks, quartz veins, and igneous rocks), Sorby (1858) demonstrated that the study of cavities containing fluid (i.e., fluid and melt inclusions) could be used to determine whether rocks in which they are contained were deposited from aqueous solutions or crystallized from a melt (e.g., Correns, 2003; Bodnar, 2018). Significantly earlier than Sorby's work, the investigations of Scrope (1825) stand out (Fig. 3a). Through the study of volcanoes and volcanic products in south Italy (Sicily, and the Pontine and Aeolian Archipelagos), Central France (Auvergne, the Velay, and Vivarais areas), and north Germany (Eifel volcanic region), Scrope (1825) proposed that a wide range of igneous rocks could originate from the differentiation of a primary magma by "some chemical process in an internal reservoir" (see below). Moreover, this author further provided other speculations on the viscosity of magmas and their relationship with crystal and vapour composition. Years later, the chemist Robert von Bunsen also investigated the origin of granites and their water content. Through his investigations on the volcanic rocks of Iceland, von Bunsen (1851) showed evidence of the coexistence between felsic and mafic lavas. As Bunsen considered it impossible that potassic granites could be formed from the same primary magma as basalt, he proposed the existence of two primary magmas, basic and acid. Years later, Bunsen's reflections would also become part of the debate on the origin of granite from a geochemical point of view (von Bunsen, 1851, 1861) (see below).

The geochemical research on granites of the British metallurgist, geologist, and mining engineer John Arthur Phillips during the last decades of the 19th century was also remarkable. Phillips published several results of granite geochemical analyses, among other rock types and soils, and became one of the first scientists, together with H. C. Sorby, to use the polarizing microscope (Phillips, 1849, 1871, 1873, 1880, 1882). This instrument had been invented by W. H. Fox Talbot in 1834 by making use of William Nicol's single-vision 1828 calcite prism, and the work of Dana (1877) and Rosenbusch (1877a, 1877b) contributed to its popularity (Fig. 4a) (see below). The most complete wholerock geochemical analyses published in the 19th century that we have been able to find are those by Phillips (1880), which included data for the Silica, Alumina, Phosphoric anhydride, Titanic, Ferric oxide, Ferrous oxide, Manganous oxide, Lime, Magnesia, Potassa, and Soda contents of granites cropping out in several quarries in England, Scotland, and Ireland. Other 19th century examples of whole-rock geochemical analyses of granites are the studies of the microscopic structure and chemical composition of Irish granites (Hull, 1874), as well as the structural, mineralogical, and geochemical investigations of granite quarries throughout England (Harris, 1888). On the other hand, the first known geochemical analysis of specific granite-forming minerals, namely feldspar, black mica, and beryl, were those presented to the Royal Irish Academy by Samuel Haughton, professor of geology at Trinity College (Dublin) (Haughton, 1853, 1855, 1862). In the United States, since the establishment of the United States Geological Survey (USGS) in 1879, several projects also reported the properties and the major element compositions of granites (among other rock types) and their mineral assemblages throughout the country, with Frank Wigglesworth Clarke as the chief chemist (e.g., Clarke, 1868, 1887; Clarke and Chatard, 1884; Kinahan, 1887; Van Hise, 1890; Turner, 1899; Mathews, 1900).

Several other works described granitic rocks around the world

during the last decades of the 19th century. Although it would be impossible to address all of these in this review, those focused on the magmatism vs. transformism debate, are referred to below. Notable is also the pioneering work of McMahon (1887) on the Gneissose-Granite of the Himalayas, the summary of the geology of southern India by Newbold (1850), and the geological interpretations and petrographical descriptions of granitoid rocks made by G. F. Scott Elliot along the Nile River summarized by Raisin (1893) (Fig. 6).

### 4. From the late 19th century, part I: global flood of knowledge

"The "facts" of today are the hypotheses of yesterday" Reginald Aldworth Daly, 1914. Igneous Rocks and Their Origi

With the development of geochemistry, research on granitic rocks shifted from being purely observational to analytical. The observational aspect, however, also flourished thanks to the emergence and optimisation of the polarizing microscope and the concomitant increase of petrographical studies during the late 19th and early 20th centuries (Fig. 4). As a result of the new ideas on the origin and significance of granites and other igneous rocks, the first problems with their classification became central to discussions among scientists worldwide. Until that moment, the term *granite* had been used to refer to any kind of massive crystalline rock with a broadly felsic composition. However, that broad use was no longer acceptable from the late 19th century, since interpretations of the origin of granitic rocks were directly correlated with their textural properties and chemical compositions (Fig. 4be).

Among many other works, two textbooks on petrography are believed to have changed the course of the study of igneous rocks (e.g., Cross, 1902; Pitcher, 1997a, 1997b): *Lehrbuch der Petrographie* (Zirkel, 1866) and *Die mikroskopische Physiographie der massigen Gesteine* (Rosenbusch, 1877a), with the latter considered to contain the first published global classification of igneous rocks. In this classification, several sub-types of granites and syenites were proposed in addition to the definition of granite *sensu stricto*: a rock containing quartz and orthoclase, small amounts of plagioclase (oligoclase or albite), and micas and/or amphibole. For granitic rock containing biotite but no muscovite or amphibole, Rosenbusch (1877a) used the expression "granite" (e.g., Marmo, 1967a, 1967b).

Several other references stand out for furthering the knowledge of granite petrology during the late 19th and early 20th centuries. A good example is the textbook of Cossa (1881) (Fig. 4b), which influenced the Italian school of mineralogy and petrology by providing instructions for thin section preparation and the use of the polarizing microscope, as well as a complete set of colour plates representing different graniteforming minerals and textures. Furthermore, the work of Teall (1888) (Fig. 4c, d), from the British school (understanding school as a group of individuals with a common scientific tradition, thought, and/or views about a given subject), also provided numerous petrographic descriptions, geochemical analyses, and raised valuable points about the classification of igneous rocks, which he did not attempt to apply into his textbook: "the more rocks are studied the less they seem to me to adapt themselves to any classification at all" (Teall, 1888) (Fig. 4c, d). In fact, these and many other schools of thought have played a significant role on the evolution of our knowledge of granitic rocks by enhancing the magnitude of scientific revolutions. Instead of convincing individual scholars, a set of inconsistencies, arguments, and alternative explanations may eventually become convincing enough so that a whole school topples at once and completely changes its view. In these scenarios, paradigm shifts may be harder to accomplish, but when they happen, they are larger. The concept of schools of thought may be compared to the assortative mating phenomenon of Human Evolution theory (e.g., Bons et al., 2019), i.e., the innate tendency to mate and have children with like persons, which creates regions with particular (genetic) traits and sharp boundaries between them. In science, these regions may be

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**Fig. 4.** (a) One of the first known schematic drawings of a polarizing microscope specifically designed for the study of rock thin sections (H. Rosenbusch model, *Serial Number 131*) (Dana, 1877; after Rosenbusch, 1873, and Zirkel, 1873). (b-e) Microscopic illustrations of granitoid textures published in some of the most influential works on igneous petrography of the late 19th century. (b) Cossa (1881), Tavola IV, Fig. 5: *thin section of a diorite in which chlorite (yellow) sheets are observed to be normally displaced at the flake plane* (light grey colours correspond to quartz and feldspar); magnification: 180 (location not provided). (c) Teall (1888), plate 35: *quartz (white), feldspar (grey), dark mica (biotite, yellow), chlorite (green), and apatite (brown) – Shap biotite-granite (granitite) (Devonian) intrusive in Ordovician strata (Cumbria, NW England); magnified 30 diameters.* (d) Teall (1888), plate 42: *feldspar (dark grey), white mica (brown and yellow), and quartz (light grey) – Pre-Carboniferous (Upper Ordovician) schistose Wicklow granite intrusive in Lower Palaeozoic strata (Leinster, Ireland); magnified 25 diameters.* (e) Rosenbusch (1873), Tafel VII, Fig. 40: Biotite-granite close to Schlierbach (Heidelberg); *quartz and feldspar in white, biotite in green and brown* (magnification not provided). Note that plate descriptions and indications on magnification in this caption are those originally published by the authors. See text for further references. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

understood as schools of thought and, yet not helpful for evolution to reach optimum fitness, it is a protective mechanism for preserving genomes (or, in science, common scientific traditions, thoughts, and views about a given subject).

One year later after Teall's (1888) work, another textbook published by the director of the Geological Survey of France, August Michel-Lévy, attempted to provide a new classification scheme for igneous rocks (Michel-Lévy, 1889). This new classification, the first of many following versions and revisions that would be published in the later decades, was based on previous observations and textural interpretations (Michel-Lévy, 1878a, 1878b, 1889; Fouqué and Michel-Lévy, 1879, 1882), as well as on the initial proposal of Rosenbusch (1877a, 1877b). Concerning granites, Michel-Lévy (1889) proposed a classification for the Famille des Granites based on their depth of emplacement, including: Profondeur (granites, granulites, and granites à amphibole), Filons (granulites and microgranulites), and Épanchements (among others, microgranulites, micropegmatites, porphyres globulaires) (see Section 5). The same classification scheme (i.e., deep, vein-type, and eruptive rocks) was also followed for the Famille des Syenites, the terminology of which was reduced and significantly simplified by Michel-Lévy (1889) after Rosenbusch (1877a, 1877b).

For the Russian school, the work of Feodor Yulievich Levinson-Lessing, Professor at the University of St. Petersburg, was particularly influential. Among other contributions, *Petrographisches Lexicon (Petrographic Lexicon)* (Levinson-Lessing, 1893a, 1895) and Tablitzy dlya Mikroskopicheskikh Opredeleni Porodoobraznykh Mineralov (LevinsonLessing, 1893b; translated into English by Cole and Gregory, 1893) provided Russian scientists with new techniques that were emerging in western Europe for the study and interpretation of rocks under the microscope, including granites and granite-forming minerals.

In the United States, the work of Joseph Paxon Iddings was extremely influential within different departments and faculties of geology and petrology. Iddings spent two years (1879-1880) at the University of Heidelberg with Prof. Rosenbusch conducting petrographic investigations that resulted in several publications on the significance, formation and classification of granitoids and other igneous rocks (e.g., Hague and Iddings, 1885; Rosenbusch and Iddings, 1888; Iddings, 1890, 1892a, 1892b, 1893a, 1893b). Among his works, Iddings (1893a, 1893b, 1898a, 1898b) presented the largest dataset ever published of geochemical analyses of igneous rocks from the Los Andes cordillera, the Yellowstone National Park, and Christiania region, and included a large set of compositional plots. Furthermore, The Origin of Igneous Rocks (Iddings, 1892b) became as influential as it was controversial for the scientific community since, besides providing new ideas on the origin and classification of igneous rocks, it questioned the assumptions and interpretations of granites that had previously been made by Scrope (1825), Darwin (1844), von Bunsen (1851), or Durocher (1859) (see Section 5).

The large amount of observational, analytical, and (embryonic) experimental data on igneous rocks, and specifically on granites, gave rise to several classification issues. For example, confusion stemmed from the inclusion of a wide range of granite-looking rocks within the

general term granite, but this term itself was not yet properly restricted (e.g., granites sensu stricto, granite-porphyries, granophyres, quartz porphyries, pitchstones, rhyolites, feldspar-bearing high-grade micaschists, and migmatites). The subclassification of these rocks according to their relative age or area of occurrence further added more layers of complexity to this problem (e.g., Johnstone, 1887; Harker, 1896; Daly, 1910). The emergence of petrography as a subfield of petrology also produced an increase in the detail of the investigations on granites and thus a magnification of their differences, leading to complex subclassifications often based on very slight changes in the modal abundance or texture of mineral assemblages (e.g., Michel-Lévy, 1889; Cross, 1898). Further confusion was related to the different approaches to the classification of granites proposed, adapted, and modified by petrologists from different schools in a brief period of time, among others: (1) the mode of occurrence (e.g., Dikes, Necks or Plugs, Sills, Laccoliths, Bysmaliths, and Bathyliths; Chamberlin and Salisbury, 1904), (2) their timing relative to orogenesis (e.g., Synkinematic, Late-kinematic, Post-Kinematic; Sederholm, 1891), (3) their petrogenesis and chemical and mineralogical features (e.g., the CIPW System or the Persalane, Dosalane, Salfemane, Dofemane, and Perfemane rock groups; Iddings, 1898b; Cross et al., 1902; Lewinson-Lessing, 1911; Young, 1999), or (4) a combination of the above (e.g., Daly, 1905). The minutes and publications derived from meetings of the American Academy of Arts and Sciences, the Royal Society of Edinburgh, and the Mineralogical Society of Great Britain and Ireland are illustrative of these developements. For more information about these and other historical classifications of igneous rocks, some of which still cause debate among scientists, the reader is referred to Section 6.1 and to the exceptional reviews of Cross (1898), Iddings (1898b), Tomkeieff (1939), Barbarin (1990, 1999), and Pitcher (1997a, 1997b). Regarding the magmatic, metamorphic, or metasomatic origin of granites (e.g., Walton, 1960; Marmo, 1967a, 1967b; Kresten, 1988; Clarke, 1996), this issue is addressed in more detail in the following section.

### 5. From the late 19th century, part II: magmatism vs. transformism

"Are there Granites and Granites after all?"

Herbert Harold Read, 1947. Granites and Granites. Geological Society of America Memoir 28: Origin of Granite

The research on granitic rocks saw its most polarized controversy from the late 19th to the late 20th centuries. As a result of the growing number of inconsistencies derived from field observations, structural interpretations, and geochemical investigations (named anomalies in Fig. 5), chemists, physicists, and geologists were divided among magmatists and transformists. The debate on whether granites are of magmatic, metamorphic, or metasomatic origin, and the related structural and petrogenetical implications of each potential hypothesis, governed more than one century of widespread debate in almost all the geological societies, schools of thought, and geological surveys. Given the significant number of reviews that have already addressed this granite controversy (Walton, 1960; Marmo, 1967a, 1967b; Pitcher, 1987; Clarke, 1996; Pitcher, 1997a, 1997b; Brown, 2013), we focus the present section on two of the main debated issues: the space or room problem (e.g., Keilhau, 1838), and the origin of granitic melts (e.g., Tuttle and Bowen, 1958).

About the room problem, Read (1948) wrote: "it is the room problem that lies at the heart of the granite problem, and the room problem is a matter to be dealt with by field geology". The room problem, in fact, dates back to 1837 when Charles Lyell, Marie Elisabeth Lyell, and Baltazar M. Keilhau visited Norway and provided contrasting interpretations on the origin of the granites cropping out close to the city of Oslo (e.g., Holtedahl, 1963; Marmo, 1967a, 1967b; Hestmark, 2011) (see Section 3; Figs. 1b, 6). At that time, the idea of magmatism implied that the space needed for the formation of a body of granite was exactly the same of that occupied by

the granite itself once formed. In contrast to this, room was not needed for the formation of granites according to the transformist views, which suggested that granites were formed by the replacement or transformation in the solid state of pre-existing rocks (Keilhau, 1838). The apparently unattainable dimensions required for the formation of granites as explained by magmatists led the transformists to gain traction in the late 19th century, and a wide range of arguments invoking granitization were reported to explain the formation of granitic rocks worldwide. Concepts such as pseudomorphism (variations in mineralogy, texture or composition of a granite simulating primary sedimentary or volcanic structures and/or mineral phases; Bowes, 1953; Stringham, 1953), gradation and continuity (the process by which a preexisting rock passes gradationally into granite; Anderson, 1937; Goodspeed, 1948; Eckelmann and Poldervaart, 1957), or relict stratigraphy (the occurrence of pre-existing layers within a granite that maintain a relative stratigraphic position; Misch, 1949; Pitcher, 1952) were used as arguments for a metasomatic or metamorphic origin of granitoid rocks. The formation of rapakivi granites through replacement of red sandstones in the Fennoscandian Peninsula suggested by Backlund (1938), or the formation of gabbro and diorite as mafic fronts resulting from the expulsion of matter during granitization of metagreywackes (Reynolds, 1943, 1947), are good examples of the aforementioned concepts.

However, years before granitization had gained popularity and was eventually refuted, several geologists attempted to solve the *room* problem from a magmatic point of view by means of *assimilation*. Within this framework, granites from Oslo were again the source for the interpretations of Kjerulf (1855, 1879). This author agreed with previous interpretations of Lyell (1838, 1841) regarding a plutonic origin for the Oslo granites despite acknowledging the *room* problem pointed out by Keilhau (1838). Kjerulf (1879) suggested that the space needed for the emplacement of granites was progressively gained by the assimilation of sedimentary rocks, rather than created through forceful intrusion by displacing the surrounding strata. Other contemporaneous influential petrologists from the French and Finnish schools further developed the ideas of assimilation and anatexis (i.e., partial melting of rocks) (e.g., Michel-Lévy, 1888, 1893; Sederholm, 1897, 1907; Lacroix, 1898).

Authors from the French school based their interpretations on pioneering investigations of the mid-19th century that addressed the origin of granites from an extreme transformist perspective. For example, Boué (1824) first proposed igneous liquefaction as a process for turning metamorphic schists into igneous rocks, and the French chemist Sainte-Claire Deville first suggested the existence of agents minéralisateurs (i.e., mineralizing agents) that chemically transform the rocks through which they diffuse (see Holmes, 1945). Virlet D'Aoust (1844a, 1844b) coined the term imbibition, a precursor of granitization, to describe the process by which igneous materials could soak into metasedimentary rocks further transforming them into granites. Later, investigations by Delesse (1861) further provided insights for the French school of transformism: "plutonic rocks are formed from metamorphic rocks; they represent the extreme term of general metamorphism; they are the effect rather than the cause of metamorphism". This author further suggested that granitized plutonic rocks may become so mobile that they could be squeezed towards higher levels of the crust to form intrusive granites.

Occupying the middle ground between the endmembers of extreme magmatism and transformism, Michel-Lévy (1888, 1893) suggested that intrusive granite masses pervasively altered their surrounding strata, promoting the formation of more granite rocks. According to this interpretation, which was based on the ideas of Virlet D'Aoust (1844b), the upwelling rate of a given granitic intrusion was determined by the rate of absorption of the roof and walls of the reservoir, where granitic material assimilates the surrounding rocks that progressively sink and are transformed into granite gneiss, gneissose granite and, finally, granite (Michel-Lévy, 1888, 1893). Furthermore, the investigations of Lacroix (1898) in the granitic rocks from the Pyrenees (Fig. 1f) emphasized the contact zones with the surrounding schists. From field and mineralogical observations along the Garonne River in the HauteAriège, Lacroix obtained strong evidence regarding the occurrence of granitic intrusions and their relationship to the assimilation and metamorphism of surrounding rocks. This author attributed strong feldspatisation of metamorphic aureoles to the migration of a *granite substance* into the country rock, in the presence of mineralizing agents, as well as the possible removal of certain elements in the exomorphic contact zone under the same intrusion conditions (Lacroix, 1896, 1898; Daly, 1898). From this work, the Pyrenees turned out to be the flagship for the *transformist* ideas of the French school, being defined as a district in which remarkable rock transformations could be clearly seen. A geological excursion to the Pyrenees led by Prof. Lacroix within the framework of the Eighth International Geological Congress (Paris, year 1900) (see summary in Adams, 1901) further fostered the discussion on the origin of the Pyrenean granites, which remained open until the late 20th century (e.g., Guitard, 1960, 1970; Autran et al., 1970).

In Finland, the spread of transformism was mainly lead by the head of the Finnish Geological Survey, Jakob Johannes Sederholm (Sederholm, 1907, 1911, 1926). From observations on Precambrian rocks of southern Finland, Sederholm encountered difficulties in explaining different outcrop features in terms of classical magmatism, which led him to introduce the terms palingenesis (i.e., rebirth) and anatexis (i.e., to melt down) to account for peculiar mixed rocks, migmatites (Sederholm, 1907): "because the mixture of different constituents is the characteristic feature of all these different rocks, I would suggest for them the name migmatites". The term palingenesis was also used by Sederholm (1907) to explain the inception of the granitization phenomenon, where the (re-)melting of older granites formed granite melts that were later on intruded at higher levels while granitizing their adjacent rocks, thereby not requiring room. Sederholm (1926) also defined migmatites as an intermediate rock type between igneous and metamorphic end members, which were commonly found in migmatitic terrains together with other rock varieties such as arterites, agmatites, eruptive breccias, bandedgneisses, ptygmatitic migmatites, and diktyonites (see comparisons between Finnish and Scottish rocks in Read, 1925).

The interpretation by French and Finnish schools regarding the formation of granites and their relationship with migmatites did not necessitate a choice between either magmatic emplacement through forceful intrusion or extreme transformism by mysterious replacement reactions driven by fluids. A coupled process under high-temperature conditions, where minor intrusions of granitic magma caused the transformation of surrounding strata into more granitic rocks, was more acceptable given the knowledge of that epoch, and furthermore reduced the *room* problem by several orders of magnitude. The growing inconsistencies between field observations (Michel-Lévy, 1888; Lacroix, 1898; Read, 1925; Sederholm, 1926) and the tenets of magmatism questioned the theories involving forceful intrusions as the origin of granitic rocks. "*Granites are very big things, not hand specimens*", noted Read (1948) several years later, when introducing the *room* problem in the well-known review *Granites and Granites*.

Despite maintaining traditional views on magmatism, some geologists were forced to acknowledge the existence and importance of metasomatic granites during the early 20th century. This was the case of the Finnish geologist Pentti Eskola. About Sederholm, Eskola wrote (see Eskola's prologue in Sederholm, 1967): "one may ask how he, a pupil of Harry Rosenbusch, the extreme magmatist, became the announcer of granitization". Sederholm's (1891) classification of granitic rocks (see Section 4) according to their time of formation relative to orogenesis was, nevertheless, adapted by Eskola (1932). In this classification, Groups I, II, and III of Sederholm (1891) were defined in Eskola (1932) and later on refined in Eskola (1955) as: (I) Synkinematic granites, of metasomatic origin with gneissose texture, which are associated with greatly differentiated complexes varying through all calci-alkalic members from peridotites to granites, (II) Late-kinematic granites, of metasomatic origin with aplitic texture, which were formed at a later stage of the orogeny and thus crosscut the country rock and some of its deformation structures, and (III) Post-kinematic granites, of magmatic origin and

completely devoid of gneissosity. This latter type of granite, often characterized in Finland by rapakivi textures, was the only *true* magmatic granite in Eskola's opinion (Eskola, 1955): "*the rapakivi plutons are the purest magmatic granites in our country. From them we can follow all gradations to metasomatic granites*" (see Marmo, 1967a, 1967b; Pitcher, 1997a, 1997b; Eklund et al., 2008).

At the other end of the spectrum are the observations reported by advocates of magmatism of the English and German schools regarding the origin of granites. One of the best examples are the geochemical investigations of Rosenbusch (1877b) on the contact metamorphic slates surrounding the Barr-Andlau Pluton in the northern Vosges (France). From a series of geochemical analyses along sampling profiles from distal towards proximal areas of the margin of the granite, the author concluded that metamorphism adjacent to granitic bodies always took place without significant chemical changes. Another exceptional example refuting the existence of granitization were the observations of Callaway (1894) in a paper entitled Is granite ever metamorphic? In this work, the author investigated contact zones between granites and surrounding schists in Great Britain, which were believed to show transitional or gradational contacts. At three investigated localities, Callaway (1894) refuted the former interpretations about a metamorphic or metasomatic origin of the exposed granites by showing evidence of granitic veins or intrusive and "net" contacts between granites and sedimentary country rocks. However, Callaway (1894) did not account for the room problem. Hans Cloos, from the German school, attempted to solve the room problem and suggested that large batholiths posing controversial space problems were in truth large intrusive sheets (Cloos, 1923). This interpretation involved less room needed to form such structures, although vertical or horizontal crustal extension was necessary to accommodate the intrusive sheets following crustal discontinuities.

Previous work by Daly (1912), who also advocated for the forceful intrusion of granitic magmas, was influential to the understanding of the room problem. Based on the occurrence of intrusion breccias observed along the margins of large granitic masses along the North American Cordillera, Daly (1912) coined the term magmatic stoping and defined it as "the successive engulfment of suites of blocks broken out of the roof and walls of the batholith". Pitcher et al. (1958) later proposed lateral magmatic widening as the mechanism of emplacement for the relatively large Donegal granite in northwest Ireland. Numerous investigations on this area had already been published, suggesting either a magmatic (Haughton, 1862) or metamorphic (Scott, 1861) origin. However, Pitcher et al. (1958) provided new meaningful observations and detailed geological mapping of different granite units, acknowledging that a "metamorphic" origin appeared reasonable for one part, and an "igneous" origin for another. Pitcher et al. (1958) found that discontinuous metasedimentary horizons in the country rocks were also present within the granite as a *relict* or *ghost* stratigraphy, a phenomenon that had classically been used as an argument for granitization. However, they attributed it to a syn-plutonic origin for these horizons, which developed a strong marginal lineation during compression caused by sequential pulses of a plexus of wedge-like granite sheets (see a review in Cobbing, 2000).

Contrasting positions were taken by geologists since the late 19th century, but the *room* problem within the magmatism vs. transformism debate was not polarized on a geographical basis. There were as many *extreme magmatists* in the French and Finnish schools as there were *extreme transformists* in the English and German schools. For example, investigations by Currier (1937, 1947), member of the United States Geological Survey, and Reynolds (1943, 1947, 1958), Fellow of the Royal Society of Edinburgh, promoted a metamorphic or metasomatic origin of granites (e.g., by Lewis, 2021).

The origin of granitic magmas was the second major question in the magmatism vs. transformism debate and, unlike the *room* problem, was solely pursued within the magmatic school of thought. If granites were magmatic, then the origin of granitic magmas and the conditions at which they formed, including their water content, were key issues to be understood. The origin of this discussion dates back to the birth of geochemistry and experimental petrology, which marked the onset of numerous investigations on the order of crystallization of mineral phases and their physicochemical conditions of growth (see Section 3). The presence of water in granitic magmas was acknowledged following the work by Sorby (1858) on fluid inclusions in granites, and by Scheerer (1862) on melting of hydrous salts as a magma analogue. However, before reaching this stage, the controversy on the origin of granitic magmas was focused on the fusion of rocks: dry, or wet? (see Chapter 6 in Young, 2018). The leading advocates for the dry fusion were the Frenchmen Baptiste Xavier Fournet and Joseph Marie Durocher (e.g., Fournet, 1844, 1845; Durocher, 1859). Wet fusion (involving the presence of H<sub>2</sub>O and other volatile phases) was accepted later during the second half of the 19th century. This theory had major implications for the understanding of how granitic magmas form and, more specifically, for the physicochemical laws governing the process of magmatic differentiation.

The concept of magmatic differentiation was proposed and subsequently investigated by, among others, Scrope (1825), von Bunsen (1851, 1861), Iddings (1893b, 1896), and Harker (1896, 1909) (see Section 3). However, it only started to gain advocates since the petrographic and geochemical insights provided in the seminal work of Brögger (1894) (see Pitcher, 1997a, 1997b). From field observations and geochemical analyses of the granites cropping out in southern Norway, Brögger (1894) suggested that the sequence of mineral phase crystallization reflected the differentiation of the original magma, from basic to acid: "the oldest rocks are the most basic, the youngest are the most acid, and between the two extremes I have found a continuous series". At that time, it was already known that, under wet conditions, a granitic melt could be formed at much lower temperatures than an equivalent anhydrous melt, which had major implications for the research on the order of crystallization of mineral phases, or "crystallization series" (e.g., Fouqué and Michel-Lévy, 1882; Brögger, 1894; Iddings, 1896). Understanding of the specific order in which mineral phases are formed during the cooling and solidification of the magma, was based on the relative melting points of minerals. It was known that the crystallization temperature of micas was lower than that of feldspar, and still lower than that of quartz (see Chapter 6 in Young, 2018). On this basis, works from the late 18th century had already suggested that quartz, the phase with the highest melting temperature, should crystallize before the more fusible feldspar and, therefore, euhedral quartz crystals should show imposed faces on the subsequent feldspar crystals rather than the opposite (e.g., Kirwan, 1793, 1800). By the end of the 19th century, however, it was known that that the actual order was precisely the reverse, as demonstrated by von Bunsen (1861): "No one seems to have considered that the temperature at which a body solidifies for itself is never the temperature at which it becomes solid from its solutions in other bodies".

Later on, the order of crystallization in granitic magmas was further investigated by Guthrie (1884), who first suggested the idea of eutectics, Vogt (1884, 1888), regarded as the founder of the modern experimental petrology, and Rosenbusch (1910), who maintained that molten material with a composition of primary magma exists within the Earth's interior, where it continually differentiates. Moreover, according to Rosenbusch (1910), variations in differentiation could occur because of variations in temperature and pressure during crystallization. This latter idea served as the basis for the seminal experimental work of Norman L. Bowen on the crystallization of silicate melts and its relation to natural associations of igneous rocks (see a review in Grove and Brown, 2018). Bowen investigated crystallization processes in different silicate melts and binary, ternary, or composite systems. Among other findings, those related to different binary and ternary systems, as well as the invaluable insights into the fractional crystallization of basalt under anhydrous conditions, represented a milestone for igneous petrology in the early 20th century (e.g., Bowen, 1912, 1913, 1916). Later, Bowen started to work with Frank Tuttle, who during the 1940s conceptualized and built

an apparatus to simulate varying conditions in the Earth's crust. The hotseal hydrothermal quenching apparatus, painstakingly described in Tuttle (1948) and nowadays stored at the Carnegie Institution for Science Earth and Planets Laboratory Archives (Washington DC, USA), represented the inception of a plethora of investigations on mineral systems including volatile components, which had been sparse until that moment (e.g., Goranson, 1936). At that time, the origin of granite was viewed differently by field researchers with strong backgrounds in structural geology and metamorphic petrology, who favoured a non-magmatic origin of granite, and those with backgrounds in experimental petrology. Two quotations from leading researchers in these fields are illustrative: "the best geologist is he who has seen the most rocks" (Read, 1939), and "look as one will at the rocks he cannot see them in process of formation" (Bowen, 1948). In this framework, the seminal work of Tuttle and Bowen (1958) on the haplogranite system NaAlSi<sub>3</sub>O<sub>8</sub>-KAlSi<sub>3</sub>O<sub>8</sub>-SiO<sub>2</sub>-H<sub>2</sub>O (Ab-Or-Qtz-H<sub>2</sub>O) demonstrated that granitic rocks could be formed from granitic magmas. Only a few years later, Winkler and von Platen (1961) experimentally tested high-grade metamorphism and anatexis of different greywackes, finding that tonalitic or granodioritic compositions were a final stage in the anatectic process. This milestone offered an even more convincing hypothesis for the formation of granitic melts and ultimately granitic rocks, significantly contributing to the understanding of their genesis. Granite, therefore, turned out to be strictly magmatic for most experimental petrologists (see Bailey and Macdonald, 1976), whereas other geologists maintained opposite views on the basis of field and petrographic evidence (see Walton, 1960; Marmo, 1967a; Vernon, 1986; Pitcher, 1987; Clarke, 1996; Pitcher, 1997a, 1997b; Brown, 2013).

### 6. The granite concept since the Plate Tectonics theory

"Some granites are cooler than others" Vhairi Mackintosh. Cosmos Magazine - 2 Jul 2018

Alfred Wegener's continental drift theory (Wegener, 1912) is widely understood as the precursor to the Plate Tectonics theory. However, it is worth noting that, long before that moment, Benjamin Franklin already pointed out that the movement of continents across the Earth's surface should be driven by some internal force: "Thus the surface of the globe would be a shell, capable of being broken and disordered by the violent movements of the fluid on which it rested" (Franklin, 1782). The initial ideas of Franklin and Wegener were initially met with great scepticism due to the lack of plausible mechanisms (Oreskes, 1999), which were not discovered and accepted until the 1960s-1970s. Processes such as mantle convection, seafloor spreading, and slab pull and ridge push, provided reliable explanations for the large-scale movement of the Earth's lithosphere, and agreed with observations from seafloor mapping, paleomagnetic analysis, geochemical compositions, and largescale tectonics (Dietz, 1961; Hess, 1962; Vine and Matthews, 1963; Wilson, 1965, 1975; McKenzie and Parker, 1967; Ringwood, 1974; Anderson, 1982; Tatsumi, 1986; McKenzie, 1989). Plate tectonics is thus a scientific revolution not only for granite science but for the understanding of the past, present and future behaviour of our planet; its influence on the knowledge of granitic rocks has therefore been crucial during the last 60 yr. (Fig. 5). One of the most important contributions was, for example, to distinguish granites from different tectonic settings and sometimes with different sources, mineralogy or compositions, such as in orogenic, anorogenic, and mid-ocean ridge environments (e.g., Chappell and White, 1974; Pitcher, 1983, 1987; Pearce et al., 1984; Maniar and Piccoli, 1989; Brown, 1994).

However, the Plate Tectonics theory did not directly end the magmatism vs. transformism debate, since many inconsistencies concerning granitic magma generation, segregation and emplacement were still debated (Fig. 5). These inconsistencies are exceptionally well documented in the work of Pitcher (1987), where it is first acknowledged that "the advent of the Plate Tectonic Hypothesis has emphasised the relationship between cause, process and the regional geological environment" but, later, it is stated that "whilst many granites are essentially magmatic and intrusive others are metamorphic and replacive". However, Plate Tectonics did contribute to increasing the explanatory power of magmatism and decreasing that of transformism (Fig. 5). The magmatism vs. transformism debate lasted until the late 20th century, with some investigations published based on studies of different regions of eastern Europe as the last witnesses of this long-lasting discussion (Korzhinskii, 1962; Dobretsov and Shafeev, 1991; Stumbea, 2003; Sedova et al., 2013; Aranovich, 2017). The magmatism paradigm prevailed, and granite research entered progressively, for a second time (see Section 3), the current paradigm-governed, stasis period of normal science (Kuhn, 1970) (Fig. 5). Accordingly, most current investigations into granitic rocks (see below in this Section) are part of a reductionist science trend, based on detailed studies about specific fields and sub-fields that offer their findings as part of a wider knowledge gap (Tex, 1990; Clarke, 1996; Petford et al., 1997).

The last half-century of research on granitic rocks, since the acceptance of the Plate Tectonics theory and the establishment of the magmatism paradigm, has mostly consisted, in Kuhn's words, of "puzzle solving"; which is, for clarity, to address problems that are believed in advance to have a solution (Fig. 5). This period has further fostered the proposal of tens of genetic and non-genetic classifications, some of them still used and some others already proven to be only valid for the context where they were proposed (see Section 6.1). Publications that categorically support a non-magmatic origin for igneous-looking rocks are almost absent in the modern literature. Noteworthy abundant are, however, investigations that tentatively attribute a non-magmatic origin to granite-looking rocks, suggesting that: (1) rocks showing very different field appearances (gneiss, trondhjemite, and mica schist among others) may be derived from similar precursors by mineralogical and chemical changes induced by deformation and metasomatism, and (2) deformation and metasomatism can also give a very similar appearance to previously different rocks (e.g., Hanor and Duchač, 1990; Almond et al., 1997; Elburg et al., 2001, 2012; Martin, 2006; Zirner et al., 2015). These investigations reflect how neglected ideas may eventually become central for explaining inconsistencies in new paradigms that have more explanatory power (Fig. 5). This phenomenon can be compared, as with the concept of the schools of thought (see Section 4), with Evolution theory. In this case, reusing old, neglected ideas to solve new problems is analogous to the role of refuges during evolution (Fig. 5), where old genomes may persist in isolated places without following the main evolutionary trend and may subsequently join the mainstream again.

It is currently widely accepted by the scientific community that granites *sensu stricto* are plutonic (Hutton, 1794; Walton, 1960; Le Maitre et al., 2002) silica-rich felsic rocks that contain modal quartz (Q), alkali feldspar (A), and plagioclase (P) in the following proportions: 20 < Q < 60 and A: P = 90: 10 to 35: 65, and that they may contain one or more types of mica, amphibole, and many other accessory minerals (Streckeisen, 1974; Le Maitre et al., 2002). However, metasomatic *pseudo-granites* are also proposed to still exist, as recently demonstrated for the Himalayan banded tourmaline leucogranites, formed by metasomatism of a psammitic country-rock (Dyck and Larson, 2023), or the granites in the Mt. Painter Inlier that are alteration products of metasedimentary rocks of the Radium Creek Group (Weisheit et al., 2013).

Current research on granitic rocks can roughly be divided in two major approaches (see below): (1) observational and analytical and (2) experimental and numerical modelling. Although it is beyond the scope of this work to thoroughly review the bountiful literature of these two approaches, below we provide a summary of their main milestones and current debates (see Sections 6.2 and 6.3).

### 6.1. The classification(s) of granites

Naming and categorizing igneous rocks, including granites, has been highly controversial for geologists for the past two centuries. Since the acceptance of the Plate Tectonics theory, more than 30 non-petrogenetic and petrogenetic classifications of granitic rocks have been proposed (see the reviews of Barbarin, 1990, 1999; Frost et al., 2001; Bonin et al., 2020; Garcia-Arias, 2020). However, widespread agreement on the most suitable classification has not been yet reached. As pointed out by Garcia-Arias (2020), this long-lasting pursuit probably lies in the intrinsic "imperfection" of the existing chemical classification systems, which, if perfect, should provide names that convey all relevant information about a given sample: modal mineralogy, texture, chemical composition and genesis. For example, the widely accepted non-genetic classification of Streckeisen (1974) can be easily applied in the field and does not require determination of the other minerals (besides quartz and feldspars) that are present in the rock, nor how these rocks formed. However, this is also a limiting factor because the classification does not consider the abundance and composition of other mineral phases that may have significant petrological implications. Another example is the petrogenetic classification of Chappell and White (1974) that, although widely adopted, is not always straightforward when indicator minerals are absent, such as cordierite for S-type, amphibole for I-type, and orthopyroxene (or pigeonite or fayalite) for C-type granites (Ague and Brimhall, 1988; Barbarin, 1990; Kilpatrick and Ellis, 1992; Frost et al., 2001). Other drawbacks of these "alphabet" genetic classification schemes are that they often use a single parameter to infer a unique source. For example, peraluminous granites are equated to S-types, and therefore deemed to have a sedimentary source, whereas other origins and pathways may lead to the same composition (Alonso-Perez et al., 2009). Moreover, the classification of granites according to these multiple schemes is sometimes seen as a goal in itself, which is an approach that contributes little to our understanding of granite petrogenesis.

In this section, most used classifications are summarized, including details of their major strengths and potential pitfalls. Classifications of granitic rocks here reviewed are based on a wide number of variables including, among others, mineralogical modal compositions, petrographic features, accessory mineral phases, nature and abundance of enclaves, major or trace element compositions, and isotope geochemistry. The principal non-petrogenetic classifications of granitic rocks are presented in Table 1 (Streckeisen, 1974; De la Roche et al., 1980; Debon and Le Fort, 1988; Middlemost, 1994; Frost et al., 2001; Barton and Young, 2002; Frost and Frost, 2008; Enrique and Esteve, 2019; Glazner et al., 2019). They are based on either mineralogical modal abundances or major element compositions, and have the major advantage of being non-genetic. It is explicitly stated in Le Maitre et al. (2002) that the Subcommission on Igneous Petrology of the IUGS was based on Streckeisen (1974) and Le Bas and Streckeisen (1991) to provide nomenclature and classification to all igneous and igneous-looking rocks. Most authors now agree to use the modal classifications recommended by the IUGS (e.g., Streckeisen, 1974; Middlemost, 1994; Le Maitre et al., 2002), which use subdivisions and nomenclatures based on a limited number of variables, and are generally widely accepted (e.g., Frost et al., 2019; Bonin et al., 2020; Garcia-Arias, 2020). However, other petrologists have recently questioned and reformulated these schemes (e.g., Glazner et al., 2019).

The principal petrogenetic classifications of granitic rocks are summarized in Table 2 (Didier and Lameyre, 1969; Capdevila and Floor, 1970; Capdevila et al., 1973; Tauson and Kozlov, 1973; Chappell and White, 1974, 2001; Orsini, 1976, 1979; Ishihara, 1977, 2004; White and Chappell, 1977; Loiselle and Wones, 1979; Lameyre, 1980; Pupin, 1980, 1985; Keqin et al., 1982; Lameyre and Bowden, 1982; Yang, 1982; Pitcher, 1983, 1987; Pearce et al., 1984; Nachit et al., 1985; Rossi and Chevremont, 1987; Maniar and Piccoli, 1989; Barbarin, 1990, 1999; Castro et al., 1991; Didier and Barbarin, 1991; Kilpatrick and Ellis, 1992; Tischendorf and Förster, 1992) and have already been exhaustively discussed in Barbarin (1999) and Frost et al. (2001). These classification schemes gained popularity with the advent of modern geochemical methods, which provided evidence of the relationships between the mineralogical, chemical, and/or isotopic composition of granitoid rocks and their petrogenesis or tectonic setting (e.g., Pearce et al., 1984;



Fig. 5. Schematic timeline of the historical development of the granite concept through time, after the ideas of Kuhn (1970) and Clarke (1996). Each new paradigm about the origin of granites is settled (i.e., established) through a "scientific revolution" after "great scientific achievements". Scientific revolutions are followed by a period of "normal" science (paradigm-governed), where specific investigations or case studies are carried out through "puzzle solving" (i.e., investigating problems that are believed in advance to have a solution; stasis periods according to Bak, 1996). As more inconsistencies and new ideas emerge (anomalies falling outside the explanatory power of a given theory), the paradigms experience a crisis and collapse progressively, which forces a new revolutionary change ("avalanche" in the sandpile analogy of Bak, 1996). Ideas from preceding paradigms are often used to foster the explanatory power of the newly settled ones. See text for further explanations.

Maniar and Piccoli, 1989). In some cases, petrogenetic granite classifications were devised as bimodal schemes that progressively gained complexity by means of new sub-divisions and/or petrogenetic constrains. The two original I- and S-type granite endmembers, initially proposed by Chappell and White (1974) and subsequently complemented by the A-, M-, C- and H-types (Chappell and White, 2001; White and Chappell, 1977; Loiselle and Wones, 1979; Castro et al., 1991), are a good example of this growing complexity.

## 6.2. Observational and analytical approaches for investigating granitic rocks

Observational and analytical approaches of current research on granitic rocks encompass field-based, mineralogical, geophysical, and geochemical investigations (e.g., Liesa et al., 2021; Dahlquist et al., 2022; Milani et al., 2022; Song and Xu, 2022; Turner et al., 2022), commonly related to geothermal energy (e.g., Singh et al., 2020; Klee et al., 2021; Chandrasekharam et al., 2022) and the exploration for ore deposits (e.g., Harlaux et al., 2021; Lehmann, 2021; Zhang et al., 2021; Wang et al., 2022a, 2022b). Global investigations are generally represented by multidisciplinary approaches at the interface between structural geology and petrology. The observational approach for the study of granitoid rocks often starts with geological-structural mapping through classical fieldwork (e.g., Cobbing, 2000; IGME-BRGM, 2009; Casini et al., 2015; Tavazzani et al., 2017; Errandonea-Martin et al., 2018; Secchi et al., 2021; Russo et al., 2022) and, in high-altitude or inaccessible areas, remote sensing-based technologies (e.g., Watts and Harris, 2005; Haselwimmer et al., 2011; Karimzadeh and Tangestani, 2021). Recently, the historical importance and future directions of geological mapping have been reviewed by Butler et al. (2024). Another

fundamental and widely applied observational mesoscale approach for the study of granites is the mapping, measurement, and quantification of fabrics (Bouillin et al., 1993; Bouchez, 1997; Moyen et al., 2003; Carreras et al., 2004). At the microscopic scale, beyond optical petrography, Electron Back Scatter Diffraction (EBSD) analyses with Scanning Electron Microscope (SEM) have also become a popular observational approach in the study of tectonic fabrics and sub-fabrics present in granitic rocks (e.g., Peternell et al., 2010; Mamtani and Renjith, 2015; Ávila et al., 2022), which may provide insights into the stress, strain, temperature, and pressure conditions of granitoid deformation (Hirth and Tullis, 1992; Stipp et al., 2002; Cross et al., 2017; Gomez-Rivas et al., 2020).

Geophysical imaging of granite bodies may also be classified as an observational approach. Among others, electric (e.g., Olhoeft, 1981; Jover, 1986), magnetic (e.g., Schwarz, 1991; Maré and Thomas, 1997; Kadioglu et al., 1998), and seismic surveys (Wenzel et al., 1987; Améglio and Vigneresse, 1999) have been conducted during the past decades. Moreover, high-resolution seismic reflection profiles and the 3D inversion of gravity data have provided valuable insights into the deep geometry and the internal structure of granitoid masses (e.g., Mair and Green, 1981; Evans et al., 1994), as well as the large-scale crustal architecture of different orogens worldwide (Muñoz, 1992; Mortimer et al., 2002; Korsch et al., 2012; Ayarza et al., 2021). Gravity survey data and methods such as 2D forward modelling and 3D inversion can be considered as the most used and insightful geophysical imaging techniques for granitic rocks (e.g., Vigneresse, 1990, 1995, 1999; Améglio et al., 1997). Since differences in gravity signatures are based on density contrasts between granitoids (generally with lower density, thus showing negative gravity signatures) and surrounding rocks (Vigneresse, 1990; Améglio and Vigneresse, 1999; Wang et al., 2011),

#### Table 1

Summary of the main non-petrogenetic classifications of granitoid rocks, including the parameters or properties utilised for categorizing, a broad summary of the resulting classes, and their potential problems and pitfalls.

Reference	Parameters considered	Classification at glance	Potential pitfalls
(chronological order)			
Lacroix (1933)	Major element composition	Calc-alkaline hyperaluminic Calc-alkaline Alkaline	In disuse – Terminology and sub-classes have been improved and/or updated
Shand (1948); Barton and Young (2002)	Major element composition (Aluminium Saturation Index)	Peralkaline; Metaluminous; Weakly/Strongly peraluminous	In use - Terminology and sub-classes have been improved and/or updated
Streckeisen (1974)	Major element composition (volumetric % normative composition of quartz, K- felspar, plagioclase feldspar)	Widely used, applicable in the field, inexpensive. Recommended by the International Union of Geological Sciences	Other minerals than quartz and feldspars are not considered.
De la Roche et al. (1980)	Major element composition (Multicationic chemical parameters R1 and R2; Oxide percentages converted to millications)	Incorporates all of the major elements that are relevant to both the rock mineralogy and petrology. Similar nomenclature as in <u>Streckeisen (1974)</u>	Not suitable for granitic rocks because K-feldspar and albite plot at the same point (see Batchelor and Bowden, 1985)
Debon and Le Fort (1988)	Major element composition (Oxide percentages converted to millications)	Applicable in two steps: sample nomenclature and identification through cationic values, and sample association to magmatic three possible groups (Aluminous / Alumino-Cafemic / Cafemic). First step with similar nomenclature as in Streckeisen (1974)	In disuse - Terminology and sub-classes have been improved and/or updated. According to Bonin et al. (2020), it should be resurrected and used as the most ideal classification
Middlemost (1994)	Major element composition (% alkali and silica content)	Recommended by the International Union of Geological Sciences – Similar nomenclature as in Streckeisen (1974)	Other minerals than quartz and feldspars are not considered
Frost et al. (2001)	Major element composition (Fe- number, Modified alkali–lime index, and Aluminium Saturation Index)	Three-tiered classification scheme that uses familiar chemical parameters to successively distinguish among (i) magnesian and ferroan, (ii) alkaline, alkali-calcic, calc-alkaline and calcic, and (iii) peraluminous, metaluminous and peralkaline types	Not (yet) widely applied. Terminology and sub- classes have been improved and/or updated - See Frost and Frost (2008)
Frost and Frost (2008)	Major element composition (Alkalinity Index & Aluminium Saturation Index)	Introduction of the alkalinity index and feldspathoid silica-saturation index allows to apply the three-tiered classification scheme of Frost et al. (2001) to all alkaline rocks, including feldspathoid-bearing rocks	Not (yet) widely applied.
Enrique and Esteve (2019)	Major element composition (SiO <sub>2</sub> , CaO, and $K_2O$ )	Chemical approximation to the normative QAP diagram. Similar nomenclature as in Streckeisen (1974)	Other minerals than quartz, feldspars, and feldspathoids are not considered. Sodium not considered.
Glazner et al. (2019)	Major element composition	Based on the QAP system but adding quantitative data to rock names	See numerous comments in Frost et al. (2019) and Hogan (2019)

these techniques have allowed the identification of the shape and extension at depth of outcropping granite bodies in, e.g., the Congo Craton (Shandini and Tadjou, 2012), the Pennine area (Kimbell et al., 2010), the Pyrenees (Ayala et al., 2021), and the Arabian Shield (Mukhopadhyay et al., 2021). Furthermore, they have allowed the detection of concealed subsurface igneous masses of contrasting size and composition in many regions where, for example, granitic basement rocks are located beneath thick sedimentary sequences (Cooper Basin in Central Australia, Meixner et al., 2014; Middle and High Atlas in northern Africa, Elabouyi et al., 2022; or Northland in New Zealand, Stagpoole et al., 2016), and even in the subsurface of the Moon (Andrews-Hanna et al., 2013, 2014) and Mars (Kiefer, 2004).

Regarding analytical approaches, whole-rock geochemical data are commonly used for the petrogenetic and non-petrogenetic classification of granitic rocks (see Tables 1 and 2 and Section 6.4). Moreover, geochemical investigations on accessory minerals of granites have gained popularity during the recent decades due to their potential use for correlation, dating, and unravelling the temperature and pressure conditions of granite formation. Zircon, for example, is an outstanding U-Pb geochronometer due to, among others, its high closure temperature and the long-lasting stability of its atomic structure (e.g., Davis et al., 2003; Harley and Kelly, 2007). U-Pb zircon geochronology has therefore been applied to granitic rocks worldwide to obtain insights into major orogenic events (e.g., Aguilar et al., 2014; Heilbron et al., 2017). However, in some cases, the robustness of zircon can be a hindrance to determine the magmatic age of granites, as the mineral is not always reset during anatexis (Bea et al., 2021). In these cases, U-Pb dating of monazite and titanite (Elburg et al., 2003) can be more informative. The trace element and isotopic composition of igneous zircons, as well as of micro-zircon inclusions in accessory minerals, are furthermore considered as petrogenetic indicators (e.g., Belousova et al., 2002; Bell and Kirkpatrick, 2022; Guo et al., 2022; Yang et al., 2024). Recently, zircon oxygen isotopic signature and water content have also been suggested as a potential tool to distinguish I- and S-type granites (Mo et al., 2023).

Beyond zircon, other accessory minerals are also targets of geochronological investigations on granites. U-Pb dating of apatite crystals can provide accurate geochronological data, although it is recommended to apply it in combination with other geochronological data for comparison and data validation (e.g., Ferreira et al., 2019; Van Daele et al., 2020). In fact, two main possible pitfalls must be considered when carrying out U-Pb dating of apatite-bearing granitoid rocks. First, its closure temperature of ~450–550 °C (e.g., Schoene and Bowring, 2007), which is significantly lower than that of zircon ( $\sim <900$  °C; e.g., Cherniak and Watson, 2003). And second, its potential crystal-plastic deformation and recrystallization, which causes chemical changes in the trace element composition substantially decreasing the U/Pb ratio (Ribeiro et al., 2020). Another useful method generally applied to zircon and apatite crystals to gain insights into the exhumation paths of granitic rocks is the measurement of fission-tracks, i.e., regions of high density of crystalline defects formed due to the damage inferred by fission decay of <sup>238</sup>U (Reiners and Brandon, 2006). Other methods applied to diverse accessory minerals for constraining the crystallization age of granitic rocks are, for example, cassiterite U-Pb and muscovite Ar-Ar dating on granite-related ore deposits, such as skarn- and greisen-type deposits (Chen et al., 2018; Tichomirowa et al., 2019; Bui et al., 2022), in-situ Lu-Hf geochronology of apatite through laser ablation tandem inductively coupled mass spectrometry (LA-ICP-MS/MS), as well as the

### Table 2

Summary of the main petrogenetic classifications attempted for granitoid rocks (updated after Barbarin, 1999), including the parameters or properties employed for categorizing, a broad summary of the resulting classes, and their potential problems and pitfalls.

Reference	Parameters considered	Classification at glance	Potential pitfalls
(chronological order)			
Didier and Lameyre (1969); Didier and Barbarin (1991)	Mineralogy and associated enclaves	C-type (crustal granitoids; leucogranites) M-type (mixed or mantle granitoids; monzonites and granodiorites)	In disuse - Diagnostic minerals and enclave types may not be present
Capdevila and Floor (1970); Capdevila et al. (1973)	Petrographic features	Mesocrustaux / Mixtes / Basicrustaux	In disuse - Diagnostic minerals and textures may not be present
Tauson and Kozlov (1973)	Trace element composition [Discrimination diagrams]	Plumasitic leucogranites / Ultra-metamorphic granites / Palingenic granites / Plagio-granites / Agpaitic leucogranites	In disuse - Outdated terminology
Chappell and White (1974, 2001); White and Chappell (1977); Loiselle and Wones (1979); Castro et al. (1991); Kilpatrick and Ellis (1992)	Major element composition and Mineralogy	I-type (Igneous source) / A-type (Anorogenic) / S- type (Sedimentary source) / M-type (Mantle-derived) / C-type (Charnockitic) / H-type (Hybrid source)	Granitoids rarely come from single sources. Diagnostic minerals may not be present for some sources. See text for further pitfalls and other comments in Frost et al. (2001)
Orsini (1976, 1979)	Petrographic features	Aluminic subalkaline / Hypoaluminic Subalkaline / Calc-alkaline	In disuse - Diagnostic minerals and textures may not be present
Ishihara (1977, 2004)	Fe-Ti oxides content, associated mineralization, $\delta^{34}S$ and $\delta^{18}O$ values	Depending on the occurrence or absence of magnetite (magnetite-andillmenite-series, respectively)	Processes that can control the stability of magnetite in granitic rocks are not fully considered. See comments in Frost et al. (2001)
Lameyre (1980); Lameyre and Bowden (1982)	Mineralogy	Based on the QAP system and the four major series, differentiates among: Leucogranites (Crustal fusion) / Calc-alkaline series / Tholeiitic series / (Per)Alkaline series	In disuse - not accepted nowadays to link QAP compositions with possible magmatic sources
Pupin (1980, 1985)	Zircon morphology, petrographic features, accessory minerals	Types 1–7 grouped in Crustal, Mixed, and Mantle granites on the basis of zircon typological populations and their "Typological Evolutionary Trend" (TET).	In disuse - Diagnostic minerals and zircon morphology may not be present
Keqin et al. (1982)	Associated mineralization	Transformation type (Continental Crust) / Syntexis type (transitional Crust) / Mantle-derived type	In disuse - outdated terminology
Yang (1982)	Petrographic features	Metamorphic type / Crustal type / Mixed source type / Mantle-derived type	In disuse - Diagnostic minerals and textures may not be present
Pitcher (1983, 1987)	Geodynamic setting	Hercynotype / Caledonian type / Andinotype / West Pacific type / Nigeria type	In disuse - outdated terminology
Pearce et al. (1984)	Trace element composition [Discrimination diagrams]	Distinction among tectonic environments on the basis of (among others) Rb-Ta-Nb-Y relationships: Ocean ridge granites / Volcanic arc granites / Within-plate granites / Collisional granites	Not applicable to post-collisional granitoids because they plot over the range of classification fields
Nachit et al. (1985)	Biotite composition [Discrimination diagrams]	Distinction among magmatic series on the basis of Al and Mg contents of biotite: Alumino-potassiques / Calcoalcalines et subalcalines / Alcalines et hyperalcalines	In disuse - Processes that can control the composition of biotite in granitic rocks are not fully considered
Rossi and Chevremont (1987)	Mineralogy (mafic minerals)	Depending on the mafic mineral content, distinction among: Aluminopotassique / Monzonitique / Calcoalcaline / Tholeitique / (Per)Alcaline	Diagnostic minerals and textures may not be present
Maniar and Piccoli (1989)	Major element composition [Discrimination diagrams]	Distinction among tectonic environments (Island arc granitoids, Continental arc granitoids, Continental collision granitoids, Postorogenic granitoids, Rift- related granitoids, Continental epeirogenic uplift granitoids, and Oceanic plagiogranites	Not applicable to post-collisional granitoids because they plot over the range of classification fields
Barbarin (1990, 1999)	Petrographic features, Mineralogy, Major element composition	Three main groups that are further subdivided and correlated to major tectonic settings: Crustal / Mixed / Mantle	It compresses the wide range of bulk compositions of granitoids into six rock types. See comments in Frost et al. (2001)
Tischendorf and Förster (1992)	Major and trace element composition [Discrimination diagrams]	Distinction among granitoids tectonic environment on the basis of K2OxRb / MgO – Na2OxZrxY diagram (Ocean ridge, Volcanic arc, Within plate, and Collisional)	In disuse – Processes such as magma mixing and wall rock assimilation are not considered

recently developed in-situ triple quadrupole inductively coupled plasma mass spectrometry (LA-QQQ-ICPMS) Rb—Sr dating of mica (Zack and Hogmalm, 2016; Wang et al., 2022a, 2022b).

Isotopic investigations also constitute a wide category within the analytical approaches addressing the origin and significance of granitic rocks. Investigations on whole-rock radiogenic isotope ratios, such as <sup>87</sup>Sr/<sup>86</sup>Sr, <sup>143</sup>Nd/<sup>144</sup>Nd, and the various Pb isotope ratios, currently play a major role on granite petrogenetic investigations, including the origin and composition of magma sources, the mantle-crust interactions, crustal melting processes, and the identification of the geotectonic setting (e.g., McCulloch and Chappell, 1982; Foden et al., 2002; Farina et al., 2014; Zametzer et al., 2022). Moreover, the study of radiogenic isotopes in single mineral grains or mineral separates, such as <sup>87</sup>Sr/<sup>86</sup>Sr on K-feldspar (e.g., Siebel et al., 2005) and <sup>176</sup>Hf/<sup>177</sup>Hf on zircon (e.g.,

Chagondah et al., 2023), also provide very meaningful insights into the assimilation of wallrock material and magma mixing processes during granite petrogenesis. In this framework, analytical profiles through zircon and feldspar crystals, namely crystal isotope stratigraphy (Davidson et al., 1998), may help to reveal both the magmatic history and post-magmatic subsolidus re-equilibration of a given granite.

Reconstructions of the temperature and pressure conditions of granite formation have also a strong analytical basis that targets many accessory minerals. Among others, the two-feldspar, hornblende-plagioclase, pyroxene-ilmenite, pyroxene-biotite, garnet-hornblende, muscovite-biotite, and garnet-biotite methods are based on mineral compositions and activity models to infer the temperature and/or pressure of crystallization, although they are strongly dependent on other factors (see Anderson, 1996; Anderson et al., 2008). When present

in granitoid rocks, garnet also represents a useful tool in petrogenetic studies for unravelling the pressure (e.g., garnet-biotite-plagioclasequartz mineral assemblage) and temperature (e.g., garnet-biotite mineral assemblage) of crystallization (e.g., Dahlquist et al., 2007). Furthermore, titanite, a common accessory mineral in mainly metaluminous granites, can also be used as an empirical geobarometer when crystallized under near-solidus conditions together with a particular mineral assemblage (Erdmann et al., 2019). Additionally, significant advances in barometric and thermometric estimations of granite formation have been made from diffusion modelling coupled to determinations of titanium concentrations in zircon (Schiller and Finger, 2019) and (more commonly) in quartz, the latter also known as Titanium-in-quartz thermobarometry (Wark and Watson, 2006; Huang and Audétat, 2012). A noteworthy finding from Titanium-in-quartz thermobarometry is the significantly low formation temperature (474-561 °C) obtained by Ackerson et al. (2018) for the Tuolumne Intrusive Suite in the Sierra Nevada Batholith, California. Although lowtemperature magma storage had already been inferred from volcanic rocks, for example in the Taupo Volcanic zone (Rubin et al., 2017), results from Ackerson et al. (2018) may represent one of the paradigmshifts of the 21st century regarding the knowledge of granitic rocks, with major implications on their petrogenesis. However, these findings have also been suggested to be problematic (Clemens et al., 2020). Furthermore, beyond investigations of granite petrogenesis, the postmagmatic evolution on granitoid rocks has also been investigated through studies on hydrothermal minerals. This wide topic includes chemical and stable isotope analyses on quartz veins and related ores hosted in granitic rocks (Sun and Eadington, 1987; Bhattacharya et al., 2014), as well as microthermometry and microanalysis of fluid inclusions (Wagner et al., 2016; Tang et al., 2021).

### 6.3. Experimental and numerical modelling approaches for investigating granitic rocks

Approaches that address the origin and significance of granitic rocks from an experimental or numerical modelling background may be classified in three different perspectives. Chronologically ordered, they correspond to the origin of melt (i.e., fluid present or absent partial melting and melt segregation), its ascent and emplacement (i.e., continuous or episodic ascent, and mode of emplacement in different tectonic settings), and the post-magmatic tectono-thermal evolution (i. e., brittle and ductile deformation, metamorphism, and metasomatism) (e.g., Petford et al., 2000). After the pioneering experimental work of Tuttle and Bowen (1958) on the haplogranite system Qz-Ab-Or-H<sub>2</sub>O, experimental melting of upper crustal rocks and crystallization of granite became very popular and was carried out at a wide range of physicochemical conditions. This includes low (1–2 kbar; e.g., Clemens et al., 1986; Rusak et al., 2021) to high (10-15 kbar and even up to 40 kbar; e.g., Green and Lambert, 1965; Huang and Wyllie, 1975) pressures and from granite solidus (500-800 °C; e.g., Johannes, 1984; Shchekina et al., 2020; Bartoli and Carvalho, 2021) to high (1400 °C; e.g., Clemens et al., 1986) temperatures, in runs that span from few hours (e.g., Green and Lambert, 1965) to several weeks or months (e.g., Maaløe and Wyllie, 1975). Such experiments have also shown that equating peraluminous felsic melts to a crustal (S-type) origin is flawed, as medium-pressure (1.2-0.7 GPa) fractionation of metaluminous mafic melts may lead to peraluminous derivative liquids (Alonso-Perez et al., 2009). The origin of granitic rocks by partial melting of the continental crust was also addressed in the models of Clemens and Vielzeuf (1987) that showed that the wide range of melting temperatures at which a given melt can be produced is a function of the composition of the rocks that are being melted. For example, the presence of micas and amphiboles is known to partially drive, through their high temperature breakdown, the amount of melt that is segregated at a given temperature (Clemens et al., 1986; Petford et al., 2000; Gao et al., 2016 and the references thereof). This fact explains, in turn, the wide range of temperature needed to produce a

given melt fraction during fluid-absent melting, depending on the rock type (e.g., Petford, 1995; Petford et al., 2000). For further information, the reader is referred to the experimental results on melt origin, composition, and formation conditions of Clemens et al. (1986) on the anorogenic, alkaline A-type Watergums Granite in southeastern Australia, Clemens and Wall (1988) on the influence of pressure-temperature conditions, volatile fugacity and melt chemistry on the mineral assemblages of S-type granites with original melt chemistry, and the reviews of Castro (2014, 2020) on I-type granites from active continental margins (Cordilleran I-type) and intracontinental orogens (Caledonian I-type). Numerically, the computer programs and calculation packages *Perplex* (Connolly, 1990), *Thermocalc* (Andersson et al., 2002), *RCrust* (Mayne et al., 2016), and *Melts* (Ghiorso and Sack, 1995) are also commonly used for calculating granitic melt volumes and/or compositions.

Other relevant studies addressing the origin of granitic magmas are those focused on melt segregation in migmatitic terrains (e.g., Vigneresse et al., 1996; Vigneresse and Burg, 2000; Leitch and Weinberg, 2002; Bons et al., 2004; Cruden and Weinberg, 2018), which are thought to be the main sources of crustal magmas. Segregation is understood as a melt transport mechanism but, unlike magma ascent and emplacement (see below), it operates at the small scale (centimetres to decimetres) and mostly within the magma source region. The structures that indicate melt segregation are diverse in size and morphology, and can be formed by processes operating from  $\sim 10 \ \mu m$  to form initial melt pockets up to several decimetres to develop granitic patches, layers, pods, or veinlets in a metamorphic host (Brown et al., 1995; Clemens and Petford, 1999; Rosenberg and Handy, 2005; Brown, 2013). Numerous experiments have attempted to replicate melt segregation in migmatitic terrains under both hydrostatic conditions (e.g., Laporte, 1994) or during pure shear (Butler, 2010 and references therein) and simple shear (Katz et al., 2006 and references therein) deformation. The experiments that were carried out until the mid-1990s are exceptionally well summarized in the review of Rushmer (1996). From that moment on, however, many other works have provided insights into this issue (e.g., Bagdassarov et al., 1996; Pickering and Johnston, 1998; Rosenberg and Handy, 2005; Druguet and Carreras, 2006). For further reviews on this topic, readers with a strong background in experimental petrology are referred to the works of Brown (2001), Clemens and Stevens (2016), and Collins et al. (2020).

The experimental and numerical modelling approaches of granitic magma ascent and/or emplacement dynamics have been central for perhaps the most controversial debate involving granites during the last decades: dykes vs. diapirs as the main magma ascent mechanism. This debate is summarized in the work of Petford and Clemens (2000): "How are granite magmas transported from their source regions in the deep crust, through 10-40 km of overlying solid rock and emplaced close to the Earth's surface?" Before the 1990s, the widely accepted mechanism for magma ascent from melt extraction and segregation zones was diapirism (see review by Cruden and Weinberg, 2018). This process was believed to be driven by the density difference between a melt volume and its surrounding country rocks and, for decades, represented a feasible explanation for the shape, size, and deformation of both granite batholiths themselves and surrounding rocks. Pioneering investigations on modelling the magma ascent through diapirism were those of Grout (1945). Using a glass tank containing viscous material as analogue for magma, and soft wet clays, corn syrup, or syrup diluted with water as analogue of the surrounding rocks, Grout (1945) showed the diverse shape of the rising masses and deformation structures in the surrounding material were related to the amount, type, and temperature of analogue materials. Subsequently, numerous experiments were undertaken to investigate the diapiric ascent of granitic magmas (Berner et al., 1972; Ramberg, 1972; Cruden, 1988, 1990; Weinberg, 1992; Weinberg and Podladchikov, 1994), establishing diapirism as a plausible mechanism for the ascent and emplacement of granitic magmas (see the reviews by Weinberg, 1996; Cruden and Weinberg, 2018). Noteworthy, this diapirdriven ascent paradigm fostered the publication of several works where granites were drawn in cross-sections as large, downwards-extending or balloon-shaped masses at different depths of the crust. In some cases, granites are nowadays still illustrated like this (Peng et al., 2015; Gamal El Dien et al., 2019; Xie et al., 2019; Salih and Rahman, 2021; de Lira Santos et al., 2022; He et al., 2023). However, several problems related to this type of graphical representation of granites had already been pointed out by Lane (1931) and further debated until recent times (e.g., Clemens and Mawer, 1992; Hutton, 1992; Petford, 1996; Cruden, 1998; Vigneresse and Clemens, 2000). In fact, numerous works have suggested that magmatic ascent by diapirism cannot overcome the high effective viscosity of the crust (which would need to be at least one order of magnitude lower than that predicted; Miller et al., 1988; Cruden and Weinberg, 2018), nor the crustal brittle-ductile transition (~10-20 km depending on the temperature of the crust). These arguments thus restricted diapirism, if present, to the lower crust (Barnichon et al., 1999; Vigneresse, 2004). However, it is worth noting that diapir-driven ascent dynamics are still invoked nowadays for the exhumation of UHP eclogite rocks to the surface of deep-seated subduction zones (e.g., Little et al., 2011: Chatterjee and Jagoutz, 2015: Massonne and Fockenberg, 2022).

Another plausible mechanism of "bulk" ascent of granitic magmas through the crust is the magmatic stoping (Daly, 1912). This mechanism has been suggested to operate at all crustal levels and to be mainly controlled by thermal fracturing, and involves the incorporation of blocks of country rock into the magma and the development of irregular contacts. Some of these blocks may remain within the molten material as stoped blocks and xenoliths, whereas others may be assimilated by the magma, changing its geochemical composition (e.g., Dumond et al., 2005; Žák et al., 2006; Pignotta and Paterson, 2007; Clarke and Erdmann, 2008). However, theoretical work by Marsh (1982, 1984, 2007) has shown that large-scale magma transport cannot be achieved through magmatic stoping because of the rapid cooling and freezing of granite magma, restricting magmatic stoping to a more local heat and mass transfer mechanism (e.g., Cruden and Weinberg, 2018). It has also been argued that the interplay between multiple space-making mechanisms could provide a solution for the room problem, including magmatic stoping, ductile shortening of host rocks, and downward return flow (Paterson and Miller, 1998a, 1998b; Paterson et al., 2008). From the 1990s, experimental and numerical modelling of granitic magma ascent in igneous-volcanic plumbing systems shifted the paradigm to dykedriven, channelled magma ascent dynamics (see the reviews of Petford, 1996; Petford et al., 2000; Petford and Clemens, 2000; Vigneresse and Clemens, 2000; Menand, 2011). Emplacement of granitic magma that ascends through dykes is nowadays attributed to dyke arresting and sill-like horizontal propagation followed by vertical inflation of the pluton (Cruden, 1998; Cruden and Weinberg, 2018; Galland et al., 2018). Under this paradigm, debate has also been raised on the interplay between the steps of initial melt formation, segregation, accumulation, and its final emplacement. The assumption of complex fracture networks where small, tributary fractures feed larger ones, known as the 'rivuletsfeeding-rivers' model, has been perhaps the dominant theory (Brown and Solar, 1998; Weinberg, 1999), although some theoretical arguments and field evidence speak against it (e.g., Bons et al., 2009). For example, the connected melt networks expected for the 'rivulets-feeding-rivers' model (e.g., the classical Port Navalo migmatites, France) have been suggested to more likely represent different generations of melt-filled veins and dykes (Bons et al., 2009). Within this framework, the discontinuous connectivity of melt-filled fractures in a stepwise accumulation model (Maaløe, 1987; Bons and van Milligen, 2001; Bons et al., 2001) has more explanatory power in terms of spatial and temporal scales involving the formation of a granitic pluton. For further reviews that address the origin and propagation of granitic magmas through dykes or dyke-diapir interaction, the reader is referred to the work of Clemens (1998), Bons et al. (2009), Cao et al. (2016), Cruden and Weinberg (2018) and references therein. Moreover, given the bountiful

literature on the topic and the multiple possible explanations for the ascent of granitic magmas through the crust, we think it is appropriate to ask here: has the *room* problem been resolved?

Investigations on the physical and chemical processes acting during the post-magmatic evolution of granites are generally focused on metamorphism, deformation, and hydrothermal and metasomatic alteration, as well as on the interplay and/or overprinting of these processes. These phenomena, although unrelated to granite genesis, are important because they control the final aspect and geochemical composition that geologists find while studying or sampling granites (e. g., Elburg et al., 2001). Among others, the metamorphic, metasomatic, hydrothermal, and tectonic post-magmatic evolution of a given granite at subsolidus conditions are included in this scope. Since granites are an important rock type being considered to host nuclear waste storage facilities (e.g., Mccarthy et al., 1978; Metz et al., 2005) and can act as reservoir storage rocks in geothermal and petroleum systems (Landes et al., 1960; Zheng et al., 2021), a significant number of experiments on fracture initiation and propagation have been conducted in granites (see a review in Zhuang and Zang, 2021). Moreover, other experimental studies on the post-magmatic evolution of granites are those focused on hydrothermal alteration and metasomatic reactions. Hydrothermal alteration in granitic rocks can occur as a result of either the interaction with magmatic fluids exsolved during crystallization processes (e.g., Berni et al., 2020), or with hydrothermal fluids of metamorphic or meteoric origin (e.g., Savage et al., 1987). Significant knowledge of the interplay between ore genesis, element mobility, and hydrothermal fluid flow has been gained from experimental studies on granite-hosted ore deposits such as, among others, tungsten (Wang et al., 2021) and gold (Pokrovski et al., 2014) (see a review in Candela, 1992). Furthermore, of significant interest for potential application in geo-energy are the experiments of Truche et al. (2021), where it is demonstrated that the production of molecular hydrogen (H2) during the alteration of peralkaline granites can be similar to, if not faster than, those formerly suggested during the serpentinization of ultramafic rocks.

Experimental and numerical modelling studies focused on the postmagmatic evolution of granitoid rocks can be roughly classified in two main research subfields that commonly appear intimately related in the literature: structural geology and metamorphic petrology. Within the structural geology subfield, rock mechanics testing of deformational (micro- to macro-) structures are divided according to the temperature and pressure conditions under which they act. Generally, at shallow crustal levels, granites behave as a brittle material where a given stress state can produce fractures, i.e., discrete planar discontinuities along which cohesion and continuity are lost (Lajtai, 1998; Gudmundsson, 2011; Parisio et al., 2019). Testing fracturing processes affecting granitic rocks is generally carried out for geo-energy applications and ore deposits research (Zheng et al., 2021; Chandrasekharam et al., 2022). In this framework, numerical approaches that evaluate the brittle behaviour of granites under different injection flow rates and pressure and temperature ranges in enhanced geothermal systems are common (e.g., Shao et al., 2015; Guo et al., 2018; Cheng et al., 2021). Furthermore, other studies have also attempted to simulate the influence of the orientation, geometry, aperture, and roughness of natural fracture systems on the fluid flow and heat transfer mechanisms in natural granitic geothermal reservoirs (e.g., Liu et al., 2020; Chabani et al., 2021). Within the ore deposits field, artificial (i.e., human-induced) hydrofracturing is often used in mining operations as a rock preconditioning method aimed at the fragmentation of ore bodies (Katsaga et al., 2015; He et al., 2016; Bons et al., 2022), and natural fractures are furthermore key structures because they exert a critical control on hydrothermal fluid flow and the formation of orebodies. Mineral precipitates that fill fractures in granites, forming veins, thus constitute very important ore sources of, among others, key elements for the energy transition such as antimony, tungsten, lithium, or cobalt (Alderton et al., 1980; Li et al., 2018; Chauvet, 2019; Slack et al., 2022). For a comprehensive review of the structural controls on ore deposits related to granitoid rocks from a

numerical modelling approach, the reader is referred to Zhang et al. (2011).

At deep crustal levels, where temperature and lithostatic pressure are higher than those related to the brittle behaviour, ductile flow controls deformation processes affecting most rock types, including granitic rocks (e.g., Gomez-Rivas et al., 2020 and references therein). Numerical modelling of ductile structures has provided meaningful insights to unravel whether granite microstructures reflect igneous histories, postmagmatic deformations, or a combination of both (e.g., Paterson et al., 1989; Holness et al., 2018). Within this framework, numerical modelling platforms such as ELLE (Jessell et al., 2001; Bons et al., 2008) and other computational approaches (e.g., Bäckström et al., 2008; Zhang et al., 2019) have contributed to the understanding of tectonic structures and microstructures commonly found in granites (e.g., Griera et al., 2011; Llorens et al., 2013). Moreover, Finch et al. (2020) simulated the development of C' shear bands, which dip in the direction of shear at an angle of 15–35° to the shear zone boundary, and described how changes in the proportions of the weak phase influence the formation of S-C fabrics, C' and C" shear bands. Finch et al. (2022) also modelled and described the wide range of interference patterns that can be developed when co-planar shearing overprints ductile simple shear in the opposite direction, providing insights into strain accommodation during the formation of these hybrid structures in granitic rocks.

In the metamorphic petrology subfield, numerical modelling of postmagmatic processes affecting granitic rocks is mainly related to the theoretical determinations of pressure-temperature-deformation-time (P-T-D-t) paths. Metamorphism is commonly recorded in granites that, after undergoing this process, become named metagranites or orthogneisses (Bucher and Frey, 2002; Schmid et al., 2004; Fettes and Desmons, 2011) and are often present as gneiss-dome systems or metamorphic core complexes (see reviews in Teyssier and Whitney, 2002; Yin, 2004). P-T-D-t estimations on metagranites allow relating the spatial and temporal evolution of pressure and temperature with a given number of deformation phases. Although the temperature and pressure may be estimated from observational (e.g., mineral assemblages and reaction textures; Eskola, 1920; Okrusch and Frimmel, 2020) and analytical approaches (e.g., geothermometry and geobarometry; Anderson et al., 2008), numerical modelling is required for the calculation of pseudosections and thermal modelling. However, these approaches are outside the scope of the present work since they are not different from the P-T-D-t paths of other rock types. For examples of P-T-D-t determinations using pseudosection and thermal modelling in metagranites, the interested reader is referred to the recent works of Massonne (2015), Jung et al. (2019) and Schorn (2022). Mineral transformations, grain coarsening, and grain recrystallization are also common textural re-equilibration processes that may take place during the cooling or metamorphism of granitic rocks. Although these processes may not alter the geochemical composition of rocks, they cause important textural changes that could be confused with primary magmatic features (see reviews by Higgins, 2011; Holness and Vernon, 2015).

### 6.4. Research on granites during the Digital Age

The aforementioned observational, analytical, experimental, and numerical modelling approaches for the study of granitic rocks have greatly benefited during the past few years from Open Access, FAIR (Findable, Accessible, Interoperable, Reusable) initiatives for data management (Wilkinson et al., 2016; Stall et al., 2019). However, until recently, geochemical data of granitic rocks were typically published with non-standardized approaches and, in some cases, lack of sample metadata. This has significantly reduced the data value and longevity in large areas of Australia, Europe, South America, and easternmost Asia (Figs. 6, 7) (e.g., Chamberlain et al., 2021).

These FAIR geochemical initiatives, which offer to the end user a standardized template for data download, can be roughly divided among

geographically constrained compilations and worldwide-focused databases (Fig. 6; Table 3). Examples of the regional compilations are the North American Volcanic and Intrusive Rock Database (NAVDAT, http ://www.navdat.org/; Walker et al., 2006), the Finnish Lithogeochemical Rock Geochemistry Database (RGDB; Rasilainen et al., 2007), and the Geochemical Database of Japanese Islands for Basement Rocks (DODAI; Haraguchi et al., 2018) (Fig. 7; Table 3). Worldwidefocused geochemical databases offer a user-friendly search engine through which users can filter data by sample location, chemical composition, or tectonic setting. Among others, the Geochemistry of Rocks of the Oceans and Continents database (GEOROC; Georg-August-Universität Göttingen, http://georoc.eu) and the community-driven Petrological Database (PetDB; EarthChem; www.earthchem. org/petdb) are data-rich examples (Figs. 6, 7; Table 3). An excellent compilation of the aforementioned databases was carried out by Gard et al. (2019), which has allowed to graphically represent the worldwide distribution and the geochemical composition of more than 25,000 whole-rock analyses on fresh granitoid samples (Figs. 6, 7; Table 3) in a user-friendly, Geographical Information System (GIS)-based resource; the Global Granite Geochemistry and Outcrop Database (GGGOD; González-Esvertit et al., 2024). This tool allows the visualization, plotting, and spatial analysis of all granitoid compositions available Open Access.

When comparing the GGGOD Database with the global distribution of granitic rocks (modified from the Open Access Global Lithological Map, GLIM, Universität Hamburg; Hartmann and Moosdorf, 2012) (Figs. 6, 7; Table 3), high granite sample concentrations stand out in northern Canada and Alaska (Fig. 7a) and Japan (Fig. 7e). In contrast, Mainland Australia (Fig. 7b), Europe (Fig. 7c), South America (Fig. 7d), and easternmost Asia (Fig. 7e) show limited data. Most (if not all) the granitic rocks cropping out in these areas have, however, already been studied, analyzed, and/or interpreted in the literature. Therefore, the efforts for sample collection and analysis of these granites can be considered as futile in terms of their subsequent reusability in later investigations, such as the present and other potential review works on granitic rocks (e.g., Piwowar et al., 2007).

### 7. Concluding remarks

The origin and significance of granitic rocks have been two of the most controversial and long-lived topics in the history of Earth Sciences. Here we have reviewed the role that granitic rocks have played on our knowledge of the behaviour of the Earth's system for more than two centuries, including the Neptunism, Plutonism, Uniformitarianism, Transformism, and Magmatism paradigms. Consecutive stasis periods of reductionist science, or "puzzle solving" periods, have successively led to the accumulation of inconsistencies, or anomalies, generally related to the field occurrence, mineralogy, geochemistry, petrology, and geophysical imaging of granitic rocks. These anomalies have produced, in turn, successive scientific revolutions towards new paradigms with more explanatory power than the preceding ones. During some periods of the history of Earth Science, two paradigms have coexisted leading to an increase in the degree of scientific discussion and to radicalization of the main arguments. Major innovations, such as the invention of the petrographic microscope, the birth of geochemistry, and the development of experimental petrology, have fostered the accumulation of anomalies, driving research on granitic rocks towards new challenges and settling new uncertainties. These anomalies have been mostly derived from observational approaches, highlighting the importance of field-based geological investigations. During the current paradigm, tens of petrogenetic and non-petrogenetic classification schemes have been proposed and, for the first time, a worldwide-scale approach for the study of granites is being underpinned by the Open Access and FAIR initiatives for sharing spatial and geochemical data. Yet today, there are still several controversies to be resolved, which will require new analytical methods and more complex theoretical frameworks. The role



**Fig. 6.** Global distribution of Open Access, FAIR geochemical data from fresh granitoid samples compared to the global distribution of granitoid outcrops. Type localities and influential works published during the 18th and 19th centuries are indicated. Geochemical data is compiled from FAIR, Open Access initiatives (dataset modified after Gard et al., 2019). Granitic rock outcrops redrawn from the Global Lithological Map (GLIM, Universität Hamburg; Hartmann and Moosdorf, 2012). Basemap: *ESRI® Ocean.* See outcrop photographs of selected outcrops in Fig. 1 and detail maps of selected sectors in Fig. 7.

of the *schools*, or the places where granite scientist have been trained, will be important for the generation of new hypotheses as well as the anomalies that will break them down.

involved in the design, execution, and dissemination of research in a given area, are a common practice to be tackled.

Finally, it is worth noting from the references and quotations included in this review that, as in many other areas of inquiry, granite research has historically been dominated by wealthy, mostly male researchers in high income, geopolitically stable countries from Europe and North America. From the 20th century onwards, a diversification trend can be observed in terms of who is involved in the research on granites and where. However, this diversification has not succeeded in eradicating past and present economical, geopolitical, and gender imbalances. The evolution of the knowledge of granites reflects, in fact, the general evolution of geoscientific knowledge. Yet today, terms as *parachute science* or *helicopter research*, where local communities are not

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Fig. 7. Distribution of Open Access, FAIR geochemical data from fresh granitic rock samples compared to the global distribution of granitoid outcrops in (a) Alaska and northwestern Canada, (b) Mainland Australia, (c) Europe and north Africa, (d) southern South America and northern Antarctic Peninsula, and (e) east Asia (Korea Peninsula, Japan, eastern China, and eastern Russia). See locations in Fig. 6 (note that the Geographic Reference System has been changed into various Projected Coordinate Reference Systems that best represent land surface, indicated in each region). Geochemical data is compiled from the Open Access, FAIR initiatives listed in Table 3 (dataset modified after Gard et al., 2019). Granitoid outcrop distribution has been redrawn from the Global Lithological Map (GLIM, Universität Hamburg; Hartmann and Moosdorf, 2012). Basemap: *ESRI® Ocean*.

#### Table 3

Summary of the FAIR, Open Access databases that host geochemical data from fresh granite samples formerly published in scientific literature (n = 26,459). See Figs. 6 and 7 for the spatial distribution of samples (raw dataset curated after Gard et al., 2019).

Database Acronym	Host institution	Reference	n*				
Geographically constrained databases:							
DODAI [Japan]	Journal of the Geological Society of Japan (Open Access)	Geochemical database of Japanese islands for basement rocks; Haraguchi et al. (2018). doi:https://doi.org/10.5575/geosoc.2018. 0027	1606				
OZCHEM [Australia] USGS [USA] CDoGS [Canada] ABJ [Antarctica] NLGA [Canada and PM] PETLAB	Geoscience Australia, Australian Government	Geoscience Australia's national whole-rock geochemical database. http://pid.geoscience.gov.au/dataset/ga/65464	298				
	United States Geological Survey (USGS)	USGS National Geochemical Database. https://usgs.gov	85				
	Natural Resources Canada, Government of Canada	Canadian Database of Geochemical Surveys. https://geochem.nrc an.gc.ca/	8				
	British Antarctic Survey	A. Burton-Johnson compilation; published in Gard et al. (2019). doi: https://doi.org/10.5281/zenodo.2592822	160				
	Dept. of Energy, Industry and Technology, Government of Newfoundland and Labrador	Newfoundland and Labrador Geoscience Atlas. https://geoatlas.gov. nl.ca/	4219				
	Geological and Nuclear Sciences Limited (GNS), New Zealand	New Zealand's national rock, mineral and geoanalytical database; Strong et al. (2016). https://pet.gns.cri.nz/	567				
Non-geographically constrained databases:							
EarthChem - USGS		EarthChem Portal - USGS National Geochemical Database. htt	10,023				
EarthChem – PetDB and NAVDAT	Lamont-Doherty Earth Observatory of Columbia University, City College of	ps://earthchem.org					
	New York, Kansas University, and South Dakota School of Mines and Technology. Funded by US National Science Foundation.	and North American Volcanic and Intrusive Rock Database (NAVDAT). https://earthchem.org	5006				
EarthChem - others GEOROC		EarthChem Portal. https://earthchem.org	23				
	Göttingen University - Formerly hosted at Max Planck Institute for Chemistry in Mainz	Geochemistry of Rocks of the Oceans and Continents; Lehnert et al. (2000). https://georoc.eu/	4087				
Gard et al. (2019) – new samples	Zenodo Data Repository (Open Access)	Gard, M., Hasterok, D., and Halpin, J. (2019). Global whole-rock geochemical database compilation (Version 1.0.0). doi:https://doi.org/10.5281/zenodo.2592823	377				

See Figs. 6 and 7 for sample location and distribution.

\* n refers to the number of available fresh granitoid samples with associated whole-rock geochemical data.

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### Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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### Data availability

Data is available in the Global Granite Geochemistry and Outcrop Database (GGGOD) - https://doi.org/10.20350/digitalCSIC/16505

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